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GEOPHYSICAL CHEMICAL

TENORE TO

DAWOOD MINES LED.

MORTH VARCOUVER, B. C.

ON THE

TOTE MINERAL CLAIMS

MRAS MERKITT, B. C.

31

SHRWIN F. KRLLY, F. ENG.

GEOLOGIST AND GEOPHYSICIST

Median, B. C.

JULY 15, 1970

Geophysical-Geochemical Geophysical Geophysical Geochemical Ltd.

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Department of Mines and Petroleum Resources ASSESSMENT REPORT

NO. 2468 MAP

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DATE OF BUILDING

Geophysical and geochemical surveys were carried out on 16 mineral claims of the DOTE group, in the old Aspen Grove copper camp, British Columbia, in the spring of 1970. The techniques employed were soil sampling (samples analyzed for copper), a magnetometer survey and an electromagnetic survey employing a Sonka Salé VIF unit. The work was carried out under the direction of George Cressy, Jr. and Jack Samson, both of Ferritt. These men have been known to me for several years and I recognise them as competent, conscientious and reliable operators.

The secretary of Damoed Mines, Jack Damson, requested me to analyze the results, give my interpretations and recommendations and prepare the geophysical-geochemical report. The mork performed was to estimate the assessment requirements on the claims DOTS \$1 to \$9 and \$11 to \$37. This report is to support the submission of the surveys for assessment work.

The base maps were drawn by Jack Pauson. I then prepared the contour maps of the magnetic and soil survey results.

DOTE NO.37	DOTE NO. 36	,				
DOTE NO.35	DOTE NO.34	E A A 17 7, BC. /				
DOTE NO.33	DOTE NO.32					
DOTE No.31	DOTE No. 30		DOTE NO.8	DOTE NO.6	DOTS No. 4	DOTE NO.2
DOTE No.29	DOTE NO.28		DOTE NO.7	DOTE NO.5	DOTE NO.Z	DOTE NO.1
DOTE NO.22	DOTE No.23	A S P E N & Roys B.C.	DOTE NO. 17	DOTE NO. 15	DOTE NO.13	DOTE NO. 11 DOTE NO.9
DOTE NO.21	DOTE NO.25	DOTE NO.19	DOTE NO. 18	DOTE NO.16	DOTS NO. 14	POTE WO.12
DOTE NO.26	DOTE NO.27	DOTE NO.20	DOTE NO.2/		DOT	MINES LTD. (N.P.L.) E GROUP GROVE, B.C.
	Fig. 1 Claim Map		B.C. HWY. No. 5	SCALE	"= ISOO!	

CIAIIS

There are 36 claims in the DOTE group, which straidles the Ferritt-Princeton Highway (Highway #5) at Aspen Grove. See Fig 1. The claims DOTE #1 to #33 (there is no DOTE #10) were staked in Pay, 1966, by Jack Dawson as agent for Dawsod Mines Ltd., then of Kelenma. DOTE #34 to #37 he staked in the same capacity, in Pay, 1968. The data are as follows:

MAKE OF CLADA		YALI	D T	2
DOTE #1 to #3	30233/20	Lay	13,	1970
DOTE #9	30223	lay	16,	1970
DOTE ALL to ALS	30225/32	ay	16,	1970
DOTE #19 to #33	30579/93	ay	30,	19 70
DOTE 194 to 137	36904/7	ay	27,	1970

All the above claims were grouped into one group on June 26, 1968, by Notice to Group #1215.

These data are on file in the office of the Mining Mecorder in Marritt.

The validity dates on the DOTS claims will be extended one year, subject to the filling and acceptance of this report.

LOCATION AND ACCESS

The DOTE group of 36 mineral claims is located in the Micola Mining Division, on the interior plateau, at an elevation of around 3500 to 4000 ft. The group straigles Highway No. 5 (Merritt-Princeton Highway) at the small settlement of Aspen Grove, some 16 miles southeast of Merritt. The co-ordinates for Aspen Grove, located about the middle of the group, are approximately:- longitude 1200 37% West, latitude 490 56% North.



The area is shown near the middle of the northerly portion of the Tulameen topographic sheet, 92 H/NE, at the scale of 1 inch to 2 miles. See Fig. 2.

Access is via Highway #5 from Princeton or Merritt. Humarous ranching and logging roads turn off from this highway, including some near Aspen Grove, which give easy access to the area of these claims.

GEDICGICAL SETTING

The bedrock of the claim area, and for many miles to the north and south, consists of the Ricola Series of volcamics and sediments, of Triassic age. The series is predominantly of volcamic origin, comprising flows from rhyolitic to baseltic in composition, with interbedded tuffs. There are some sedimentary beds in the series, usually conglomerates, shales and limestones. Prevailing strikes are northerly to northeasterly. A short distance north of the DOTE group, on the Porcupine claims of Amalgamated Resources Ltd., the volcamics are moderately thin bedded, usually a few feet thick, and dip southeasterly at about 30°.

The Nicola beds were intruded in Jurassic time by batholiths as well as by plugs of lesser dimensions, varying in composition from gabbre to granite. Diorites, quarts districts and granodicrites predeminate, however. Many of these are mineral-bearing, usually carrying copper and sometimes molybdenum. They are believed to be the source rocks of mineral deposits frequently found in Micola host rocks, as at Copper Mountain and at the Craigment wine. In the Highland Valley type of occurrence, the huge Guichen Batholith is both source rock and host rock for the great copper ore bodies being developed in that Valley.

In the Aspen Grove area, there are several Jurassic intrusives not far from the DOTE claims. The southwestern and of the Penask batholith lies some six or seven miles

to the northeast. Five or six miles to the southwest, the Allison Lake intrusive lies along the highest. The geological map accompanying Memoir 203, "Geology and Mineral Deposits of the Princeton Map-Area, British Columbia" by H. H. A. Rice, Department of Mines and Resources, Ottawa, 1947, shows these bodies as well as several smaller plugs of granite or granoiderite, intruding the Micola beds near the claim group.

The Micola rocks are favorable host rocks for the reception of mineralizing solutions. Of the numerous Jurassic intrusives which have penetrated them, some were source rocks of mineralizing solutions. The band of Micola series, stretching from the U. S. border to Kawloope Lake, therefore presents a favorable hunting ground for copper deposits.

Early prespectors working in this area, discovered copper mineralization from Lake Missesula north to Micola Mtn., in the beginning of this century. As a result, the area around Aspen Grove became a center of activity and was known as the Aspen Grove Camp. The early prospectors were looking for shipping ore, which they did not find. The Aspen Grove region is nevertheless dotted with old trenches, pits, adite, and shafts which reveal copper mineralization assaying up to three or four percent copper. The mineralization thus uncovered, yields useful information which can help to guide the interpretation of geophysical data. Such information concerns both the type of mineralization and the structural or formational controls of mineral deposition.

The sulphide minerals usually occur in disseminated form in the beds of the Micola volcanies. The dissemination may be dust-fine particles to moderate-eight blebs and crystals. These are found in both tuffs and flows. In some cases, the mineral particles are so fine as to be almost invisible to the naked eye. Copper sulphide predominates and pyrite, where it occurs, is usually sparse; pyrchotite is almost

non-existent. Expertite does not seem to have accompanied the mineralization, but is found as a primary constituent in some of the flows (such as in basalt). To a depth of around 100 feet, secondary beautic and chalcocite frequently are developed. Palow about 100 feet, although the depth will vary from place to place, the predeciment sulphide is the primary copper sulphide, chalcopyrite.

The absence of regnetite as an ecompanisent of mineralization, means that magnetic surveys will yield little or no useful information on the occurrence of copper sulphides. Such surveys may, however, yield useful data on the occurrences, transa, and dislocations of volcanic flows with distinctive magnetic reactions.

Since the sulphide mineralization is largely disseminated, spontaneous polarization and the usual electromagnetic nothers are not likely to produce distinctive anscalies. The induced polarization and the V.L.F. (very low frequency) methods are probably better adapted to the type of occurrences found in the Aspen Grove area.

The structural and formational controls of mineral deposition still present most questions. There is evidence favouring the hypothesis that mineralizing solutions have been introduced along faults cutting the bedding of the Sicola series. There is also a good deal of evidence that mineralization has favoured particular beds of the Nicola series and hence may be bedding-controlled. Possibly, mineralizing solutions following fault somes and shears have selectively perseated certain beds. In any case, it became important to trace the trends of the underlying Sicola beds and to locate areas of faulting, shearing, dislocation. Thus it is in this field of enquiry that magnetic methods may be helpful.

SEED LAYOUT

A grid was laid out to cover the claims on the east side of Higheny #5, namely DOTS #1 to #18, not including #9 (there is no DOTE #10). A north-south base line was run approximately slong the location line between DOTS #4 and #6 on the north to between #10 and #16 on the south. This line was 4800 feet long. East-west grid lines were turned off at 400 foot intervals, beginning with line 0400 on the north and ending with 48400 on the south. These lines are of varying lengths, but usually run about 2500 feet west of the base line and 2000 feet east of it. Grid lines and base lines were picketed at 100 foot intervals, the stakes being numbered east and west from the base lines. See Figs. 3 and 4.

The surveys conducted on this grid were magnetic, V.L. and soil sampling.
Soil samples were taken and observations were made with the magnetometer and the
V.L. . instrument, at 100 foot intervals along the grid lines. Soil samples were
also taken at 100 foot intervals along the base line. The resultant observations
were entered on appropriate plan maps, as described below.

ACTORIC SUCCES

Fagnetic observations were made on the vertical component of the earth's respectic field, using a flumpate Esgnetometer. The instrument exployed was a Sharpe NF-1 Flumpate Esgnetometer, social #511377. It is a hand-held instrument and needs only coarse leveling and is independent of extentation. Scales available are, for full scale deflection, 1,000, 3,000, 10,000, 30,000, 100,000 gammas. The constitutive at the 1,000 gamma range is 20 gammas per scale division, making it possible to estimate to about 5 gammas. The temperature stability is 1 gamma per degree fabrenheit.

As a base and control point, the station at 2100 W on line 40+00 S was chosen. It was arbitrarily given a value of 1,000 games and all readings were referred to that as datum. Control points were also established at stations 2100 W, line 24+00 S and at 2100 W, line 12+00 S. The survey was carried out in loops, taking the readings at 100 ft. intervals along all grid lines, returning to a control station three times per day in order to correct for diurnal variations. At some locations where details were desired on anomalous somes, the reading interval was reduced to 50 ft.

The recorded values, corrected for disrnal variation, were entered on a plan map of the grid area and the values contoured thereon. See Fig. 3. The claim lines are also shown on this map.

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as well as along the base line. Holes were dug, with mattock and shovel to the "B" horizon, or to the "C" horizon wherever the "B" was lacking. Soil samples were then taken with a wooden spoon and placed in Kraft sample bags, duly Mentified, and shipped to Bendar-Tegg and Company in Borth Vancouver for determining copper values. This company uses hot nitric acid for extraction of the metal, whose quantities are then determined by the stock absorbtion process.

The values returned by Bondar-Clegg and Company were entered on a plan map of the survey grid and the values were then contoured. See Fig. 4. The claim lines are also shown on this map.

The technique known as the VIF-E.U. method, was used on the same grid, covering a part of the PCTE mineral claim group. The method relies on measurements of the broadcasting field examating from one of the U.S. Mavy stations set up to communicate with naval vessels, especially submarines. The station to which the field instrument was tuned on this survey, was the one at Jim Crook, near Scattle, Washington, broadcasting with 300 Kilowatts power on a frequency of 18.6 kHz. The stations mentioned utilize vertical antennae, with the result that the electromagnetic fields examating therefrom, are horizontal.

When the field radiating from such an enterms encounters a conductive fermation, it generates a secondary field therein. This secondary field distorts the primary one, and is not in phase with it. The resultant distortions and phase relationships are deduced from observations nade on the vertical and horizontal components of the resultant electromagnetic field. The instrument explayed to measure these parameters, was a Sonka 28-16, namufactured by Geomics Ltd., of Toronto, serial number 78.

seadings are taken of the real (in-phase) and quadrature (out-of-phase) portions of the vertical component of the secondary field. They are plotted as profiles on the plan of the grid lines, each profile being laid out along the grid line on which the readings were taken. See Fig. 5. A sharp cross-over of the real component, from above to below the line, usually indicates a steeply dipping, sheet type conductor. It reads in the direction in which the instrument was facing, in this case west, so the lines should be read from right to left (east to west). A high ratio of in-phase to quadrature component, indicates a good conductor of large size.

The VL method is a sensitive one, which presents both advantages and disadvantages. It will respond not only to sulphide mineralization, but also to conductive faults, contacts, shear somes, etc. Thus, it provides a useful tool for suppling such geological

structures. On the other hand, it may senetimes be difficult to distinguish anomalies caused by such features from those arising from sulphide deposits. Consequently, it is advisable to correlate results from the VIF method with those from other techniques, such as geochemical and magnetic surveys. Further checks by other electrical methods may also be desirable.

LICUSSION OF RESIDES

lemetic

The contours of the magnetic intensities recorded on this group of claims, reveal an interesting contrast between the east and the west portions of the grid area. See Fig. 3. The base line more or less marks the boundary between the two contrasting areas. West of this line, the magnetic relief is generally slight and the values are low. The contoured area is marked by mamerous somes of low values, outlined by depression contours. There are only a few peaks of magnetic intensity, rarely attaining a maximum value of 1500 gammas. In general, each is confined to a single grid line, at least in the upper values.

These phenomena of low values, and only weak, scattered peaks, indicate an unionlying bedrock of generally low magnetic susceptibility and lacking in marked contrasts.

Under these discussioness of low magnetic relief, the bedrock is probably not one of strongly basic constitution, i.e. not one high in iron-magnesium minerals. It could consist of acidic (high in cilica) flows or intrusives, or of sedimentary beds. The general trend of the contours implies a northerly strike for underlying flows or beds.

East of the base line, the magnetic pattern is characterised by mumerous peaks with values as high as 3000 gausses. Depression contours are less numerous than in the western portion and in general do not indicate values quite as low. Furthernore, they

are more constricted and tend to a narrow, elongate chape, whereas in the western portion many of the depression contours are bread and diffuse in outline.

The high value contours are also narrow and clongeted with a striking, north-couth trend. There is an indication of a curvature in this trend, with a general tendency to swing from a north-couth orientation in the south-castern portion of the grid area to a slightly west of north trend in the north-central portion. This could indicate a slight bend in the unierlying beis, but needs careful checking in the field as there is also a possibility that this apparent change in direction could be due to topography.

The topography of the western portion of the survey area is reported to be flatter than in the eastern portion. A north-south ridge structure spreads castuard from the area of the base line. Under such circumstances, the positions of the underlying beds or flows at bedrock surface, will be affected by the bedrock topography. If, as seems probable, the dips are to the east, then the bedrock surface expression of a given bed will lie further to the west at a high elevation than at a lower one. Thus, it will take careful correlation in the field between the topography and magnetic contours to discover whether the contours reflect only the topography or, on the other hand, if they are indicative of slight folding or of dislocation by faulting. This information will be useful in tracing either specific beds carrying mineralization, or beds lying in some definite relationship to mineralized beds.

Elongate magnetic loss between nurrow bands of highs, may indicate relatively acidic flows or tuffs, or possibly even sedimentary bads, intercalated in a series of flows of predominantly baseltic or near-baseltic character. Preminent loss interrupting a given band of highs, may signify a discontinuity caused by faulting. It could also be the result of the alteration of some of the magnetite in the flow, due to the action of circulating hydrothermal solutions.

There may be other causes capable of producing these effects, but the above are of most immediate interest as possibilities.

Soil Analyses

An interesting feature of the contours outlining copper soil enomalies, is that they seem to avoid the magnetic highs. For example, the central portion of the area east of the base line, where magnetic peaks abound, is notably deficient in soil anomalies. See Fig. 4.

In general, there is little contrast between the soil anomalies in the western and eastern portion of the survey grid. In both areas there are long anomalies and short anomalies, as well as weak ones and strong ones. There are occasional, isolated anomalies, confined to one station and which show high intensities. Noticeable examples occur on line 36+00 3 where an anomaly of 900 ppm lies 500 ft. west of the base line and another of 410 ppm lies 1600 ft. east of the base line. High values such as these, severely restricted in area, lead to the suspicion that they correspond to an isolated pocket of mineralized material, such as a boulder, buried in the overburden. With this in mind, they should nevertheless be investigated.

Of greater interest are the anomalies which extend across several grid lines, such as those on claims Dote \$7, on Dote \$1, \$2, \$3, and \$4, and on Dote \$16 and \$13. These areas are characterised by elongated anomalies with fairly high peak values.

The background value for copper in this area, is about 25 ppm and the first contoured value is 3 times background, or 75 ppm. Contours of 100, 200, 300, etc. ppm are drawn more heavily and correspond to values of 4, 8, 12, etc., times background. Anomalies of three times background and higher, are considered to be of definite interest.

There is no consistent correlation between soil anomalies and magnetic anomalies, except as previously mentioned that the soil amenalies seem to avoid the magnetic highs. Otherwise, some of them coincide with magnetic lows, some lie on the flanks of magnetic highs and some lie at the ends of magnetic highs or in saddles between magnetic highs.

The avoidance of magnetic highs by the copper soil anomalies, implies that the more basic flows, such as basaltic types, presented unfavorable host rocks for copper deposition. The tendency to follow lows, or low saddles between highs, indicates at least two possible controls of copper deposition. The elongate lows between ridges of highs, may correspond to less basic, more acidic flows or tuffs, which presented favorable conditions for copper deposition. The saddles which interrupt a given band of magnetic highs, may be due to transverse faults or shears; when occupied by copper soil anomalies, they imply a favorable environment for copper deposition in the shear none. Also, such interruptions by magnetic lows of a band of highs, could be due to alteration of magnetite by the hydrothermal solutions from which the copper minerals are precipitated. The alteration of the magnetite, (to hematite, for example) would destroy the magnetic effect of the formation at that locality.

There are several alignments of interesting soil anomalies, with a generally east-of-morth trend. One of the most persistent starts in the south-west corner of the grid area, on claim Dote \$18. Scattered, weak to moderate soil anomalies occur on nearly every grid line, in a band extending to the north boundary of the grid area, 400 ft. west of the base line, on line 0+00. A long individual anomaly in this band, extends from 1100 feet west of the base line on line 20+00 S, to 800 ft. west of the base line on line 12+003. It corresponds to a magnetic low and is separated by a magnetic high from what might be its continuation, at 1400 feet west of the base line on line 28+00 S.

Just west of the above described, long anomaly, a moderately strong soil anomaly at 800 ft. west of the base line, on line 4400 S, shows a parallel orientation and points toward a very strong high on line 12400 S, 1300 ft. west of the base line. This latter copper anomaly, however, has an independent north-south orientation. Other anomalies in this survey-area, also show the same orientation, as will be discussed later in this chapter.

A second band of soil enoughly highs, with a similar east-of-morth trend, lies some 600 ft. east of the first one described. In this one, the anomalies are generally stronger but more scattered. It starts on line 40+00 S, 900 ft. west of the base line and extends north-easterly to the vicinity of the intersection of the base line and line 4+00 S. The anomalies roughly correspond with magnetic lows, except that at the last mentioned point, near the north end of the base line, the anomaly lies on the flank of a magnetic high.

A strong high with no continuation to the north or south, as already noted lies on line 36*00 S, 500 ft. west of the base line. It occurs between a high of the trend just described and a high on the next parallel one to the east. Jack Dawson stated that a large block of fresh district appeared in the roots of an uprooted tree in the locality of this high. No mineralisation was evident. The manner of occurrence of this high and the description of the appearance of the presumed bedrock, tend to reinforce the suspicion that the reading here is due to an erratic packet of mineralisation, possibly a boulder, in the overturden.

The next trend of highs is discontinuous and is more preminent in the south portion of the grid. It starts with the anomaly between 300 ft. and 800 ft. west of the base line on line 48+00 S, and extends north-easterly to the intersection of the base line and line 32+00 S. The prolongation of that trend would then lead

through a crowded nest of magnetic highs, devoid of soil copper anomalies. If the trend is projected however, it encounters a strong anomaly 900 ft. east of the base line on line 16+00 S; a moderate one 1400 ft. east of the base line on line 8+00 S and another 1600 ft. east of the base line on line 4+00 S. These latter anomalies, however, lie in an assemblage of soil anomalies of north-south orientation. Their correlation with the north-easterly trending band is therefore questionable.

East of the base line, there are several soil anomaly highs, rather widely scattered and not too prominent, most of which exhibit the prevalent north-easterly trend. In the north-east corner of the survey grid, however, there are several soil anomalies which have a north-south strike. The regnetic contours in the same area show a similar strike. For the magnetic highs, this is not so anomalous a direction as it is for the soil copper highs. The evidence may indicate a swing in this vicinity, to a more northerly strike of the underlying beds. The area is also reported to be one of ridges and troughs with a north-south direction, and at least some of the anomalies coincide with north-south gulleys. Careful evaluation of these anomalies is essential, because they may represent a trapping and accumulation of copper ions by the clay soil which tends to form in low areas of poor drainage. Under these conditions, very strong copper anomalies could develop which are totally unrelated to the immediately subjacent bedrock. The copper thus might have drained from mineralized material higher up on the hillside.

Before undertaking intensive investigation of any soil anomalies, they should be carefully compared with the topography in order to determine whether or not a situation such as described above, could be responsible for the anomalous highs. If so, it is highly probable that the strong anomalies will be of considerably loss interest than the weaker ones.

The trend of the soil anomalies is generally east of north, which is not entirely in accordance with the trend of the magnetic highs, which generally run north-south. The latter orientation is more definite in the east-central portion of the survey area, where there are no strong soil anomalies, than in the western portion. In the latter area there are some magnetic highs and lows trending east of north, with which soil anomalies showing the same trend, are associated. Consequently, there may be a slight change in strike of the bals in the eastern area, compared to the western portion.

It is surprising that there are almost no anomalous soil readings in the vicinity of the shafts, near the eastern and of line 4400 3. There are a couple of weak anomalies trending east of north, just south of these shafts. Information in the area indicates that the soil is shallow, which may mean fairly quick drainage and no thick blankst of soil in which copper could accumulate. Mineralization is evident in the shafts, and a sample taken by George Cressy, Jr. and Jack lawson from sheared wall rock in the west shaft, assayed 0.895 copper. A fragment from the dump assayed 1.355 copper. Amother fragment from the dump of the eastern shaft, assayed 0.195 copper. Assurite, malachite, and obslicecite are reported evident in the dump and in the walls of the shafts. Further investigation should be carried out in this vicinity to check the prior determinations. If they are found valid, an affort should be made to discover the reason for the lack of soil anomalies here. If it is found that good copper mineralization exists here and it fails to give soil anomalies, the validity of soil surveys in this area would be called in question. For this an evident explanation will need to be found which could be applied to other, similar circumstances.

Possibly the lack of soil anomalies near the shaft is due, as previously mentioned, to rapid drainage and that the drainage has accumulated the copper ions in low ground

to the west, forming the strong enously which extends north and south, close to the nearby claim posts.

VL Survey

The profiles recorded in the VLF survey are shown on Fig. 5. The "cross-overs" that are thought to be of interest, have been indicated by extra heavy cross lines drawn through the cross-over. The usual indication of an underlying, steeply dipping sheet of conductive material (sulphide mineralization) is given by the in-phase profile line (solid line) descending and crossing, or closely approaching the line representing the quadrature (out-of-phase) component, which is shown as a dashed line.

The cross-overs that have been emphasized, are those which are closely associated with soil anomaly highe, or which are obviously on strike with such a high. In both the western and eastern portions of the area, they correspond very well with the soil anomalies and reinforce the supesition of an east-of-north strike.

Worthy of note is the fact that there are some cross-overs in the vicinity of the shafts. Some of the cross-overs appear where there are no nearby soil anomalies, as on the east end of line 0+00. One, near the east end of line 8+00 S, is associated with a weak soil copper high, and others lie on the prolongation of its trend, to the north and south. Lack of soil anomalies around the shaft and the relatively poor response of the VLF measurements in this area, leads to some suspicion about the strength of the mineralization here. The VLF profiles and especially the soil anomalies, indicate that the area west of the base line may well be more provising than that east of it.

SUPPLIES AND CONCLUSIONS

An excellent program of geophysical and geochemical investigations by a variety of techniques, has been conducted on the DOTE group of mineral claims. These claims at Aspen Grove, near Ferritt, B. C., belong to Passoci Films 14d. The techniques employed were geochemical sell analyses for copper, plus magnetic and the VLF electromagnetic methods. These procedures should be extended over the balance of the baldings in this group.

Copper-bearing beds in the Micola velcanics, or mineralised shear sones at flat angles to the strike of the Micola beds, are indicated by the bands of soil anomalies trending east-of-north. "Cross-overs" in the VLF readings correspond, in many cases, with anomalous soil areas and also confirm the east-of-north trends. The soil anomalies generally avoid magnetic highs and are found in lows, or on the ends or flanks of highs. The magnetically strong formations, probably near-baseltic in composition, therefore seem to be unfavorable host rocks for copper deposition. Such deposition has presumably favored more acidic beds, and mones where magnetite in the basic formations may have been destroyed by the mineralizing, hydrothermal solutions.

The eastern half of the grid is crowded with provinent alignments of regnetic highs, generally with a north-south trend. Where such highs are most numerous, in the cast-central area, soil anomalies are well-nigh absent. The north-south negnetic trends indicate a possible change in strike, or dislocation, of the bed-rock formations with respect to the western half of the grid. In the western portion, the regnetic map shows a weaker and more diffuse pattern, but with some confirmation of the east-of-porth trend evident in the soil anomaly and VII readings.

The soil anomalies and confirmatory magnetic and VLF results, imply that the western portion of the grid is more premising than the eastern part, for paraletent bands of copper mineralization.

Soil anomalies are surprisingly scarce near the chafts in the northeast corner of the grid area. Scirock mineralization is nevertheless revealed by these workings. Further investigation of this area is suggested, to develop an explanation for this phenomenon.

The geophysical and geochemical exploration on these claims has produced excellent indications of copper mineralization. Further development by bulklosing and diamond drilling is desirable, especially on the more consistent indications.

Particular attention should be paid to those locations where VLF cross-every coincide with soil anomalies of seco length, which are associated with regnetic loss.

Respectfully submitted

Shermin F. Kelly, P. Eng., Geologist and Geophysicist

Adelphi Hotel Merritt, B. C. July 15, 1970

STATEMENT OF LOUIS PROPERTY.

The following sets forth the number of persons employed on the geophysical and geochemical surveys herein described, their names, the work they performed and the time devoted thereto, and the period within which the work was done.

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Twenty days. Assistant on Line-cutting for the grid and in geochemical soil suspling survey.

Aventy days. Assistant on outding and chaining lines and in geochemical soil sampling survey.

> Shorwin F. Kelly, F. Eng., Geologist and Geophysicist

Atalphi Hotel Herritt, B. C. July 15, 1970

CERTIFICATE OF CUALIFICATIONS

I. Sherwin F. Holly, P. Eng., residing at the Adelphi Hotel in Merritt, B. C., cortify that:-

- (1) I am a registered Professional Regineer in the Province of British Columbia.
- (2) I received the Jegree of B. Sc. in Mining Engineering from the University of Mansas in 1917.
- (3) I pursued graduate work in geology and mineralogy at the Sorbonne, Ecole des Mines and Museum d'Histoire Maturelle in Paris and at the University of Karsas and the University of Toronto. I also taught those two subjects at the two latter universities. I received my training in geophysics from Prof. Commad Schlumberger of the Ecole des Mines, in Paris.
- (4) I have practised as a geologist and geophysicist in Europe, North Africa, United States, Canada, Mexico, Central America, South America and the Caribbean, since 1920. Since 1936, my work has been principally as a consultant.
- claim group, is based on the field data presented to me in transcribed form, by Jack Eswson and George Creesy Vr. They had taken the geophysical observations and supervised the collecting of the geochemical samples. I have also relied on my personal knowledge of, and experience in that general area, extending over the last ten years. References used include Semoir 200, of the Geological Survey of Canada, "Geology and Mineral Deposits of the Micola Map Area, British Columbia," by W. E. Cockfield, 1948; "Geology and Mineral Deposits of the Princeton Map Area, Memoir 24 of the Geological Survey of Canada, by H. M. A. Rice, 1947; "Five Years of Surveying with the TLF-E.M. Method" by Sorman S. Paterson and Vaine Sonks, presented at the 1969 Annual International Secting of the Society of Replevation Geophysicists.
- (6) I have no interest in the claims herein reported on, or in any securities referring to them, nor have I been promised any.
- (7) I give my permission to publish this report in a prospectus, or other literature dealing with the property herein reviewed.

Respectfully ministres.

Sharwin F. Kelly, F. Eng., Geologist and Geophysicist

Adelphi Hetel Hearitt. B. C. July 15, 1970

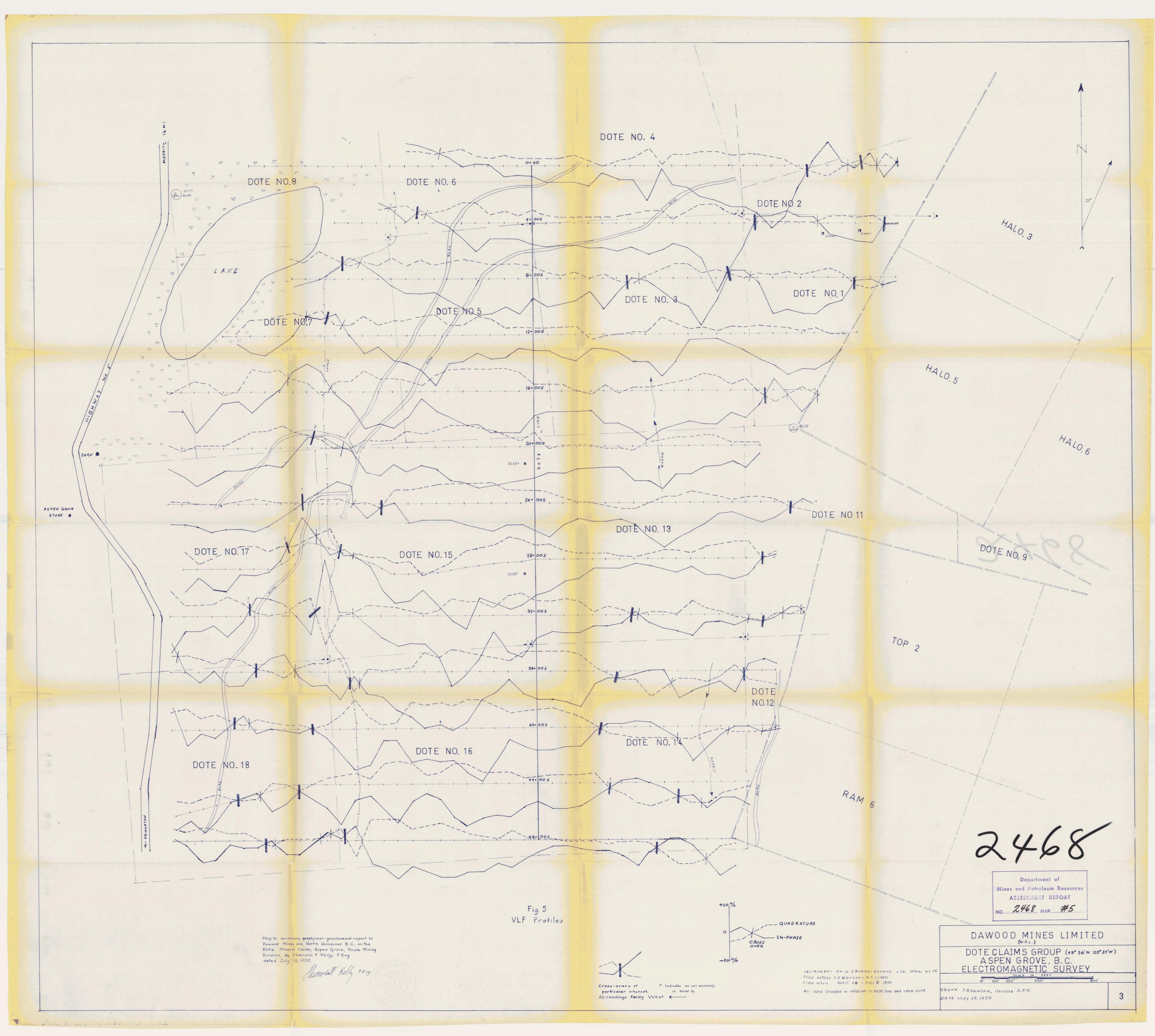
Statement Covering Expanditures incurred for Geophysical-Geochemical Surveys on DOTS Mineral Claim Group.

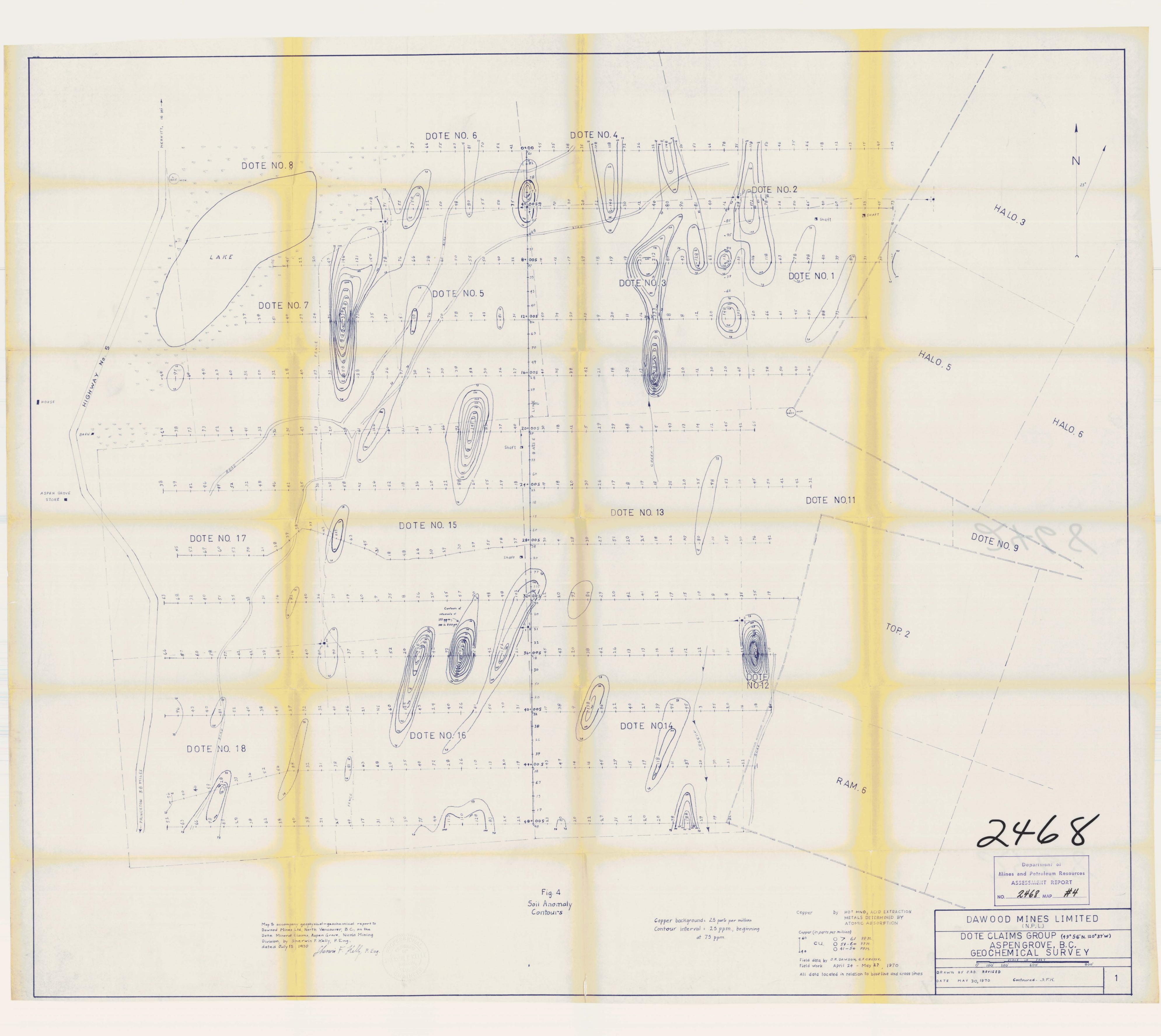
J. B. Dawson, April 24 to May 27, 1970. 1 month @ \$750 per month\$	750.00	
George F. Cressy, Jr., May 4 to May 23, 1970. Twenty days @ \$30 per day	600.00	en e
R. Dawson, May 4 to May 23, 1970.	400.00	
Twenty days \$ \$20 per day\$ Sherwin F. Kelly, P. Eng.	400.00	
Contract price for report	400.00 2,550.00	\$ 2,550.00
Rental of 4x4 Jeep for 20 days	50.00	
Rental of equipment	276.00	\$ 276.00
Cost of soil analyses	706.80	\$ 706,80
Maccillaneous		
Freight on samples	3.00 62.05 70.05	\$ 70.05
Total expenditures		\$ 3,602.85

I hereby certify that the above is a true statement of the expenses incurred for the geophysical and geochemical surveys on part of the DOTS mineral claim group, conducted in April and May of 1970.

July 15, 1970

M. R. Dawson, Secretary Treasurer and Director of Dawson Mines Ltd. (N.P.L.)





DOTE NO.4 DOTE NO. 6 LAKE DOTE NO. 7 HOUSE ASPEN GROVE DOTE NO. 17 TOP 2 DOTE NO.16 DOTE NO. 18 RAM Department of Mines and Petroleum Resources ASSESSMENT REPORT Fig. 3 NO. 2468 MAP #3 Magnetic Contours DAWOOD MINES LIMITED (NPL) Map to accompany geophysical-geochemical report to Dawood Mines Ltd., North Vancouver, B.C., on Contour interval = 100 gammas the Date Mineral Claims, Aspen Grove, DOTE CLAIMS GROUP (49° 56'N 120° 37'W) Nicola Min. Div., by Shervin F. Kelly, P. Eng. dated July 15, 1970. Therwin F. Kelly, P. Eng. Instrument - MF-1 FLUXGATE MAGNETOMETER Field data by: J.R. DAWSON, G.F. CRESSY, Field Work April 28 - May 1970 DRAWN J.R.D. REVISED DATE MAY 27, 1970 Contoured S.F.K.