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A Geological and Physical Assessment Report of the Joan, Bea and part of the Web groups of Mineral Claims Situated on Mt. Copeland North-West of Revelstoke B.C. 51°N, 118°W, S.E.

By George A. Wilson, P.Eng. on behalf of

King Resources, Ltd.

July 17 to August 31, 1965

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INTRODUCTION

In August 1964, Messrs. Ewer and Derickson located 12 mineral claims known as the Joan Group, and 11 mineral claims known as the Bea Group, on the north side of Mt. Copeland 15 miles north-west of Revelstoke, B.C. It was known that molybdenite occurred on the north slope of the mountain at an elevation of 7300', and it was thought that more could be found in the area. In addition to the molybdenite showings, there were fragments of syenite in talus near the top of the mountain with bornite mineralization and extensive malachite and azurite stain. An additional 84 mineral claims were located north, east and south of the original groups of claims in the spring of 1965.

In order to determine the need for, and extent of additional work it was decided to collect assay samples from strata thought to contain molybdenite and to conduct concurrently a geological mapping program of adjacent outcrops.. To these ends, a four man crew consisting of a geologist two prospector-samplers and a prospector-cook were moved to a camp at 7350° on mineral claim Joan 7 early on July 19, 1965. On July 27, the camp was moved down to a bench on Joan 9 at an elevation of 6400°. On July 31, the crew with the exception of the geologist and a helper was moved out. The geologist was moved to Revelstoke late on August 1. All moves were made by helicopter.

For the remainder of August the geologist managed operations from Revelstoke,, while making excursions to the field according to the availability of helicopters.

On August 21 and 22 a diamond drill was moved by helicopter to a location on Joan 7, where it commenced a drilling program.

Assistance from Dr. J.T. Fyles in the form of useful discussions and advice, and loaned air photographs and base maps is gratefully acknowledged.

LOCATION

The Joan, Bea, and Web groups of mineral claims are situated astride the summit of Mt. Copeland near its west end. Mt. Copeland part of the Gold Range of the Monashee Mountains is a ridge with a nearly east-west trend and summit elevations ranging from 7800' to 8400'. It lies between Copeland Creek to the north and Hiren Creek to the south. Both creeks flow east into the Jordan River which flows south-east to the Columbia River which it joins at Revelstoke.

TOPOGRAPHY

Mountain slopes, particularly those facing north, are steep and in places can be climbed only with considerable difficulty. Glaciers occupy most depressions on the north slopes. Snow covers much of the mineralized area on the Joan Group until late July. Late August is probably the best time to examine outcrops on the north slopes. Copeland Creek is at an elevation of 5000' in the vicinity of the claims. In six miles it descends to slightly lesss than 2200'. Hiren Creek to the South of Copeland Mountain has a similar grade.

The lower slopes of the valleys have a dense growth of trees and shrubs. Below 5500', outcrops are rare. Trees of fair size extend to 6500', but above that elevation they are stunted, Tree line is near 7300'

Many of the south slopes are meadows and it is on these slopes that foot traversing is most feasible.

ACCESS.

A logging road extends up the Jordan River to a point not far above the mouth of Hiren Creek. From there a trail is said to lead up Copeland Creek. Freight can be moved 7 or 8 miles up the Jordan River by road, and from there to the property by helicopter. A road could be built up either Copeland Creek to the north side of the property or up Hiren Creek and one of its tributaries to the south side of the property.

METHOD OF WORK

Geological mapping and channel sampling of mineralized zones were done concurrently during the last two weeks of July. Extensive weathering accompanied by deposition of iron oxides on the rocks obscured mineral content in many places. It was possible that surface alteration and leaching might affect the quality of samples collected. Therefore, the crew was equipped with a portable gas drill. The samplers on the crew drilled and blasted cuts across selected zones where alteration was extensive, to permit sampling of fresh rock. Samples 40 to 60 pounds in weight were collected at several places. The assay results are listed in the appendix.

After the departure of thepprospector-samplers, the time was spent in processing data, drawing base maps, making foot traverses over specific areas, and in attempting to secure a diamond drill.

Late on August 23 a diamond drill commenced drilling. Results from this part of the program are not available at the time this report is being written.

GEOLOGY

Stratigraphy

All names applied to rock units in this report are based on field examination with a lOx hand lens and may be regarded with some skepticism. This is particularly true of names such as sillimanite gneiss, and nepheline syenite. The term sillimanite is used for a porphyroblastic grey to colourless silicate common in some brown weathering gneisses and which may be sillimanite. Nepheline syenite is applied to a grey syenite with a mineral resembling nepheline. Wheeler (1) applied the term to a crystalline rock in an area which may coincide with that to which the term is applied in this report.

With the exception of some late dark green basic dykes, all of the rocks in this area are part of the Shuswap metamorphic complex. All, except some quartzites are foliated to some degree and many are quite schistose. In the area of the Joan, Bea, and Web groups, schistosity is parallel with the original stratigraphic units where such units are recognisable.

Tops, and therefore order of superposition could not be established on any of the units examined on the Joan, Bea, and Web groups. Relative positions of the units are shown on the map and stratigraphic sections which accompany this report. Stratigraphic positions of control stations are shown on the stratigraphic sections but the lines of the sections were ommitted from the map in order to avoid cluttering it.

Descriptions of the rock units are furnished on the sections mentioned above, and will not be repeated here, although some general remarks are in order.

The quartzite which occurs on both sides of the axis of a fold below control station GW-37 on the north slope of Mt. Copeland is different from the interbedded quartzite and carbonate which occurs above station GW-37. The proportion of carbonate in the interbedded quartzite and carbonate varies from place to place, partly due to original deposition and partly due to

(1) Wheeler, J.O., 1965: Big Bend map-area, British Columbia; Geol. Surv. Can., Paper 64-32. squeezing out of carbonate during folding.

The proportion of dolomite to limestone varies from place to place. This variation is due to a combination of original deposition and to metamorphic history. The relative importance of the two agencies is difficult to asses at this stage of the investigation.

Another variation in depositional character is apparent. From station GW-43 at least as far west as GW=21, the carbonate and lime silicate band which is the locus of mineralization is separated from the biotite-amphibole schist by a succession of quartzose clastic beds as much as 100' thick. These siliceous clastic rocks are quartzite in places and hard sandstone in others. These sandstones and quartzites include numerous thin lenticular beds of carbonate. Almost everywhere the siliceous clastic rocks are mineralized with pyrite and pyrhotite and in places with magnetite. On fresh surfaces they are grey to yellowish grey. Weathered surfaces are almost universally yellowish brown to a dark red-brown. The unit clearly tapers east to station GW-8 where it is only a few feet thick. East of there it is absent between the biotite-amphibole schist and the carbonate but similar rocks occur on the opposite side of the carbonate unit.

Changes in texture attributable to original deposition occur within the carbonate succession. One such change occurs at GW-20. There, a wedge of finely crystalline dolomite thickens in a westerly direction from O' to nearly 10' in approximately 50'. It occupies the position of mineralized beds to the East but is only sparsely mineralized, the molybdenite occurring above and below it. This could not be fully investigated due to the steepness of the face on which it occurs.

Adjacent to the carbonate beds and on either side of them, are what probably were quartzose limestones or quartzose dolomitic limestones, now converted to a lime-silicate skarn with coarse to very coarse diopside, fine brown garnet, and coarse brown mica with patches of coarse calcite and dolomite. These are the beds of greatest economic interest. Carbonate beds are thin to absent in association with the lime-silicate skarn west of the steeply sloping gully between GW-24 and GW-43. West of the gully,nepheline syenite is in contact with the skarn.

Advanced metamorphism has produced numerous masses of crystalline rock, the composition of which is a function of the composition of the beds from which they were derived. These masses of crystalline rock trend parallel with recognizable stratigraphic units and grade into them along strike. This is particularly noticeable in the biotite-amphibole schist. At GW-5 on the south side of the syncline-like fold, there is a one foot interval of grey banded syenite. Eastward this unit expands so that in the vicinity of GW-10 it occupies some 50' of the section. A similar wedge of a different syenite occurs south of the carbonate in the biotite-amphibole schist. Its width increases from 6", 30' east of GW-5 to 10' some 45' west of GW-5. The south margin of the wedge of syenite conforms fairly well to bedding and schistosity but the south margin crosses both at a small angle. Foliation within the syenite matches foliation in the biotite-amphibole schist.

Mineralization

Pyrite, pyrhotite, and magnetite occur to some degree in all of the strata adjacent to the carbonate beds or the limesilicate skarn. Their presence has already been mentioned with respect to the sandstones and quartzites north of the limesilicate beds. Magnetite is a common constituent of the medium to coarse grained pegmatite stringers and lenticles in the various gneisses. Near the summit of Mt. Copeland at GW-25, magnetite occurs as crystals up to 3/4" across with coarse feldspar. Magnetite also occurs in disseminated form in fair quantity in syenites just south of the crest of Mt. Copeland.

The important mineral, molybdenite, is almost entirely confined to the metamorphosed quartzose carbonate adjacent to the carbonate beds. This unit is now a skarn composed of calcite, dolomite, fine-grained garnet, medium- to coarse-grained mica, and coarse- to very coarse-grained diopside in large lenticular masses and clusters of coarse crystals. The molybdenite seems to be associated with the coarse diopside. It has the form of blebs and masses up to 2" inddiameter, but it also occurs as lenticular masses sub-parallel with bedding. In places, these masses cover fracture planes over areas as much as 2' across with a thickness of 1/8". Exceptionally high grade molybdenite is found in intervals from 6" to 1' thick, and of lower grade over as much as 20' of beds in finely disseminated blebs and specks. Molybdenite forms a rim around diopside crystals in many places. In others, it is scattered at random in the skarn. The high grade skarn occurs fairly consistently a few feet from the syenite from GW-15 west. This may not have much significance as far as the syenite is concerned because equally good, or even better, uccurrances as at GW-5 do not seem to be directly related to crystalline rocks.

Molybdenite was found in the syenite at only one place, that is just west of GW-3. In that locality, blebs and masses of molybdenite are disseminated in a thin tongue of syenite which extends into a biotite-amphibole schist

It is not known what the genesis of the molybdenite is. There is no clear relationship between it and the crystalline rock although the processes which formed the crystalline rock may have changed its distribution. Other possibilities to consider are, that the molybdenite is hydrothermal and that it accumulated at some early stage, or that it is an altered syngenetic deposit. The ability of molybdenite to resist weathering could be cited as support for the last hypothesis.

Considerable further work must be done before one of the

above or some other hypothesis should be selected as the most likely for the accumulation of the molybdenite.

Structure

An interpretation of geologic structure is necessary in the development of an exploration policy to evaluate the mineral deposit. Unfortunately much of the data necessary to understand structure will be obtained by exploration and may be available at a rather late stage.

The strata which include the molybdenite seem to have the form of a syncline whose axial plane dips to the south west as shown on the map accompanying this report. This idea is supported by the configuration of some of the beds and their repetition in reverse order at the east end of the feature. Examination of the north slopes of Mt. Copeland from a helicopter and from the north side of Copeland valley reveals that there are two sets of folds. One set is isoclinal and recumbant with a nearly horizontal axial plane. A fold of this set is well exposed at about 6000'on the west end of the north side of Mt. Copeland. There, it outlined by quartzites. Another member of the same set is visible to the East of the Joan Group at a similar elevation. Unfortunately it was not in a place where the helicopter could land to permit a closer examination. However, it did appear to have a nearly horizontal axial plane.

The other set of folds has a south-east trend and is more open. Dr. J.T. Fyles of theB.C. Dept. of Mines thinks that the isoclinal recumbent folds are older (personal communication).

George .Eng. Expiry Date: August 4, 1966







Sandstone; grey to yellow-grey, weathering yellow-broan to brown, fine grained, grainular to mosaic tabric, beds 1' to 2' thick with 1/10 to 1" laminae; probably had calc stale partings now altered to line silicates and garnet nica marble Small bedding plane Faults. Ty & to throughout. Hosz in disseminated blebs to 1/8" sparse in Upper part as drawn.

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Sundatone; yellow-arey, weathering yellow brown, calcareous, Schistose, intotheds bio a wythikale gaeiss

Sandstone; grey to dark grey, weathers dark brown, hard, quartzose, fine grained, thinky laminated, sparse interbeds biotite-dimptibole gneiss

Biotite-amphibole schist; dark green weathering similar Atundant purphyroblasts sillimanite? Within this unit there is a light grey banded gneiss of approximate syonite composition, which widens trom 6" sume 30' cast of the line to 10' sume 45' wost of line of section. Banding within the tongue of crystalline rock conforms to bedding in surrounding beds. The upper contact as plotted constotus to bedding; the lower and crosses beds at a low angle.

Section measured from GWB south across south limb of syncline and plotted so that south end of section is at bottom and north end is at top.

Note: In second unit from top there is a tongue of grey banded enystalline gness just above molybdenum bads as plutled. This tongue expands to east so that it includes some 50' of beds below white dolomite. It is similar to dionitic type in out crops down slope and to east of 600

Scale 1" to 10' -saw.

SECTION 3









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Appendix 3.

PHONE 469-2391 (24-HOUR SERVICE)

Atlas Testing Laboratories Ltd.

INDUSTRIAL AND RESEARCH CHEMISTS TESTING AND INSPECTION SERVICES

7911 ARGYLL ROAD EDMONTON, ALBERTA

August 2, 1965.

Geological Consultants Ltd., 104 Waterloo Drive, CALGARY, Alberta.

Lab No. 1189

Attention: Mr. George Wilson, P.Eng.,

Dear Sir:

The ore samples received July 26, 1965, assayed as follows:

	Sample No.	Molybdenum	Copper	Lead	<u>Zinc</u>
	E - 1	.09	.02	trace	.65
-	E - 2	.18	.03	trace	.45
	E - 3	•49	.05	trace	.30
	E - 4,	.04	.03	trace	.50
,	E - 5	.11	.02	trace	.50
	E – 6	.02	.02	trace	•45
ł	E - 7	.07	.03	trace	.40
	E - 8	5.03	.03	trace	-75

Yours very truly, ATLAS TESTING LABORATORIES LTD.

J.O. Garner, Head Assayer:

JOG/bb

FORM NO. B			PHONE: 8 4111 CABLE ADDRESS "ELDRICO"
Mr. George A. Wilson,	Clertificate of Assau	Member Canadian Testing Association	FILE NO. A.3-W.1-65 18599
104 Waterleo Drive, Calgary, Alta.	COAST ELDRIDGE		DATE August 9, 1965
- -	125 EAST 4TH AVE. VANCOUVER 10. CANADA		

	G	OLD	SILVER	Lead(Pb)	Zinc(Zn)	Copper(Cu)	Molybdenum		
MARKED	OUNCES PER TON	VALUE PER TON	OUNCES PER TON	PER Cent	PER CENT.	PER () CENT.	105 ₂) per 2 cent.	PER CENT.	PÉF CEN
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76862		-	Trace	0.05	0.12	0.04	0.26		
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76865				Trace	0.07	0.04	0.05		
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Gold calculated at \$_____per ounce

Note. Rejects retained one week. Pulps retained one month. Pulps and rejects may be stored for a maximum of one year by special arrangement.

Unless it is specifically stated otherwise, gold and silver values reported on these sheets have not been adjusted to compensate for losses and gains inherent in th e assay process.

Sharples Provincial Assayer

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FORM NO. B		ACCTLC	PHONE: 6 4111 CABLE ADDRESS "ELDRICO"
Mr. G.A. Wilson		Canadian Teshing Association	FILE NO. A. 3-W.1-65 18875
104 Waterloo Dr.,			DATE August 19, 1965
Calgary, Alta.	ENGINEERS & CHEMISTS LTD.		
	125 EAST 4TH AVE. VANCOUVER 10. CANADA		

Lead (Pb)

PER

CENT.

Trace

Trace

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Zinc (Zn)

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Note. Rejects retained one week.

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Pulps retained one month. Pulps and rejects may be stored for a maximum of one year by special arrangement.

Unless it is specifically stated otherwise, gold and silver values reported on these sheets have not been adjusted to compensate for losses and gains inherent in the reassay process.

H. Sharpen Provincial Assayer

Molybdenum

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white Quartzite weatherin massive, sparse inferbed, of paragmeiss. Paragneiss; grey with brown tint; weathering grey as below interbedded with guartzite du abave. Paragneiss; grey with brown tint, weathering grey, alternate laminatorich in gtz or mica, brown, Quartzite; white to grey, weathering brown-grey, massive. Covered; partly underlain by conterted quartaite Quartaite; white, weathering light grey with interbedded paragness. Paragness. similar to unit 5 1 This unit at base of section is at or very mean unial plane of syncline-like fold. The quartaites are dyplicated down slopio. Units 25 to 39 are probably same beds as unit 20 on apposite sides of a syncline lise told with axial plane near centre of unit 24. P. a. 1.1 Scale 1" to 100' -Ga.W. To accompany geological report by George A. Wilson, P. Eng., on the Jean, Bea and Web groups on Mr. Copeland, Revelstoke, M.O., Dated August 28/65

