ASSESSMENT REPORT ON
DIAMOND DRILLING AND GEOLOGICAL MAPPING
on the

SILVER FOX AND MOLLY BLUE
MINERAL CLAIMS

by<br>D.A. Bending

NTS 93M/14W
$127^{\circ} 25^{\prime} \mathrm{W} 55^{\circ} 44^{\prime} \mathrm{N}$
$\$$
situated on Goathead Creek in the Omineca Mining Division
owned by
TEXASGULF CANADA LTD. (now KIDD CREEK MINES LTD)
work by
TEXASGULF INC. \&
KIDD CREEK MINES LTD.
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## INTRODUCTION

Location, Access, Terrain and Climate
The Kisgegas molybdenite prospect is located in the Atna Range near the headwaters of Goathead Creek, 58 kilometres north of Hazelton, British Columbia (Figures 1 and 2).

Direct access to the property is by helicopter. Chartered helicopters are available in Smithers, 125 km to the south. Equipment and supplies can be flown from logged areas near the confluence of the Skeena and Babine Rivers, 15 km west, or from farms near Kispiox, 30 km southwest.

Regional topography is characterized by isolated peaks separated by broad wooded valleys. Peaks above 2000 metres are surrounded by snow and ice fields. Relief on the property is moderate to extreme. The showings occur in a north facing cirque between elevations of 1500 and 1600 metres. Local tree line is about 1300 metres.

The region has a cool temperate climate with moderate rainfall. Much of the property is covered by a small glacier. Snow on and peripheral to this glacier persists until late summer. Accumulation of snow does not begin until October, although sporadic snowfall can occur at higher elevations during most of the summer. History

The history of the property has been reviewed by DeLancey (1979). The property, originally known as the 0le Group, was held by Canex from 1961 to 1963. Initial exploration was focused on low grade Mo-Cu mineralization in rusty hornfels adjacent to a granodiorite stock.

ALIAX personnel observed higher grade Mo mineralization in a quartz vein stockwork in the stock and scheelite in adjacent skarns.
$1)$
Figure 2

| Texasgulf lnc. |  |  |  |
| :---: | :---: | :---: | :---: |
| LOCATION MAP |  |  |  |
| MT. THOMLINSON $\boldsymbol{a}$ KISGEGAS PEAK MO PROSPECTS |  |  |  |
| WORK AY | ORAW | OATE | DANS MO |
| PR.D. | E.R. | 5-10-79 |  |
| -. |  |  |  |

The Fog and Frost Claims were staked to cover these occurrences. During 1964, 1965 and 1966, Amax carried out progranmes of geological mapping, trenching, rock chip sampling and one diamond drillhole 453 metres deep (location shown in figure 4): Although the upper part of the hole was reported to contain molybdenite, the overall results were apparently not encouraging enough to warrant further work, and the property was allowed to lapse.

In 1977 John Bot, an independent prospector from Smithers, staked the Molly Blue Claim (Figure 3). He optioned the property to Texasgulf on May 16, 1979. P.R. DeLancey (see DeLancey, 1979) spent four days on the property in August, 1979. He produced a sketch map of the geology on a scale of 1:5000, reported the style of the molybdenite and scheelite occurrences, and noted the presence of high grade $\mathrm{MoS}_{2}$ in angular float near the edge of the glacier (location shown in Figure 4). He proposed a drill hole to test the apparent source of the mineralized float, and staked the Silver Fox Claim (4 units) to cover parts of the stockwork outside the Molly Blue Claim.

The general retreat of snow and ice between 1966 and 1979 provided better exposure than that previously available. Amax's sample location markings on rock faces three to five metres above the ice indicate a significant retreat of the glacier during these fourteen years.

On the afternoons of September 13 and 14, 1980, the author and assistant examined the showings. A summary of observations made during the brief examination was produced for assessment credit (Bending, 1981).

## Property Status

The principal mineral occurrences are covered by the Molly Blue and Silver Fox claims (Figure 3). The Molly Blue Claim

was transferred to Texasgulf Canada Ltd. according to the terms of the option agreement signed on May 16, 1979 (Bill of Sale, September 27, 1979). The Silver Fox Claim is covered by terms of the option concerning peripheral ground. The Goat 1, 2, 3 and 4 Claims were staked in September 1981 to cover scheelite occurrences north and west of the Molly Blue Claim. The Goat 4 Claim and part of the Goat 1 Claim lie within the area of influence of the option agreement for the Molly Blue Claim, as shown in Figure 3. Fexasgulf Canada tetd. changed its name to Kidd Creek Mines Ltd. effective December 30, 1981.

TABLE 1
Claỉm Administration Data - Kisgegas Property

| Claim | Units | Date Staked | Record Number |
| :---: | :---: | :---: | :---: |
| Molly Blue | 8 | June 16, 1977 | 624 |
| Silver Fox | 4 | August 20, 1977 | 2118 |
| Goat 1 | 20 | September 4, 1981 | 4308 |
| Goat 2 | $10^{2}$ | September 4, 1981 | 4309 |
| Goat 3 | 6 | September 4, 1981 | 4310 |
| Goat 4 | 4 | September 20, 1981 | 4311 |

Summary of Work Completed, 1981
The work done in 1981 consisted of construction of a base camp and drillsites, prospecting, geologic mapping on a scale of 1:2500 and 719.3 metres of $B Q$ diamond drilling in two holes.

The camp was prepared in July by a crew contracted from BEMA Industries. Texasgulf personnel began geological work and drillsite preparation August 22. Longyear personnel were on the property August 29, and drilling commenced August 37. Drilling progressed steadily, without significant delays, and continued until September 22. The drill was dismantled and demobilized September 23. The camp was winterized and all Texasgulf personnel demobilized on September 25.


Mineralized core was split, and all of the core was logged and photographed. Three metre sample intervals were assayed for $\mathrm{MoS}_{2}$. The samples were grouped into fifteen metre composites and analysed for $\mathrm{Cu}, \mathrm{F}$, and Mn. The core is stored in camp.

Work Distribution
Most work was carried out on the Molly Blue and Silver Fox Claims. Two days of reconnaissance prospecting were done on the Goat Claims.

## DRILLING

The purpose of the 1981 drilling programme was to test, at depth, the extension of the exposed mineralization and alteration. Two BQ diamond drill holes, totaliing 712.3 metres, were completed. The locations of these holes are shown in Figure 4, assays and analyses are listed in Appendix 1, summary geological logs are given in Appendix 2, and all 3 metre intervals with assays greater than $0.1 \%$ $\mathrm{MOS}_{2}$ are listed in Table 2.

TABLE 2
Assay Intervals with Greater than $0.1 \% \mathrm{MoS}_{2}$

| Hole | Interval | Length(m) | \%MoS2 |
| :--- | :---: | :---: | :---: |
|  | $51.0-54.0$ | 3.0 | 0.143 |
| K-1-81 | $267.0-270.0$ | 3.0 | 0.125 |
| K-1-81 | $342.0-373.0$ | 33.0 | 0.193 |
| including | $342.0-345.0$ | 3.0 | 0.152 |
|  | $345.0-348.0$ | 3.0 | 0.400 |
|  | $351.0-354.0$ | 3.0 | 0.179 |
|  | $354.0-357.0$ | 3.0 | 0.295 |
|  | $357.0-360.0$ | 3.0 | 0.145 |
|  | $360.0-363.0$ | 3.0 | 0.285 |
|  | $369.0-372.0$ | 3.0 | 0.409 |
|  | $108.0-111.0$ | 3.0 | 0.107 |

K-1-87
DDH K-1-81 was drilled to test the mineralized stockwork, beneath the glacier, for a possible source of the angular high grade molybdenite bearing float found near the west edge of the ice. Figure 4 shows the location of the holes and the float occurrences, and Figure 6 shows the rock types, veining, and alteration encountered. Hole K-1-81 was drilled at Az. $156^{\circ} /-50^{\circ}$, to a depth of 421.3 metres. It penetrated variably altered and mineralized granodiorite, and was stopped in fresh granodiorite with sparse quartz veins. Most of the hole contained scattered molybdenite along quartz veins, and as fine disseminations in narrow zones of potassic alteration. The best intersection averaged $0.193 \% \operatorname{MoS}_{2}$ across 33 metres at a depth of 342 to 375 metres. Within this intersection was a three metre section which assayed $0.409 \% \mathrm{MoS}_{2}$.

K-2-81
DDH K-2-81 (291.0 metres Az.225 $/-51^{\circ}$ ) was drilled (a) to test the interpretation that the angular high grade float. was locally deriyed talus and not transported by the glacier; and (b) to test the stockwork exposed along the West Ridge. The hole cut granodiorite, with local compositional variations from quartz monzonite to quartz diorite, and locally abundant felsic dykes. Figure 7 is a geological section showing the veining,alteration, and rock types encountered. The first 160 metres intersected a well developed quartz vein stockwork, with moderate to intense potassic alteration and widespread traces of $\mathrm{MoS}_{2}$. . The best assay interval was 108.0 - 111.0 metres, with $0.107 \% \mathrm{MoS}_{2}$ across three metres. The quartz veining and alteration became progressively less abundant below 200 metres, and the hole was stopped in fresh, unmineralized granodiorite at 291.0 metres.

## GEOLOGY

## Regional Setting

The Kisgegas property lies within the intermontane structural belt, in the southeast corner of the Bowser Basin. Most of the region is underlain by argillites, siltstones and minor carbonates of the Jurassic to Cretaceous Bowser Lake Group. These are intruded by a northwest trending group of roughly contemporaneous granodiorite and quartz monzonite stocks called the Bulkley Intrusions. The Bulkley intrusions have radiometric apparent ages of 70 to 84 Ma . (Carter, 1976). The Kisgegas molybdenite and scheelite occurrences are related to one such Late Cretaceous granodiorite stock.

## Property Geology

The geology of the property has been discussed by Delancey (1979) and briefly by Bending (1981). The known mineralization is within and peripheral to an elongate east - west trending stock 1500 metres long and 600 metres wide. Molybdenite, chalcopyrite, and pyrite occur in a weakly developed quartz vein stockwork and dis-, seminated in altered areas within the granodiorite. Pyrite, pyrrhotite, and lesser ammounts of scheelite, chalcopyrite and molybdenite, occur in hornfels near the eastern contact. Scheelite occurs in sparse veins and along fractures in fresh argillite, and in garnet - epidote skarn developed in calcareous beds near the granodiorite contacts.

## Sedimentary Rocks

Although the Bowser Lake Group sedimentary rocks exposed in the area have not been mapped in detail, four distinct assemblages were recognized during the 1981 programme. An unknown thickness of argillite and siltstone is overlain by a fifty metre thick section of interbedded argillites and greywackes. This is overlain by an interval characterized by locally calcareous argillites interbedded with one to two metre thick limestone subunits. The
highest peaks are capped by massive chert pebble conglomerates. The only fossils found were pelecypod fragments in the limestones. Igneous Rocks

Granodiorite porphyry
The dominant igneous rock underlying the property is granodiorite, with local compositional variations that range from quartz diorite to quartz monzonite. The rock is generally medium grained, with 2 to 3 cm . zoned phenocrysts of pink potash feldspar, and 3-5 mm quartz subhedra ('quartz-eyes') in a groundmass of plagjoclase, minor orthoclase and biotite; some hornblende may also be present.

Where fresh, this rock is weakiy magnetic. Hornfelsed argillite inclusions occur along intrusive contact zones and are particularly abundant near the west end of the stock. Pyritic, biotite rich schlieren that probably represent assimilated argillite inclusions occur near intrusive contacts and occasionally in the central portion of the stock.

The chronology of the igneous rocks, veins and alteration is summarized in Table 3. The following discussions classify each type of dyke by composition, texture and position in this chronology, from oldest to youngest.

Granodiorite dykes
Some granodiorite dykes intruding argillite near the north contact are clearly contemporaneous with the stock itself, but others crosscut the granodiorite, and are slightly later. These dykes are similar in composition to the granodiorite but are finer grained. with textures varying from a medium grained ground mass with 2-3 cm. $k$-feldspar phenocrysts, to a very fine grained groundmass with phenocrysts smaller than 1 cm .

TABLE 3
GENERAL PARAGENESIS - KISGEGAS PROPERTY

## Intrusive Rocks

(01dest)
Granodiorite Porphyry
Granodiorite Dykes

Brown-Pink Aplite Dykes I
Pale Grey Aplite Dykes

> Pink Pegmatitic Veins(rare)
> Early Potassic Veins $\left(\mathrm{MoS}_{2}\right)$

## Brown-Pink Aplite Dykes II

Pink and Buff Felsic Dykes

Irregular mafic intrusives with abundant inclusions
Intermediate and Mafic Dykes veins
West Ridge Veining and Alteration
Deep Pink Potassic Veins (rare)
Sheeted Veins
Vuggy Quartz, K-Feldspar, Pyrite Veins
Large Vuggy Quartz Veins

Veins
Alteration

Pervasive, light green alteration

Pink K-feldspathization Pink K-feldspathization along selvages $1-2 \mathrm{~cm}$ thick

Thin pale green selvages
Silicification, potassic alteration
Dark pink
K-feldspathization
Weak pink K-feldspathization
Weak pink K-feldspathization None

Argillic and late green alteration

Brown-Pink aplites I and II
Many brown-pink finely crystalline dykes 0.3 to 20 cm . wide (generally 0.5 to 2 cm .) outcrop along the ridge west of the glacier, throughout hole K-2-81, and near the bottom of hole K-1-81. Occasionally these dykes contain 1-3 mm 'quartz-eyes'. These dykes are cut by the early potassic veins. A second generation of brown-pink aplite dykes (II) cut the early potassic veins.

Pale grey a.plites
Pale grey aplite dykes up to one metre thick are exposed along the West Ridge and in hole $\mathrm{K}-2-81$. They vary in texture from finely crystalline to porphyritic, and contain distinct'quartz-eyes' up to 5 mm in diameter.

Pink and buff felsic dykes
Pink (occasionally pale buff) aphanitic to medium grained felsic dykes cut the early potassic veins. These dykes vary from less than 2 cm . wide to irregular masses greater than five meters in width. The larger dykes are medium grained, with $3-4 \mathrm{~mm}$. 'quartz-eyes' in an equigranular groundmass of quartz and feldspar (dominantly potash feldspar).

As outlined in Table 3, several generations of yeins separate the previous? described intrusive events from later intermediate and mafic intrusives.

Irregular mafic intrusives with abundant inclusions
Several small exposures of mafic, fine grained intrusive rock with inclusions of fresh and altered granodiorite and felsite occur within the granodiorite stock. The bodies themselves are irregular in shape. They crosscut the common vein types. The exposures are adjacent to the 1965 Amax drillsite and adjacent to the west edge of the glacier along the West Ridge.

Intermediate and mafic dykes
One porphyritic dyke of intermediate composition is exposed along the West Ridge. This dyke cuts all veins and alteration, and is itself unaltered and weakly magnetic. One mafic porphyritic dyke is exposed along the West Ridge. It cuts all veins and alteration, contains finely disseminated pyrite, and is moderately magnetic. Contact Effects

The contact zone adjacent to the granodiorite generally lacks extensive metamorphism or intense deformation. Contact effects, where present, include drag folds, contact metamorphism, felsite dykes and pods, and small lenses of intense silicification. Contaminated border zones are locally present in the granodiorite.

The bedding of the sedimentary country rocks is generally subhorizontal and not affected by the stock, but along the north contact the bedding of the argillite has been locally deformed in response to the intrusion.

The argillites alter along intrusive contacts to form a hornfels zone that varies in thickness from less than 1 metre to several tens of metres. The hornfels contains disseminations, fracture fillings, and local irregular concentrations of pyrite, pyrrhotite and, in some cases,chalcopyrite, molybdenite, and scheelite. Limestone beds adjacent to intrusive contacts are altered to skarn with garnets, epidote, diopside, pyrite, calcite, and scheelite. No systematic sampling has been performed, but some float found in the cirque has been visually estimated to contain up to $2 \%$ scheelite. The extent of the skarn zones has not been determined. In most cases these contacts are not readily accessible due to extreme topography. Numerous one metre thick skarn beds are exposed in the headwall above the glacier; this is apparently the source of much of the mineralized skarn and hornfels float.

The north contact zone is locally characterized by the presence of numerous angular inclusions of argillite and granodiorite in tan felsite. This zone, which is 2-3 metres thick, has the superficial appearance of a breccia. The felsite crosscuts both argillite and granodiorite, with sharp contacts and narrow but distinct chilled margins.

The 'North Boundary Felsite' consists of a pale tan, sugary textured siliceous rock with diffuse feldspar relics and occasional traces of pyrite and molybdenite. The superficial appearance of this rock is of a felsic intrusive, but it is probably an alteration zone in the granodiorite. Gradational contacts separating fresh granodiorite porphyry from 'felsite' can occur within a single outcrop.

Veins, Alteration and Mineralization
Molybdenite and lesser scheelite mineralization occur in quartz veins cutting the granodiorite pyritic hornfels peripheral to the stock, and in garnet pyroxene skarn within calcareous beds adjacent to the intrusive contact. The principal focus of the 1981 programme was the quartz vein stockwork within the granodiorite.

The mineralogy and chronology of the veins and alteration types are summarized in Table 3. Eight types of veins, each categorized on the basis of structure, mineralogy, alteration, texture and paragenetic position, occur in the stock. Some of these vein types have uncertain temporal relationships. Several types of veins that show intense pink potassic alteration may be contemporaneous or otherwise closely related, but are herein separated for discussion. Pre-intrusive white quartz veins, present in the argillite country rocks, are not included in Table 3. Significant amounts of molybdenite occur in veins of two ages: the early potassic veins and grey quartz veins. These veins are separated in time by the brown-pink aplite
and pale grey felsite dykes. Traces of molybdenite occur in two younger vein types in some localities. These are called 'deep pink potassic veins' and the 'West Ridge veining and alteration'. Some vuggy quartz veins with K-feldspar contain minor pyrite.

Peryasive, light green alteration
Much of the granodiorite has undergone weak, pervasive alteration which produces a diffuse green colour. The green (sericitic?) alteration is the first hydrothermal event; it is crosscut by all veins and felsic dykes. This alteration destroys the weak magnetism present in the granodiorite. Traces of pyrite in biotite flakes, and finely disseminated chalcopyrite, are present where this green alteration is strongest. The north, west, and east fringes of the stock are less altered than the central area.

- Pink pegmatitic veins

Two pink pegmatitic veins are exposed in the cirque near the north contact. These 2-3 thick veins are composed of pink potash feldspar, minor quartz, traces of biotite and rare molybdenite. The margins of these veins display $1-2 \mathrm{~cm}$ selvages of pink potassic alteration.

Early potassic veins
The early potassic veins generally contain quartz, pyrite, K-feldspar, and minor molybdenite. They are characterized by pale pink 0.5-1 cm thick selvages of potash feldspathization, frequently accompanied by finely disseminated molybdenite and pyrite. The veins vary in thickness from less than one millimetre to a maximum of 3 centimetres. Fluorite, gypsum, stibnite and sphalerite are present in some early potassic veins in hole k-1-81.

In some short intervals of drill holes K-1-81 and K-2-81 the veins are so closely spaced that the potassic alteration appears pervasive. More commonly, the potassic selvages are separated by fresh granodiorite, and the intensely altered bands represent only five to ten percent of the rock.

In DDH K-2-87, the alteration mineralogy of the early potassic veins varies with depth. Near the collar the selvages are the characteristic pale pink $K$-feldspar zones. At a depth of about 100 metres the outer margins of the selvages are lined with 2-3 mm bands of pale green sericitic alteration. The relative proportions of these two types of alteration vary with depth so that at 250 metres the sericitic alteration is predominant.

The mineralogy of the potassic selvages varies with host lithology. This is best demonstrated by following an individual vein through different rock types. Most granodiorite alters to form pale pink selvages, while the more mafic intrusive phases alter to a green colour. The argillite country rocks alter to form prominent pale grey-green sericitic selvages l-2 cm thick.

Near the toe of the glacier, 300 metres northeast of drillsite K-1-81, :are large angular blocks of granodiorite float cut by potassic veins bearing scheelite and powellite. No scheelite was observed in the core, and tungsten analyses of the core were uniformly low ( 2 to 3 ppm ). This float may be evidence that a tungsten zone exists in peripheral parts of the early potassic vein system.

Grey guartz $\left(\mathrm{MoS}_{2}, \mathrm{CuFeS}_{2}\right)$ veins
The most prominent concentrations of molybdenite, chalcopyrite and pyrite, occur in grey quartz veins. These veins vary from 2 mm to 25 cm in width but are generally less than 3 cm wide. Molybdenite occurs as smears along the vein margins and between thin quartz bands within some larger veins. The margins of these veins are characterized by narrow $2-5 \mathrm{~mm}$ selvages of silicification and weak green alteration. The relative lack of alteration adjacent to the veins contrasts with the early potassic veins. The grey quartz veins are most abundant in the outcrops in the vicinity of drill site K-1-81.

West Ridge veining and alteration
Figure 4 shows a cluster of felsic dykes in an area of intensive quartz veining, and silicic and potassic alteration cropping out along West Ridge. The silicic and potassic alteration is associated with traces of molybdenite and chalcopyrite. In some exposures silicification and potash feldspathization are so intense that very little remains of the primary igneous fabric of the granodiorite. This alteration probably postdates the early potassic veins, but because the grey quartz (MoS2) veins have not been observed in this area the paragenesis is not established.

Deep pink potassic veins
Holes $\mathrm{k}-1-81$ and $\mathrm{K}-2-81$ intersected some quartz- $k$-feldspar veins, characterized by $k$-feldspar, with a deep brown-pink colour and 1-2 cm wide potassic selvages; these veins postdate the grey molybdenite bearing veins. Some of these deep pink potassic veins contain traces of molybdenite. Although the relative chronology of the Hest Ridge veining and alteration is not well established, textural and mineralogical similarity suggests that it is related to these deep pink potassic veins.

Sheeted veins
Much of the west end of the stock is penetrated by a swarm of quartz veins (and subordinate K-feldspar) with weak potassic selvages and traces of pyrite. These veins are 0.5 to 2.0 cm thick, $10-20 \mathrm{~cm}$ apart, and occur uniformly spaced across exposures more than fifty metres wide. They postdate the early potassic veins.

The sheeted veins exposed in the West Ridge do not contain molybdenite. Some large angular blocks of talus that are possibly derived from the north flank of the zone, about 100 metres below the outcrops examined, contain smears of molybdenite along veins that are similar in habit to the sheeted vein system.

Vuggy quartz, K-feldspar, pyrite veins
Veins containing vuggy quartz, $k$-feldspar and pyrite are scattered throughout the granodiorite and the hornfels. They vary from 1 cm. to 10 cm . wide but are generally less than 5 cm . Many show weak pink potassic alteration.

Large vuggy quartz veins
Vuggy white quartz veins 10 cm .to 50 cm .wide occur in the granodiorite and in the hornfelsed argillite. These veins have no distinct alteration. They are notably continuous; an individual vein near the west contact can be traced for more than 200 metres. All of these vuggy quartz veins strike about $045^{\circ}$ and dip $60^{\circ}-85^{\circ}$ northwest.

Argillic alteration
Shear zones show intense argillic alteration of feldspars, producing a soft, friable light grey or pale grey-green aggregate. Thase intensely argillized shears are bounded by-2-3 metre wide zones of very intense deep green alteration characterized by destruction of biotite, removal of quartz, conversion of plagioclase to a green intergrowth of clay and epidote (or chlorite?) and incipient argillization of K -feldspar megacrysts.

This type of altered granodiorite weathers easily and is not generally observed in outcrop. Some is exposed along the West Ridge above drill site $\mathrm{K}-2-81$, and in the fault zone in the ciiffs east of the glacier. Short intervals of sheared and argillized rocks were intersected by holes K-1-81 and K-2-81.

Structure
The structural geology of the Kisgegas property can be considered in terms of the geometry of the various igneous phases and the orientations of the veins, fractures and faults.

## Intrusive rocks

The granodiorite stock is elongate along an east-west axis. Several dykes of granodiorite that are contemporaneous with the stock dip steeply to the north and strike between $090^{\circ}$ and $095^{\circ}$.

Figure 3-1 is a stereographic plot of the poles to all felsic dyke orientations measured. It shows a random orientation for the felsic dykes throughout the property, except for the West Ridge, where the dykes show a northeast preferred trend (see the geological map, Figure 4).

Early potassic yeins
Figure 3-2 is a stereographic plot of poles to orientations of the early potassic veins. It shows that they have a generally random orientation. In the central part of the stock, near drillsite K-1-81, these veins have a very strong preferred orientation; they strike northeast and dip $55^{\circ}-75^{\circ}$ northwest. The angles of intersection between the potassic veins and the core axis in DDH K-1-81 show a gradual decrease from $85^{\circ}$ near the collar to $35^{\circ}$ at depth. The orientations of the potassic veins in hole K-2-81 are less systematic; they form a stockwork of veins that intersect the core axis at $20^{\circ}-80^{\circ}$.

Grey quartz $\left(\mathrm{MoS}_{2}, \mathrm{CuFeS}_{2}\right)$ veins
The grey quartz ( $\mathrm{MoS}_{2}, \mathrm{CuFeS}_{2}$ ) veins display a strong NE-SW trend (Figure 3-3). The dips of these veins vary from steeply southeast to steeply northwest, including some that are subhorizontal or have shallow northwesterly dips. The shallow dips are localized in the central part of the intrusive, near drillsite K-1-87. Figure 5 shows the orientation of the grey veins in this hole. In the top 100 metres, most of the veins dip about $30^{\circ}$ to the $N W$ and intersect the core axis at about $80^{\circ}$. They are oblique to the most prominent set
of potassic veins. The dip of the grey veins becomes progressively steeper with depth, until at 350 metres they intersect the core axis at about $40^{\circ}$. The grey quartz $\left(\mathrm{MoS}_{2}\right)$ veins intersected by DDH K-2-81 are subparallel to the core axis. The apparent general pattern of the grey veins across this mineralized, central part of the stock, is nearly horizontal over the axis of the zone (in the central part of the stock) and steeply dipping along the fringes.

Other veins
The poles to measured orientations of the sheeted vein. system exposed in the West Ridge are plotted in Figure 3-4. The poles are tightly grouped and show a very strong preferred orientation, trending east-northeast and dipping $70^{\circ}-80^{\circ}$ northwesterly. This is a much stronger grouping than that in any previous vein set.

Figure $3-5$ is a composite plot of poles to orientations of late stage vein types. The large vuggy quartz veins show a strong preferred orientation parallel to the sheeted vein system. The other vein types have a random orientation.

> Faults

The stock is cut by numerous local faults. The only significant fault apparent in outcrop is exposed near the east end of the property. This is a reverse fault, oriented $052^{\circ} / 40^{\circ} \mathrm{NW}$, with at least 100 metres of vertical displacement. The trace of this fault is marked by a slickensided zone with $20-30 \mathrm{~cm}$ of rusty clay-rich gouge. This fault displaces and postdates all the types of veins exposed at the east end of the cirque.

K-1-81 intersected numerous shear zones showing argillic alteration. These zones intersect the core axis at about $30^{\circ}$ and appear to be parallel to each other. Comparison of shear and vein orientations in core to adjacent outcrops indicates that these shears probably trend northeast and dip about $80^{\circ}$ northwest. DDH K-2-81 has a higher proportion of altered shear zones than K-1-81. Most of the shears in $\mathrm{K}-2-81$ intersect the core axis at about $15^{\circ}$, but some are up to $45^{\circ}$ to the core axis.

## GEOCHEMISTRY

Fifteen metre composite samples from the drill holes were analysed by geochemical methods for $\mathrm{Sn}, \mathrm{W}, \mathrm{F}, \mathrm{Mn}$ and Cu . These elements were selected in an attempt to identify primary dispersion patterns related to the molybdenite mineralization. Sn and W analyses were discontinued when the first shipment of samples contained no detectable tin and a uniformly low (2-3ppm) tungsten content.

The $\mathrm{Cu}, \mathrm{Mn}$, and F values are plotted along with corresponding $\mathrm{MoS}_{2}$ assays in Figures 7 and 8. The data show no obvious pattern, but in the context of the geology of the holes some conclusions are possible. Each vein and alteration type has a distinct geochemical signature, and rigorous examination of primary dispersion patterns would require sampling of individual vein and alteration systems. Hole $\mathrm{K}-1-81$ has generally lower Cu values than Hole K-2-81, and a much higher Mo/Cu ratio. Cu-values in hole K-1-81 vary from 670 ppm to 188 ppm , with a mean Cu value of 350.8 ppm and a mean $\mathrm{MoS}_{2} / \mathrm{Cu}$ ratio. of 1.27. Cu values in hole K-2-81 range from a maximum of 850 ppm , near the collar, to 80 ppm at the bottom, with a mean of 373.6 ppm and a mean $\mathrm{MoS}_{2} / \mathrm{Cu}$ ratio of 0.67 . Hole k-2-81 penetrated the south margin of the stockwork, whereas DDH K-1-81 was closer to the centre of the system. The difference between these holes may indicate the presence of a Cu halo peripheral to the mineralized zone.

DISCUSSION
The Kisgegas Mo(W) property has extensive, locally intense potassic alteration, widespread quartz veining, and locally attractive grades of molybdenite mineralization in vein systems of two distinct ages. However, most of the quartz veins do not contain molybdenite and the most attractive mineralization is not related to the prominent potassic alteration.

The early potassic vein and alteration system contains molybdenite in veins and as disseminations in thin selvages. Because these veins are generally widely spaced, they make a minor contribution to the overall molybdenite grades, but intervals in which they are well developed and closely spaced (so about $50 \%$ of the interval has been subjected to potassic alteration) grade up to $0.15 \% \mathrm{MoS}_{2}$. The molybdenite content, the width of the selvages, and the width of these veins appear to be greatest in the central part of the stock, in the area of drill site K-1-81.

The grey quartz-MoS2 veins contain the richest molybdenite concentrations in the property. The best mineralization ( 0.2 to $0.407 \% \mathrm{MoS}_{2}$ across 3 metre intervals) intersected by hole K-1-81 occurs in grey veins, with a small proportion in early potassic veins. The grey vein system is exposed from the kest Ridge to the east side of the glacier, but in most exposures these veins are very narrow and widely scattered. The best exposures of this vein system are in the vicinity of $\mathrm{K}-1-81$, where the veins are more abundant and contain proportionately more molybdenite than elsewhere in the property. The fact that the best thirty metres in hole K-1-81 was richer and more continuous than any exposed mineralization serves to focus attention on the potential of this system at depth and beneath the glacier.

The area of intense veining and quartz-k-spar alteration exposed along the West Ridge contains only traces of molybdenite. The intensity of the alteration and the presence of traces of molybdenite suggest a possible relationship to more attractive mineralization. The 1965 AMAX drill hole was apparently planned to test this system; it was directed below the centre of the area of intense veining and alteration. The upper part was reported to contain very
low grade mineralization similar to that in the upper parts of K-1-87. The lower part, which penetrated the rock below the exposed quartz-K-spar alteration, contained only traces of molybdenum.

The outcrops of the sheeted vein system do not contain molybdenite. It is possible that this well developed vein system is the apical expression of a mineralized zone at depth, however, the available evidence in support of this idea is not strong enough to warrant drilling at this time.

The next steps in evaluating the property should be further drilling to test the mineralized zone intersected in hole K-1-81 at greater depth, and along strike to the east. Proposed hole K-3-81 is an attempt to penetrate the K-1-81 mineralization at a greater depth to further evaluate the grade and geometry of the system. Proposed hole K-4-82 is situated to test the eastern strike extension of the same stockwork.

The tungsten potential of the skarn and hornfels was not evaluated in 1981. Poor weather and limited manpower prevented systematic mapping and sampling of the cliffs where the skarns are exposed. As noted by DeLancey. (1980), some float in the cirque contains attractive quantities of scheelite. The beds that are the most probable sources of this mineralization are exposed along the cliffs east of the glacier and in the headwall of the cirque. They are generally less than one metre thick and not uniformly mineralized. More work will be necessary to evaluate the tungsten potential of this contact zone.

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$\bigcirc$

## APPENDIX 1

HOLE NO.: _K-7-81 PAGE 1 of 5
 LONGITUDE: _ 598,830E* DIP: $\quad-53 \quad$ INCLINATION: $156^{\circ} /-55$ at 415.1 ELEVATION: 1764*_____ at

| SAMPLE | METRES |  | $\mathrm{MoS}_{2}$ | Sn | F |  | W | Mn | Cu |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| No. | FROM | T0 | ASSAYS | ppm | ppm |  | ppm | ṕpm | ppm |
| 18226 | 1.2 | 3.0 | 0.020 |  |  |  |  |  |  |
| 18227 | 3.0 | 6.0 | 0.017 |  |  |  |  |  |  |
| 18228 | 6.0 | 9.0 | 0.030 | ND | 500 |  |  | 210 | 268 |
| 18229 | 9.0 | 12.0 | 0.028 |  |  |  |  |  |  |
| 18230 | 12.0 | 12.0 | 0.019 |  |  |  |  |  |  |
| 18076 | 15.0 | 18.0 | 0.043 |  |  |  |  |  |  |
| 18077 | 18.0 | 21.0 | 0.035 |  |  |  |  |  |  |
| 18078 | 21.0 | 24.0 | 0.055 | ND | 470 |  | 2 | 240 | 220 |
| 18079 | 24.0 | 27.0 | 0.027 |  |  |  |  |  |  |
| 18080 | 27.0 | 30.0 | 0.030 |  |  |  |  |  |  |
| 18081 | 30.0 | 33.0 | 0.020 |  |  |  |  |  |  |
| 18082 | 33.0 | 36.0 | 0.042 |  |  |  |  |  |  |
| 18083 | 36.0 | 39.0 | 0.033 | ND | 410 |  | 2 | 190 | 280 |
| 18084 | 39.0 | 42.0 | 0.057 |  |  |  |  |  |  |
| 18085 | 42.0 | 45.0 | 0.062 |  |  |  |  |  |  |
| 18086 | 45.0 | 48.0 | 0.030 |  |  |  |  |  |  |
| 18087 | 48.0 | 57.0 | 0.025 |  |  |  |  |  |  |
| 18088 | 57.0 | 54.0 | 0.143 | ND | 550 |  | 3 | 210 | 405 |
| 18089 | 54.0 | 57.0 | 0.040 |  |  |  |  |  |  |
| 18090 | 57.0 | 60.0 | 0.063 |  |  |  |  |  |  |
| 18091 | 60.0 | 63.0 | 0.050 |  |  |  |  |  |  |
| 18092 | 63.0 | 66.0 | 0.065 |  |  |  |  |  |  |
| 18093 | 66.0 | 69.0 | 0.032 | ND | 520 |  | 3 | 255 | 360 |
| 18094 | 69.0 | 72.0 | 0.063 |  |  |  |  |  |  |
| 18095 | 72.0 | 75.0 | 0.050 |  |  |  |  |  |  |
| 18096 | 75.0 | 78.0 | 0.037 |  |  |  |  |  |  |
| 18097 | 78.0 | 81.0 | 0.047 |  |  |  |  |  |  |
| 18098 | 87.0 | 84.0 | 0.068 | ND | 550 |  | 2. | 220 | 358 |
| 18099 | 84.0 | 87.0 | 0.037 |  |  |  |  |  |  |
| 18100 | 87.0 | 90.0 | 0.067 |  |  |  |  |  |  |
| 18101 | 90.0 | 93.0 | 0.022 |  |  |  |  |  |  |
| 18102 | 93.0 | 96.0 | 0.014 |  |  |  |  |  |  |
| 18103 | 96.0 | 99.0 | 0.077 | ND | 550 |  | 3 | 235 | 298 |
| 18104 | 99.0 | 102.0 | 0.013 |  |  |  |  |  |  |
| 18105 | 102.0 | 105.0 | 0.073 |  |  |  |  |  |  |

HOLE No.: _ K-1-81 PAGE 2 of 5 LATITUDE: ___ AZIMUTH: INCLINATION: $\qquad$ at LONGITUDE:___ DIP:

INCL INATION: $\qquad$ at
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| SAMPLE <br> №. | METRES |  | $\mathrm{MoS}_{2} \quad 1 \mathrm{Sn}$ |  |  | F |  | 1 W | Mn | CH |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | EROM | T0 | ASSAYS |  | ppin | ppm |  | ppm. | ppm | ppm |
| 18106 | 105.0 | 108.0 | 0.013 |  |  |  |  |  |  |  |
| 18107 | 108.0 | 117.0 | 0.012 |  |  |  |  |  |  |  |
| 18108 | 111.0 | 114.0 | 0.033 |  | ND | 500 |  | 3 | 220 | 670 |
| 18109 | 114.0 | 117.0 | 0.020 |  |  |  |  |  |  |  |
| 18110 | 117.0 | 120.0 | 0.017 |  |  |  |  |  |  |  |
| 18111 | 120.0 | 123.0 | 0.023 |  |  |  |  |  |  |  |
| 18172 | 123.0 | 126.0 | 0.024 |  |  |  |  |  |  |  |
| 18113 | 126.0 | 129.0 | 0.022 |  | ND | 550 |  | 3 | 220 | 515 |
| 18114 | 129.0 | 132.0 | 0.080 |  |  |  |  |  |  |  |
| 18115 | 132.0 | 135.0 | 0.033 |  |  |  |  |  |  |  |
| 18176 | 135.0 | 138.0 | 0.043 |  |  |  |  |  |  |  |
| 18177 | 138.0 | 141.0 | 0.050 |  |  |  |  |  |  |  |
| 18118 | 141.0 | 144.0 | 0.025 |  | ND | 550 |  | 3 | 225 | 515 |
| 18179 | 144.0 | 147.0 | 0.045 |  |  |  |  |  |  |  |
| 18120 | 147.0 | 150.0 | 0.058 |  |  |  |  |  |  |  |
| 18121 | 150.0 | 153.0 | 0.090 |  |  |  |  |  |  |  |
| 18122 | 153.0 | 156.0 | 0.044 |  |  |  |  |  |  |  |
| 18123 | 156.0 | 159.0 | 0.025 |  | ND | 500 |  | 3 | 300 | 341 |
| 18124 | 159.0 | 162.0 | 0.018 |  |  |  |  |  |  |  |
| 18125 | 162.0 | 165.0 | 0.003 |  |  |  |  |  |  |  |
| 18126 | 165.0 | 168.0 | 0.040 |  |  |  |  |  |  |  |
| 18127 | 168.0 | 171.0 | 0.043 |  |  |  |  |  |  |  |
| 18128 | 171.0 | 174.0 | 0.042 |  | ND | 710 |  | 2 | 290 | 302 |
| 18129 | 174.0 | 177.0 | 0.075 |  |  |  |  |  |  |  |
| 18130 | 177.0 | 180.0 | 0.027 |  |  |  |  |  |  |  |
| 18137 | 180.0 | 183.0 | 0.027 |  |  |  |  |  |  |  |
| 18132 | 183.0 | 186.0 | 0.007 |  |  |  |  |  |  |  |
| 18133 | 186.0 | 189.0 | 0.073 |  | ND | 580 |  | 2 | 275 | 238 |
| 18134 | 189.0 | 192.0 | 0.018 |  |  |  |  |  |  |  |
| 18135 | 192.0 | 195.0 | 0.013 |  |  |  |  |  |  |  |
| 18136 | 195.0 | 198.0 | 0.023 |  |  |  |  |  |  |  |
| 18137 | 198.0 | 207.0 | 0.027 |  |  |  |  |  |  |  |
| 18138 | 207.0 | 204.0 | 0.017 |  | ND | 710 |  | 2 | 270 | 235 |
| 18739 | 204.0 | 207.0 | 0.018 |  |  |  |  |  |  |  |
| 18140 | 207.0 | 210.0 | 0.017 |  |  |  |  |  |  |  |

PROPERTY: Kisgegas
HOLE No.: _K-1-81 PAGE 3 of 5

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PROPERTY: __Kisgegas _
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HOLE NO.: K-1-81 PAGE 5 of 5

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| SAMPLE No. | METRES |  | MoS? | Sn | $F$ | 1.6 | $\mathrm{Mn} \quad \mathrm{Cu}$ |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | FROM | TO | ASSAYS | ppm | R0n | ppm | ppm. | ppm |
| 18211 | 420,0 | 421.2 | 0.007 | ND | 410 |  | 160 | 188 |
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HOLE NO.: __K-2-8] PAGE $\quad$ __ of 3
LATITUDE: 6,179,750N* AZIMUTH: $226^{\circ}$ LONGITUDE: $\quad$ 598,735E* DIP: __ $-51^{\circ}$ INCLINATION: $224,-51$ at 137.2 m INCLINATION: $229 / 51^{\circ}$ at 288.1 m
 $\qquad$ /__at $\qquad$


PROPERTY: _Kisgegas
HOLE No.: __K-2-8I PAGE 2 of 3 LATITLUDE: $\quad$ AZIMUTH:
LONGITUDE: $\qquad$ DIP:
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PROPERTY: Kisgegas
HOLE No.: K-2-87
PAGE 3 of 3

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| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | SAMPLE |  |  | $\mathrm{MoS}_{2}$ | Sn | F | W | Mn | Cu |
|  | No. | FROid | T0 | ASSAYS | ppin | .ppm | ppin | ppm | npm. |
|  | 1937 | 210.0 | 213.0 | 0.002 |  |  |  |  |  |
|  | 1938 | 213.0 | 216.0 | 0.002 | ND | 320 |  | 230 | 114 |
|  | 1939 | 216.0 | 219.0 | 0.002 |  |  |  |  |  |
|  | 1940 | 219.0 | 222.0 | 0.003 |  |  |  |  |  |
|  | 1941 | 222.0 | 225.0 | 0.003 |  |  |  |  |  |
|  | 1942 | 225.0 | 228.0 | 0.007 |  |  |  |  |  |
|  | 1943 | 228.0 | 231.0 | 0.003 | ND | 300 |  | 200 | 149 |
|  | 1944 | 231.0 | 234.0 | 0.007 |  |  |  |  |  |
|  | 1945 | 234.0 | 237.0 | 0.012 |  |  |  |  |  |
|  | 1946 | 240.0 | 243.0 | 0.003 |  |  |  |  |  |
|  | 1947 | 240.0 | 243.0 | 0.003 |  |  |  |  |  |
|  | 1948 | 243.0 | 246.0 | 0.003 | ND | 290 |  | 270 | 80 |
|  | 1949 | 246.0 | 249.0 | 0.008 |  |  |  |  |  |
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## APPENDIX 2

SUMMARY GEOLOGICAL LOGS

DDH K-1-81
DDH K-2-81



| TE | SGU | 1 NC | DRILL HOLE LOG ${ }^{\text {HOLE NO. }}$ N-T-81 ${ }^{\text {K }}$ |
| :---: | :---: | :---: | :---: |
| DEPTH ${ }^{(m)}$ |  | REC'Y | DESCRIPTION |
| FROM | TO |  |  |
| 198.0 | 213.2 | 98-100\% | Variably altered granodiorite porphyry with_locally pervasive pink potassic |
|  |  |  | alteration along early potassic yeins $65^{\circ}$ to $80^{\circ}$ to the core axis and sparse |
|  |  |  | grey quartz, MoS 2 veins $70^{\circ}$ to $80^{\circ}$ to core axis. |
| 213.2 | 220.1 | 100\% | Granodiorite dyke; similar in texture to the granodiorite but darker in colour |
|  |  |  | (with more abundant biotite). Chilled margins 0.8 metre thick are dark orey, unaltered, |
|  |  |  | and weakly magnetic. The inner phase of the dyke is weakly propyllitized and not |
|  |  |  | magnetic. Note that the early potassic veins cut the dyke but seem to be less |
|  |  |  | abundant than in the enclosing granodiorite. |
| 220.1 | 227.3 | 99\% | Variably altered granodiorite with pink felsite and grey-green intermediate |
|  |  |  | dykes. The granodiorite is characterized by weak to moderate pervasive propyllitic |
|  |  |  | alteration, $\mathrm{MoS}_{2}$ bearing early potassic veins, occasioned MoSp-bearing grey quartz |
|  |  |  | veinlets, later 1 to 3 cm . thick quartz k-feldspar pyrite veins, and vugay white quartz |
|  |  |  | veins. The pink dykes cut the early potassic veins but show no distinct crosscutting |
|  |  |  | relationships to the late vuggy quartz veins. |
| 227.3 | 224.0 |  | Variably altered granodiorite porphyry; pink potassic yeins and alteration are |
|  |  |  | abundant and represent about $30 \%$ of the rock. These early potassic veins cut the core |
|  |  |  | axis atangles of $35^{\circ}$ to $80^{\circ}$ in a stockwork. Later MoS ${ }_{2}$ bearing grey quartz yeins |
|  |  |  | cut the core at about $70^{\circ}$. These later veins are sparse, with an average of two per |
|  |  |  | metre. |
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| TEXASGULF INC. |  |  | DRILL HOLE LOG $\left.\quad$HOLE NO. <br> $\mathrm{K}-2-81$ \right\rvert\, $\begin{gathered}\text { PAGE } \\ \end{gathered}$ |
| :---: | :---: | :---: | :---: |
| DEPTH(m) |  | REC'Y | DESCRIPTION |
| FROM | T0 |  |  |
|  |  |  | potassic veins show zoned selvages, with 0.5 cm . green sericitic rims outside the |
|  |  |  | more typical pale pink selvages. These early potassic veins occur $30^{\circ}$ to subparallel |
|  |  |  | to the core axis. A dark grey, biotite - rich, partly assimilated inclusion occurs at |
|  |  |  | 116.4 m. . The interval $125-128$ contains numerous 1 cm . thick brown - pink dykes |
|  |  |  | subparallel to the core axis, that predate the early potassic veins. A 25 cm . thick |
|  |  |  | flesh brown quartz - eye felsite dyke at 128.3, cuts the core axis at $50^{\circ}$. |
|  |  |  | This dyke cuts the early potassic veins. |
| 133.5 | 142.5 | 100\% | Weakly to intensely altered granodiorite as above. Some later, deep pink potassic |
|  |  |  | alteration occurs along quartz, K-feldspar, pyrite veins that cut the early potassic |
|  |  |  | veins. The interval 134.8 to 135.5 is characterized by shearing and intense argillic |
|  |  |  | alteration that postdates the veins. |
| 142.5 | 757.5 | 97\% | Sheared, intensely argillized and propyllitized granodiorite with relics of early |
|  |  |  | potassic veins and white quartz veinlets. The interval 145.7-148.1 contains numerous |
|  |  |  | white quartz veins and deep brown - pink potassic atteration, overprinted by shearing |
|  |  |  | and green clay - rich alteration. The slickensided shears cut the core at $15^{\circ}-30^{\circ}$. |
| 157.5 | 161.0 | 100\% | Weakly to intensely altered granodiorite porphyry with some $1-2 \mathrm{~cm}$. pale tan - buff |
|  |  |  | felsite dykes $20^{\circ}$ to subparallel to the core axis, and some short dark biotite - rich |
|  |  |  | intervals that may represent assimilated country rocks. The interval is nearly |
|  |  |  | fresh, with a very weak green colour, and is cut by a sparse stockwork of early potassic |
|  |  |  | veins (most cut the core at $20^{\circ}-30^{\circ}$ ) and two vuggy white quartz veins. |
| 161.0 | 163.5 | 100\% | Pale tan aplite dyke with pinheads of $\mathrm{MOS}_{2}$ and $\mathrm{FeS}_{2}$. Locally altered (argillic, |
|  |  |  | green) shears. |
| 163.5 | 177.0 | 100\% | Variably altered granodiorite porphyry. Most is very |
|  |  |  | weakly altered to a diffuse green colour with traces of $\mathrm{FeS}_{2}$ and $\mathrm{CuFeS}_{2}$ in corroded |
|  |  |  | biotite flakes. A sparse set of early potassic veins with traces of $\mathrm{MoS}_{2}$ |


| TEX | ASGU | INC | DRILL HOLE LOG $\quad$ HOLE $\mathrm{NO}_{\mathrm{K}-2-81} \mathrm{PAGE}_{4} \mathrm{NO}$ |
| :---: | :---: | :---: | :---: |
| DEPTH |  | REC'Y | DESCRIPTION |
| FROM | T0 |  |  |
|  |  |  | cuts the core at about $35^{\circ}$. Note 21 cm . wide grey quartz, $\mathrm{MoS}_{2}$ veins, cutting the core |
|  |  |  | at $15^{\circ}$, in the interval 172.0-173.0. These are accompanied by 1 cm . thick green |
|  |  |  | selvages. |
| 177.0 | 216.0 | 100\% | Weakly to moderately altered granodiorite to quartz monzonite porphyry with |
|  |  |  | pervasive, very weak green alteration, abundant tan - pink aplite dykes and a weak |
|  |  |  | stockwork of early potassic veins with traces of MoS ${ }_{2}$. The selvages of the early |
|  |  |  | potassic veins in and below this interval are generally pink along, the veins and grade |
|  |  |  | into greenish sericitic alteration, 0.5 cm . wide, away from the veins. Less than $5 \%$ of |
|  |  |  | the interval has been subjected to intense alteration. Note some quartz, pyrite, K-feld- |
|  |  |  | spar and quartz, sphaterite, pyrite, K-feldspar veinlets, cutting the core at $20^{\circ}-80^{\circ}$, |
|  |  |  | with green - pink potassic selvages, in the interval 197.0-198.0. The dykes that |
|  |  |  | characterize much of the interval are 0.5 to 15 cm . thick, with textures that vary from |
|  |  |  | sucrosic to medium grained with quartz eyes. These dykes predate the veins. |
| 216.0 | 220.0 | 100\% | Variably altered granodiorite porphyry as above, without dykes. This interval |
|  |  |  | is characterized by 0.2-0.3 metre sections with diffuse concentrations of biotite |
|  |  |  | in the granodiorite. |
| 220.0 | 235.3 | 100\% | Variably altered granodiorite porphyry; generally yery weakly altered, with a faint |
|  |  |  | green tint and traces of pyrite. Sparse early potassic yeins that cut the core at $35^{\circ}$ - |
|  |  |  | $40^{\circ}$ contain traces of $\mathrm{MoS}_{2}$. Several 1-2 cm. thick quartz, pyrite veins with weak |
|  |  |  | potassic selvages cut the core at $40^{\circ}$. The interval $225 .-227$ is characterized |
|  |  |  | by pervasive weak green sericitic alteration. The interyal $234-235$ shows locally |
|  |  |  | intense, orange - pink potassic alteration that postdates the other vein types. |
| 235.3 | 236.7 | -99\% | . Intensely altered_granodiorite__ Pervasive light_pink potassic_alteration_along_ |
|  |  |  | closely spaced early potassir veins $15^{\circ}-35^{\circ}$ to the core axis, with finely disseminated |
|  |  |  | $\mathrm{MoS}_{2}$. |



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| :---: | :---: | :---: | :---: |
| TEX | SGUL | INC. |  |
|  |  | REC'Y | DESCRIPTION |
| FROM | T0 |  |  |
|  |  |  | propyllitic alteration. The dykes predate all veins and alteration. <br> Quartz diorite with gradational contacts with the granodiorite. Generally fresh and |
| 277.7 | 285.1 | 100\% |  |
|  |  |  | moderately magnetic, richer in biotite than the granodiorite and lacking $k$-feldspar |
|  |  |  | phenocrysts. This is cut by a tan - pink aplite dyke in the interval 279.1-280.0, |
|  |  |  | and numerous 1-2 cm. brown - pink dykes, at $30^{\circ}-40^{\circ}$ to the core axis. This interval |
|  |  |  | is cut by many $0.5-1 \mathrm{~cm}$. quartz, carbonate, pyri.te veins, $45^{\circ}$ to $65^{\circ}$ to the core axis, |
|  |  |  | with prominent green sericitic selvages. Individual veins of this type, where traced into |
|  |  |  | the felsic dykes, change markedly in mineralogy: in the felsite they are quartz, $k$-feldspar |
|  |  |  | veins with pink feldspathic selvages. |
| 285.1 | 291.0 | 100\% | Granodiorite to quartz monzonite; generally fresh, weakly magnetic. Cut by medium to |
|  |  |  | fine grained brown - pink felsic dykes at $50^{\circ}$ to the core axis, and quartz, k-feldspar |
|  |  |  | veins with deep pink potassic selvages. |
|  |  |  | 291.0 metres: End of Hole $\quad$, |
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## APPENDIX 3

Stereographic Plots of Structural Data

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Fig. 3-1 Kisgegas property: stereographic plot of poles to felsic dyke orientations. 32 measurements.

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Fig. 3-2 Kisgegas property: stereographic plot of poles to measured early potassic vein orientations. 111 measurements.
$\bigcirc$


Fig. 3-3 Kisgegas property: stereographic plot of poles to measured grey quartz $\left(\mathrm{MoS}_{2}, \mathrm{CuFeS}_{2}\right.$, $\mathrm{FeS}_{2}$ ) vein orientations. 20 measurements. ${ }^{\text {' }}$


Fig. 3-4 Kisgegas property: stereographic plot of poles to local composites of measurements of sheeted quartz veins. Each point is based on five to ten measurements in a small area. Total 72 measurements.

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Fig. 3-5 Kisgegas property: sterographic plot of poles to measured orientations of late stage veins.

LEGEND
Quartz, K-feldspar, pyrite veins

- 39 measurements Large vuggy quartz veins
- 6 measurements
Other late veins
10 measurements


## APPENDIX 4

Statement of Expenditures

## STATEMENT OF EXPENDITURES

KISGEGAS-82 GROUP

## SALARIES AND FRINGE BENEFITS - TEXASGULF INC.

G.R. Peatfield - P.Eng. Sept. 10, 11; 2 days @ $\$ 220$.
P.R. DeLancey - P.Eng. Sept. 10, 11; 2 days @ $\$ 200$. 400.00
D.A. Bending - Geologist

Period Aug. 21 - Sept.23;
32 days @ $\$ 140$.
4,480.00
E. Potsepp - Cook

Period Sept. 19-25; 6 days @ $\$ 705$.
J. Etzkorn - Cook

Period Aug.22-Sept.21; 29 days @ $\$ 80$. 2,320.00
G. Cooper - Geologist

Period Sept. 15-25; 9 days © $\$ 95$. 855.00
D. Piroshco - Geologist

Period Sept. 19-25; 5 days @ $\$ 75$. 375.00
P. Mouldey - Assistant

Period Aug. 21-29; 8 days © $\$ 60$. 480.00
M. Stanley - Assistant

Period Aug.21-25; 4 days @ $\$ 55.220 .00$
R. Larsen - Assistant

Period Sept. 13-25; 12 days @ $\$ 55 . \quad 660.00$
J. Leigh - Assistant

Period Sept.11-25; 14 days @ \$45.
630.00
$\$ 17,490.00$
$\$ 11,490.00$
ROOM AND BOARD
Tg Personne 1 - 123 days @ \$90. \$17,070.00
Longyear personnel - 108 days $0 \$ 90$. $\frac{9,720.00}{\$ 20,790.00}$
$\$ 20,790.00$

## HELICOPTER SUPPORT

| Invoice totals Highland 206 B | \$7,130.00 |  |
| :---: | :---: | :---: |
| Invoice totals Okanagan 206 B | 12,758.61 |  |
| $206 \mathrm{~L}-1$ | 13,377.40 |  |
| Texasgulf leased A-Star | 13,37.40 |  |
| 75 hours @ \$550. | 41,250.00 |  |
|  | \$74,576.07 | \$74,516.01 |

DIAMOND DRILLING

| Longyear Canada invoices for drilling, <br> survey, core boxes, supplies and <br> equipment, moving, mob. and demob. |  |
| :--- | ---: |
| Rental of Sperry-Sun survey instrument | 1,480.50$\$ 57,103.76$ |

## ANALYTICAL COSTS

| 224 MoS2 assays @ $\$ 8.00$ | $1,792.00$ |  |
| :--- | ---: | ---: |
| 224 sample preparation | 168.00 |  |
| (composite) @ $\$ 0.75$ | 75.00 |  |
| 20 W analyses @ $\$ 3.75$ | 460.00 |  |
| $46 \mathrm{Sn}, \mathrm{F}, \mathrm{Cu}, \mathrm{Mn}$ analyses @ $\$ 10.00$ | $\$ 2,495.00$ | $2,495.00$ |

REPORT PREPARATION
G.R. Peatfield, P.Eng. 1/2 day @ \$220. 110.00
D.A. Bending 5 days @ $\$ 140$. 700.00

Contract drafting 1,344.82
Inhouse drafting 300.00
Secretarial
Reproduction, etc.
250.00 150.00 \$2,854.82

2,854.82

## MISCELLANEOUS

Office and technical supplies 100.00
Pro-rated share of travel 640.00

Shipping and storage 800.00

Communications (radio, etc.)


2,540.00

0

## APPENDIX 5

Statement of Qualifications

## STATEMENT OF QUALIFICATION

## D.A. Bending - Geologist

D.A. Bending holds a B.Sc. degree in Geology from the University of Oregon (1976), and is presently completing an M.Sc. degree at the University of Toronto. He was employed by Texasgulf from May I, 1980 to February 1982, when he returned to the University of Toronto.








