| LOG NO: $10-10$ |
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Report \#1088

# DIGHEMIII SURVEY <br> FOR 

KRL RESOURCES CORP.
STEWART, B.C.

NTS 103P/14, 104N/3, 4


DIGHEM SURVEYS \& PROCESSING INC. MISSISSAUGA, OMTARIO July 25, 1990

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Ministry of
Energy, Mines and Petroleum Resources

## AssEsSMENT REPORT TITLE PAGE AND SUMMARY



## SUMMARY GEOLOGY (liznology, ape, structure, alteration, mineralization, ize, and mtitude):

The propertjes are underlain by north trending moder ate to shallow west dipping volcanic and sedimentary rocks of the Hazelton Group and Bowser Group. The MM property is intruded by bodies of granodiorite and monzonite ánd disrupted by north trending•faults:

[^0]

1. MM Property
MM 100 ..... 1594 ..... 18
MM 2 ..... 33111
MM 3 ..... 3312 ..... 1
MM 5 3313 ..... 1
MM 1 Fraction ..... 3314 ..... 1
MM 4 Fraction ..... 3315 ..... 1
MM 6 Fraction ..... 3316 ..... 1
Lake 16 ..... 3139 ..... 1
Lake 17 3140 ..... 1
Buck 709 ..... 3138 ..... 3
Buck ..... 803412
Dunwell 4 Fraction ..... 5871 ..... 1
2. Rufus Property
21401
Rufus 1 ..... 525 ..... 1
Rufus 2 ..... 526 ..... 1
Rufus 3 ..... 2141 ..... 1
Rufus 4 ..... 527
Rufus 6 ..... 528
Baby Rufus Fraction ..... 529 ..... 1
Wide Fraction ..... 530
Silver Fraction ..... 531 ..... 1
Long Fraction ..... 5321
Argyle Fraction ..... 520 ..... 11Argyle 15341
Argyle 2 ..... 535
Argyle 21
536
Argyle 3
537
Argyle 4
538
Argyle 5
539
Duke Fraction ..... 540
Comet 4 ..... 522
Veteran ..... 523
Veteran 3 ..... 524
3. Med Big Property
Med 1 ..... 799720
Med 2 ..... 7998 ..... 20
Med 3 ..... 7999 ..... 20
Med 4 ..... 8000 ..... 20
Med 5 ..... 8001 ..... 20
Med 6 ..... 8002 ..... 20
Med 7 ..... 8003 ..... 20
Med 8 ..... 8004 ..... 20
Med 9 ..... 8005 ..... 20
Med 10 ..... 8006 ..... 20
Med 11 ..... 8007 ..... 20
Med 12 ..... 8008 ..... 20
4. Med Little Property
20
Med 17 ..... 8013
Med 18 ..... 8014 ..... 20
Med 19 8015 ..... 20
Med 20 8016 ..... 20

(3)



## SUMHARY

This report describes the logistics and results of a DIGHEMII airborne geophysical survey carried out for KRL Resources Corp., over four properties in the Stewart area, B.C.

The purpose of the survey was to detect zones of conductive mineralization and to provide information that could be used to map the geology and structure of the survey areas. This was accomplished by using a DIGHEMIII multicoil, multi-frequency electromagnetic system, supplemented by a high sensitivity Cesium magnetometer and a two-channel VLF receiver. The information from these sensors was processed to produce maps which display the magnetic and conductive properties of the survey areas. An electronic navigation system, operating in the UHF band, ensured accurate positioning of the geophysical data with respect to the base maps.

The EM survey detected several anomalies of possible bedrock origin. Most of the inferred bedrock conductors appear to warrant further investigation using appropriate surface exploration techniques. Areas of interest may be assigned priorities on the basis of supporting geophysical, geochemical and/or geological information. After initial investigations have been carried out, it may be necessary to


FIGURE 1
THE SURVEY AREAS

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## IFHRODUCTION

A DIGHEMIII electromagnetic/resistivity/magnetic/VLF survey was flown for KRL Resources Corp. from May 17 to May 26, 1990, over four survey blocks in the Stewart area, B.C. They are located on NTS map sheets $103 \mathrm{P} / 14$ and $104 \mathrm{~A} / 3$, 4 . (See Figure 1).

Survey coverage consisted of approximately 115 line-km in the MM100 area, 88 line-km in the Rufus area, 146 line-km in the Med (Little) area, and 376 line-km in the Med (Big) area. Traverse lines were flown in an azimuthal direction of $90^{\circ} / 270^{\circ}$ for all areas. All survey blocks were flown with a line separation of 200 m . Tie lines were flown for all areas perpendicular to the traverse line direction.

The survey employed the DIGHEM ${ }^{\text {III }}$ electromagnetic system. Ancillary equipment consisted of a magnetometer, radar altimeter, video camera, analog and digital recorders, a VLF receiver and an electronic navigation system. Details on the survey equipment are given in Section 2.

The instrumentation was installed in an Aerospatiale AS350B turbine helicopter which was provided by Canadian Helicopters Corporation. The helicopter flew at an average
airspeed of $64 \mathrm{~km} / \mathrm{hr}$ with an EM bird height of approximately 30 m.

Section 2 also provides details on the data channels, their respective sensitivities, and the navigation/flight path recovery procedure. Noise levels of less than 2 ppm are generally maintained for wind speeds up to $35 \mathrm{~km} / \mathrm{h}$. Higher winds may cause the system to be grounded because excessive bird swinging produces difficulties in flying the helicopter. The swinging results from the $5 \mathrm{~m}^{2}$ of area which is presented by the bird to broadside gusts.

## SURVEY EOUIPMENT

This section provides a brief description of the geophysical instruments used to acquire the survey datas

## Electromagnetic System

| Model: | DIGHEM III |  |
| :---: | :---: | :---: |
| Type: | Towed bird, symmetric dipole configuration operated at a nominal survey altitude of 30 metres. Coil separation is 8 metres for the 900 Hz and 7200 Hz coil pairs. |  |
| Coil orientat | -ns/frequencies: | coaxial / 900 Hz coplanar/ 900 Hz coplanar/7200 Hz |
| Channels recor | ded: | 3 inphase channels <br> 3 quadrature channels <br> 1 monitor channel |
| Sensitivity: |  | 0.2 ppm at 900 Hz 0.4 ppm at 7200 Hz |
| Sample rate: |  | 10 per second |

The electromagnetic system utilizes a multi-coil coaxial/coplanar technique to energize conductors in different directions. The coaxial transmitter coil is vertical with its axis in the flight direction. The coplanar coils are horizontal. The secondary fields are sensed simultaneously by means of receiver coils which are maximum


#### Abstract

coupled to their respective transmitter coils. The system yields an inphase and a quadrature channel from each transmitter-receiver coil-pair.


## Magnetometer

| Model: | Picodas PMAG 3340 |
| :--- | :--- |
| Type: | Optically pumped Cesium vapour |
| Sensitivity: | 0.01 nT |
| Sample rate: | 10 per second |

The magnetometer sensor is towed in a bird 15 m below the helicopter.

## Base Station Magnetometer

Model: Geometrics G826A
Type: Digital recording proton precession
Sensitivity: 0.5 nT
Sample rate: 0.2 per second

Both digital and analog recorders are operated in conjunction with the base station magnetometer to record the diurnal variations of the earth's magnetic field. The clock of the base station is synchronized with that of the airborne system to permit subsequent removal of diurnal drift.

## VLF System



The VLF receiver measures the total field and vertical quadrature components of the secondary VLF field. Signals from two separate transmitters can be measured simultaneously. The VLF sensor is towed in a bird 10 m below the helicopter.

## Radar Altimeter

Manufacturer: Honeywell/Sperry
Type:
AA 220
Sensitivity: 1 ft

The radar altimeter measures the vertical distance between the helicopter and the ground. This information is used in the processing algorithm which determines conductor depth.

## Digital Data Acquisition System/Analog Recorder

## Manufacturer: RMS Instruments

Type:
DGR33
Tape Deck: RMS TCR-12, 6400 bpi, tape cartridge recorder

The analog profiles were recorded on chart paper in the aircraft during the survey. Table 2-1 lists the geophysical data channels and the vertical scale of each profile.

The digital data were used to generate several computed parameters. Both measured and computed parameters were plotted as "digital profiles" during data processing. These parameters are shown in Table 2-2.

In Table 2-2, the log resistivity scale of 0.06 decade/mm means that the resistivity changes by an order of magnitude in 16.5 mm . The resistivities at 0,33 and 67 mm up from the bottom of the digital profile are respectively 1 , 100 and 10,000 ohm-m.

Table 2-1. The Analog Profiles

| Channel <br> Name | Parameter | Scale units/mm | Designation on digital profile |  |
| :---: | :---: | :---: | :---: | :---: |
| 1X9I | coaxial inphase ( 900 Hz ) | 2.5 ppm | CXI | ( 900 Hz ) |
| $1 \mathrm{X9Q}$ | coaxial quad ( 900 Hz ) | 2.5 ppm | CXQ | ( 900 Hz ) |
| 2P9I | coplanar inphase ( 900 Hz ) | 2.5 pqm | CPI | ( 900 Hz ) |
| 2P9Q | coplanar quad ( 900 Hz ) | 2.5 pgm | CPQ | ( 900 Hz ) |
| 3P7I | coplanar inphase ( 7200 Hz ) | 5 pqm | CPI | ( 7200 Hz ) |
| 3P7Q | coplanar quad ( 7200 Hz ) | 5 pqm | CPQ | (7200 Hz) |
| ALIR | altimeter | 3 m | ALIT |  |
| VF1T | VIF-total: primary stn. | $5 \%$ |  |  |
| VF10 | VLF-quad: primary stn. | 5\% |  |  |
| VF2T | VLF-total: secondary stn. | $5 \%$ |  |  |
| VF2Q | VLF-quad: secondary stn. | 5\% |  |  |
| CMGC | magnetics, coarse | 20 nT | MAG |  |
| CMGF | magnetics, fine | 2.0 nT |  |  |
| CXSP | coaxial spherics monitor |  | CXS |  |

Table 2-2. The Digital Profiles

| Channel Name (Frea) | Observed parameters | Scale units/nm |
| :---: | :---: | :---: |
| MAG | magnetics | 10 nT |
| ALT | bird height | 6 m |
| CXI ( 900 Hz ) | vertical coaxial coil-pair imphase | 2 ppm |
| CXQ ( 900 Hz ) | vertical coaxial coil-pair quadrature | 2 ppm |
| CXS | ambient noise monitor (coaxial receiver) |  |
| CPI ( 900 Hz ) | horizontal coplanar coil-pair inphase | 2 ppm |
| CPQ ( 900 Hz ) | horizontal coplanar coil-pair quadrature | 2 ppm |
| CPI ( 7200 Hz ) | horizontal coplanar coil-pair inphase | 4 ppm |
| CPQ ( 7200 Hz ) | horizontal coplanar coil-pair quadrature | 4 ppm |
|  | Computed Parameters |  |
| DFI ( 900 Hz ) | difference function inphase from CXI and CPI | 2 ppm |
| DFQ (900 Hz) | difference function quadrature from CXQ and CPQ | 2 ppm |
| CDT | conductance | 1 grade |
| RES ( 900 Hz ) | log resistivity | . 06 decade |
| RES ( 7200 Hz ) | log resistivity | . 06 decade |
| DP $(900 \mathrm{~Hz})$ | apparent depth | 6 m |
| DP ( 7200 Hz ) | apparent depth | 6 m |

## Tracking Camera

| Type: | Panasonic Video |
| :--- | :--- |
| Model: | AG 2400 |

Fiducial numbers are recorded continuously and are displayed on the margin of each image. This procedure ensures accurate correlation of analog and digital data with respect to visible features on the ground.

## Navigation System

| Model: | Del Norte 547 |
| :--- | :--- |
| Type: | UHF electronic positioning system |
| Sensitivity: | 1 m |
| Sample rate: | 0.5 per second |

The navigation system uses ground based transponder stations which transmit distance information back to the helicopter. The ground stations are set up away from the survey area and are positioned such that the signals cross the survey block at an angle between $30^{\circ}$ and $150^{\circ}$. After site selection, a traverse line is flown at right angles to a base line drawn through the transmitter sites to establish an arbitrary coordinate system for the survey area. The onboard Central Processing Unit takes any two transponder distances and determines the helicopter position relative to these two ground stations in cartesian coordinates.
The cartesian coordinates are transformed to UTM coordinates during data processing. This is accomplished by correlating a number of prominent topographical locations with the navigational data points. The use of numerous visual tie points serves two purposes: to accurately relate the navigation data to the map sheet and to minimize location errors which might result from distortions in uncontrolled photomosaic base maps.

## PRODUCTS AND PROCESSIHG TECEHIOUES

The following products are available from the survey data. Those which are not part of the survey contract may be acquired later. Refer to Table 3-1 for a summary of the maps which accompany this report, some of which may be sent under separate cover. Most parameters can be displayed as contours, profiles, or in colour.

## Base Maps

Base maps of the survey areas have been prepared from published topographic maps although photomosaics can also be used. Topographic maps provide an accurate, distortion-free base which facilitates correlation of the navigation data to the UTM grid. Photomosaics are useful for visual reference and for subsequent flight path recovery, but usually contain scale distortions. Orthophotos are ideal, but their cost and the time required to produce them, usually precludes their use as base maps.

## Electromagnetic Anomalies

Anomalous electromagnetic responses are selected and analysed by computer to provide a preliminary electromagnetic anomaly map. This preliminary EM map is used, by the

Table 3-1 Procucts Available from the Survey

| NO. OF | ANOMALY | PROFILES |  | OURS | HADOW |
| :---: | :---: | :---: | :---: | :---: | :---: |
| MAP SHEEIS | MAP | ON MAP | INK | COLOLR | MAP |
| Electromagnetic Anomalies 3 | 20,000 | N/A | N/A | N/A | N/A |
| Probable Bedrock Conductors | - | N/A | N/A | N/A | N/A |
| Resistivity ( 900 Hz ) | N/A | - | - | - | - |
| Resistivity ( 7200 Hz ) 3 | N/A | - | 20,000 | 20,000 | - |
| EM Magnetite | N/A | - | - | - | - |
| Total Field Magnetics 3 | N/A | - | 20,000 | 20,000 | - |
| Enhanced Magnetics | N/A | - | - | - | - |
| Vertical Gradient Magnetics 3 | N/A | - | 20,000 | 20,000 | - |
| 2nd Vertical Derivative Magnetics - | N/A | - | - | - | - |
| Magnetic Susceptibility - | N/A | - | - | - | - |
| Filtered Total Field VLF 3 | N/A | - | 20,000 | 20,000 | - |
| EM Profiles ( 900 Hz ) - | N/A | - | N/A | N/A | N/A |
| EM Profiles ( 7200 Hz ) | N/A | - | N/A | N/A | N/A |
| Overburden Thickness | N/A | - | - | - | N/A |
| Digital Profiles | Horksheet profiles |  |  |  | - |
|  | Interpreted profiles |  |  |  | 20,000 |

[^1]geophysicist, in conjunction with the computer-generated digital profiles, to produce the final interpreted EM anomaly map. This map includes bedrock, surficial and cultural conductors. A map containing only bedrock conductors can be generated, if desired.

## Resistivity

The apparent resistivity in ohm-m may be generated from the inphase and quadrature EM components for any of the frequencies, using a pseudo-layer halfspace model. A resistivity map portrays all the EM information for that frequency over the entire survey area. This contrasts with the electromagnetic anomaly map which provides information only over interpreted conductors. The large dynamic range makes the resistivity parameter an excellent mapping tool.

## EM Magnetite

The apparent percent magnetite by weight is computed wherever magnetite produces a negative inphase EM response.

## Total Field Magnetics

The aeromagnetic data are corrected for diurnal variation using the magnetic base station data. The regional

IGRF gradient is removed from the data, if required under the terms of the contract.

## Enhanced Magnetics

The total field magnetic data are subjected to a processing algorithm. This algorithm enhances the response of magnetic bodies in the upper 500 m and attenuates the response of deeper bodies. The resulting enhanced magnetic map provides better definition and resolution of nearsurface magnetic units. It also identifies weak magnetic features which may not be evident on the total field magnetic map. However, regional magnetic variations, and magnetic lows caused by remanence, are better defined on the total field magnetic map. The technique is described in more detail in Section 5.

## Magnetic Derivatives

The total field magnetic data may be subjected to a variety of filtering techniques to yield maps of the following:
vertical gradient
second vertical derivative
magnetic susceptibility with reduction to the pole upward/downward continuations


#### Abstract

All of these filtering techniques improve the recognition of near-surface magnetic bodies, with the exception of upward continuation. Any of these parameters can be produced on request. Dighem's proprietary enhanced magnetic technique is designed to provide a general "all-purpose" map, combining the more useful features of the above parameters.


VLE

The VLF data are digitally filtered to remove long wavelengths such as those caused by variations in the transmitted field strength.

## pigital Profiles

Distance-based profiles of the digitally recorded geophysical data are generated and plotted by computer. These profiles also contain the calculated parameters which are used in the interpretation process. These are produced as worksheets prior to interpretation, and can also be presented in the final corrected form after interpretation. The profiles display electromagnetic anomalies with their respective interpretive symbols. The differences between the worksheets and the final corrected form occur only with respect to the EM anomaly identifier.

## Contour, Colour and Shadow Map Displays

The geophysical data are interpolated onto a regular grid using a cubic spline technique. The resulting grid is suitable for generating contour maps of excellent quality.

Colour maps are produced by interpolating the grid down to the pixel size. The parameter is then incremented with respect to specific amplitude ranges to provide colour "contour" maps.

Monochromatic shadow maps are generated by employing an artificial sun to cast shadows on a surface defined by the geophysical grid. There are many variations in the shadowing technique which may be applied to total field or enhanced magnetic data, magnetic derivatives, VLF, resistivity, etc. Of the various magnetic products, the shadow of the enhanced magnetic parameter is particularly suited for defining geological structures with crisper images and improved resolution.

## SITRYEY RESUITRS

## GENERAL DISCUSSION

The survey results are presented on separate map sheets for each parameter at a scale of 1:20,000. Tables 4-1, 4-2, 4-3 and 4-4 summarize the EM responses in the survey areas, with respect to conductance grade and interpretation.

The anomalies shown on the electromagnetic anomaly maps are based on a near-vertical, half plane model. This model best reflects "discrete" bedrock conductors. Wide bedrock conductors or flat-lying conductive units, whether from surficial or bedrock sources, may give rise to very broad anomalous responses on the EM profiles. These may not appear on the electromagnetic anomaly maps if they have a regional character rather than a locally anomalous character. These broad conductors, which more closely approximate a half space model, will be maximum coupled to the horizontal (coplanar) coil-pair and should be more evident on the resistivity parameter. Resistivity maps, therefore, may be more valuable than the electromagnetic anomaly maps, in areas where broad or flat-lying conductors are considered to be of importance. Contoured resistivity maps, based on the 7200 Hz coplanar data, are included with this report.

- 4-2 -

TABLE 4-1

## EM ANOMALY STATISTICS

## MM100 AREA

## CONDUCTOR GRADE

7
6
5
4
3
2
1
*
TOTAL

CONDUCTOR
MODEL
D
B
S
H
L

TOTAL

CONDUCTANCE RANGE SIEMENS (MHOS)
$>100$
50-100
20-50
10 - 20
5 - 10
1-5
$<1$
INDETERMINATE

MOST LIKELY SOURCE

DISCRETE BEDROCK CONDUCTOR DISCRETE BEDROCK CONDUCTOR

NUMBER OF RESPONSES
NUMBER OF RESPONSES

0
0
4
5
11
4
13
37 CONDUCTIVE COVER 17 ROCK UNIT OR THICK COVER 4 CULTURE

5
37

## CONDUCTOR <br> GRADE

7
6
5
4
3
2
1
*
TOTAL

CONDUCTOR MODEL

CONDUCTANCE RANGE SIEMENS (MHOS)

$$
\begin{aligned}
&>100 \\
& 50-100 \\
& 20- \\
& 10= \\
& 5- \\
& 1- \\
& 10 \\
&< 1 \\
& \text { < } 1 \\
& \text { INDETERMINATE }
\end{aligned}
$$

MOST LIKELY SOURCE

DISCRETE BEDROCK CONDUCTOR CONDUCTIVE COVER

NUMBER OF RESPONSES

0
0
0
0
0
1
2
3
6

NUMBER OF RESPONSES

B
S
TOTAL

2
4
6

TABLE 4-3

## EM ANOMALY STATISTICS

MED (LITTLE) AREA

CONDUCTOR GRADE

7
6
5
4
3
2
1

TOTAL

## CONDUCTOR <br> MODEL

CONDUCTANCE RANGE
SIEMENS (MHOS)
' $>100$
50-100
20-50
$10-20$
5 - 10
1 - 5
$<1$
INDETERMINATE

MOST LIKELY SOURCE

DISCRETE BEDROCK CONDUCTOR DISCRETE BEDROCK CONDUCTOR CONDUCTIVE COVER

4
5 ROCK UNIT OR THICK COVER 5
EDGE OF WIDE CONDUCTOR
TOTAL
D
B
S
H
E
1

73

NUMBER OF RESPONSES

NUMBER OF RESPONSES

## TABLE 4-4

## EM ANOMALY STATISTICS

MED (BIG) AREA

| CONDUCTOR | CONDUCTANCE RANGE | NUMBER OF |
| :---: | :---: | :---: |
| GRADE | SIEMENS (MHOS) | RESPONSES |
| 7 | $>100$ | 0 |
| 6 | 50-100 | 0 |
| 5 | 20-50 | 0 |
| 4 | 10-20 | 4 |
| 3 | $5-10$ | 3 |
| 2 | $1-5$ | 46 |
| 1 | $<1$ | 42 |
| * | INDETERMINATE | 79 |
| TOTAL |  | 174 |
| CONDUCTOR | MOST LIKELY SOURCE | NUMBER OF |
| MODEL |  | RESPONSES |
| D | DISCRETE BEDROCK CONDUCTOR | 10 |
| B | DISCRETE BEDROCK CONDUCTOR | 23 |
| S | CONDUCTIVE COVER | 134 |
| H | ROCK UNIT OR THICK COVER | 6 |
| E | EDGE OF WIDE CONDUCTOR | 1 |
| TOTAL |  | 174 |

(SEE EM MAP LEGEND FOR EXPLANATIONS)

Excellent resolution and discrimination of conductors was accomplished by using a fast sampling rate of 0.1 sec and by employing a common frequency ( 900 Hz ) on two orthogonal coil-pairs (coaxial and coplanar). The resulting "difference channel" parameters often permit differentiation of bedrock and surficial conductors, even though they may exhibit similar conductance values.

Anomalies which occur near the ends of the survey lines (i.e., outside the survey areas), should be viewed with caution. Some of the weaker anomalies could be due to aerodynamic noise, i.e., bird bending, which is created by abnormal stresses to which the bird is subjected during the climb and turn of the aircraft between lines. Such aerodynamic noise is usually manifested by an anomaly on the coaxial inphase channel only, although severe stresses can affect the coplanar inphase channels as well.

The EM anomalies resulting from this survey appear to fall within three general categories.

The first type consists of discrete, well-defined anomalies which Yield marked inflections on the difference channels. These anomalies are usually attributed to conductive sulphides or graphite and are generally given a
"B", "T" or "D" interpretive symbol, denoting a bedrock source.


#### Abstract

The second class consists of moderately well-defined quadrature responses which coincide with low amplitude or poorly defined inphase responses. In these areas, where EM responses are evident primarily on the quadrature components, zones of poor conductivity are indicated. Where these responses are coincident with magnetic anomalies, it is possible that the inphase component amplitudes have been suppressed by the effects of magnetite. In areas where magnetite causes the inphase components to become negative, the apparent conductance and depth of $E M$ anomalies may be unreliable.


Depending on the strength and definition of the quadrature responses, anomalies in this category may be given interpretive symbols varying from "B?" to "S?".

The third class of anomalies comprises broad responses which exhibit the characteristics of a half space and do not yield well-defined inflections on the difference channels. Anomalies in this category are usually given an "S" or "H" interpretive symbol. The lack of a difference channel
response usually implies a broad or flat-lying conductive source such as overburden.

The effects of conductive overburden are evident over parts of the survey areas. Although the difference channels (DIFI and DIFQ) are extremely valuable in detecting bedrock conductors which are partially masked by conductive overburden, sharp undulations in the bedrock/overburden interface can yield anomalies in the difference channels which may be interpreted as possible bedrock conductors. Such anomalies usually fall into the "S?" or "B?" classification but may also be given an "E" interpretive symbol, denoting a resistivity contrast at the edge of a conductive unit.

It is recommended that an attempt be made to compile a suite of geophysical "signatures" over areas of interest. Anomaly characteristics are clearly defined on the computerprocessed geophysical data profiles which are supplied as one of the survey products.

Even weak conductors may be of economic significance in the survey areas. A complete assessment and evaluation of the survey data should be carried out by one or more qualified professionals who have access to, and can provide a
meaningful compilation of, all available geophysical, geological and geochemical data.

A Geometrics G826A proton precession magnetometer was operated at the survey base to record diurnal variations of the earth's magnetic field. The clock of the base station was synchronized with that of the airborne system to permit subsequent removal of diurnal drift.

The background magnetic levels have been adjusted to the mean IGRF value for the survey areas. However, the IGRF gradient across the survey blocks has not been removed.

The total field magnetic data have been presented as contours on the base maps using a contour interval of 10 nT where gradients permit. The maps show the magnetic properties of the rock units underlying the survey areas.

## MM100 AREA

## Magnetics

The magnetic data suggest that this area is structurally complex. These structural complexities are evident on the contour map as variations in magnetic intensity, irregular patterns, and as offsets or changes in strike direction.

There is evidence on the magnetic contour map that suggests that the area has been subjected to moderately strong deformation and/or alteration.

The highest magnetic values are situated within a complex magnetic high which extends from fiducial 1528 on line 10090 to fiducial 1400 on line 10160.

The magnetic results, in conjunction with the other geophysical parameters, should provide valuable information which can be used to map the geology and structure of the survey area.

## Resistivity

A resistivity map, which displays the conductive properties of the survey area, was produced from the 7200 Hz coplanar data.

In general, the resistivity patterns show limited agreement with magnetic trends.

A resistivity low, which trends north-northeast from the southwest corner of the survey block to fiducial 870 on line 10200, is coincident with the Bear River. Other resistivity


#### Abstract

lows also seem to be related to topographic features, although there are several resistivity features which seem to be associated with bedrock conductors.


## VLF

The VLF method is quite sensitive to the angle of coupling between the conductor and the propogated EM field. Consequently, conductors which strike towards the VLF station will usually yield a stronger response than conductors which are nearly orthogonal to it.

The VLF parameter does not normally provide the same degree of resolution available from the EM data. Closelyspaced conductors, conductors of short strike length or conductors which are poorly coupled to the VLF field, may escape detection with this method. Erratic signals from the VLF transmitters can also give rise to strong, isolated anomalies which should be viewed with caution. The filtered total field VLF contours are presented on the base maps with a contour interval of one percent.

VLF results were obtained from the transmitting stations at Seattle, Washington (NLK - 24.8 kHz ) and Lualaulei, Hawaii
(NPM - 23.4 kHz ). The VLF map shows the contoured results of the filtered total field from Seattle for all survey lines.

In general, the VLF trends over this survey area are moderately strong and well defined. Most anomalies exhibit a general north/south strike.

## CONDUCTORS IN THE MM100 AREA

The electromagnetic anomaly maps show the anomaly locations with the interpreted conductor type, dip, conductance and depth being indicated by symbols. Direct magnetic correlation is also shown if it exists. The strike direction and length of the conductors are indicated when anomalies can be correlated from line to line. When studying the map sheets, consult the anomaly listings appended to this report.

Most of the interpreted anomalies in Block A are probably indicative of conductive surficial sources, although there are several moderately strong anomalies which may be possible bedrock responses.

Anomalies 10040A, 10040B, 10060A, 10071B and 10071C

These anomalies reflect moderately strong, possible bedrock sources. All are situated within a resistivity low. Anomalies 10040B, 10060A and 10071B show very weak magnetic correlation. These anomalies are all associated with a north-northeast/south-southwest trending VLF anomaly.

Anomalies 10060B, 10060C and 10071D

These anomalous responses are indicative of moderately weak, possible bedrock sources. They are situated within a resistivity low which is approximately outlined by the 100 ohm-metre contour. These anomalous responses exhibit no direct magnetic or VLF correlation.

## RUFUS AREA

## Magnetics

The magnetic contour patterns suggest that this area is structurally complex. Truncations, and offsets in the magnetic contour patterns give many features a complex, segmented appearance.

The digital profiles show that several of the more magnetic units in this survey block contain magnetite.

## Resistivity

A resistivity map, which displays the conductive properties of the survey area, has been produced from the 7200 Hz coplanar data. This survey block is generally resistive, exhibiting resistivity values of greater than 5000 ohm-metres over most of the survey block. One area of moderate conductivity, which extends east/west over lines 20020 through 20040, is coincident with a river valley. There do not appear to be any bedrock conductors within this survey block.

VLF

VLF results were obtained from transmitting stations at Seattle, Washington (NLK - 24.8 kHz ), Lualualei, Hawaii (NPM - 23.4 kHz ), and Annapolis, Maryland (NSS - 21.4 kHz ). The VLF map displays the contoured results of the filtered total field data from Seattle for all survey lines.

In general, the VLF trends in this survey area are
moderately strong and well-defined. Most trends seem to strike approximately north-northwest/south-southeast.

## CONDUCTORS IN THE RUFUS AREA

Few anomalies of either surficial or bedrock origin are located within this survey block.

Anomaly 20110A

This single-line anomaly reflects a weak, possible bedrock source, although the conductance value may be erroneous because of the presence of magnetite. This anomaly is situated on the west flank of a highly magnetic, magnetite-rich zone. It exhibits a moderately strong magnetic correlation, which suggests it is coincident with a local magnetic peak.

Anomaly 20120A

This single line anomaly also reflects a weak, possible bedrock source. It is situated on the east flank of a magnetite-rich feature. It is associated with a weak resistivity low, approximately defined by the 5000 ohm-metre contour.

## MED (LITHLE) AREA

## Magnetics

The magnetic data suggest that this area is structurally complex. These structural complexities can be seen as variations in magnetic intensity, or as offsets or changes in strike direction. Several apparent linear structural features, which trend northeast/southwest are evident as discontinuities in the magnetic contour patterns.

The highest magnetic values in this block are associated with a thin magnetic high which trends northeast from the west end of line 30070 to fiducial 1100 on line 30120. This magnetic trend seems to transect the local magnetic strike.

Most of the bedrock anomalies interpreted from the survey data within this block are situated in the eastern half of the area. They are associated with relatively nonmagnetic features.

## Resistivity

A resistivity map, which displays the conductive properties of the survey block, was produced from the 7200 Hz coplanar data. The resistivity patterns show little


#### Abstract

agreement with magnetic trends, although several structural features, inferred from the magnetic data, are evident as discontinuities in the resistivity contours.


A highly conductive zone is situated in the eastern half of the survey block. The limit of this zone is approximately defined by the 1000 ohm-metre contour. This zone displays resistivity values of lower than 50 ohm-metres. Most of the bedrock conductors interpreted from the electromagnetic data are situated within this conductive zone.

## VLF

VLF results were obtained from the transmitting stations at Seattle, Washington (NLK - 24.8 kHz ), Cutler, Maine (NAA24.0 kHz ) and Annapolis, Maryland (NSS - 21.4 kHz ). The VLF map displays the contoured results of the filtered total field data from Seattle except for lines 30110 through 30210 which used Annapolis. In general, VLF trends in the survey area are moderately weak and poorly defined. Most trends appear to strike north-northwest/south-southeast to north/south. Most trends are of limited strike length and do not extend over more than 2 or 3 lines.

## CONDUCTORS IN THE MED (LITTLE) AREA

Most of the bedrock conductors defined by the survey in this area are located in the eastern half of the block. They are located east of a resistivity contrast extending north from fiducial 692 on line 30010 to fiducial 2878 on line 30210. This contrast is approximately defined by the 1000 ohm-metre contour on the 7200 Hz resistivity map.

Within the zone located to the east of this resistivity contrast, several highly conductive zones contain numerous bedrock conductors.

Zone A

This zone is situated near the eastern ends of lines 30150 through 30210. The approximate limit of this zone is defined by the 100 ohm-metre contour on the 7200 Hz resistivity map.

This zone contains several moderately weak conductors of limited strike length, and several single line anomalies. No direct magnetic correlation is exhibited by any of the anomalies in this zone. A possible structural linear feature trends northeast/southwest through this zone.

Conductors 30150A-30170A and 30150B-30160B

These conductors are situated west of zone A. Both reflect moderately weak, possible bedrock sources of limited strike length. Conductor 30150A-30170A possibly reflects a thin source.

Both trends give rise to moderately well-defined resistivity lows. Neither trend exhibits magnetic correlation.

Zone B

The approximate limit of this zone is defined by the 250 ohm-metre contour. It contains several weak, probable bedrock conductors, and weak single line responses. There is no direct magnetic correlation with this zone, as it appears to be associated with an area of low gradient magnetics.

MED (BIG) AREA

## Magnetics

The magnetic data for this area are characterized by broad, low gradient features.

A thin, moderately magnetic linear feature trends northwest from the east end of line 40210 to fiducial 4660 on line 40410 . This trend displays magnetic values approximately 60 nT above background.

This magnetic feature is coincident with a resistivity low.

Resistivity

A resistivity map, which portrays the conductive properties of the survey area, was produced from the 7200 Hz coplanar data.

In general, the resistivity patterns show little agreement with the magnetic data. One resistivity low, however, extending northwest from the east end of line 40220 to fiducial 4344 on line 40400, is coincident with a thin magnetic high. This resistivity low may be adversely affected by conductive surficial material as it coincides with an area of low topography associated with Nelson Creek.

Other resistivity zones also seem to be associated with topographic features. Several resistivity highs are coincident with topographic peaks.

Several of the bedrock anomalies interpreted from the survey data give rise to moderately well-defined resistivity lows.

VLF

VLF results for this survey area were obtained from the transmitting stations at Seattle, Washington (NLK - 24.8 kHz ) and Annapolis, Maryland (NSS - 21.4 kHz ). The VLF map shows the contoured results of the filtered total field from Seattle for all survey lines.

In general, the VLF trends in the survey area are moderately strong and well-defined. Most trends strike approximately north/south to north-northwest/south-southeast.

A weak VLF trend extending northwest from the east end of line 40210 to fiducial 4346 on line 40400 is coincident with a resistivity low and a linear magnetic high. This trend may be due to a combination of surficial and bedrock sources, as it is coincident with a creek.

## CONDUCTORS IN THE MED (BIG) AREA

Few bedrock conductors have been interpreted from the survey data for this area. All possible bedrock responses are situated at the western ends of line 40410 through 40500.

Conductors 40450A-40470B, 40460A-40490C, 40480A-40490A and 40490E-40500A

These conductors are all situated within a resistivity low whose approximate limit is defined by the 200 ohm-metre contour on the 7200 Hz resistivity map.

Conductor 40450A-40470B reflects a north/south trending, thin bedrock source. This trend is indicative of a moderately weak source. It is coincident with a weak VLF anomaly. There is no direct magnetic correlation with this trend.

Conductors 40460A-40490C, 40480A-40490A and 40490E40500A all strike approximately northwest/southeast.

Conductor 40480A-40490A reflects a strong, thin bedrock source, whereas conductors 40460A-40490C and 40490E-40500A are indicative of weak bedrock sources.

No magnetic correlation is apparent for any of these conductors, although they appear to be related to a magnetic low.

## Conductor 40420B-40440B, and Anomalies 40440C, 40470C and 40470D

These bedrock features are situated within a resistivity low whose limit is approximately defined by the 150 ohm-metre contour.

Conductor 40420B-40440B reflects a weak, thin, possible bedrock source of limited strike length. It trends northwest/southeast.

Anomalies 40440C and 40470C reflect weak, possible bedrock responses, whereas anomaly 40470 D reflects a thin bedrock source which exhibits a possible dip to the east. These anomalies are also associated with a magnetic low.

## BACKGROUND INFORMATION


#### Abstract

This section provides background information on parameters which are available from the survey data. Those which have not been supplied as survey products may be generated later from raw data on the digital archive tape.


## ELECTROMAGNETICS

DIGHEM electromagnetic responses fall into two general classes, discrete and broad. The discrete class consists of sharp, well-defined anomalies from discrete conductors such as sulfide lenses and steeply dipping sheets of graphite and sulfides. The broad class consists of wide anomalies from conductors having a large horizontal surface such as flatly dipping graphite or sulfide sheets, saline water-saturated sedimentary formations, conductive overburden and rock, and geothermal zones. A vertical conductive slab with a width of 200 m would straddle these two classes.

The vertical sheet (half plane) is the most common model used for the analysis of discrete conductors. All anomalies plotted on the electromagnetic map are analyzed according to this model. The following section entitled Discrete Conductor Analysis describes this model in detail, including
the effect of using it on anomalies caused by broad conductors such as conductive overburden.

The conductive earth (half space) model is suitable for broad conductors. Resistivity contour maps result from the use of this model. A later section entitled Resistivity Mapping describes the method further, including the effect of using it on anomalies caused by discrete conductors such as sulfide bodies.

## Geometric interpretation

The geophysical interpreter attempts to determine the geometric shape and dip of the conductor. Figure 5-1 shows typical DIGHEM anomaly shapes which are used to guide the geometric interpretation.

## Discrete conductor analysis

The EM anomalies appearing on the electromagnetic map are analyzed by computer to give the conductance (i.e., conductivity-thickness product) in siemens (mhos) of a vertical sheet model. This is done regardless of the interpreted geometric shape of the conductor. This is not an unreasonable procedure, because the computed conductance increases as the electrical quality of the conductor increases, regardless of its true shape. DIGHEM anomalies
Conductor
location
Channel DIFI
Ratio of
amplitudes
CxI/ CPI

Fig. 5-1 Typical DIGHEM anomaly shapes
are divided into seven grades of conductance, as shown in Table 5-1 below. The conductance in siemens (mhos) is the reciprocal of resistance in ohms.

Table 5-1. EM Anomaly Grades

| Anomaly Grade | siemens |
| :---: | :---: |
| 7 | $\gg 100$ |
| 6 | $20-100$ |
| 5 | $10-50$ |
| 4 | $5-20$ |
| 3 | $1-10$ |
| 2 | $<$ |
| 1 |  |

The conductance value is a geological parameter because it is a characteristic of the conductor alone. It generally is independent of frequency, flying height or depth of burial, apart from the averaging over a greater portion of the conductor as height increases. Small anomalies from deeply buried strong conductors are not confused with small anomalies from shallow weak conductors because the former will have larger conductance values.

Conductive overburden generally produces broad EM responses which may not be shown as anomalies on the EM maps. However, patchy conductive overburden in otherwise resistive areas can Yield discrete anomalies with a conductance grade (cf. Table 5-1) of 1,2 or even 3 for conducting clays which
have resistivities as low as 50 ohm-m. In areas where ground resistivities are below $10 \mathrm{ohm}-\mathrm{m}$, anomalies caused by weathering variations and similar causes can have any conductance grade. The anomaly shapes from the multiple coils often allow such conductors to be recognized, and these are indicated by the letters $S, H$, and sometimes $E$ on the electromagnetic anomaly map (see EM map legend).

For bedrock conductors, the higher anomaly grades indicate increasingly higher conductances, Examples: DIGHEM's New Insco copper discovery (Noranda, Canada) yielded a grade 5 anomaly, as did the neighbouring copper-zinc Magusi River ore body; Mattabi (copper-zinc, Sturgeon Lake, Canada) and Whistle (nickel, Sudbury, Canada) gave grade 6; and DIGHEM's Montcalm nickel-copper discovery (Timmins, Canada) yielded a grade 7 anomaly. Graphite and sulfides can span all grades but, in any particular survey area, field work may show that the different grades indicate different types of conductors.

Strong conductors (i.e., grades 6 and 7) are characteristic of massive sulfides or graphite. Moderate conductors (grades 4 and 5) typically reflect graphite or sulfides of a less massive character, while weak bedrock conductors (grades 1 to 3 ) can signify poorly connected graphite or heavily disseminated sulfides. Grades 1 and 2
conductors may not respond to ground EM equipment using frequencies less than 2000 Hz .

The presence of sphalerite or gangue can result in ore deposits having weak to moderate conductances. As an example, the three million ton lead-zinc deposit of Restigouche Mining Corporation near Bathurst, Canada, yielded a well-defined grade 2 conductor. The 10 percent by volume of sphalerite occurs as a coating around the fine grained massive pyrite, thereby inhibiting electrical conduction.

Faults, fractures and shear zones may produce anomalies which typically have low conductances (e.g., grades 1 to 3). Conductive rock formations can yield anomalies of any conductance grade. The conductive materials in such rock formations can be salt water, weathered products such as clays, original depositional clays, and carbonaceous material.

On the interpreted electromagnetic map, a letter identifier and an interpretive symbol are plotted beside the EM grade symbol. The horizontal rows of dots, under the interpretive symbol, indicate the anomaly amplitude on the flight record. The vertical column of dots, under the anomaly letter, gives the estimated depth. In areas where anomalies are crowded, the letter identifiers, interpretive
symbols and dots may be obliterated. The EM grade symbols, however, will always be discernible, and the obliterated information can be obtained from the anomaly listing appended to this report.

The purpose of indicating the anomaly amplitude by dots is to provide an estimate of the reliability of the conductance calculation. Thus, a conductance value obtained from a large ppm anomaly (3 or 4 dots) will tend to be accurate whereas one obtained from a small ppm anomaly (no dots) could be quite inaccurate. The absence of amplitude dots indicates that the anomaly from the coaxial coil-pair is 5 ppm or less on both the inphase and quadrature channels. Such small anomalies could reflect a weak conductor at the surface or a stronger conductor at depth. The conductance grade and depth estimate illustrates which of these possibilities fits the recorded data best.

Flight line deviations occasionally yield cases where two anomalies, having similar conductance values but dramatically different depth estimates, occur close together on the same conductor. Such examples illustrate the reliability of the conductance measurement while showing that the depth estimate can be unreliable. There are a number of factors which can produce an error in the depth estimate, including the averaging of topographic variations by the
altimeter, overlying conductive overburden, and the location and attitude of the conductor relative to the flight line. Conductor location and attitude can provide an erroneous depth estimate because the stronger part of the conductor may be deeper or to one side of the flight line, or because it has a shallow dip. A heavy tree cover can also produce errors in depth estimates. This is because the depth estimate is computed as the distance of bird from conductor, minus the altimeter reading. The altimeter can lock onto the top of a dense forest canopy. This situation yields an erroneously large depth estimate but does not affect the conductance estimate.

Dip symbols are used to indicate the direction of dip of conductors. These symbols are used only when the anomaly shapes are unambiguous, which usually requires a fairly resistive environment.

A further interpretation is presented on the EM map by means of the line-to-line correlation of anomalies, which is based on a comparison of anomaly shapes on adjacent lines. This provides conductor axes which may define the geological structure over portions of the survey area. The absence of conductor axes in an area implies that anomalies could not be correlated from line to line with reasonable confidence.

DIGHEM electromagnetic maps are designed to provide a correct impression of conductor quality by means of the conductance grade symbols. The symbols can stand alone with geology when planning a follow-up program. The actual conductance values are printed in the attached anomaly list for those who wish quantitative data. The anomaly ppm and depth are indicated by inconspicuous dots which should not distract from the conductor patterns, while being helpful to those who wish this information. The map provides an interpretation of conductors in terms of length, strike and dip, geometric shape, conductance; depth, and thickness. The accuracy is comparable to an interpretation from a high quality ground EM survey having the same line spacing.

The attached EM anomaly list provides a tabulation of anomalies in ppm, conductance, and depth for the vertical sheet model. The EM anomaly list also shows the conductance and depth for a thin horizontal sheet (whole plane) model, but only the vertical sheet parameters appear on the EM map. The horizontal sheet model is suitable for a flatly dipping thin bedrock conductor such as a sulfide sheet having a thickness less than 10 m . The list also shows the resistivity and depth for a conductive earth (half space) model, which is suitable for thicker slabs such as thick conductive overburden. In the EM anomaly list, a depth value of zero for the conductive earth model, in an area of thick
cover, warns that the anomaly may be caused by conductive overburden.

Since discrete bodies normally are the targets of $E M$ surveys, local base (or zero) levels are used to compute local anomaly amplitudes. This contrasts with the use of true zero levels which are used to compute true EM amplitudes. Local anomaly amplitudes are shown in the EM anomaly list and these are used to compute the vertical sheet parameters of conductance and depth. Not shown in the EM anomaly list are the true amplitudes which are used to compute the horizontal sheet and conductive earth parameters.

## Questionable Anomalies

DIGHEM maps may contain EM responses which are displayed as asterisks (*). These responses denote weak anomalies of indeterminate conductance, which may reflect one of the following: a weak conductor near the surface, a strong conductor at depth (e.g., 100 to 120 m below surface) or to one side of the flight line, or aerodynamic noise. Those responses that have the appearance of valid bedrock anomalies on the flight profiles are indicated by appropriate interpretive symbols (see EM map legend). The others probably do not warrant further investigation unless their locations are of considerable geological interest.

## The thickness parameter

DIGHEM Can provide an indication of the thickness of a steeply dipping conductor. The amplitude of the coplanar anomaly (e.g., CPI channel on the digital profile) increases relative to the coaxial anomaly (e.g., CXI) as the apparent thickness increases, i.e., the thickness in the horizontal plane. (The thickness is equal to the conductor width if the conductor dips at 90 degrees and strikes at right angles to the flight line.) This report refers to a conductor as thin when the thickness is likely to be less than 3 m , and thick when in excess of 10 m . Thick conductors are indicated on the EM map by parentheses " ( )". For base metal exploration in steeply dipping geology, thick conductors can be high priority targets because many massive sulfide ore bodies are thick, whereas non-economic bedrock conductors are often thin. The system cannot sense the thickness when the strike of the conductor is subparallel to the flight line, when the conductor has a shallow dip, when the anomaly amplitudes are small, or when the resistivity of the environment is below $100 \mathrm{ohm}-\mathrm{m}$.

## Resistivity mapping

encountered during surveys. In such areas, anomalies can be generated by decreases of only 5 m in survey altitude as well as by increases in conductivity. The typical flight record in conductive areas is characterized by inphase and quadrature channels which are continuously active. Local EM peaks reflect either increases in conductivity of the earth or decreases in survey altitude. For such conductive areas, apparent resistivity profiles and contour maps are necessary for the correct interpretation of the airborne data. The advantage of the resistivity parameter is that anomalies caused by altitude changes are virtually eliminated, so the resistivity data reflect only those anomalies caused by conductivity changes. The resistivity analysis also helps the interpreter to differentiate between conductive trends in the bedrock and those patterns typical of conductive overburden. For example, discrete conductors will generally appear as narrow lows on the contour map and broad conductors (e.g., overburden) will appear as wide lows.

The resistivity profiles and the resistivity contour maps present the apparent resistivity using the so-called pseudo-layer (or buried) half space model defined by Fraser $\left(1978^{1}\right.$. This model consists of a resistive layer overlying

1 Resistivity mapping with an airborne multicoil electromagnetic system: Geophysics, v. 43, p.144-172
a conductive half space. The depth channels give the apparent depth below surface of the conductive material. The apparent depth is simply the apparent thickness of the overlying resistive layer. The apparent depth (or thickness) parameter will be positive when the upper layer is more resistive than the underlying material, in which case the apparent depth may be quite close to the true depth.

The apparent depth will be negative when the upper layer is more conductive than the underlying material, and will be zero when a homogeneous half space exists. The apparent depth parameter must be interpreted cautiously because it will contain any errors which may exist in the measured altitude of the EM bird (e.g., as caused by a dense tree cover). The inputs to the resistivity algorithm are the inphase and quadrature components of the coplanar coil-pair. The outputs are the apparent resistivity of the conductive half space (the source) and the sensor-source distance. The flying height is not an input variable, and the output resistivity and sensor-source distance are independent of the flying height. The apparent depth, discussed above, is simply the sensor-source distance minus the measured altitude or flying height. Consequently, errors in the measured altitude will affect the apparent depth parameter but not the apparent resistivity parameter.

The apparent depth parameter is a useful indicator of simple layering in areas lacking a heavy tree cover. The DIGHEM system has been flown for purposes of permafrost mapping, where positive apparent depths were used as a measure of permafrost thickness. However, little quantitative use has been made of negative apparent depths because the absolute value of the negative depth is not a measure of the thickness of the conductive upper layer and, therefore, is not meaningful physically. Qualitatively, a negative apparent depth estimate usually shows that the EM anomaly is caused by conductive overburden. Consequently, the apparent depth channel can be of significant help in distinguishing between overburden and bedrock conductors.

The resistivity map often yields more useful information on conductivity distributions than the EM map. In comparing the EM and resistivity maps, keep in mind the following:
(a) The resistivity map portrays the absolute value of the earth's resistivity, where resistivity $=$ 1/conductivity.
(b) The EM map portrays anomalies in the earth's resistivity. An anomaly by definition is a change from the norm and so the EM map displays anomalies, (i) over narrow, conductive bodies
and (ii) over the boundary zone between two wide formations of differing conductivity.

The resistivity map might be likened to a total field map and the EM map to a horizontal gradient in the direction of flight ${ }^{2}$. Because gradient maps are usually more sensitive than total field maps, the EM map therefore is to be preferred in resistive areas. However, in conductive areas, the absolute character of the resistivity map usually causes it to be more useful than the EM map.

## Interpretation in conductive environments

Environments having background resistivities below 30 ohm-m cause all airborne EM systems to yield very large responses from the conductive ground. This usually prohibits the recognition of discrete bedrock conductors. However, DIGHEM data processing techniques produce three parameters which contribute significantly to the recognition of bedrock conductors. These are the inphase and quadrature difference channels (DIFI and DIFQ), and the resistivity and depth channels (RES and DP) for each coplanar frequency.

2 The gradient analogy is only valid with regard to the identification of anomalous locations.

The EM difference channels (DIFI and DIFQ) eliminate most of the responses from conductive ground, leaving responses from bedrock conductors, cultural features (e.g., telephone lines, fences, etc.) and edge effects. Edge effects often occur near the perimeter of broad conductive zones. This can be a source of geologic noise. While edge effects Yield anomalies on the EM difference channels, they do not produce resistivity anomalies. Consequently, the resistivity channel aids in eliminating anomalies due to edge effects. On the other hand, resistivity anomalies will coincide with the most highly conductive sections of conductive ground, and this is another source of geologic noise. The recognition of a bedrock conductor in a conductive environment therefore is based on the anomalous responses of the two difference channels (DIFI and DIFQ) and the resistivity channels (RES). The most favourable situation is where anomalies coincide on all channels.

The DP channels, which give the apparent depth to the conductive material, also help to determine whether a conductive response arises from surficial material or from a conductive zone in the bedrock. When these channels ride above the zero level on the digital profiles (i.e., depth is negative), it implies that the EM and resistivity profiles are responding primarily to a conductive upper layer, i.e., conductive overburden. If the DP channels are below the
zero level, it indicates that a resistive upper layer exists, and this usually implies the existence of a bedrock conductor. If the low frequency DP channel is below the zero level and the high frequency $D P$ is above, this suggests that a bedrock conductor occurs beneath conductive cover.

The conductance channel CDT identifies discrete conductors which have been selected by computer for appraisal by the geophysicist. Some of these automatically selected anomalies on channel CDT are discarded by the geophysicist. The automatic selection algorithm is intentionally oversensitive to assure that no meaningful responses are missed. The interpreter then classifies the anomalies according to their source and eliminates those that are not substantiated by the data, such as those arising from geologic or aerodynamic noise.

## Reduction of geologic noise

Geologic noise refers to unwanted geophysical responses. For purposes of airborne EM surveying, geologic noise refers to EM responses caused by conductive overburden and magnetic permeability. It was mentioned previously that the EM difference channels (i.e., channel DIFI for inphase and DIFQ for quadrature) tend to eliminate the response of conductive overburden. This marked a unique development in airborne EM
technology, as DIGHEM is the only EM system which yields channels having an exceptionally high degree of immunity to conductive overburden.

Magnetite produces a form of geological noise on the inphase channels of all EM systems. Rocks containing less than 18 magnetite can yield negative inphase anomalies caused by magnetic permeability. When magnetite is widely distributed throughout a survey area, the inphase EM channels may continuously rise and fall, reflecting variations in the magnetite percentage, flying height, and overburden thickness. This can lead to difficulties in recognizing deeply buried bedrock conductors, particularly if conductive overburden also exists. However, the response of broadly distributed magnetite generally vanishes on the inphase difference channel DIFI. This feature can be a significant aid in the recognition of conductors which occur in rocks containing accessory magnetite.

## EM magnetite mapping

The information content of DIGHEM data consists of a combination of conductive eddy current responses and magnetic permeability responses. The secondary field resulting from conductive eddy current flow is frequency-dependent and consists of both inphase and quadrature components, which are
positive in sign. On the other hand, the secondary field resulting from magnetic permeability is independent of frequency and consists of only an inphase component which is negative in sign. When magnetic permeability manifests itself by decreasing the measured amount of positive inphase, its presence may be difficult to recognize. However, when it manifests itself by yielding a negative inphase anomaly (e.g., in the absence of eddy current flow), its presence is assured. In this latter case, the negative component can be used to estimate the percent magnetite content.

A magnetite mapping technique was developed for the coplanar coil-pair of DIGHEM. The technique yields a channel (designated FEO) which displays apparent weight percent magnetite according to a homogeneous half space model. ${ }^{3}$ The method can be complementary to magnetometer mapping in certain cases. Compared to magnetometry, it is far less sensitive but is more able to resolve closely spaced magnetite zones, as well as providing an estimate of the amount of magnetite in the rock. The method is sensitive to 1/4\% magnetite by weight when the EM sensor is at a height of 30 m above a magnetitic half space. It can individually resolve steep dipping narrow magnetite-rich bands which are

3 Refer to Fraser, 1981, Magnetite mapping with a multi-coil airborne electromagnetic system: Geophysics, v. 46, p. 1579-1594.


#### Abstract

separated by 60 m . Unlike magnetometry, the EM magnetite method is unaffected by remanent magnetism or magnetic latitude.


The EM magnetite mapping technique provides estimates of magnetite content which are usually correct within a factor of 2 when the magnetite is fairly uniformly distributed. EM magnetite maps can be generated when magnetic permeability is evident as negative inphase responses on the data profiles.

Like magnetometry, the EM magnetite method maps only bedrock features, provided that the overburden is characterized by a general lack of magnetite. This contrasts with resistivity mapping which portrays the combined effect of bedrock and overburden.

## Recognition of culture

Cultural responses include all EM anomalies caused by man-made metallic objects. Such anomalies may be caused by inductive coupling or current gathering. The concern of the interpreter is to recognize when an EM response is due to culture. Points of consideration used by the interpreter, when coaxial and coplanar coil-pairs are operated at a common frequency, are as follows:

1. Channel CPS monitors 60 Hz radiation. An anomaly on this channel shows that the conductor is radiating power. Such an indication is normally a guarantee that the conductor is cultural. However, care must be taken to ensure that the conductor is not a geologic body which strikes across a power line, carrying leakage currents.
2. A flight which crosses a "line" (e.g., fence, telephone line, etc.) yields a center-peaked coaxial anomaly and an m-shaped coplanar anomaly. 4 When the flight crosses the cultural line at a high angle of intersection, the amplitude ratio of coaxial/coplanar response is 4. Such an EM anomaly can only be caused by a line. The geologic body which yields anomalies most closely resembling a line is the vertically dipping thin dike. Such a body, however, yields an amplitude ratio of 2 rather than 4. Consequently, an m-shaped coplanar anomaly with a CXI/CPI amplitude ratio of 4 is virtually a guarantee that the source is a cultural line.
3. A flight which crosses a sphere or horizontal disk yields center-peaked coaxial and coplanar anomalies with a CXI/CPI amplitude ratio (i.e., coaxial/coplanar) of

1/4. In the absence of geologic bodies of this geometry, the most likely conductor is a metal roof or small fenced yard. 5 Anomalies of this type are virtually certain to be cultural if they occur in an area of culture.
4. A flight which crosses a horizontal rectangular body or wide ribbon yields an m-shaped coaxial anomaly and a center-peaked coplanar anomaly. In the absence of geologic bodies of this geometry, the most likely conductor is a large fenced area. 5 Anomalies of this type are virtually certain to be cultural if they occur in an area of culture.
5. EM anomalies which coincide with culture, as seen on the camera film or video display, are usually caused by culture. However, care is taken with such coincidences because a geologic conductor could occur beneath a fence, for example. In this example, the fence would be expected to yield an m-shaped coplanar anomaly as in case \#2 above. If, instead, a center-peaked coplanar anomaly occurred, there would be concern that a thick

5 It is a characteristic of EM that geometrically similar anomalies are obtained from: (1) a planar conductor, and (2) a wire which forms a loop having dimensions identical to the perimeter of the equivalent planar conductor.
geologic conductor coincided with the cultural line.
6. The above description of anomaly shapes is valid when the culture is not conductively coupled to the environment. In this case, the anomalies arise from inductive coupling to the EM transmitter. However, when the environment is quite conductive (e.g., less than 100 ohm-m at 900 Hz$)$, the cultural conductor may be conductively coupled to the environment. In this latter case, the anomaly shapes tend to be governed by current gathering. Current gathering can completely distort the anomaly shapes, thereby complicating the identification of cultural anomalies. In such circumstances, the interpreter can only rely on the radiation channel CPS and on the camera film or video records.

## MAGNETICS

The existence of a magnetic correlation with an EM anomaly is indicated directly on the EM map. In some geological environments, an EM anomaly with magnetic correlation has a greater likelihood of being produced by sulfides than one that is non-magnetic. However, sulfide ore bodies may be non-magnetic (e.g., the Kidd Creek deposit near Timmins, Canada) as well as magnetic (e.g., the Mattabi deposit near Sturgeon Lake, Canada).

The magnetometer data are digitally recorded in the aircraft to an accuracy of one $n T$ (i.e., one gamma) for proton magnetometers, and 0.01 nT for cesium magnetometers. The digital tape is processed by computer to yield a total field magnetic contour map. When warranted, the magnetic data may also be treated mathematically to enhance the magnetic response of the near-surface geology, and an enhanced magnetic contour map is then produced. The response of the enhancement operator in the frequency domain is illustrated in Figure 5-2. This figure shows that the passband components of the airborne data are amplified 20 times by the enhancement operator. This means, for example, that a 100 nT anomaly on the enhanced map reflects a 5 nT anomaly for the passband components of the airborne data.

The enhanced map, which bears a resemblance to a downward continuation map, is produced by the digital bandpass filtering of the total field data. The enhancement is equivalent to continuing the field downward to a level (above the source) which is $1 / 20 \mathrm{th}$ of the actual sensorsource distance.

Because the enhanced magnetic map bears a resemblance to a ground magnetic map, it simplifies the recognition of trends in the rock strata and the interpretation of geological structure. It defines the near-surface local


Fig. 5-2 Frequency response of magnetic enhancement operaior for a sample Interval of 50 m .
geology while de-emphasizing deep-seated regional features. It primarily has application when the magnetic rock units are steeply dipping and the earth's field dips in excess of 60 degrees.

Any of a number of filter operators may be applied to the magnetic data, to yield vertical derivatives, continuations, magnetic susceptibility, etc. These may be displayed in contour, colour or shadow.

## VLIF

VLF transmitters produce high frequency uniform electromagnetic fields. However, VLF anomalies are not EM anomalies in the conventional sense. EM anomalies primarily reflect eddy currents flowing in conductors which have been energized inductively by the primary field. In contrast, VLF anomalies primarily reflect current gathering, which is a non-inductive phenomenon. The primary field sets up currents which flow weakly in rock and overburden, and these tend to collect in low resistivity zones. Such zones may be due to massive sulfides, shears, river valleys and even unconformities.


Fig. 5-3 Frequency response of VLF operator.

The VLF field is horizontal. Because of this, the method is quite sensitive to the angle of coupling between the conductor and the transmitted VLF field. Conductors which strike towards the VLF station will usually yield a stronger response than conductors which are nearly orthogonal to it.

The Herz Industries Ltd. Totem VLF-electromagnetometer measures the total field and vertical quadrature components. Both of these components are digitally recorded in the aircraft with a sensitivity of 0.1 percent. The total field yields peaks over VLF current concentrations whereas the quadrature component tends to yield crossovers. Both appear as traces on the profile records. The total field data are filtered digitally and displayed as contours to facilitate the recognition of trends in the rock strata and the interpretation of geologic structure.

The response of the VLF total field filter operator in the frequency domain (Figure 5-3) is basically similar to that used to produce the enhanced magnetic map (Figure 5-2). The two filters are identical along the abscissa but different along the ordinant. The VLF filter removes long wavelengths such as those which reflect regional and wave transmission variations. The filter sharpens short wavelength responses such as those which reflect local geological variations.

## CONCLUSIONS AND RECOMMENDATIONS

This report provides a very brief description of the survey results and describes the equipment, procedures and logistics of the survey.

The various maps included with this report display the magnetic and conductive properties of the survey areas. It is recommended that the survey results be reviewed in detail, in conjunction with all available geophysical, geological and geochemical information. Particular reference should be made to the computer generated data profiles which clearly define the characteristics of the individual anomalies.

The anomalies defined by the survey should be subjected to further investigation, using appropriate surface exploration techniques. Anomalies which are currently considered to be of moderately low priority may require upgrading if follow-up results are favourable.

Colour maps have been prepared from the survey data. These maps result from additional processing of the geophysical data which was carried out in order to extract the maximum amount of information from the survey results. These maps provide valuable information on structure and
lithology, which may not be clearly evident on the contour maps. Current processing techniques can yield images which define subtle, but significant, structural details.

Respectfully submitted, DIGHIEM SURVEYS \& PROCESSING INC.

for Ruth A. Pritchard
Geophysicist

RAP/sdp
A1088JUL. 91 R

## APPENDIX A

## LIST OF PERSONNEL

The following personnel were involved in the acquisition, processing, interpretation and presentation of data, relating to a DIGHEM ${ }^{I I I}$ airborne geophysical survey carried out for KRL Resources Corp., over four properties in the Stewart area, B.C.

Steve Kilt
Robert Gordon
Maurice Bergstrom John Jess

Gordon Smith
David E. Pritchard
Ruth A. Pritchard
Reinhard Zimmerman
Susan Pothiah

Vice President, Operations Geophysical Operator and Geophysical Operator Pilot (Canadian Helicopters Corporation)
Computer Processing Supervisor
Computer Processor
Geophysicist/Interpreter
Draftsperson
Word Processing Operator

The survey consisted of 725 km of coverage, flown from May 17 to May 26, 1990.

All personnel are employees of Dighem Surveys \& Processing Inc., except for the pilot who is an employee of Canadian Helicopters Corporation.

DIGHEM SURVEYS \& PROCESSING INC.


Ruth A. Pritchard
Geophysicist
RAP/sdp
Ref: Report \#1088
A1088JUL. 91R

## APPENDIX $B$

## STATEMENT OF COST

## Date: July 25, 1990

IN ACCOUNT WITH
DIGHRM SURVEYS \& PROCESSING INC.

To: Dighem flying of Agreement dated April 2, 1990, pertaining to an Airborne Geophysical Survey in the Stewart area, B.C.

## Survey Charges

| Mobilization |  |
| :--- | :--- |
| 725 km of data | $\$$$3,500.00$ <br>  |
| $\underline{72,500.00}$ |  |

## Allocation of Costs

- Data Acquisition (60\%)
- Data Processing
(20\%)
- Interpretation, Report and Maps
(20\%)

DIGHEM SURVEYS \& PROCESSING INC.


Ruth A. Pritchard
Geophysicist

RAP/sdp

A1088JUL. 91 R

## APPENDIX C

## STATEMENT OF QUALIFICATIONS

I, Ruth A. Pritchard of the City of Brampton, Province of Ontario, do hereby certify that:

1. I am a geophysicist, residing at 31 Barrington Crescent, Brampton, Ontario, L6Z 1N2.
2. I am a graduate of York University, Downsview, Ontario, with a Specialized Honours B.Sc. Earth Sciences - Geophysics (1986).
3. I have been actively engaged in geophysical exploration since 1986.
4. The statements made in this report represent my best opinion and judgement.

for Ruth A. Pritchard Geophysicist

APPENDIXD

EM ANOMALY LIST

| COAXIAL | COPLANAR | COPLANAR | VERTICAL | HORIZONIAL CONDUCIIVE | MAG |
| ---: | ---: | ---: | :---: | ---: | :--- |
| 900 HZ | 900 HZ | 7200 HZ | DIKE | SHEET | EARIH |
| CORR |  |  |  |  |  |

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LINE 10080 (FLIGFP 3)

| A | 1221 | L | 1 | 2 | 1 | 2 | 2 | $4 \cdot$ | - | - | $\cdot$ | - | - | - |
| :--- | :---: | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- | :--- |
| B | 999 | B? | 1 | 6 | 1 | 7 | 26 | $23 \cdot$ | 0.8 | 0 | $\cdot$ | 1 | 43 | 225 |
| C | 859 | S | 0 | 5 | 1 | 6 | 25 | $23 \cdot$ | 0.7 | 0 | 0 | 1 | 73 | 213 |

LTNE 10090 (FLTGFP 3)

| A | 1448 | S? | 1 | 2 | 1 | 2 | 2 | 4 | $\cdot$ | - | - | $\cdot$ | - | - |
| :--- | :--- | :--- | :--- | :--- | :--- | ---: | ---: | ---: | ---: | ---: | ---: | ---: | ---: | ---: |
| B | 1563 | S? | 1 | 1 | 1 | 4 | 17 | 16 | $\cdot$ | 1.0 | 0 | - | 1 | 74 |
| C | 1616 | S? | 2 | 3 | 1 | 3 | 12 | 19 | . | 0.6 | 0 | . | 1 | 51 |

$\begin{array}{llllllllllllllllllll}\text { C } 1616 & \text { S? } & 2 & 3 & 1 & 3 & 12 & 19 & 0.6 & 0 & 0 & 1 & 51 & 310 & 23 & 0\end{array}$
.* ESTITMATED DEPIH MAY BE UNRELIABLE BECAUSE THE STRONGER PART

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- LIINE, OR BECAUSE OF A SHALLOW DIP OR OVERBURDEN EFFECTS.
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- IINE, OR|BECAUSE OF A SHALLOW DIP OR OVERBURDEN EFFECTS.

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| $\begin{aligned} & \text { LINE } 20010 \\ & \text { A } 2600 \mathrm{~S} \end{aligned}$ | $\begin{gathered} \text { (FLIGHT } \\ 12 \end{gathered}$ | 9) | 2 | 2 | 4 | - | - | - | - | - | - |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| LINE 20110 | (FLIGHP | 8) |  |  |  |  |  |  |  |  |  |  |
| A 1292 B? | 02 | 0 | 2 | 0 | 4 | - | - | - | - | - | - | 130 |
| LINE 20120 | (FLIGHIP | 8) |  |  |  |  |  |  |  |  |  |  |
| A 916 B? | 02 | 0 | 3 | 0 | 10 | 0.5 | 14 | 1 | 159 | 1035 | 0 | 0 |
| LINE 20131 | (FLIGHiP | 7) |  |  |  |  |  |  |  |  |  |  |
| A 3155 S ? | 00 | 0 | 1 | 0 | 10 | 0.1 | 0 | 1 | 111 | 8112 | 8 | $0$ |
| LINE 20150 | (FLITGFI | 7) |  |  |  |  |  |  |  |  |  |  |
| A 1662 S? | 01 | 0 | 2 | 0 | 15 | 1.7 | 59 | 1 | 157 | 1035 | 0 | $0$ |
| LINE 20161 | (FLIGHT | 7) |  |  |  |  |  |  |  |  |  |  |
| A 1464 S? | 01 | 0 | 0 | 0 | 4 | - | - | - | - | - | - | $0$ |

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- LINE, OR BECAUSE OF A SHALIOW DIP OR OVERBURDEN EFFECIS.

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| LTNE 30010 | (FLIGHIT | 13) |  |  |  |  |  |  |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| A 714 B? | 12 | 1 | 2 | 2 | 4 | - | - |  | - | - | - | - | 0 |
| B 729 B? | 28 | 7 | 14 | 53 | 33 | 2.3 | 0 |  | 1 | 56 | 207 | 11 | 0 |
| C 734 D | $9 \quad 18$ | 7 | 16 | 64 | 73 | 4.1 | 0 |  | 1 | 38 | 160 | 0 | 0 |
| LINE 30020 | (FLIGHT | 13) |  |  |  |  |  |  |  |  |  |  |  |
| A 110 B? | 59 | 1 | 10 | 43 | 64 | 2.4 | 2 |  | 1 | 32 | 548 | 0 | 0 |
| B 53 B ? | 49 | 4 | 12 | 37 | 14 | 3.0 | 0 |  | 1 | 33 | 161 | 0 | 0 |
| C 49 B ? | 12 | 1 | 2 | 2 | 4 | - | - |  | - | - | - | - | 4 |
| LINE 30040 | (FLIGHT | 12) |  |  |  |  |  |  |  |  |  |  |  |
| A 2596 B? | 57 | 1 | 6 | 22 | 22 | 3.7 | 7 |  | 1 | 70 | 757 | 0 | 0 |
| B 2588 B? | 25 | 6 | 6 | 22 | 4 | 4.3 | 20 |  | 1 | 92 | 78 | 52 | 7 |
| C 2565 D | 918 | 10 | 22 | 79 | 65 | 4.3 | 0 |  | 1 | 38 | 151 | 2 | 90 |
| LINE 30050 | (FLIGHT | 12) |  |  |  |  |  |  |  |  |  |  |  |
| A 2414 B | 714 | 6 | 19 | 80 | 72 | 3.5 | 0 |  | 1 | 44 | 109 | 9 | 0 |
| B 2422 D | 89 | 10 | 12 | 27 | 18 | 7.2 | 3 |  | 1 | 52 | 105 | 14 | 0 |
| C 2428 D | 67 | 10 | 12 | 22 | 26 | 6.8 | 15 |  | 1 | 64 | 217 | 18 | 0 |
| LINE 30060 | (FLTGHFI | 12) |  |  |  |  |  |  |  |  |  |  |  |
| A 1894 B? | 23 | 3 | 7 | 37 | 34 | 2.8 | 6 |  | 1 | 59 | 122 | 17 | 0 |
| B 1888 D | 12 | 1 | 2 | 2 | 4 | - | - |  | - | - | - | - | 0 |
| LITE 30070 | (FLIGET | 12) |  |  |  |  |  |  |  |  |  |  |  |
| A 1533 B? | 27 | 2 | 9 | 28 | 38 | 1.8 | 4 |  | 1 | 48 | 531 | 0 | 0 |
| B 1540 B? | 7 | 3 | 5 | 23 | 23 | 4.0 | 6 |  | 1 | 64 | 161 | 19 | 0 |
| C 1546 B? | 37 | 3 | 8 | 32 | 24 | 2.6 | 9 |  | , | 53 | 151 | 13 | 0 |
| D 1552 D | 28 | 3 | 11 | 38 | 53 | 1.9 | 4 |  | 1 | 50 | 260 | 7 | 0 |
| LINE 30080 | (FLIGHT | 12) |  |  |  |  |  |  |  |  |  |  |  |
| A 771 B? | 58 | 1 | 7 | 32 | 36 | 3.5 | 8 |  | 1 | 59 | 289 | 10 | 0 |
| B 760 B? | 68 | 2 | 9 | 37 | 52 | 3.4 | 5 |  | 1 | 58 | 151 | 16 | 13 |
| LINE 30090 | (FLIGFIT | 1) |  |  |  |  |  |  |  |  |  |  |  |
| A 2080 D | 12 | 1 | 2 | 2 | 4 | - | - |  | - | - | - | - | 0 |
| B 2071 S? | 36 | 1 | 8 | 28 | 27 | 1.8 | 0 |  | 1 | 98 | 154 | 43 | 0 |
| C 2043 B? | 23 | 2 | 4 | 14 | 13 | 1.0 | 0 |  | 1 | 58 | 179 | 33 | 0 |
| LINE 30101 | (FLIGHP | 1) |  |  |  |  |  |  |  |  |  |  |  |
| A 1602 B? | 06 | 0 | 8 | 0 | 16 | 8.0 | 60 |  | 1 | 103 | 860 | 27 | 60 |
| B 1589 B ? | 02 | 0 | 2 | 0 | 4 | - | - |  | - | - |  | - | 80 |
| .* ESTTMAIED DEPTH MAY BE UNRELIABLE BECAUSE THE SIRTCNGER PART OF THE CONDUCIOR MAY BE DEEPER OR TO ONE SIDE OF THE FLIGHT LINE, OR BECAUSE OF A SHALION DIP OR OVERBURDEN EFFECIS. |  |  |  |  |  |  |  |  |  |  |  |  |  |



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| ITNE 30101 |  | TLIGHT | 1) |  |  |  |  |  |  |  |  |  |  |  |
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| C 1518 B | 3 | 6 | 5 | 8 | 30 | 39 |  | 3.9 | 31 | 1 | 81 | 172 | 37 | 0 |
| D 1452 S | 1 | 2 | 1 | 2 | 2 | 4 |  | - | - | - | - | - | - | 0 |
| ITNE 30110 |  | TICHT | 10) |  |  |  |  |  |  |  |  |  |  |  |
| A 674 D | 5 | 8 | 9 | 12 | 34 | 40 |  | 5.5 | 0 | 1 | 0 | 259 | 0 | 0 |
| B 659 D | 4 | 9 | 2 | 8 | 22 | 53 |  | 2.2 | 0 | 1 | 0 | 122 | 0 | 0 |
| C 646 B ? | 4 | 6 | 5 | 9 | 26 | 14 |  | 3.8 | 0 | 1 | 0 | 75 | 0 | 0 |
| LINE 30120 |  | LIGHP | 10) |  |  |  |  |  |  |  |  |  |  |  |
| A 1365 B | 3 | 6 | 6 | 8 | 25 | 16 |  | 4.2 | 55 | 2 | 138 | 46 | 103 | 0 |
| LTNE 30130 |  | TIGHT | 10) |  |  |  |  |  |  |  |  |  |  |  |
| A 1676 B? | 1 | 2 | 0 | 2 | 2 | 4 |  | - | - | - | - | - | - | 0 |
| B 1641 B? | 6 | 7 | 8 | 9 | 27 | 16 |  | 7.0 | 0 | 2 | 67 | 44 | 34 | 0 |
| LINE 30150 |  | TIGFIP | 10) |  |  |  |  |  |  |  |  |  |  |  |
| A 2576 B? | 1 | 2 | 1 | 2 | 2 | 4 |  | - | - | - | - | - | - | 0 |
| B 2554 B? | 1 | 2 | 1 | 2 | 2 | 4 |  | - | - | - | - | - | - | 0 |
| C 2541 D | 1 | 2 | 1 | 1 | 2 | 4 |  | - | - | - | - | - | - | 0 |
| D 2530 D | 1 | 2 | 1 | 2 | 2 | 4 |  | - | - | - | - | - | - | 0 |
| E 2476 S | 1 | 2 | 1 | 2 | 2 | 4 |  | - | - | - | - | - | - | 0 |
| LINE 30160 |  | TIGHT | 10) |  |  |  |  |  |  |  |  |  |  |  |
| A 3508 D | 9 | 8 | 5 | 7 | 23 | 12 |  | 8.5 | 0 | 1 | 64 | 189 | 16 | 0 |
| B 3522 B? | 2 | 1 | 3 | 5 | 14 | 12 |  | 5.3 | 33 | 1 | 86 | 76 | 47 | 0 |
| C 3534 B ? | 4 | 4 | 1 | 3 | 8 | 14 |  | 0.5 | 0 | 1 | 0 | 113 | 0 | 0 |
| D 3557 B ? | 4 | 3 | 4 | 3 | 11 | 3 |  | 10.1 | 0 | 1 | 10 | 85 | 0 | 11 |
| E 3572 B? | 1 | 2 | 1 | 2 | 2 | 4 |  | - | - | - | - | - | - | 0 |
| ITNE 30170 |  | TIGHT | 11) |  |  |  |  |  |  |  |  |  |  |  |
| A 418 D | 11 | 12 | 6 | 14 | 45 | 35 |  | 6.4 | 12 | 1 | 57 | 182 | 16 | 0 |
| B 390 B? | 1 | 2 | 1 | 2 | 2 | 4 |  | - | - | - | - | - | - | 0 |
| C 371 H | 2 | 4 | 8 | 11 | 37 | 26 |  | 3.9 | 13 | 2 | 74 | 43 | 42 | 0 |
| D 340 H ? | 1 | 3 | 5 | 11 | 32 | 20 |  | 2.5 | 11 | 1 | 58 | 139 | 18 | 0 |
| LTNE 30180 |  | TIGHT | 11) |  |  |  |  |  |  |  |  |  |  |  |
| A 1342 B | 32 | 49 | 25 | 56 | 174 | 134 |  | 7.5 | 1 | 2 | 42 | 39 | 18 | 0 |
| B 1356 B ? | 7 | 12 | 5 | 12 | 35 | 34 |  | 3.8 | 4 | 2 | 76 | 51 | 43 | 0 |
| C 1370 B? | 10 | 7 | 8 | 9 | 30 | 24 |  | 10.9 | 1 | 2 | 69 | 26 | 41 | 0 |
| LINE 30190 |  | TIIGHI | 11) |  |  |  |  |  |  |  |  |  |  |  |
| A 1665 D | 25 | 21 | 16 | 21 | 75 | 50 |  | 12.8 | 2 | 1 | 43 | 128 | 7 | 0 |

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- IINE, OR BECAUSE OF A SHALTON DIP OR OVERBURDEN EFFECIS.

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| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 900 HZ | 900 HZ | 7200 HZ | DIKE | SHEEST | EARIH | CORR |

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- LINE, OR BECAUSE OF A SHALIOW DIP OR OVERBURDEN EFFECIS.

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| 900 HZ | 900 HZ | 7200 HZ | DIKE |  | SHEET | EAR | CORR |

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- OF THE CONDUCIOR MAY BE DEEPER OR TO ONE SIDE OF THE FLIGHT
- LINE, OR BECAUSE OF A SHALLOW DIP OR OVERBURDEN EFFECTS.

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| LINE 40100 | (FLIGHP | 15) |  |  |  |  |  |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| A 1490 S | 01 | 1 | 2 | 2 | 4 | - | - | . - | - | - | - | 0 |
| B 1500 S | 01 | 0 | 2 | 2 | 4 | - | - | - - | - | - | - | 0 |
| LTNE 40110 | (FLIGHP | 15) |  |  |  |  |  |  |  |  |  |  |
| A 2032 S | 12 | 1 | 2 | 2 | 4 | - | - | - - | - | - | - | 0 |
| B 2014 S | 16 | 1 | 7 | 25 | 53 | 0.5 | 0 | 1 | 39 | 580 | 0 | 0 |
| LTNE 40120 | (FLIGHP | 15) |  |  |  |  |  |  |  |  |  |  |
| A 2121 S | 03 | 1 | 7 | 21 | 34 | 0.5 | 0 | 1 | 47 | 552 | 0 | 0 |
| B 2195 S? | 02 | 1 | 1 | 2 | 4 | - | - | - - | - | - | - | 0 |
| LINE 40130 | (FLTGHE | 16) |  |  |  |  |  |  |  |  |  |  |
| A 1114 S | 02 | 0 | 2 | 2 | 4 | - | - | . - | - | - | - | 0 |
| B 1131 S | 12 | 0 | 2 | 2 | 4 | - | - | . - | - | - | - | 0 |
| C 1146 S | 12 | 1 | 2 | 2 | 4 | - | - | - - | - | - | - | 5 |
| D 1177 S | 23 | 0 | 4 | 11 | 23 | 0.5 | 0 | . 1 | 14 | 370 | 0 | 0 |
| E 1190 S | 12 | 0 | 2 | 2 | 4 | - | - | - - | - | - | - | 0 |
| F 1211 S | 12 | 1 | 2 | 2 | 4 | - | - | - - | - | - | - | 0 |
| G 1230 S | 12 | 1 | 2 | 2 | 4 | - | - | . - | - | - | - | 0 |
| LINE 40140 | (FLIGHP | 16) |  |  |  |  |  |  |  |  |  |  |
| A 1511 S | 25 | 2 | 7 | 23 | 41 | 1.9 | 18 | 1 | 67 | 299 | 20 | 0 |
| B 1492 S | 12 | 1 | 2 | 2 | 4 | - | - | - - | - | - | - | 0 |
| LINE 40150 | (FLIGHP | 16) |  |  |  |  |  |  |  |  |  |  |
| A 1579 S | 12 | 2 | 6 | 23 | 21 | 2.1 | 0 | 1 | 24 | 381 | 0 | 0 |
| LINE 40160 | (FLIGHIP | 16) |  |  |  |  |  |  |  |  |  |  |
| A 1991 S | 11 | 1 | 2 | 2 | 4 | - | - | - | - | - | - | 0 |
| B 1945 S | 02 | 1 | 4 | 14 | 18 | 0.8 | 0 | 1 | 47 | 272 | 22 | 0 |
| C 1872 S | 12 | 1 | 2 | 2 | 4 | - | - | . - | - | - | - | 0 |
| LINE 40170 | (FLIGHT | 16) |  |  |  |  |  |  |  |  |  |  |
| A 2079 H | 36 | 3 | 8 | 30 | 29 | 2.5 | 2 | 1 | 47 | 163 | 7 | 0 |
| B 2119 S | 14 | 1 | 8 | 22 | 50 | 0.9 | 6 | 1 | 49 | 352 | 5 | 0 |
| C 2127 S | 18 | 0 | 11 | 20 | 88 | 0.6 | 11 | 1 | 50 | 545 | 6 | 0 |
| D 2150 S | 11 | 1 | 2 | 2 | 4 | - | - | - | - | - | - | 0 |
| E 2168 S | 12 | 1 | 2 | 2 | 4 | - | - | , - | - | - | - | 0 |
| LINE 40180 | (FLIGET | 16) |  |  |  |  |  |  |  |  |  |  |
| A 2566 S | 57 | 4 | 8 | 23 | 20 | 4.3 | 0 | 1 | 65 | 109 | 24 | 0 |


| COAX | COPLANAR | COPIANAR | EERTICAL | HRRIZONIAL CO | MAG |
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|  | 900 Hz | 7200 HZ |  | SHEET PAPTH |  |

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COAXIAL COPLANAR COPLANAR - VERRIICAL . HORIZGNIAL CONDUCTIVE MAG 900 HZ 900 HZ 7200 HZ . DIKE . SHEET EARIH CORR ANOMALY/ REAL QUAD REAL QUAD REAL QUAD . COND DEPIH*. COND DEPIH RESIS DEPIH FID/INIERP PPM PPM PPM PPM PPM PPM .SIEMEN M .SIEMEN M OHM M M M MT

| LINE 40280 <br> B 793 S | $\begin{aligned} & \text { (FLIGHTP } \\ & 5 \quad 29 \end{aligned}$ | 17) | 53 | 177 | 233 | 1.6 | 0 | 1 | 23 | 144 | 0 | 12 |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| LINE 40290 | (FLuGHir | 17) |  |  |  |  |  |  |  |  |  |  |
| A 1153 S | 822 | 12 | 40 | 135 | 111 | 3.1 | 0 | 1 | 29 | 91 | 0 | 0 |
| LINE 40300 | ( FLIGHIT | 17) |  |  |  |  |  |  |  |  |  |  |
| A 1319 H | 12 | 1 | 2 | 2 | 4 | - | - | - | - | - | - | 0 |
| B 1307 S | 27 | 3 | 12 | 44 | 46 | 1.3 | 0 | 1 | 11 | 457 | 0 | 5 |
| LINE 40310 | (FLIGHP | 17) |  |  |  |  |  |  |  |  |  |  |
| A 1478 S | 14 | 1 | 6 | 18 | 36 | 0.5 | 0 | 1 | 63 | 677 | 0 | 0 |
| B 1574 H | 12 | 1 | 2 | 2 | 4 | - | - | - | - | - | - | 0 |
| LTNE 40320 | (FLIGHT | 17) |  |  |  |  |  |  |  |  |  |  |
| A 2064 S? | 2 | 1 | 2 | 2 | 4 | - | - | - | - |  | - | 0 |
| B 1936 H | 9 | 7 | 18 | 48 | 52 | 3.0 | 1 | 1 | 39 | 114 | 5 | 40 |
| LINE 40330 <br> A 2229 H | $\begin{aligned} & \text { (FLIGHTP } \\ & 3 \quad 11 \end{aligned}$ | $\begin{gathered} \text { 17) } \\ 6 \end{gathered}$ | 18 | 55 | 108 | 2.5 | 4 | 1 | 25 | 322 | 0 | 40 |
| $\begin{aligned} & \text { LINE } 40340 \\ & \text { A } 2605 \mathrm{~S} \end{aligned}$ | $\begin{aligned} & \text { (FITGHP } \\ & 1 \quad 15 \end{aligned}$ | 17) | 28 | 79 | 114 | 0.7 | 0 | 1 | 23 | 314 | 0 | 0 |
| $\begin{aligned} & \text { LTNE } 40350 \\ & \text { A } 2884 \text { S } \end{aligned}$ | ${ }_{1}^{\text {(FITGHT }}$ | 17) | 15 | 45 | 88 | 0.5 | 0 | 1 | 19 | 586 | 0 | 30 |
| LINE 40360 | (FITGHP | 17) |  |  |  |  |  |  |  |  |  |  |
| A 3171 S | 16 | 2 | 12 | 42 | 75 | 1.0 | 1 | 1 | 15 | 529 | 0 | 20 |
| B 3151 S | 28 | 2 | 14 | 45 | 63 | 1.2 | 0 | 1 | 12 | 505 | 0 | 0 |
| LINE 40370 | (FLIGHT | 17) |  |  |  |  |  |  |  |  |  |  |
| A 3408 S | 12 | 1 | 2 | 2 | 4 | - | - | - | - | - | - | 0 |
| B 3443 S | 25 | 1 | 7 | 18 | 43 | 1.5 | 6 | 1 | 33 | 580 | 0 | 0 |
| C 3455 s | 11 | 1 | 2 | 2 | 4 | - | - | - | - | - | - | 0 |
| D 3468 S | 24 | 2 | 6 | 25 | 26 | 2.2 | 0 | 1 | 23 | 422 | 0 | 0 |
| LINE 40380 | (FLIGHPI | 17) |  |  |  |  |  |  |  |  |  |  |
| A 3806 S | 12 | 1 | 5 | 20 | 31 | 0.7 | 0 | 1 | 48 | 195 | 24 | 0 |
| B 3743 S | 28 | 3 | 14 | 38 | 67 | 1.2 | 0 | 1 | 18 | 474 | 0 | 0 |
| LINE 40390 <br> A 3875 S? | $\begin{gathered} \text { (FLIGHT } \\ 3 \end{gathered}$ | 17) | 4 | 13 | 23 | 0.6 | 0 | 1 | 50 | 193 | 27 | 0 |

.* ESTIMATED DEPPIH MAY BE UNRETTABLE BECAUSE THE STRONGER PART •

- OF The CONDUCIOR MAY BE DEEPER OR TO ONE SIDE OF THE FLIGHIT .
- LINE, OR BECAUSE OF A SHAILOW DIP OR OVERBURDEN EFFECTS.


ANOMALY/ REAL QUAD REAL QUAD REAL QUAD . COND DEPPIH*. COND DEPIH RESIS DEPIH FID/INIERP PPM PPM PPM PPM PPM PPM .SIEMEN M .SIEMEN M OHM-M M NT

.* ESTITMALED DEPTH MAY BE UNRELIABLE BECAUSE THE SIRONGER PART

- OF IHE CORDUCTOR MAY BE DEEPER OR TO ONE SIDE OF THE FLIGHT
- LINE, OR BECAUSE OF A SHALION DIP OR OVERBURDEN EFFECCIS.

| COAXIAL | COPLANAR | COPLANAR | VERTITCAL | HORIZGNTAL CONDUCTIVE | MAG |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 900 HZ | 900 HZ | 7200 HZ | DIKE | SHEET | EARTH | CORR |

ANOMALY/ REAL QUAD REAL QUAD REAL QUAD . COND DEPPH*. COND DEPIH RESIS DEPIH FID/INIERP PPM PPM PPM PPM PPM PPM .SIENEN M.SIEMEN M OHM-M M

.* ESTIMATED DEPTH MAY BE UNRELIABLE BECAUSE THE SIRONGER PART

- OF THE CONDUCIOR MAY BE DEEPER OR TO ONE SIDE OF THE FLIGHI .
- LINE, OR BECAUSE OF A SHALION DIP OR OVERBURDEN EFFECIS. .


ANOMALY/ REAL QUAD REAL QUAD REAL QUAD . COND DEPPIH*. COND DEPPIH RESIS DEPTH FID/INIERP PPM PPM PPM PPM PPM PPM .SIEMEN M .SIEMEN M OFM-M M NT

.* ESTITMAIEED DEPPTH MAY BE UNRETLIABTE BECAUSE THE SIRCNGER PART

- OF THE CONDUCIOR MAY be dexeper or to one side of the flitgit
- LINE, OR BECAUSE OF A SHALIOW DIP OR OVERBURDEN EFFECIS.

| COAXIAL | COPLANAR | COPLANAR | VERTICAL | HORIZONIAL CONDUCTIVE | MAG |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 900 HZ | 900 HZ | 7200 HZ | DIKE | SHEET | EARIH | CORR |

ANOMALY/ REAL QUAD REAL QUAD REAL QUAD • COND DEPIH*. COND DEPIH RESIS DEPIH FID/INIERP PPM PPM PPM PPM PPM PPM .SIEMEN M .SIEMEN M OHM-M M NT

| LINE 49010 | (FLIGHT | 19) |  |  |  |  |  |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| A 1823 B? | 39 | 2 | 9 | 37 | 35 | 2.1 | 0 | 1 | 62 | 561 | 0 | 0 |
| B 1811 B? | 26 | 0 | 5 | 8 | 39 | 1.4 | 16 | 1 | 68 | 216 | 24 | 0 |
| LINE 49020 | (FLIGHP | 19) |  |  |  |  |  |  |  |  |  |  |
| A 2444 S | 02 | 0 | 2 | 2 | 4 | - | - | - | - | - | - | 7 |
| LINE 49030 | (FLIGHP | 19) |  |  |  |  |  |  |  |  |  |  |
| A 2111 S | 25 | 3 | 7 | 36 | 22 | 2.4 | 1 | 1 | 25 | 256 | 0 | 0 |

.* ESTIIMATED DEPIH MAY BE UNREUTABLE BECAUSE THE STRONGER PART .
: OF THE CONDUCIOR MAY BE DEEPER OR TO ONE SIDE OF THE FLIGFFT .

- LITE, OR BECAUSE OF A SHALIOW DIP OR OVERBURDEN EFFECTS.


[^0]:    REFERENCES TO PREvIOUS WORK. Assess: Repts: $\cdot$ Nos: •11422; -17660, . 18096
    BCDM Annual Repts: .1909, .1916, 1917, .1919, .1920, .1922, .1924, .1925, .1926, .1927, . 1928 . GSC Mem. Nos. 159, 175

[^1]:    N/A Not available
    Highly recammended due to its overall information content
    ** Recommended

    * Qualified recammendation, as it may be useful in local areas
    - No reccamendation

    20,000 Scale of delivered map, i.e, 1:20,000

