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GEOLOGICAL, GEOCHEMICAL AND GEOPHYSICAL REPORT ON THE OVP 1/36 AND MK 1/60 CLAIMS

53°35' 127°20'

by

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for

Aston Resources Limited

TABLE OF CONTENTS

page

page

LIST OF ILLUSTRATIONS

INTRODUCTION

SUMMARY

The property is located on the east flank of the Coast Mountains in north-central British Columbia. Variable amounts of chalcopyrite, pyrite and molybdenite occur in several ages of intrusive rocks and in altered volcanic rocks. A major ore control is a north - northwesterly trending structure that is occupied by a complex swarm of dikes, many of which are porphyritic. A large granodiorite stock is an apparent lithologic ore control, and is a favourable host rock where traversed by the north - northwesterly structural elements. Some of the dikes are mineralized, whereas others are clearly post-mineral in age.

The present investigation was designed to study the geological, geochemical and geophysical interrelationships of the various igneous rock units.

LOCATION

The claim group is near the west end of Troitsa Lake, approximately 90 miles south of Smithers, B.C. Latitude and longitude are 53°35' and 127°20' respectively. An all-weather road from Houston, B.C. to the east end of Tahtsa Lake ends ten air miles north of the claim group. Present access requires the charter of aircraft from Smithers, B.C.

HISTORY

Mineralization was discovered by two independent veteran prospectors (Messers. G.Bleiler and F.Giauque) in the field season of 1966. These men were undoubtedly attracted by a prominent gossan that occurs near the south shore of Troitsa

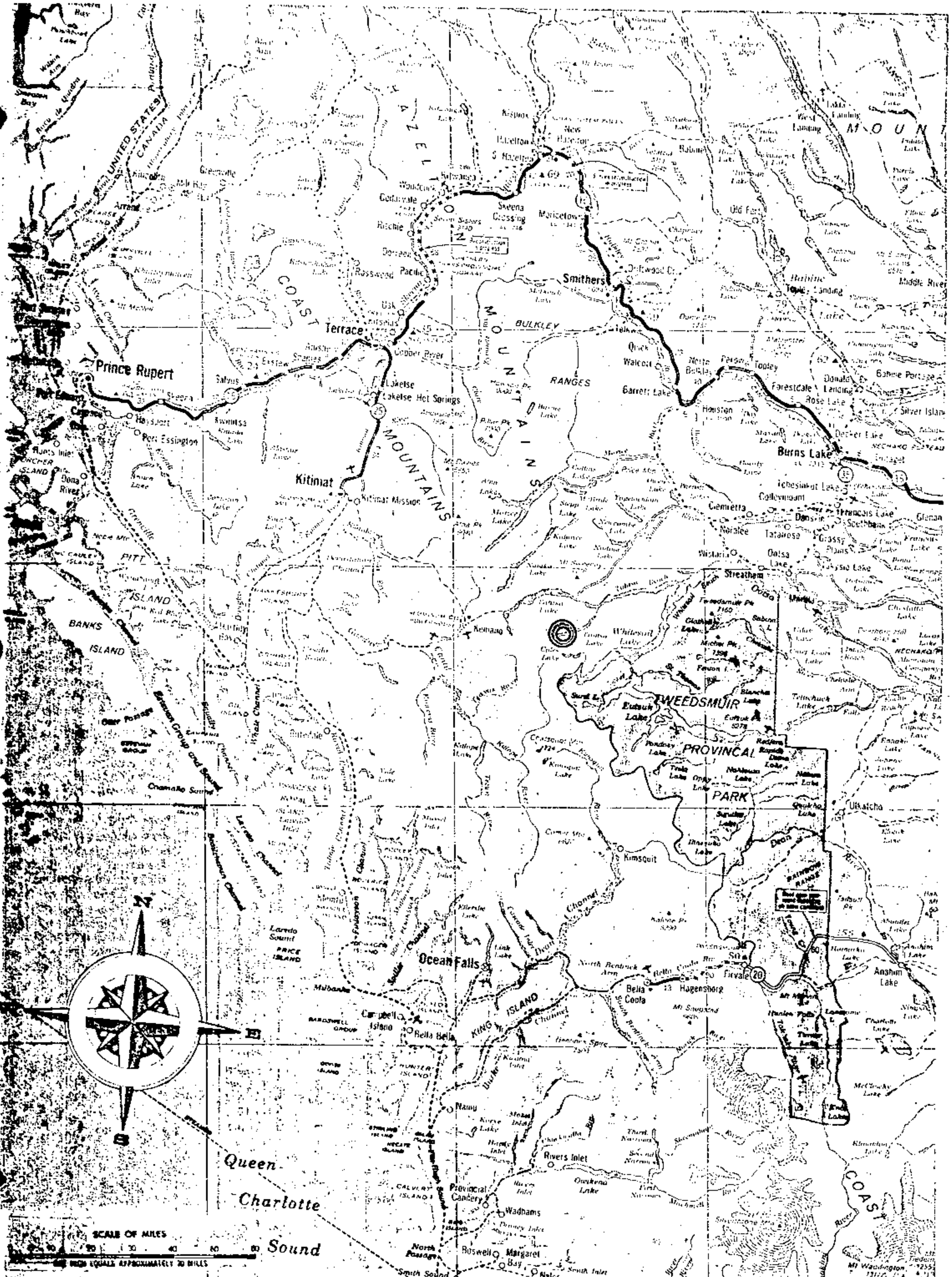


FIGURE I LOCATION MAP

Lake. The prospectors staked thirty six claims (OVP 1/36) in August, and added an additional forty claims (OVP 37/76) in September, 1966. Silver Standard Mines optioned the property in the late summer of 1966, and some mapping and sampling was carried out.

In the winter of 1966-1967 an additional forty claims (FAG 1/40) were staked to cover a gossan south of the present claim group. During the summer of 1967, three diamond drill holes totalling 1215 feet were drilled, as well as some additional geological mapping, trenching, and a geophysical survey on a small part of the property. A group of sixty claims (MK 1/60) was staked to cover newly discovered mineralization in an upland valley south of Troitsa Lake.

In 1968, two holes totalling 1187 feet were drilled. Silver Standard subsequently relinquished their option on the property.

In the summer of 1969 Aston Resources Ltd. and the original prospectors entered into an agreement to further explore the area.

PURPOSE OF THE STUDY

The present program involved a detailed geological, geochemical and geophysical study within an area of possible composite igneous activity near the north end of the porphyry dike swarm. In addition, an area of approximately thirty square miles was tested by airborne geophysics.

Detailed geologic work involved checking and compiling previous work, as well as additional studies on a scale of 1" = 500' in areas where the interrelationship of igneous map units was considered significant. Previous mapping had been done on scales of 1" = 500' and 1" = 200'. This data was compiled and reduced to a common scale (Map 1). In addition some gaps of information existed in areas pertinent to the

present interpretation.

A 200' x 500' grid covering an area 8000' by 3000' was laid out on the north end of the dike swarm (Map 2). This will subsequently be referred to as the "Grid Area". This control was used in the detailed geologic studies, for the collection of material for rock geochemical studies, and for a ground E.M. survey. B.C. Government aerial photographs (1" = 1/2 mile) were used for control in reconnaissance studies along the dike swarm. These photos were also used as control for the airborne magnetometer and E.M. survey. Flight lines 1/4 mile apart covered an area of approximately 30 square miles centered on the dike swarm.

Evergreen Explorations Limited contracted the work involved in setting up camp, and in providing the personnel and all equipment for the geological, geochemical and geophysical field work. A base camp was established in a saddle at an elevation of 4800 feet, some 1 1/2 miles south of the west end of Troitsa Lake. All field work was done under the supervision of R. Woolverton P.Eng. D.A. Davidson P.Eng. interpreted the geological and geochemical work. R. Woolverton P.Eng. compiled and interpreted the geophysical data.

PHYSIOGRAPHY

The area of interest is in the Tahtsa Ranges near the north-western edge of the Interior Plateau and on the eastern flank of the Coast Mountains. Mountains in the Tahtsa Range are mostly in east-west trending ranges, and the major valleys are commonly occupied by lakes.

Rugged serrate peaks and steep ridges rise to more than 7500 feet. Alpine glaciers, often in well developed cirques, are present on most of the higher peaks. Morrainal deposits related to alpine glaciation are evident in many of the valleys.

Relief on the ridges and peaks covered by the present work is marked. The ridge immediately south of Troitsa Lake rises from 2947 feet at the lake to 6500 feet in approximately one mile. A mile wide upland valley at an elevation of 4000 - 4800 feet separates this ridge from a series of steep ridges and peaks to the south that rise to 6700 feet. The valley of Coles Lake lies to the south of these mountains.

Timber-line varies from 4000 to 4500 feet. No data is available on annual precipitation, but the proximity to the Coast Mountains would suggest a moderate to abundant amount of rain and snow. In the claim group, snow remains in the sheltered upland valley into late June.

GEOLOGY

REGIONAL GEOLOGY

The area of interest is underlain mostly by volcanic, pyroclastic and sedimentary rocks of the Hazelton Group approximately three miles east of the eastern contact of the Coast Range intrusive complex. The area is also traversed by the Skeena Arch, a major geological feature that trends northeasterly beneath the Mesozoic volcanic and sedimentary rocks, and includes rocks of the eastern portion of the Coast Crystalline Belt and the western portion of the Cassiar Crystalline Belt. Granitic apophyses of probable late Mesozoic age are common near the eastern contact of the Coast Range intrusive rocks. The arch is further characterized by numerous small stocks and bosses that occupy a well defined zone between the two major "Crystalline Belts". Recent work has shown that many of these small intrusive bodies are Tertiary in age, and that many are closely associated in time and space with mineralization.

LOCAL GEOLOGY

GENERAL STATEMENT

The local geology is generalized in the Whitesail Lake Map-Area (G.S.C. Memoir 299) by S. Duffell.

The oldest rocks in the area are a series of Hazelton volcanic, and pyroclastic rocks of intermediate composition. Several discontinuous, lenticular (?) layers of argillaceous rocks occur within the Hazelton map unit. These rocks trend north-northwesterly and dip at moderate to steep angles to the west. A large "tear-drop" shaped mass of granodiorite of probable late Mesozoic age has intruded the Hazelton rocks and underlies much of the map area. The long axis of this unit trends northerly.

The prominent gossan on the ridge immediately south of Troitsa Lake is underlain mostly by an irregular ovoid complex of pyritiferous rhyolite, porphyritic rhyolite and aplite. These rocks occur in and near a re-entrant on the west side of the granodiorite and appear to be in part discordant to the Hazelton rocks. Some aplitic and porphyritic rhyolite dikes have been observed cutting the granodiorite. This evidence suggests that the aplitic - rhyolitic complex is younger than the granodiorite.

All of the previously mentioned rock units are cut by a north-northwesterly trending dike swarm that has been traced for a strike length exceeding four miles. Recent mapping has shown that the dike swarm becomes more complex and exhibits a crude sub-radial habit near the main rhyolite mass. Previous mapping suggested that the dikes widen and may coalesce immediately south of Troitsa Lake. The writer has interpreted all of these features as representing composite high level phases of a buried intrusive center. The rhyolite - aplite unit is considered to represent an early, chilled phase, whereas the dikes, some of which are mineral related, represent a later pulse of intrusive activity.

Most of the structural features recognized in the mapped area trend north-northwest parallel to the regional trend of the Coast Mountains. These features include primary structures in the Hazelton rocks, secondary foliations, some jointing, weak shearing and faulting. These features are believed to have been formed in the structural environment set up during the emplacement of the Coast Range batholith and the attendant orogeny, and the stresses have persisted through emplacement of the dike swarm.

ROCK TYPES

HAZELTON GROUP

Hazelton group rocks have been essentially undifferentiated in both the previous and recent mapping. In the

area of study, this map unit consists mostly of a series of fine to medium grained volcanic and pyroclastic rocks to intermediate composition. These rocks include grey green to dark green andesite flows, tuffs, lapilli tuffs and tuff breccias, as well as minor amounts of dark reddish brown to dark brown basaltic flows and tuffs. Some of the volcanic rocks are porphyritic.

Several discontinuous and presumably lenticular layers of fine grained dark brown-black to black rock with an incipient concoidal fracture have been mapped as argillite. These rocks may be a slightly hornfelsed, extremely fine grained volcanic or pyroclastic rock.

The Hazelton rocks trend north-northwesterly and dip steeply to the west. These rocks do not appear to have been appreciably deformed by the younger intrusive rocks on the property.

GRANODIORITE

A large "tear-drop" shaped mass of granodiorite underlies much of the map area. This stock is approximately 3 miles long and 1 1/2 miles wide and is elongate in a north south direction. The granodiorite is considered to be an apophysis of the Coast Range batholithic complex that lies two miles to the west.

The granodiorite is a greenish grey to pinkish green grey, medium grained hypidiomorphic - granular rock. Locally the rock is weakly porphyritic (feldspar). A typical specimen is composed of up to 15% quartz, 20% pink potash feldspar, 55% plagioclase feldspar and 10% mafic minerals. The mafic minerals consist mostly of hornblende, biotite, ragged, fine grained chlorite (after hornblende?) and some fresh anhedral to euhedral black biotite of probable secondary origin. With the exception of the alteration of the ferromagnesian minerals, this rock is generally fresh in appearance.

RHYOLITIC - APLITIC COMPLEX

Most of this map unit occurs in an irregular mass

in a re-entrant on the west side of the granodiorite near the south shore of Troitsa Lake. Several distinctive phases of this unit have been recognized, but as yet no attempt has been made to differentiate these in mapping.

Much of the mass consists of light grey, light buff-grey aphanitic rhyolite. In the eastern portion of the mass, porphyritic rhyolite and fine grained aplite become more abundant. The porphyritic phases contain up to 10% subhedral watery quartz and/or white feldspar grains to 3/16 of an inch set in a light grey to light buff-grey aphanitic matrix. The aplitic rocks consist of very fine grained (<0.5mm) equigranular quartz and white feldspar. All phases in this unit may contain variable amounts of fine grained pyrite or its weathered equivalent.

Aplite and porphyritic rhyolite dikes cut the granodiorite. Identical rocks are found within the main mass of this map unit.

Bleaching of the Hazelton rocks and/or contamination of the rhyolitic rocks has resulted in mottlings of greens and browns in aphanitic phases near the western contact.

DIKE SWARM COMPLEX

A complex swarm of dikes of mostly intermediate composition occurs in a north-northwesterly trending zone that has been traced from Coles Lake to Troitsa Lake. Recent mapping has shown that these dikes are much more numerous and widespread than had been indicated by previous work (Map 1,2). In addition, some of the areas mapped as altered feldspar porphyry dikes in granodiorite, are simply well defined zones of sheeted vein and fracture controlled alteration in the granodiorite.

Compositionally, most of the dikes would be classified as monzonitic through dioritic and/or their more silicic equivalents. Some dikes have been mapped as granitic and syenitic. Texturally, some of the older dikes are porphyritic, with subhedral grains of grey feldspar and dark green horn-

blende to 3/8 inch set in an aphanitic to fine grained grey to green matrix. These may be cut by green porphyritic and non-porphyritic andesite dikes. No attempt was made to differentiate the dikes in the present study.

The age relationships of the various members of the dike swarm are complex, however, some of the older porphyritic dikes are mineralized, whereas younger andesitic dikes have been observed to be later than mineralization.

Many of the dikes trend parallel to the predominant north-northwesterly structural grain in the area. A sub-radial trend appears to be well developed in the dikes recently mapped within the rhyolitic - aplitic complex (Map 2). The focus of this trend would be near the "Main Showing". It is not known if the apparent marked increase in the number of dikes within the rhyolitic - aplitic rocks is real. These dikes are difficult to trace in the darker Hazelton rocks and in the granodiorite. If real, the apparent localization of these dikes would have to be explained.

BRECCIA

A breccia that contains some volcanic fragments similar to Hazelton rocks, and some small angular light grey rhyolitic fragments occur in an irregular area within and near the southwest end of the main mass of the rhyolite - aplite complex. The matrix of the breccia is a green grey, somewhat porphyritic (feldspar) andesite and/or fine grained diorite. Some specimens that are reported to represent the matrix of the breccia are similar to some phases of the porphyritic dikes.

BASIC DIKES

Several dark brown to black fine grained to aphanitic dikes occur within the map area. These are believed to be the youngest intrusive rocks on the property.

MINERALIZATION

Chalcopyrite is the chief mineral of economic significance. The principle modes of occurrence are as follows:

1. Chalcopyrite - pyrite in scales and grains on fracture planes and in small quartz veinlets in both the main granodiorite mass, in the rhyolitic - aplitic complex and in the Hazelton rocks. Fracture plane mineralization of this type has been reported in some dikes. Chalcopyrite also occurs as fine grained disseminations in alteration haloes in wall rocks adjacent to some of the mineralized fractures and veins. Many of the more prominent mineralized fractures and veinlets trend north - northwest and dip steeply to the west.
2. Fine grained disseminations of chalcopyrite and pyrite in altered granodiorite. Most of the mineralized altered zones appear to be related to fractures, veins and/or zones of dislocation that commonly trend north-northwesterly and dip steeply west.
3. Silver Standard geologists have reported fine grained to coarse grained disseminated pyrite and chalcopyrite in dike rocks. These sulphides occur as accessories and in places as replacements of mafic phenocrysts.

Minor amounts of bornite have been found associated with chalcopyrite. Some fine grained flakes of molybdenite have been noted both in small veins and on fracture planes. A few fine to coarse grains of molybdenite have been noted as disseminations in the aplitic - rhyolitic rocks.

Some small lead - zinc veins have been reported in the Hazelton rocks east of the body of granodiorite.

Most of the physical work done to date has been localized in three areas along the dike swarm (Map 1).

The original "Lake Showing" is located in a sharp creek valley immediately above the shore line at the southwest end

of Troitsa Lake. Outcrop is restricted to the creek channel in this area.

At least five closely spaced sets of quartz - pyrite - chalcopyrite veins and mineralized fractures occur in granodiorite and altered volcanic rocks over 100 feet. These host rocks and the stockworks mineralization have been cut by at least three ages of andesitic and dioritic dikes. Previous trenching indicated average grades in the order of .60% Cu in the stockworks mineralization, and values of .30% to .40% Cu in one of the late dikes. Minor amounts of molybdenite were noted in some veins and fractures.

Approximately 1 3/4 miles south of the lake shore drilling and trenching has tested a zone of strongly altered granodiorite that averages 1.49% Cu across a true width of 40 feet ("Main Showing") in a creek canyon. An additional 30 feet of altered granodiorite on the east side of the altered zone averages .40% Cu. Lesser amounts of mineralization and alteration were noted for some distance along the creek both east and west of the "Main Showing". Most of this mineralization and the related alteration is associated with a sheeting of veins and mineralized fractures that trend 150° and dip steeply to the west. The veins vary in width up to 5 inches, and consist mostly of medium to coarse grained quartz, chalcopyrite, pyrite and minor amounts of molybdenite. Scales and plates of pyrite and chalcopyrite are present in variable amounts on fracture planes. In the highly altered granodiorite, pyrite and chalcopyrite occur in replacement aggregates scattered throughout the rock. The altered granodiorite is cut by an unaltered two foot porphyritic andesite dike that trends north-northwest and dips steeply west.

At an elevation of 5300 feet, approximately 2 1/2 miles south of the lake shore, altered mineralized zones in the granodiorite are exposed near the toe of a glacier in a small cirque ("Cirque Showing"). Mineralization and alteration are

similar in character but weaker in intensity than that in the "Main Showing". Silver Standard took a total of 44 ten foot chip samples over a zone 140 feet wide. Of this width, altered granodiorite formed a total of 55 feet and averaged .33% Cu. The unaltered granodiorite averaged .13% Cu.

ALTERATION

Minor hornfelsing and possible bleaching of the volcanic and pyroclastic occur adjacent to the rhyolitic - aplitic complex and/or the granodiorite stock. Epidote and pyrite are locally abundant in Hazelton rocks, but it is not known if these are related to a hydrothermal event.

Within the granodiorite stock, some of the primary ferromagnesian minerals have been altered to ragged, fine grained aggregates of chlorite. Some fresh subhedral to euhedral grains of black biotite appear to be secondary in nature. Locally, feldspar grains have been sericitized or kaolinized. Some planar zones of moderate to strong alteration occur within the granodiorite. These appear to be controlled by zones of dislocation that trend north-northwesterly. Much of this alteration is related to mineralization. The strongest alteration of this type is found in the zone that contains the "Main Showing". Here, most of the ferromagnesian minerals have been destroyed, and the feldspars are strongly kaolinized.

Much of the alteration in the dike rocks and in the adjacent host rocks is considered to be of a deuteric nature relative to the emplacement of the dikes. Some of the andesitic dikes display weak to moderate propylitic alteration. In the porphyritic dikes, feldspar phenocrysts may be argillized or sericitized, and primary ferromagnesian minerals may progressively alter to rusty, ragged aggregates of fine grained chlorite.

ORE CONTROLS

The major ore controls of the known mineralization are structural and lithologic. Much of the veining and fracture

plane mineralization is controlled by the north-northwesterly structural elements that also controlled emplacement of dikes in the dike swarm. Exposure of the true stockworks veining at the "Lake Showing" do not fit this previous trend. In the very limited exposure available, the most prominent vein sets trend 235 and dip steeply north. The stockwork may be related to structural elements involved in the emplacement of the rhyolitic - aplitic intrusive complex. Note the scattered mineral occurrences along the north contact of this mass on Map 1.

The granodiorite, and to a lesser extent the Hazelton volcanic rocks, appear to be a lithologic control. However, relative timing of mineralization with respect to the younger intrusive rock units has yet to be established. The rhyolite and at least some of the dikes in the dike swarm are locally weakly mineralized.

GEOCHEMISTRY

GENERAL STATEMENT

A program of detailed rock geochemistry was carried out on a 500 x 200 foot grid near the north end of the dike swarm (Maps 3,4). The purpose of this work was to study the geochemical interrelationships of several ages of intrusive rocks, at least one of which is known to be closely associated in time and space with mineralization. Most of the study area was underlain by patches of outcrop separated by fine talus. "Logging" of rock chips in all individual samples with subsequent comparison to geological mapping shows excellent correlation. Therefore, mechanical dispersion of the talus is considered to be slight.

The extension of the sample grid to the east on the steep north slope above Troitsa Lake was hampered by sharp snow filled canyons and precipitous cliffs.

The 500 x 200 foot sample grid was laid out by tape and compass. Sample sites were marked with a two foot piece of 1 1/2 inch cedar lath on which details of the location were recorded. A total of 49,500 feet of line was sampled.

SAMPLE COLLECTION AND PREPARATION

Approximately one half pound of rock sample chips were taken at each sample site. This material consisted of either fine talus material (generally minus 1.5 inches), or of a crude "chip" sample taken from outcrop. Individual samples were labeled and bagged in 6 mil 8" x 11" plastic bags.

The samples were crushed and pulverized in Smithers, B.C. Approximately 7 oz. of pulp was forwarded to Crest Laboratories, 1068 Homer Street, Vancouver 3, B.C.

LABORATORY ANALYSIS

All samples were analysed for Cu and Mo. A pH determination was made on every fifth sample. Composite samples will be analysed for WO_3 and precious metals at a later date.

INTERPRETATION

STATISTICAL CONSIDERATIONS

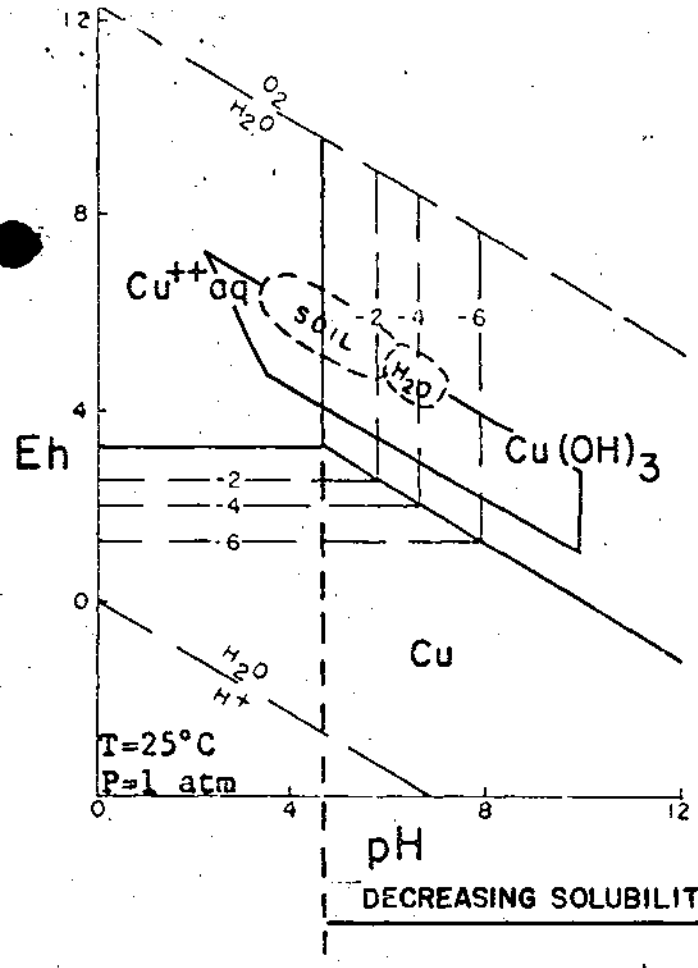
Geochemical results have been statistically studied in order to define background and anomolous values. Frequency distribution diagrams have been plotted for each element determined in the laboratory. The mode, or curve peak is arbitrarily taken as the background value. Using the law of log normal distribution, the threshold or minimum concentration level of significance has been arbitrarily defined as 2 times background. The range of weak, moderate, and strongly anomolous values are taken as 2 to 4, 4 to 8, and 8 to 16 times background respectively. Extremely anomolous values exceed 16 times background. The following tabulation illustrates the significant values (in ppm) determined for each element in the study:

Element	Mode or Background	Threshold 2xB.G.	Weak 2-4xB.G.	Moderate 4-8xB.G.	Strong 8-16xB.G.	Extreme 16xBG
Cu	10	20	20-39	40-79	80-160	>160
Mo	2	4	4-7	8-15	16-32	> 32

EFFECT OF pH ON DISPERSION

Experimental studies have shown that pH may have an important effect upon the mobility of metal ions when minerals oxidize.

Figure 2 shows the stability relationship for copper is dependant on both Eh and pH. Emperical data on weathering environments suggests that for our conditions the solubility of Cu ions is essentially pH controlled. The fields for the soluble cupric ion (Cu^{++}) and for copper hydroxide ($Cu(OH)_3$)



$Cu-H_2O$ SYSTEM

Total Copper 1 molal.

IONS

Cu^{++} - Cupric ion

OXIDES

$Cu(OH)_3$ - Copper hydroxide

ELEMENT

Cu - Native copper

DECREASING SOLUBILITY

$Mo-H_2O$ SYSTEM

Total Molybdenum 10^{-6} m.

IONS

$MoO_4^=$ - Molybdate anion

$HMoO_4^-$ - Acid molybdate anion

MoO_2^{++} - Molybdene cation

OXIDES

Mo_3O_8 - Ilsemannite

MoO_2 - Molybdenum oxide

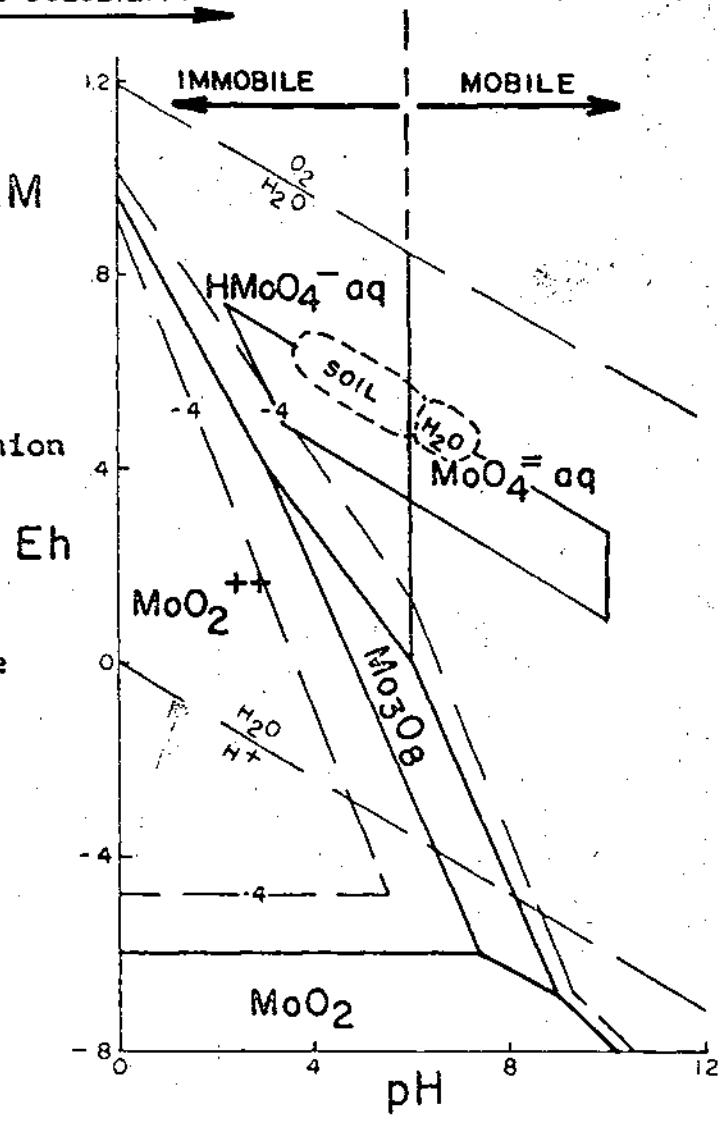


Figure 2 - STABILITY RELATIONS IN SOME SIMPLE AQUEOUS SYSTEMS, 25°C & 1 atm.

meet near a pH of 4.7 for most weathering environments. The solubility of copper hydroxide decreases with increasing pH.

Fifty-one pH determinations have been made in the present study (each fifth sample). The average pH by rock type is as follows: Hazelton rocks 6.1, rhyolitic - aplitic rocks 5.9, feldspar porphyry 6.6 and the granodiorite 6.8. In environments with pH values of this magnitude, the chemical dispersion of Cu ions would be inhibited.

The stability relationships of Mo are also shown in Figure 2. Note the inverse behaviors of Mo to Cu. For Mo, the critical pH is approximately 6.0. With pH less than 6.0, the insoluble acid molybdate anion ($HMoO_4^-$) is the stable form in most environments. With pH greater than 6.0, the soluble molybdate anion (MoO_4^-) is the stable form. In the pH conditions of the study area, chemical dispersion of Mo ions should be favoured.

These conclusions would be dependent upon a study of the pH of surficial drainage. If this is also alkaline or near neutral, then leaching of Cu from surface exposures would not be considered significant. This is borne out by field evidence as fresh chalcopyrite is found at the surface in most of the mineralized zones. Mo on the other hand would be readily removed in such an environment.

DISTRIBUTION AND INTERRELATIONSHIP OF ANOMOLIES

Map 3 and Map 4 show the distribution of Cu and Mo values respectively.

The granodiorite is decidedly anomolous in Cu with respect to the other geological map units in the Grid Area. The edge of the anomoly coincides closely with the geologic boundary of the granodiorite and little chemical or physical dispersion is indicated. The relatively high values within the granodiorite are considered to be related to more strongly mineralized zones of alteration and/or dislocation that are known to occur in these areas. Vague north-northwesterly trending "lows" that cross the rhyolitic - aplitic complex are interpreted as

being a reflection of dike related structures.

No well defined Mo anomolies exist within the Grid Area. Some weakly anomolous areas roughly coincide with areas of stronger Cu geochemical values along some dikes and near the western edge of the granodiorite stock. As discussed previously, the apparent absence of a well defined Mo anomoly could be due to unfavourable pH conditions in the weathering environment. The isolated extreme of 35ppm in the northern part of the Grid Area was obtained from a sample of dark green to black basaltic material, and is considered to be an "erratic high".

GEOPHYSICS

In 1967, Silver Standard Mines Ltd., conducted a limited Induced Polarization survey around the main showings near the camp. The survey was sufficiently encouraging that several holes were drilled and the anomalies tested were found to be caused by sulfides in a "porphyry copper environment."

The 1969 geophysical program by Aston Resources was planned to:

1) evaluate the rhyolite mass between Troitsa Lake and the camp and
2) outline other porphyry environments within or adjacent to the present claim group.

The rhyolite body was investigated by a Radem survey. One hundred and thirteen line miles of airborne magnetic and electromagnetic surveying by Lockwood Survey Corporation was completed over about 30 square miles centred around the existing claim group.

RADEM SURVEY

EQUIPMENT

The Radem unit used in the ground survey is a one man EM radio receiver utilizing the 12 to 24 kilocycle United States Naval Communications Broadcast Stations. It was

built by Crone Geophysics Limited, 979 Lakeshore Road East, Port Credit, Ontario. The instrument utilizes higher than normal EM frequencies and is capable of detecting disseminated sulfides. However, due to the high frequency, it is affected by clay and conductive overburden.

SURVEY AND RESULTS

Readings were taken facing Northwest and using the Seattle Washington station (17.8 kc) at stations along the grid lines as shown on map 5 in the pocket of this report. No strong conductors were outlined by the survey. However, the zero and 500 lines indicate a few weak conductors. These areas of weak conductivity correspond with the areas of higher rock geochem at the rhyolite-granodiorite contact. The rhyolite, however, is almost entirely featureless electricly.

AIRBORNE MAG - EM SURVEY

On July 22, 23, and 24th a 113 line mile airborne mag EM survey of about 30 square miles was completed. The survey area included the O.V.P. and MK mineral claims as shown on map 6 in the pocket. The equipment and crew were supplied by Lockwood Survey Corporation Limited of Toronto.

Readings of the in and out of phase components of the resultant electromagnetic field plus the vertical

magnetic intensity were taken from an FH-1100 helicopter platform. The equipment included a 4300 c.p.s. electromagnetic system and a Gulf Mark III magnetometer. The receiver coils and magnetometer head were carried in a 30 foot bird suspended beneath the helicopter on a 100 foot cable. Flying height was maintained at about 220 feet above the ground and the bird travelled at about 100 feet above the ground. Terrain clearance was measured by a Bonzar radio altimeter and recorded on the E.M. tapes. A Gulf recorder was used for the magnetometer and the EM system was linked to a Taylor recorder in which conductors are recorded in positive values.

Flight lines were oriented northeast-southwest approximately at right angles to the known mineralized structures on the MK claims. The lines were approximately one quarter of a mile apart. However, the steep terrain made both line direction and terrain clearance difficult to maintain. Because the survey area is largely mountainous, the flight lines were "broken off" when the helicopter was unable to continue climbing up the mountain side. Often the new line, started after the aircraft had regained a safe flying speed, was not coincident with the old one at the "break off" point. However, a continuous film record was kept of the flight lines so that they were easily plotted.

The picking of the 35mm film for location points and the data reduction was done by Versatile Drafting Ltd., 448 Seymour Street, Vancouver, B. C. The data was drafted at 1" = ¼ mile and the map is included in the pocket (see map 6) of this report.

RESULTS

The E.M. background or noise is absent because of the lack of conductive overburden in the mountainous terrain. All the conductors are therefore exceptionally distinct on the tapes.

Eight conductors were outlined by the survey. Six of these, numbers 1 to 3 and 6 to 8, are of possible economic significance.

Anomaly number one is a 40 ppm in phase and 25 ppm negative out of phase on the flank of a 140 gamma moderately sharp magnetic high. This anomaly is probably caused by a conductor which is parallel to the flight line in a possible porphyry environment.

Anomaly number two is a 70 ppm negative in phase and a very poor out of phase with a coincident 1000 gamma mag low. It is probably caused by conductive magnetite.

Anomaly number three is a 45 ppm negative in phase, poor to zero out of phase, with a coincident 600 gamma mag

high. This anomaly also probably reflects the presence of conductive magnetite.

Anomalies four and five are broad conductors with poor in phase to out of phase ratios. Number four is in a turn. Neither anomaly has associated magnetic features. They are probably spurious anomalies or ones caused by conductive overburden.

Anomaly six is also probably caused by conductive magnetite. The strong negative in phase of about 200 ppm and equally strong positive out of phase have a coincident 100 gamma mag high.

Anomalies seven and eight are very strong positive conductors and may be caused by massive sulfides. Anomaly eight went off scale on the chart. Seven has a coincident 500 gamma low whereas eight is on the flank of a 1300 gamma high. Anomalies six, seven, and eight are near the north boundary of Silver Standards' I.P. survey area. They form an EM complex on the airborne tapes from which only the more obvious conductors have been plotted.

A northeast trending shear structure is suggested by a series of magnetic lows between Seel and Troitsa Lakes. The strike of the zone parallels the long axis of Seel Lake which may be a topographic expression of the

structure. Apparently some magnetite may be associated with this structure as indicated by anomaly number two. EM anomaly two is within this structure and its electrical characteristics suggest the presence of magnetite so that the alteration of this structure apparently predates the injection of magnetite. There is also a series of northeast trending lows through the camp although they are probably dipoles adjacent to the major magnetic feature north of camp.

APPENDICES

I LISTING OF CLAIM DATA

II GEOCHEMICAL ANALYTICAL PROCEDURES

III LISTING OF GEOCHEMICAL RESULTS

IV STATEMENT OF PROJECT COSTS

APPENDIX 1

CLAIM DATA

CLAIM NAME	RECORD NUMBERS
OVP 1/36	44750/44785
MK 1/60	53881/53940

See Map 6 in pocket for claim configuration

APPENDIX II

GEOCHEMICAL ANALYTICAL PROCEDURES

APPENDIX III

LISTING OF GEOCHEMICAL RESULTS

PHONE 688-8586

CREST LABORATORIES (B.C.) LTD.

B.C. REGISTERED ASSAYERS
INDUSTRIAL and RESEARCH CHEMISTS

1068 HOMER STREET
VANCOUVER 3, B.C.

August 12, 1969

Versatile Drafting Ltd.
448 - Seymour Street
VANCOUVER, B.C.

Dear Sirs:

Re: Analytical Treatment of Rock Samples:

The samples when recieved were pulverized to approximately -100 mesh. These samples were dried at a temperature of 150° F. to remove any water condensation. One gram portions of the -100 mesh rock were placed in a 25 x 200 mm culture tubes and then digested in a mixture of perchloric and nitric acids at 425° F. for a period of three hours. The resulting digested residues were then made up to 50 milliliters volume in 10% perchloric acid.

Copper Analysis:

The respective sample solutions were then asperated into a Techtron Atomic Absorption Spectrophotometer Model 5 and the copper absorption readings were recorded. The absorption readings were then transformed into parts per million copper.

Molybdenum Analysis:

Molybdenum was determined by the colorimetric thiocyanate methods. Briefly, an aliquot (10 ml) of the sample solution is placed into a separatory funnel, and stannous chloride is added to reduce iron; ammonium thiocyanate is added to complex molybdenum, and isoamyl alcohol is added to extract the molybdenum-thiocyanate complex. The intensity of the colour of the molybdenum-thiocyanate complex in the isoamyl alchhol was measured with a Spectronic 20 colorimeter. The color of this above complex is a function of the molybdenum content.

pH Analysis:

Five grams of the -100 mesh rock was placed in a 250 ml beaker and 10 milliliters of neutral distilled water is added and mixed with a glass stirring rod until homogeneous slurry is obtained. The pH of the slurry is measured with a Radiometer Copenhagen Model 28 pH meter. A measurement is recorded after approximately two minutes or when the pH value of the meter reaches equilibrium.

Yours truly,

CREST LABORATORIES (B.C.) LTD.

Alfred A. Burgoyne

Alfred A. Burgoyne
Geologist - Geochemist

AAB/seb

PHONE 688-8586

CREST LABORATORIES (B.C.) LTD.

B.C. REGISTERED ASSAYERS
INDUSTRIAL and RESEARCH CHEMISTS

1068 HOMER STREET
VANCOUVER 3, B.C.

August 12, 1969

Versatile Drafting Ltd.
448 - Seymour Street
VANCOUVER, B.C.

Dear Sirs:

Lot 21 G: Geochemical Analyses for Molybdenum and Copper:

Mesh Size: As Received.
Digestion Method: $HClO_4$ - HNO_3
Analytical Method: Atomic Absorption & Colorimetric

Date Received: August 4, 1969
Date Analysed: August 5, 1969

Sample No.	Moly. ppm.	Copper ppm.	pH ppm.
001	< 1	4	5.5
002	< 1	36	
003	1	10	
004	< 1	6	
005	5	36	5.5
006	2	8	
007	2	36	
008	5	50	
009	2	40	
010	2	6	7.8
011	< 1	20	
012	1	8	
013	< 1	4	
014	< 1	18	
015	1	16	7.2
016	< 1	6	
017	< 1	66	
018	2	38	
019	5	16	
020	2	8	7.1
021	2	24	
022	< 1	6	
023	2	20	
024	2	12	
025	7	76	4.6
026	< 1	8	
027	2	38	
028	5	8	
029	< 1	12	
030	< 1	226	5.7
031	< 1	16	
032	2	86	
033	2	66	
034	2	210	
035	2	70	7.4
036	< 1	8	
037	2	82	
038	5	36	
039	7	52	
040	2	24	7.0
041	5	20	
042	5	76	

Continued on page 2...

Sample No.	Moly. ppm.	Copper ppm.	pH ppm.
043	7	52	
044	7	204	
045	15	220	7.1
046	5	36	
047	7	80	
048	5	24	
049	2	42	
050	5	46	6.9
051	2	88	
052	2	12	
053	5	8	
054	< 1	8	
055	2	52	5.2
056	2	6	
057	< 1	8	
058	2	20	
059	2	9	
060	2	16	5.3
061	< 1	214	
062	< 1	30	
063	< 1	44	
064	2	14	
065	30	82	7.2
066	2	10	
067	< 1	8	
068	2	26	
069	< 1	4	
070	5	24	4.6
071	2	12	
072	10	12	
073	2	36	
074	5	12	
075	< 1	36	7.1
076	< 1	30	
077	2	44	
078	< 1	68	
079	2	60	
080	2	48	5.5
081	7	56	
082	2	20	
083	< 1	16	
084	< 1	4	
085	2	6	7.0
086	< 1	10	
087	2	162	
088	2	10	
089	5	16	
090	2	48	6.4
091	< 1	13	
092	< 1	28	
093	5	60	
094	10	4	
095	< 1	7	7.2
096	< 1	9	
097	1	26	
098	1	20	
099	2	28	
100	1	26	7.1
101	1	20	
102	8	76	
103	7	78	
104	5	76	

Continued on page 3...

Sample No.	Moly ppm.	Copper ppm.	pH ppm.
105	1	28	7.4
106	1	26	
107	1	40	
108	< 1	28	
109	7	122	
110	1	108	7.0
111	2	19	
112	3	20	
113	1	8	
114	2	6	
115	2	7	6.9
116	1	6	
117	2	24	
118	3	14	
119	3	28	
120	< 1	4	7.1
121	< 1	11	
122	< 1	4	
123	2	10	
124	< 1	9	
125	1	5	7.8
126	3	13	
127	2	31	
128	3	90	
129	1	25	
130	5	32	7.1
131	1	92	
132	2	500	
133	5	96	
134	2	242	
135	3	280	7.0
136	1	550	
137	2	17	

Sample No.	Moly ppm.	Copper ppm.	pH ppm.
138	< 1	16	
139	3	7	
140	3	14	6.4
141	1	9	
142	< 1	2	
143	2	7	
144	4	8	
145	4	12	7.2
146	3	12	
147	4	7	
148	15	25	
149	2	15	
150	4	80	6.8
151	1	10	
152	3	52	
153	1	10	
154	3	8	
155	2	5	7.1
156	2	18	
157	15	16	
158	3	16	
159	6	15	
160	8	10	5.5
161	4	14	
162	2	12	
163	3	62	
164	4	46	
165	3	18	6.4
166	3	8	
167	2	22	
168	1	10	

Continued on page 4...

Sample No.	Moly. ppm.	Copper ppm.	pH ppm.
169	2	30	
170	3	9	6.8
171	1	11	
172	1	6	
173	3	12	
174	4	21	7.0
175		NO SAMPLE	
176	3	19	
177		NO SAMPLE	
178	2	12	
179	4	12	
180	2	16	6.2
181	3	20	
182	2	25	
183	7	24	
184	35	14	
185	2	10	7.2
186	1	38	
187	3	12	
188	1	22	
189	3	32	
190	2	13	6.4
191	2	15	
192	2	20	
193	1	8	
194	2	22	
195	4	84	6.4
196	2	16	
197	3	34	
198	2	22	
199	4	38	
200	7	48	5.2

Sample No.	Moly. ppm.	Copper ppm.	pH
201	3	15	
202	4	5	
203	2	7	
204	3	5	
205	5	19	7.0
206	2	10	
207	1	7	
208	1	18	
209	2	20	
210	1	20	5.5
211	1	34	
212	4	7	
213	1	13	
214	7	7	
215	6	4	5.0
216	2	12	
217	8	5	
218	3	8	
219	2	20	
220	3	4	5.5
221	2	25	
222	1	4	
223	5	7	
224	< 1	12	
225	1	16	5.0
226	< 1	7	
227	2	3	
228	1	16	
229	11	12	
230	2	21	7.0
231	1	40	
232	2	5	

Continued on page 5...

Sample No.	Moly. ppm.	Copper ppm.	pH
233	1	13	
234	1	4	
235	1	2	5.2
236	4	5	
237	4	3	
238	1	8	
239	1	4	
240	1	4	6.8
241	1	24	
242	2	20	
243	< 1	8	
244	1	5	
245	1	6	7.1
246	2	7	
247	3	14	
248	2	156	
249	1	9	
250	1	14	5.5
251	2	12	

Yours truly,

CREST LABORATORIES (B.C.) LTD.

Alfred A. Burgoyne

Alfred A. Burgoyne
Geologist-Geochemist

AAB/scb

cc: Mr. D. Davidson
Box 696
Smithers, B.C.

APPENDIX IV

STATEMENT OF PROJECT COSTS

Evergreen Explorations Ltd.

• R. WOOLVERTON
GEOLOGIST

• R. C. O'BRIEN
FIELD SUPERVISOR

• JOHN C. OSWALD and COMPANY
ACCOUNTANT
635 - 789 W. PENDER ST.
VANCOUVER 1, B.C., CANADA

CONTRACT EXPLORATION

• 5424 HALIFAX ST., BURNABY 2, B.C., CANADA, PHONE - 299-6998
• P.O. BOX 604, SMITHERS, B.C., CANADA PHONE - 847-3523

CHARGES FOR ASTON RESOURCES JULY 1969

PERSONNEL		
Operators	39 days @ $40.00	$ 1560.00
Assistants and expediting	22 days @ $25.00	550.00
Supervision and Geology	11 days @ $100.00	1100.00
EQUIPMENT		
Truck	4 days @ $20.00	80.00
Radem Rental		150.00
Field Equipment Rental		210.00
Radio and License		115.00
SUPPLIES		
Groceries, expendables, hardware and board		650.64
SERVICES PURCHASED		
Telephone		35.00
Freight		33.82
Assaying		438.00
Drafting		300.00
Helicopter Charter		1544.00
Fixed Wing Aircraft Charter		200.00
AIRBORNE GEOPHYSICS		
Includes all flying, data redution, drafting, etc.	113 line miles @ $55/mile	6215.00
TOTAL		$13,181.46

STATEMENT OF PROJECT CHARGES

Charges for Geological, Geochemical, and Geophysical Field Work per Evergreen Explorations Ltd. Billing	$ 13,181.46
Sample Preparation in Smithers, B.C.	129.00
Charges by D.A.Davidson for Sample Study and Report Compilation Ten Days @ $100/Day	1,000.00
Total	$ 14,310.46

The undersigned consider the above charges applicable as assessment work.

PROFESSIONAL ENGINEER
PROVINCE OF BRITISH COLUMBIA
R. W. WOOLVERTON

R. Woolverton P.Eng.

D.A. Davidson P.Eng.

TROITSA LAKE

W/L 2,947'

BLANKET LAKES

W/L 3000'

SNOW

CAMP

INITIAL POST MK 33 & 34
FINAL POST MK 31 & 32.

LEGEND

8		BASALT DYKE
7		QUARTZ – SERICITE ALTERATION
6		PORPHRY DYKES – ANDESITE, DIORITE, FELDSPAR PORPHYRY
5		RHYOLITE AND RHYOLITIC TUFFS
4		GRANODIORITE
3		VOLCANIC (?) BRECCIA – Relationship to Hazelton Group uncertain.
2		ARGILLITE (?)
1		HAZELTON GROUP – BASALT, ANDESITE, DACITE; BRECCIA and SKARN (?)

SYMBOLS

- Cu — Copper mineralization: Malachite, Chalcopyrite and Bornite.
- Geological contact.
- Geological contact, assumed.
- 80 — Strike and dip.
- Topographic contour at 50 foot intervals.
- Outcrop

TO ACCOMPANY A GEOLOGICAL, GEOPHYSICAL, AND GEOCHEMICAL REPORT BY ROWLOVERTON ON THE O.V.P. AND M.K. GROUPS, TROITSA LAKE, B.C., OMINECA M.D., DATED

PROFESSIONAL ENGINEER
PROVINCE OF BRITISH COLUMBIA
D. A. DAVIDSON

2026

ASTON RESOURCES PROPERTY

TROITSA LAKE PROJECT
WHITESAIL LAKE AREA

GEOLOGY MAP

SCALE: 1" = 500'

SURVEY BY: EVERGREEN EXPLORATIONS LTD., SMITHERS, B.C.
JULY, 1969

TROITSA LAKE

BLANKET LAKES

LAKE

LAKE

Camp

SYMBOLS

- OUTCROP
- GEOLOGICAL CONTACT
- GEOLOGICAL CONTACT, ASSUMED
- FAULT
- BEDDING, STRIKE AND DIP
- JOINT
- QUARTZ VEIN
- TOPOGRAPHIC CONTOURS AT 100' INTERVALS
- MINERAL OCCURRENCE

LEGEND

11 LAMPROPHYRE

10 DIORITE, SYENITE, MONZONITE, - MEDIUM GRAINED, UNALTERED, LOCALLY MINERALIZED WITH PYRITE, CALCOPYRITE.

9 GRANITE PORPHRY (MAFIC BIOTITE)

8 FELDSPAR PORPHRY

7 FELDSPAR PORPHRY (ALTERED)

6 ANDESITE PORPHRY

5 RHYOLITE WITH DISSEMINATED PYRITE

4 RHYOLITE TUFF

3 GRANODIORITE (GRADATIONAL GRANITE TO DIORITE)

2 ARGILLITE (?)

1 HAZELTON GROUP (ANDESITE, BASALT, DACITE, RHYOLITE, BRECCIA, AGGLOMERATE, TUFF, GREYWACKE, SANDSTONE, CONGLOMERATE AND LIMESTONE.

TO ACCOMPANY A GEOLOGICAL, GEOPHYSICAL AND GEOCHEMICAL REPORT BY R. WOOLVERTON ON THE O.V.P. AND M.K. GROUPS, TROITSA LAKE, B.C. OMINECA M.D. DATED

2026

ASTON RESOURCES PROPERTY

TROITSA LAKE PROJECT
WHITESAIL LAKE AREA

GEOLOGY MAP

SCALE: 1" = 500'

SURVEY BY: SILVER STANDARD MINES LIMITED, 1967

MAP 1

TROITSA LAKE

WL 2947'

BLANKET

OVP 1 | OVP 2 | OVP 3 | OVP 4 | OVP 5 | OVP 6 | OVP 7 | OVP 8 | OVP 9 | OVP 10 | OVP 11 | OVP 12 | OVP 13 | OVP 14 | OVP 15 | OVP 16 | OVP 17 | OVP 18 | OVP 19 | OVP 20 | OVP 21 | OVP 22 | OVP 23 | OVP 24 | OVP 25 | OVP 26 | OVP 27 | OVP 28 | OVP 29 | OVP 30 | OVP 31 | OVP 32 | OVP 33 | OVP 34 | OVP 35 | OVP 36

MK 1 | MK 2 | MK 5 | MK 6 | MK 7 | MK 8 | MK 9 | MK 10 | MK 35 | MK 38

73-1 | 73-2 | 73-3 | 73-4 | 73-5 | 73-6 | 73-7 | 73-8

SNOW

SYMBOLS

- ○[9] ROCK CHIP SAMPLE IN PLACE
- △[2] ROCK CHIP SAMPLE IN TALUS

LEGEND

○	0-3	NOT ANOMALOUS
●	4-7	WEAKLY ANOMALOUS
●	8-15	MODERATELY ANOMALOUS
●	16-32	STRONGLY ANOMALOUS
●	>32	EXTREMELY ANOMALOUS

TO ACCOMPANY A GEOLOGICAL, GEOPHYSICAL, AND GEOCHEMICAL REPORT BY R. GRANT-TAYLOR ON THE O.V.P. AND M.K. GROUPS, TROITSA LAKE AREA, OMINECA M.D., DATED Aug. 3/76

D.A. DAWSON
BRITISH COLUMBIA
ENGINEER

2026

ASTON RESOURCES PROPERTY

TROITSA LAKE PROJECT
WHITESAIL LAKE AREA

ROCK GEOCHEMISTRY MAP

TOTAL Mo IN PPM

SCALE: 1"= 500'

SURVEY BY EVERGREEN EXPLORATIONS LTD.
SMITHERS, B.C.

MAP 4

TROITSA LAKE

W/L 2,947'

PLANKET

SNOW

OVP 1
OVP 2
OVP 3
OVP 4
OVP 5
OVP 6
OVP 7
OVP 8
OVP 9
OVP 10
OVP 11
OVP 12
OVP 13
OVP 14
OVP 15
OVP 16
OVP 17
OVP 18
OVP 19
OVP 20
OVP 21
OVP 22
OVP 23
OVP 24
OVP 25
OVP 26
OVP 27
OVP 28
OVP 29
OVP 30
OVP 31
OVP 32
OVP 33
OVP 34
OVP 35
OVP 36

MK 1
MK 2
MK 5
MK 6
MK 7
MK 8
MK 9
MK 10
MK 36
MK 38

73-1
73-2
73-3
73-4
73-5
73-6
73-7
73-8

SYMBOLS

○[80] ROCK CHIP SAMPLE IN PLACE

△[24] ROCK CHIP SAMPLE IN TALUS

LEGEND

- 0-19 NOT ANOMALOUS
- 20-39 WEAKLY ANOMALOUS
- 40-79 MODERATELY ANOMALOUS
- 80-160 STRONGLY ANOMALOUS
- >160 EXTREMELY ANOMALOUS

TO ACCOMPANY A GEOLOGICAL, GEOPHYSICAL, AND GEOCHEMICAL REPORT BY R. W. [illegible] ON THE O.V.P. AND M.K. GROUPS, TROITSA LAKE AREA, OMINECA M.D., DATED [illegible]

PROFESSIONAL ENGINEER
PROVINCE OF BRITISH COLUMBIA
D. A. [illegible]

ASTON RESOURCES PROPERTY

TROITSA LAKE PROJECT
WHITESAIL LAKE AREA

ROCK GEOCHEMISTRY MAP

TOTAL Cu IN PPM

SCALE: 1"= 500'

SURVEY BY EVERGREEN EXPLORATIONS LTD.
SMITHERS, B.C.

2026

MAP 3

TROITSA LAKE

W/L 2947'

BLANKET LAKES

W/L 3500'

OVP 1
OVP 2
OVP 3
OVP 4
OVP 5
OVP 6
OVP 7
OVP 8
OVP 9
OVP 10
OVP 11
OVP 12
OVP 13
OVP 14
OVP 15
OVP 16
OVP 17
OVP 18
OVP 19
OVP 20
OVP 21
OVP 22
OVP 23
OVP 24
OVP 25
OVP 26
OVP 27
OVP 28
OVP 29
OVP 30
OVP 31
OVP 32
OVP 33
OVP 34
OVP 35
OVP 36

73-1
73-2
73-3
73-4
73-5
73-6
73-7
73-8

MK 1
MK 2
MK 5
MK 6
MK 7
MK 8
MK 9
MK 10
MK 35
MK 36
MK 38

2000
1500
500
0

ASTON RESOURCES PROPERTY

TROITSA LAKE PROJECT
WHITESAIL LAKE AREA

RADEM MAP

SCALE: 1" = 500'

SURVEY BY EVERGREEN EXPLORATIONS LTD.
SMITHERS, B.C.

TO ACCOMPANY A GEOLOGICAL, GEOPHYSICAL, AND GEOCHEMICAL REPORT BY R. WOOLVERTON ON THE O.V.P. AND M.K. GROUPS, TROITSA LAKE, B.C., OMINECA M.D., DATED 10 Sept 1969

R. Woolverton

2026

MAP 5

TIE-LINE

L 29 W.
L 28 E.
L 27 W.
L 26 E.
L 25 W.
L 24 E.
L 23 W.
L 22 E.
L 20 W.
L 19 E.
L 18 W.
L 17 W.
L 16A E.
L 16B W.
L 15 W.
L 14 E.
L 13 E.
L 12 E.
L 11 E.
L 10 W.
L 9 W.
L 8 W.
L 7 E.
L 6 W.
L 4 E.
L 5 W.
L 3 E.
L 2 W.
L 1 E.

TROITSA LAKE

BLANE

SEEL LAKE

COLES LAKE

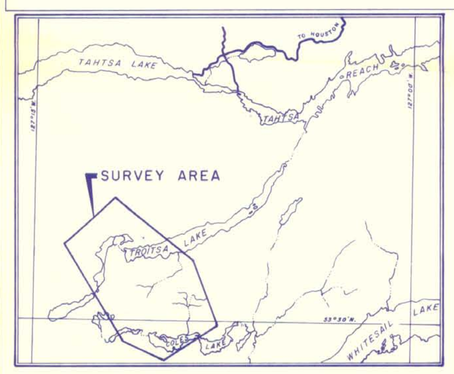
OVP

MK

EM

1 2 3 4 5 6 7 8

LOCATION MAP



SCALE : 1" = 4 miles

27°00'

TRUE NORTH
MAG. NORTH

SCALE : 1" = 1/4 MILE

LEGEND

CONDUCTORS

- Prime Conductor
- Secondary Conductor
- Surficial or Formational Conductors
- Conductor Axis

Base intensity arbitrary

Flight lines and photo identity points 882 862 851

Isomagnetic Contours 500 Gammas

100 Gammas

Magnetic Depression

TO ACCOMPANY A GEOLOGICAL, GEOPHYSICAL, AND GEOCHEMICAL REPORT BY R. WOOLVERTON ON THE O.V.P. AND M.K. GROUPS, TROITSA LAKE, B.C., OMINECA M.D., DATED 10 Sept 1969

R. Woolverton

ASTON RESOURCES PROPERTY

TROITSA LAKE PROJECT

HELICOPTER MAGNETIC AND ELECTRO MAGNETIC SURVEYS

WHITESAIL LAKE AREA, B.C.

JULY, 1969.

2026

2026

Evergreen Explorations Ltd.

• R. WOOLVERTON
GEOLOGIST

• R. C. O'BRIEN
FIELD SUPERVISOR

• JOHN C. OSWALD and COMPANY
ACCOUNTANT
635 - 789 W. PENDER ST.
VANCOUVER 1, B.C., CANADA

CONTRACT EXPLORATION

• 5424 HALIFAX ST., BURNABY 2, B.C., CANADA, PHONE - 299-6998
• P.O. BOX 604, SMITHERS, B.C., CANADA PHONE - 847-3523

Oct.21,1969

R.H. McCrimmon,
Chief Gold Commissioner,
Dept. of Mines and Petroleum Resources,
Victoria, B.C.

Dear sir,

Re: OVP 1/36 and MK 1/60 claims
Your file 166-Omenica

The EM anomolies shown on Map 6 were selected from the EM tapes supplied by Lockwood Corp. Copies of the anomolous profiles, which were described in detail in the text of the report, are enclosed for your examination. They were located on maps by use of fiducial points similar to those of the mag tapes.

Yours truly,

R. Woolverton

R. Woolverton P. Eng.

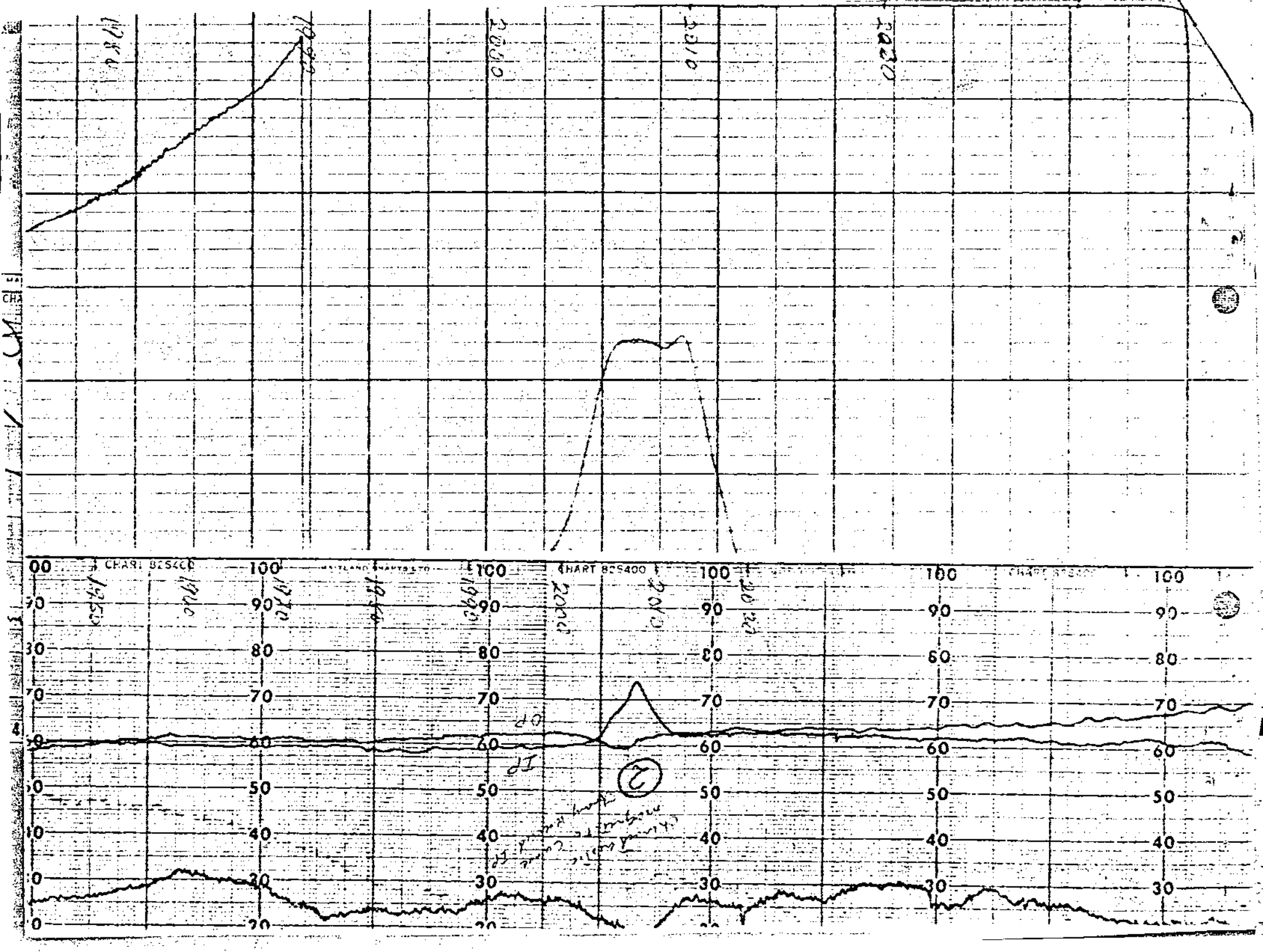
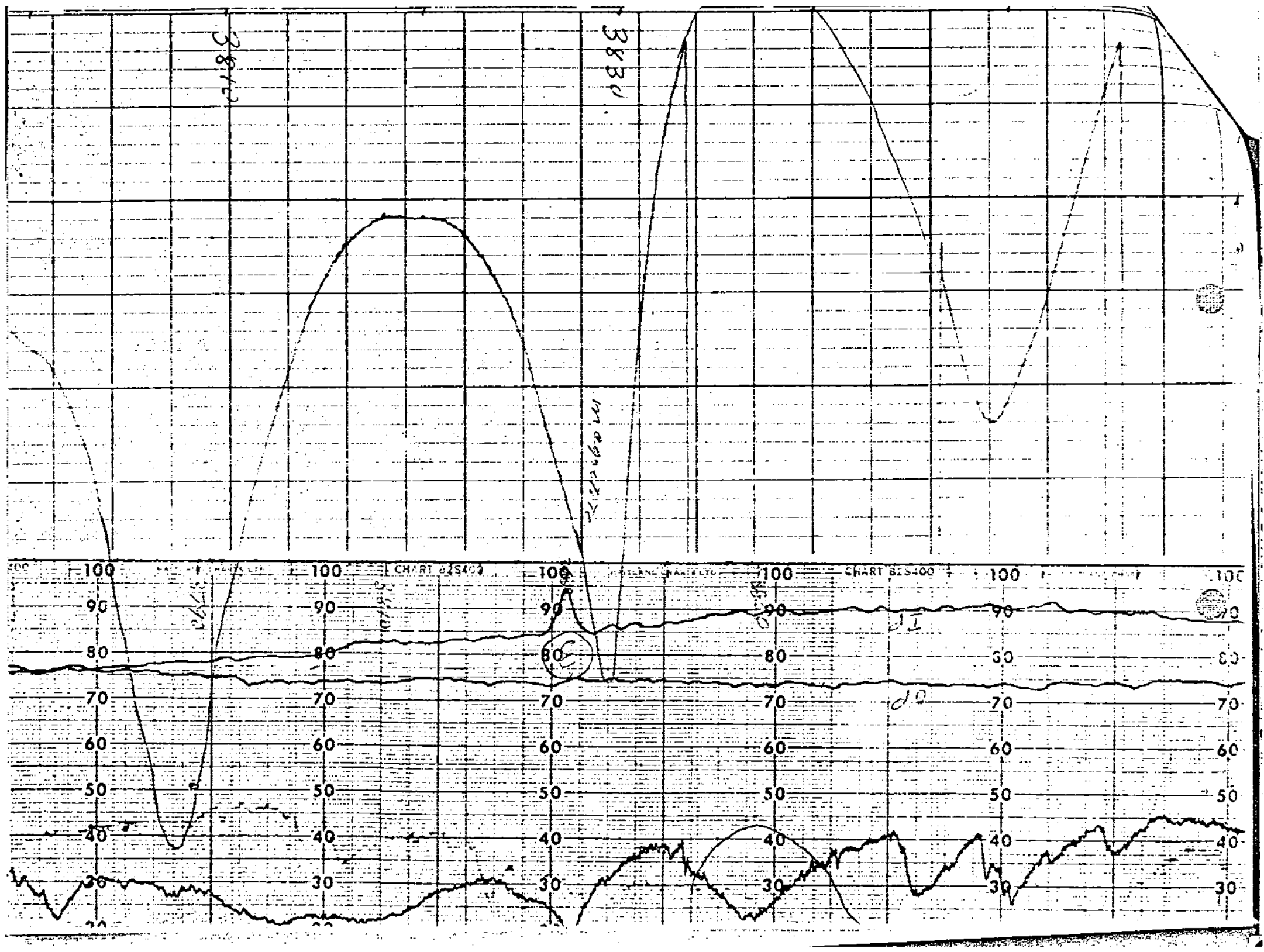


CHART 825400
DP
IP
2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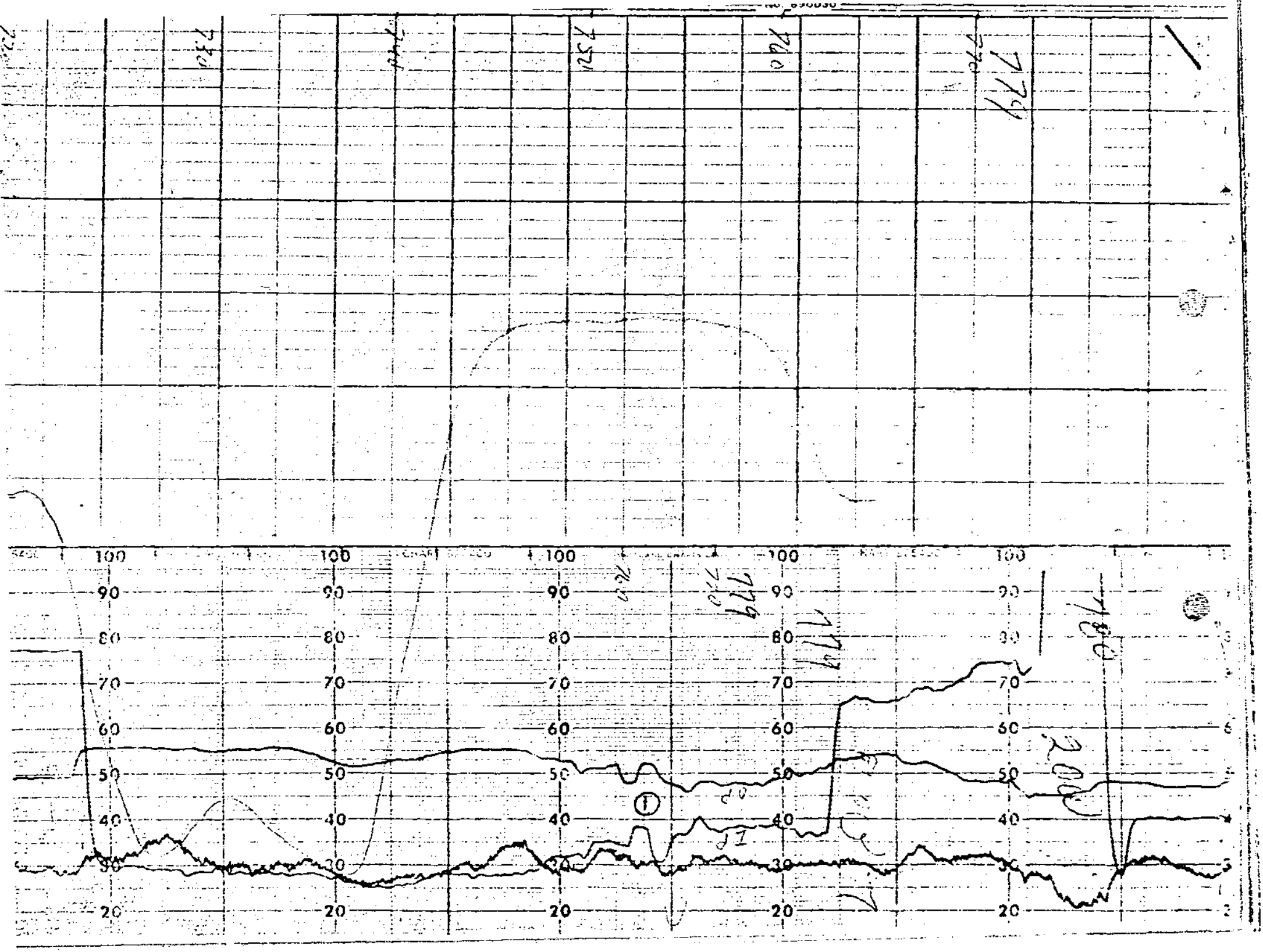
100
90
80
70
60
50
40
30


100
90
80
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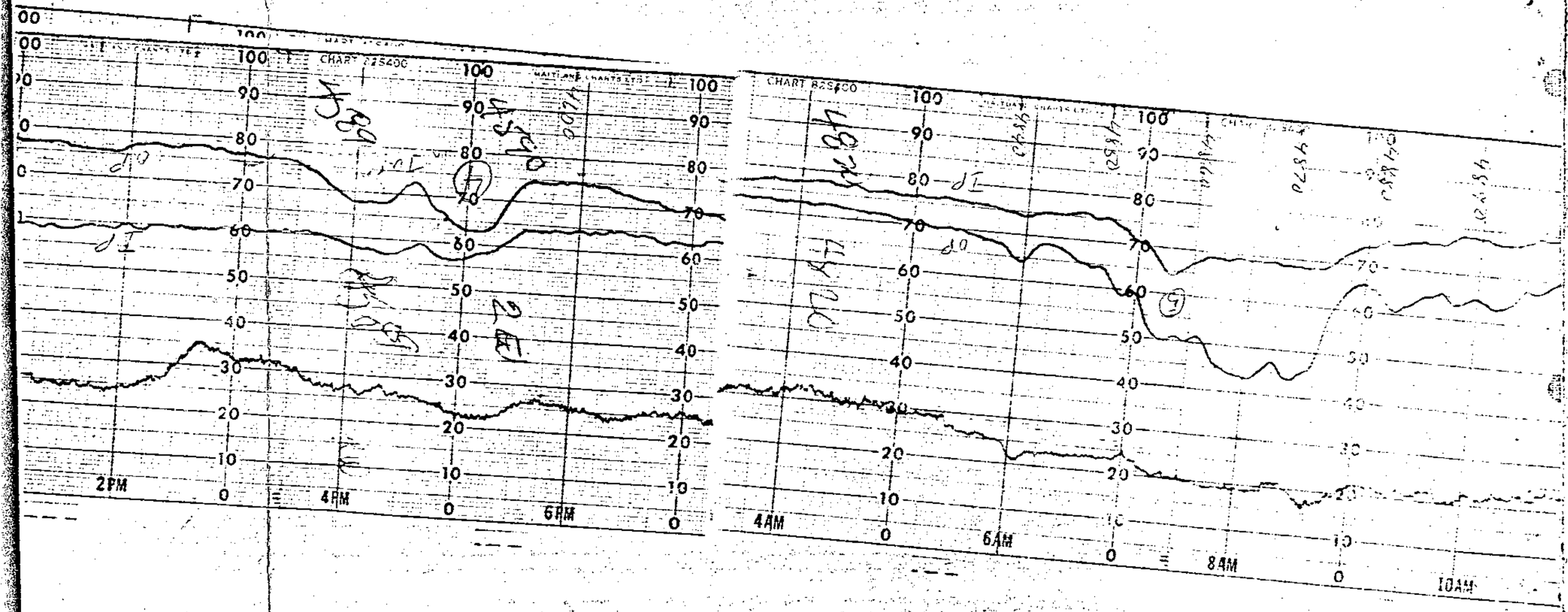
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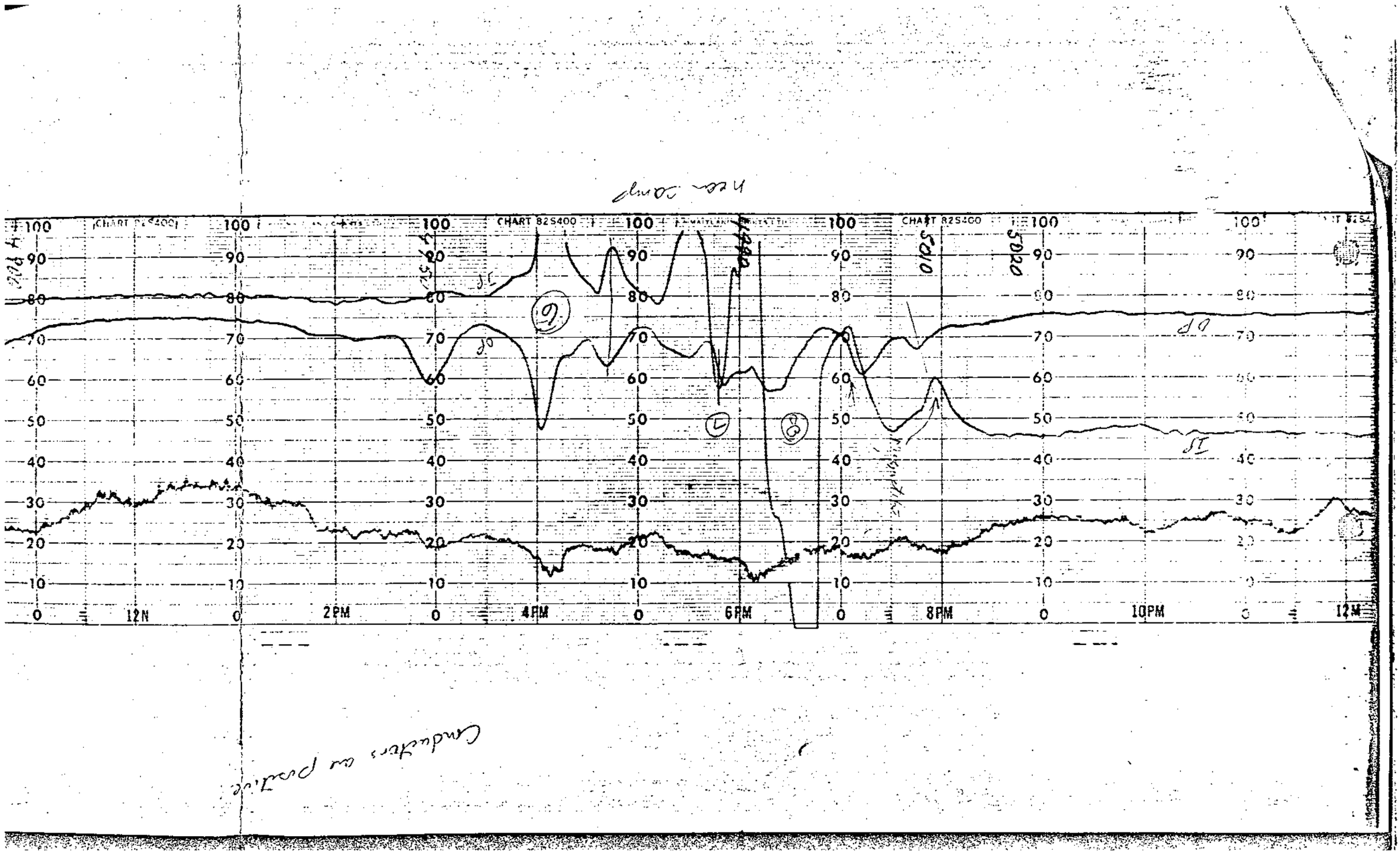
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