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DIGHEMIV SURVEY

FOR

NORANDA EXPLORATION COMPANY, LIMITED

BABINE LAKE PROJECT

BRITISH COLUMBIA

NTS 93L/16, 93 M/I

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GEOLOGICAL BRANCH ASSESSMENT REPORT

DIGHEM SURVEYS & PROCESSING INC. MISSISSAUGA, ONTARIO November 28, 1990 Paul A. Smith Geophysicist

A1090NOV.91R

Type of Report/SURVEY(S) Geo physicalN sheet only \$50,623TOTAL COST 1257500AUTHORISIRaul A. SmithSIGNATURE(S)Del MyosDATE STATEMENT OF EXPLORATION AND DEVELOPMENT FILED. 7. Now. 1990. YEAR OF WORK !PROPERTY NAME(S)MUNERALINVENTORY NUMBERISI, IF KNOWNOPT. CU.COMMODITIES PRESENTCU.COMMODITIES PRESENTCU.OPT. CU.NINER COLSPANOPT. CU.NINER COLSPANOPT. CU.OPT. CU.NINER COLSPANOPT. CU.NINER COLSPANOPT. CU.COMMODITIES PRESENTCU.COMMODITIES PRESENTCU.OPT. CU.NINER COLSPANOPT. CU.NINER COLSPANOPT. CU.OPT. CU.NINER COLSPANOPT. CU.NINE COLSPANOPT. CU.INTER COLSPANOPT. CU.CU.CU.OPT. CU.OPT. CU.OPT. CU.INTER COLSPANOPT. 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Noranda Exploration Company, Limited (no personal liability) Unit 3A 1750 Quinn Street Prince George, B.C. V2N 1X3

# noranda

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24 January 1991

Phone (604) 562-0022 Fax (604) 562-5043

Gold Commissioner Omineca Mining Division Bag 2000 Smithers, B.C. VOJ 2N0

Re: Babine Lake Dighem Airborne EM Survey, Assessment Report

Dear Sir,

Enclosed please find two copies of Mr. Paul A. Smith's report on the Dighem IV Airborne Survey. I am submitting this report to you for assessment on various claims held by Noranda Minerals Inc. or Noranda Exploration Company, Limited (No Personal Liability) in the Babine Lake area of the Omineca Mining Division.

As only the north map sheet is being submitted for assessment purposes, references to results of the south sheet have been deleted from the text. No other changes have been made to Mr. Smith's report.

The airborne survey as described in Mr. Smith's report consisted of 1797 line-kilometers (p i.) of survey at a total contract cost of \$132,600. To this I have added an additional cost of \$2,900 as follows:

> report preparation \$600 supervision \$2300

This brings the total cost claimed to \$135,500.

Rather than estimate the line-km flown on each claim or claim group, I have planimetered the survey areas of the north and south map sheets and estimated the costs on a per area basis, as follows:

N Sheet	124.390 km*km
S Sheet	208.558 km*km
ABEM survey area	332.948 km*km

ON NORTH SHEET ONLY area flown on Morris Group claims (33 units) = 4.916 km\*km area flown on Bine 1-22 claims (173 units) = 35.992 km\*km area flown on LWO claims (74 units) = 4.692 km\*km area flown off claims = 78.79 km\*km

total survey, N sheet 124.390 km\*km

unit cost = \$135,500 / 332.948 km\*km = \$406.970 / km\*km
cost of N sheet = \$406.970 km\*km \* 124.390 km\*km = \$50,623
cost on Morris Group claims = \$ 2,000.66
cost on Bine 1-22 claims = \$ 14,647.67
cost on LWO claims = \$ 1,909.50
cost off claims = \$ 32,065.17
total cost, N sheet \$ 50,623.

These costs have been applied as follows to various claim groups:

0	Morri	s	Group	(7	Nov.	90)	=	\$ 6,400
0	Lake	Ν	Group	(30	Nov.	90)	Ξ	\$ 23,300
50	Lake	$\mathbf{S}$	Group	(30	Nov.	90)	=	\$ 20,900
:0	PAC						=	\$ 23
						total`		\$ 50,623

These costs do not exceed the total cost of the N sheet. The areas applied do not exceed ten times the area of any claim group.

Please note that all but three flight lines were flown on or after 14 July 1990. The area flown on 14 July did not include any ground staked by Noranda that day.

A map is attached showing the location of the claims and the survey area.

I believe the above apportionment of costs to the various claim groups created in the ABEM survey area to be in accordance with the assessment report guidelines.

Sincerely,

Milling + Emp

Delbert E. Myers, Jr., FGAC Senior Project Geologist

encl.

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#### SUMMARY

This report describes the logistics and results of a DIGHEM<sup>IV</sup> airborne geophysical survey carried out for Noranda Exploration Company, Limited, over a property located in the Babine Lake area of British Columbia. Total coverage of the survey block amounted to 1797 km. The survey was flown from July 11 to July 22, 1990.

The purpose of the survey was to detect zones of conductive mineralization and to provide information that could be used to map the geology and structure of the survey This was accomplished by using a DIGHEM<sup>IV</sup> multiarea. coil, multi-frequency electromagnetic system, supplemented by a high sensitivity Cesium magnetometer and a two-channel VLF The information from these sensors was processed receiver. to produce maps which display the magnetic and conductive properties of the survey area. An electronic navigation system, operating in the UHF band, ensured accurate positioning of the geophysical data with respect to the base Visual flight path recovery techniques were used in maps. areas where transponder signals were blocked by topographic features.

Several bedrock conductors were detected on the property, most of which are considered to be of moderate to high priority as exploration targets. Most of these inferred bedrock conductors appear to warrant further investigation using appropriate surface exploration techniques. Areas of interest may be assigned priorities on the basis of supporting geophysical, geochemical and/or geological information. After initial investigations have been carried out, it may be necessary to re-evaluate the remaining anomalies based on information acquired from the follow-up program.



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- A. List of Personnel
- B. Statement of Cost
- C. EM Anomaly List

#### INTRODUCTION

A DIGHEM<sup>IV</sup> electromagnetic/resistivity/magnetic/VLF survey was flown for Noranda Exploration Company, Limited from July 11 to July 22, 1990, over a survey block in central British Columbia. The property, designated the Babine Lake Project, abuts the eastern shore of Babine Lake. The survey area can be located on NTS map sheets 93L/16 and 93M/1, (See Figure 1).

Survey coverage consisted of approximately 1797 line-km, including tie lines. Flight lines were flown in an azimuthal direction of 070° with a line separation of 200 metres.

The survey employed the DIGHEM electromagnetic system. Ancillary equipment consisted of a magnetometer, radar altimeter, video camera, analog and digital recorders, a VLF receiver and an electronic navigation system. Details on the survey equipment are given in Section 2.

The instrumentation was installed in an Aerospatiale AS350B turbine helicopter (Registration C-GNIX) which was provided by Questral Helicopters Ltd. The helicopter flew at an average airspeed of 120 km/h with an EM bird height of approximately 30 m. Section 2 also provides details on the data channels, their respective sensitivities, and the navigation/flight path recovery procedure. Noise levels of less than 2 ppm are generally maintained for wind speeds up to 35 km/h. Higher winds may cause the system to be grounded because excessive bird swinging produces difficulties in flying the helicopter. The swinging results from the 5  $m^2$  of area which is presented by the bird to broadside gusts.

The southwestern portion of the property contains numerous sources of cultural interference, which are related to the current operations of the Bell Copper mine. The mining operations cover a large portion of the Newman Peninsula, and have affected the geophysical responses in this area. There are many anomalies in the southwestern portion which have been attributed to possible bedrock sources, but which may be due to culture. Detailed follow-up will be required to confirm conductor sources in the presence of cultural features.

### SURVEY EQUIPMENT

This section provides a brief description of the geophysical instruments used to acquire the survey data:

Electromagnetic System

Model:

IV DIGHEM

Type: Towed. bird, symmetric dipole configuration operated at a nominal survey altitude of 30 metres. Coil separation is 8 metres for 900 Hz and 7200 Hz, and 6.3 metres for the 56,000 Hz coil-pair.

Coil	orientations/frequencies:	coaxial /	/ 900	Hz
	, -	coplanar/	/ 900	Hz
		coplanar	7,200	Hz
		coplanar/	/56,000	Hz

Channels	recorded:	4	inphase channels
		4	quadrature channels
		2	monitor channels

Sensitivity: 0.2 ppm at 900 Hz 0.4 ppm at 7,200 Hz 1.0 ppm at 56,000 Hz

Sample rate:

10 per second

The electromagnetic system utilizes a multi-coil coaxial/coplanar technique to energize conductors in different directions. The coaxial transmitter coil is vertical with its axis in the flight direction. The coplanar coils are horizontal. The secondary fields are sensed simultaneously by means of receiver coils which are maximum coupled to their respective transmitter coils. The system yields an inphase and a quadrature channel from each transmitter-receiver coil-pair.

#### Magnetometer

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Model: Picodas 3340 Type: Optically pumped Cesium vapour Sensitivity: 0.01 nT Sample rate: 10 per second

The magnetometer sensor is towed in a bird 15 m below the helicopter.

### Magnetic Base Station

Model:	Scintrex MP-3
Type:	Digital recording proton precession
Sensitivity:	0.10 nT
Sample rate:	0.2 per second

A digital recorder is operated in conjunction with the base station magnetometer to record the diurnal variations of the earth's magnetic field. The clock of the base station is synchronized with that of the airborne system to permit subsequent removal of diurnal drift.

## VLF System

Manufacturer: Herz Industries Ltd. Type: Totem-2A Sensitivity: 0.1% Stations: Seattle, Washington; NLK, 24.8 kHz Lualualei, Hawaii; NPM, 23.4 kHz

The VLF receiver measures the total field and vertical quadrature components of the secondary VLF field. Signals from two separate transmitters can be measured simultaneously. The VLF sensor is towed in a bird 10 m below the helicopter.

### Radar Altimeter

Manufacturer: Honeywell/Sperry Type: AA 220 Sensitivity: 1 ft

The radar altimeter measures the vertical distance between the helicopter and the ground. This information is used in the processing algorithm which determines conductor depth.

## Analog Recorder

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Manufacturer: RMS Instruments

Туре:	DGR33 dot-matrix graphics recorder
Resolution:	4x4 dots/mm
Speed:	1.5 mm/sec

The analog profiles were recorded on chart paper in the aircraft during the survey. Table 2-1 lists the geophysical data channels and the vertical scale of each profile.

# Digital Data Acquisition System

Manufacturer:	RMS	Instrume	ents				
Туре:	DGR	33					
Tape Deck:	RMS	TCR-12,	6400	bpi,	tape	cartridge	recorder

The digital data were used to generate several computed parameters. Both measured and computed parameters were plotted as "multi-channel stacked profiles" during data processing. These parameters are shown in Table 2-2.

In Table 2-2, the log resistivity scale of 0.06 decade/mm means that the resistivity changes by an order of magnitude in 16.6 mm. The resistivities at 0, 33 and 67 mm up from the bottom of the digital profile are respectively 1, 100 and 10,000 ohm-m.

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Channel NameParameterScale units/mmDesignation on digital profile1X91coaxial inphase (900 Hz)2.5 ppmCXI (900 Hz)1X92coaxial quad (900 Hz)2.5 ppmCXI (900 Hz)3P91coplanar inphase (900 Hz)2.5 ppmCPI (900 Hz)3P92coplanar quad (900 Hz)2.5 ppmCPQ (900 Hz)3P92coplanar quad (700 Hz)5 ppmCPQ (900 Hz)2P71coplanar inphase (7200 Hz)5 ppmCPQ (7200 Hz)2P72coplanar quad (7200 Hz)5 ppmCPQ (7200 Hz)4P51coplanar inphase(56000 Hz)10 ppmCPQ (56 kHz)4P52coplanar quad (56000 Hz)10 ppmCPQ (56 kHz)ALTRaltimeter3 mALTCMGCmagnetics, coarse25 nTMAGCMGFmagnetics, fine2.5 nTMAGVF10VLF-total: primary stn.2%VF2VF20VLF-quad: primary stn.2%CXSVF20VLF-quad: secondary stn.2%CXSCXSPcoaxial sferics monitorCXSCXPLcoaxial powerline monitorCXS				
1X9Icoaxial inphase (900 Hz)2.5 ppmCXI (900 Hz)1X9Qcoaxial quad (900 Hz)2.5 ppmCXQ (900 Hz)3P9Icoplanar inphase (900 Hz)2.5 ppmCPI (900 Hz)3P9Qcoplanar quad (900 Hz)2.5 ppmCPQ (900 Hz)3P9Qcoplanar quad (700 Hz)5 ppmCPQ (900 Hz)2P7Icoplanar inphase (7200 Hz)5 ppmCPQ (7200 Hz)2P7Qcoplanar quad (7200 Hz)5 ppmCPQ (7200 Hz)4P5Icoplanar quad (56000 Hz)10 ppmCPQ (56 kHz)4P5Qcoplanar quad (56000 Hz)10 ppmCPQ (56 kHz)ALTRaltimeter3 mALTTCMGCmagnetics, coarse25 nTMAGCMGFmagnetics, fine2.5 nTMAGVF1QVLF-quad: primary stn.2%VF2QVF2TVLF-total: secondary stn.2%CXSVF2QVLF-quad: secondary stn.2%CXSCXSPcoaxial sferics monitorCXSCXSPcoaxial powerline monitorCXS	Channel Name	Parameter	Scale units/mm	Designation on digital profile
CPPL coplanar powerline monitor CPP	1X91 1X9Q 3P91 3P9Q 2P71 2P7Q 4P51 4P5Q ALTR CMGC CMGF VF1T VF1Q VF1T VF1Q VF2T VF2Q CXSP CPSP CXPL CPPL	coaxial inphase (900 Hz) coaxial quad (900 Hz) coplanar inphase (900 Hz) coplanar quad (900 Hz) coplanar quad (900 Hz) coplanar quad (7200 Hz) coplanar quad (7200 Hz) coplanar quad (56000 Hz) coplanar quad (56000 Hz) altimeter magnetics, coarse magnetics, fine VLF-total: primary stn. VLF-quad: primary stn. VLF-quad: primary stn. VLF-quad: secondary stn. VLF-quad: secondary stn. coaxial sferics monitor coplanar sferics monitor coplanar powerline monitor	2.5 ppm 2.5 ppm 2.5 ppm 2.5 ppm 5 ppm 10 ppm 10 ppm 3 m 25 nT 2.5 nT 2.5 nT 2% 2% 2% 2%	CXI ( 900 Hz) CXQ ( 900 Hz) CPI ( 900 Hz) CPQ ( 900 Hz) CPQ ( 7200 Hz) CPQ (7200 Hz) CPQ (56 kHz) CPQ (56 kHz) ALT MAG

Table	2-1.	The	Analo	y Protiles
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Table 2-2. The Digital Profiles

Channel <u>Name (Freq)</u>	Observed parameters	Scale <u>units/mm</u>
MAG ALT CXI (900 Hz) CXQ (900 Hz) CPI (900 Hz) CPQ (900 Hz) CPI (7200 Hz) CPQ (7200 Hz) CPQ (7200 Hz) CPI (56 kHz) CPQ (56 kHz) CXS CPP	magnetics bird height vertical coaxial coil-pair inphase vertical coaxial coil-pair quadrature horizontal coplanar coil-pair inphase horizontal coplanar coil-pair quadrature horizontal coplanar coil-pair inphase horizontal coplanar coil-pair quadrature horizontal coplanar coil-pair inphase horizontal coplanar coil-pair inphase horizontal coplanar coil-pair guadrature coaxial sferics monitor coplanar powerline monitor	10 nT 6 m 2 ppm 2 ppm 2 ppm 2 ppm 4 ppm 4 ppm 10 ppm 10 ppm
	Computed Parameters	
DFI (900 Hz) DFQ (900 Hz) RES (900 Hz) RES (7200 Hz) RES (56 kHz) DP (900 Hz) DP (7200 Hz) DP (56 kHz) CDT	difference function inphase from CXI and CPI difference function quadrature from CXQ and CPQ log resistivity log resistivity log resistivity apparent depth apparent depth conductance	2 ppm 2 ppm .06 decade .06 decade .06 decade 6 m 6 m 1 grade

# Tracking Camera

Type: Panasonic Video

Model: AG 2400/WVCD132

Fiducial numbers are recorded continuously and are displayed on the margin of each image. This procedure ensures accurate correlation of analog and digital data with respect to visible features on the ground.

### Navigation System

Model:	Del Norte 547
Туре:	UHF electronic positioning system
Sensitivity:	1 m
Sample rate:	2 per second

The navigation system uses ground based transponder stations which transmit distance information back to the helicopter. The ground stations are set up well away from the survey area and are positioned such that the signals cross the survey block at an angle between 30° and 150°. The onboard Central Processing Unit takes any two transponder distances and determines the helicopter position relative to these two ground stations in cartesian coordinates. The cartesian coordinates are transformed to UTM coordinates during data processing. This is accomplished by correlating a number of prominent topographical locations with the navigational data points. The use of numerous visual tie points serves two purposes: to accurately relate the navigation data to the map sheet and to minimize location errors which might result from distortions in uncontrolled photomosaic base maps.

#### PRODUCTS AND PROCESSING TECHNIQUES

The following products are available from the survey data. Those which are not part of the survey contract may be acquired later. Refer to Table 3-1 for a summary of the maps which accompany this report, some of which may be sent under separate cover. Most parameters can be displayed as contours, profiles, or in colour.

#### Base Maps

Base maps of the survey area have been produced from published topographic maps. These provide a relatively accurate, distortion-free base which facilitates correlation of the navigation data to the UTM grid. Photomosaics are useful for visual reference and for subsequent flight path recovery, but usually contain scale distortions. Orthophotos are ideal, but their cost and the time required to produce them, usually precludes their use as base maps.

## Electromagnetic Anomalies

Anomalous electromagnetic responses are selected and analysed by computer to provide a preliminary electromagnetic anomaly map. This preliminary map is used, by the

Table 3-1	Plots	Available	from	the	Survey	
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MAP PRODUCT	NO. OF SHEETS	ANOMALY MAP	PROFILES ON MAP	CON INK	IOURS COLOUR	SHADOW MAP
Electromagnetic Anomalies	2	20,000	N/A	N/A	N/A	N/A
Probable Bedrock Conductors		-	N/A	N/A	N/A	N/A
Resistivity ( 900 Hz)	2	N/A	-	20,000	20,000	-
Resistivity ( 7,200 Hz)	2	N/A	-	20,000	20,000	-
Resistivity (56,000 Hz)	2	N/A	-	20,000	20,000	-
EM Magnetite		N/A	-	-	-	-
Total Field Magnetics	2	N/A	-	20,000	20,000	-
Enhanced Magnetics		N/A	-	-	-	-
1st Vertical Derivative Magnet	ics 2	N/A	-	20,000	20,000	-
2nd Vertical Derivative Magnet	lics	N/A	-	-	-	_
Filtered Total Field VLF	2	N/A	-	20,000	20,000	-
VLF Profiles		N/A	-	N/A	N/A	N/A
Electromagnetic Profiles( 900	Hz)	N/A	-	N/A	N/A	N/A
Electromagnetic Profiles(7200	Hz)	N/A	-	N/A	N/A	N/A
Multi-channel stacked profiles		Worksheet profiles				-
		Interpre	ted profil	les		10,000

N/A Not available

- Not required under terms of the survey contract

\* Recommended

20,000 Scale of delivered map, i.e, 1:20,000

Notes:

- Inked contour maps are provided on transparent media and show flight lines, EM anomalies and suitable registration. Three paper prints of each map are supplied, in addition to three colour plots of each contour map.

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geophysicist, in conjunction with the computer-generated digital profiles, to produce the final interpreted EM anomaly map. This map includes bedrock, surficial and cultural conductors. A map containing only bedrock conductors can be generated, if desired.

#### Resistivity

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The apparent resistivity in ohm-m may be generated from the inphase and quadrature EM components for any of the frequencies, using a pseudo-layer halfspace model. A resistivity map portrays all the EM information for that frequency over the entire survey area. This contrasts with the electromagnetic anomaly map which provides information only over interpreted conductors. The large dynamic range makes the resistivity parameter an excellent mapping tool.

#### EM Magnetite

The apparent percent magnetite by weight is computed wherever magnetite produces a negative inphase EM response.

#### Total Field Magnetics

The aeromagnetic data are corrected for diurnal

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variation using the magnetic base station data. The regional IGRF can be removed from the data, if requested.

# Enhanced Magnetics

The total field magnetic data are subjected to a processing algorithm. This algorithm enhances the response of magnetic bodies in the upper 500 m and attenuates the response of deeper bodies. The resulting enhanced magnetic map provides better definition and resolution of nearsurface magnetic units. It also identifies weak magnetic features which may not be evident on the total field magnetic map. However, regional magnetic variations, and magnetic lows caused by remanence, are better defined on the total field magnetic map. The technique is described in more detail in Section 5.

## Magnetic Derivatives

The total field magnetic data may be subjected to a variety of filtering techniques to yield maps of the following:

vertical gradient

second vertical derivative

magnetic susceptibility with reduction to the pole upward/downward continuations

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All of these filtering techniques improve the recognition of near-surface magnetic bodies, with the exception of upward continuation. Any of these parameters can be produced on request. Dighem's proprietary enhanced magnetic technique is designed to provide a general "all-purpose" map, combining the more useful features of the above parameters.

#### VLF

The VLF data are digitally filtered to remove long wavelengths such as those caused by variations in the transmitted field strength.

# Multi-channel Stacked Profiles

Distance-based profiles of the digitally recorded geophysical data are generated and plotted by computer. These profiles also contain the calculated parameters which are used in the interpretation process. These are produced as worksheets prior to interpretation, and can also be presented in the final corrected form after interpretation. The profiles display electromagnetic anomalies with their respective interpretive symbols. The differences between the worksheets and the final corrected form occur only with respect to the EM anomaly identifier.

# Contour, Colour and Shadow Map Displays

The geophysical data are interpolated onto a regular grid using a modified Akima spline technique. The resulting grid is suitable for generating contour maps of excellent quality.

Colour maps are produced by interpolating the grid down to the pixel size. The parameter is then incremented with respect to specific amplitude ranges to provide colour "contour" maps. Colour maps of the total magnetic field are particularly useful in defining the lithology of the survey area.

Monochromatic shadow maps are generated by employing an artificial sun to cast shadows on a surface defined by the geophysical grid. There are many variations in the shadowing technique. These techniques may be applied to total field or enhanced magnetic data, magnetic derivatives, VLF, resistivity, etc. Of the various magnetic products, the shadow of the enhanced magnetic parameter is particularly suited for defining geological structures with crisper images and improved resolution.

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#### SURVEY RESULTS

# GENERAL DISCUSSION

The survey results are presented on 2 separate map sheets for each parameter at a scale of 1:20,000. Table 4-1 summarizes the EM responses in the survey area, with respect to conductance grade and interpretation.

The anomalies shown on the electromagnetic anomaly maps are based on a near-vertical, half plane model. This model best reflects "discrete" bedrock conductors. Wide bedrock conductors or flat-lying conductive units, whether from surficial or bedrock sources, may give rise to very broad anomalous responses on the EM profiles. These may not appear on the electromagnetic anomaly map if they have a regional character rather than a locally anomalous character. These broad conductors, which more closely approximate a half space model, will be maximum coupled to the horizontal (coplanar) coil-pair and should be more evident on the resistivity parameter. Resistivity maps, therefore, may be more valuable than the electromagnetic anomaly maps, in areas where broad or flat-lying conductors are considered to be of Contoured resistivity maps, based on the 900, importance. 7200 and 56,000 Hz coplanar data are included with this report.

# TABLE 4-1

# EM ANOMALY STATISTICS

# BABINE LAKE PROJECT

CONDUCTOR	CONDUCTANCE RANGE	NUMBER OF
GRADE	SIEMENS (MHOS)	RESPONSES
7	> 100.0	2
6	50 - 100.0	2
5	20 - 50.0	14
4	10 - 20.0	39
3	5 - 10.0	135
2	1 - 5.0	637
1	< 1.0	220
*	INDETERMINATE	485
TOTAL		1534

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CONDUCTOR MODEL	MOST LIKELY SOURCE	NUMBER OF RESPONSES
D	DISCRETE BEDROCK CONDUCTOR	28
В	CONDUCTIVE COVER	208
S	CONDUCTIVE COVER	1050
Н	ROCK UNIT OR THICK COVER	122
Е	EDGE OF WIDE CONDUCTOR	33
L	CULTURE	93
TOTAL		1534

# (SEE EM MAP LEGEND FOR EXPLANATIONS)

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Excellent resolution and discrimination of conductors was accomplished by using a fast sampling rate of 0.1 sec and by employing a common frequency (900 Hz) on two orthogonal coil-pairs (coaxial and coplanar). The resulting "difference channel" parameters often permit differentiation of bedrock and surficial conductors, even though they may exhibit similar conductance values.

Anomalies which occur near the ends of the survey lines (i.e., outside the survey area), should be viewed with caution. Some of the weaker anomalies could be due to aerodynamic noise, i.e., bird bending, which is created by abnormal stresses to which the bird is subjected during the climb and turn of the aircraft between lines. Such aerodynamic noise is usually manifested by an anomaly on the coaxial inphase channel only, although severe stresses can affect the coplanar inphase channels as well.

#### Magnetics

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A Scintrex MP-3 proton precession magnetometer was operated at the survey base to record diurnal variations of the earth's magnetic field. The clock of the base station

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was synchronized with that of the airborne system to permit subsequent removal of diurnal drift.

The background magnetic levels have been adjusted to match the International Geomagnetic Reference Field (IGRF) for the survey area. The IGRF gradient across the survey block is left intact. This procedure ensures that the magnetic contours will match contours from any adjacent surveys which have been processed in a similar manner.

The total field magnetic data have been presented as contours on the base maps using a contour interval of 5 nT where gradients permit. The maps show the magnetic properties of the rock units underlying the survey area.

The total field magnetic data have been subjected to a processing algorithm to produce maps of the calculated first vertical derivative. This procedure enhances near-surface magnetic units and suppresses regional gradients. It also provides better definition and resolution of magnetic units and displays weak magnetic features which may not be clearly evident on the total field maps. Maps of the second vertical magnetic derivative can also be prepared from existing survey data, if requested.

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There is some evidence on the magnetic maps which suggests that the survey area has been subjected to deformation and/or alteration. These structural complexities are evident on the contour maps as variations in magnetic intensity, irregular patterns, and as offsets or changes in strike direction. Some of the more prominent linear features are also evident on the topographic base maps.

The magnetic relief on the Babine Lake Project is quite high, ranging from a low of about 57,250 nT near 10350C to a high of more than 60,800 nT on line 10551 near fiducial 1610.

The southern half of the survey block (Sheet 1) generally exhibits lower magnetic values than are evident on Sheet 2, to the north. The individual units which comprise the magnetic trends are more clearly defined on the vertical gradient maps. The latter product also identifies several faults, in addition to the geological contacts.

On the south sheet, the inferred geological strikes vary from 330° to 020°. In the north, trends usually favour directions between 310° and 360°. It is interesting to note that the rock units underlying the Newman Peninsula are of relatively low magnetic intensity, usually less than 58000

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nT. The large open pit near the western end of line 10540, however, correlates with a moderately weak, but well-defined circular magnetic anomaly about 100 nT above the background. This may be a signification factor in the search for similar copper mineralization in the area.

If a specific magnetic intensity can be assigned to the rock type which is believed to host the target mineralization, it may be possible to select areas of higher priority on the basis of the total field magnetic data. This is based on the assumption that the magnetite content of the host rocks will give rise to a limited range of contour values which will permit differentiation of various lithological units.

The magnetic results, in conjunction with the other geophysical parameters, should provide valuable information which can be used to effectively map the geology and structure in the survey area.

#### VLF

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VLF results were obtained from the transmitting station at Lualualei, Hawaii (NPM - 23.4 kHz) and Seattle, Washington (NLK - 24.8 kHz). The VLF maps show the contoured results of

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the filtered total field from Seattle. Gaps in coverage are evident between lines 10060-10130, 10470-10591 and 11100-11120. The noise spikes at the extreme western end of line 11100 and the eastern end of line 10590 should be disregarded.

The VLF method is quite sensitive to the angle of coupling between the conductor and the propogated EM field. Consequently, conductors which strike towards the VLF station will usually yield a stronger response than conductors which are nearly orthogonal to it. The general north-northwest/south-southeast strike in the survey area provides good coupling with the VLF field from Seattle.

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The VLF parameter does not normally provide the same degree of resolution available from the EM data. Closelyspaced conductors, conductors of short strike length or conductors which are poorly coupled to the VLF field, may escape detection with this method. Erratic signals from the VLF transmitters can also give rise to strong, isolated anomalies which should be viewed with caution. Regardless of these limitations, however, the VLF results have provided some additional information, particularly within the more resistive portions of the survey area. The VLF method could probably be used as a follow-up tool in most areas, although its effectiveness will be somewhat limited in areas of moderate to high conductivity. The filtered total field VLF contours are presented on the base maps with a contour interval of one percent.

#### <u>Resistivity</u>

Resistivity maps, which display the conductive properties of the survey area, were produced from the 900 Hz, 7200 Hz and 56,000 Hz coplanar data. In general, the resistivity patterns show moderately good agreement with the magnetic trends. This suggests that many of the resistivity lows are probably related to bedrock features, rather than conductive overburden. There are some areas, however, where resistivity contour patterns are obviously influenced by surficial conductivity and/or cultural sources.

The central core of the property is dominated by moderately high resistivities of more than 1000 ohm-m. This major resistivity high correlates very closly with the stronger magnetic units. Most of the zones which yield resistivities of more than 1000 ohm-m (on the 7200 Hz resistivity map) are associated with magnetic units yielding values of more than 58,250 nT.

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There is a well-defined resistivity gradient which separates the central resistive unit from a highly conductive zone to the west. The 500 ohm-m resistivity contour approximates the limit of this conductive zone. The gradient extends in a southeasterly direction from the northwest corner of the grid (at the eastern shore of Morrison Lake), along the western contact of the magnetic high to line 11170.

From this point, it swings south-southeast to line 10891 where it is offset to the west before continuing in a southeasterly direction to line 10370. At this point, it meets a second resistivity gradient which strikes almost north/south. A third resistivity gradient extends in a south-southwesterly direction from the eastern end of line 10580, forming a 'V'shaped conductive zone at the southcentral portion of the grid.

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One would normally expect to see a fairly consistent relationship between water-covered or low-lying areas and lower resistivities. This does not always appear to be the case on this property. In some areas the land portions are more conductive than the lake areas.

Note, for example, the small island at the intersection of line 10100 and the central tie-line 19020. Resistivities

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on the island are considerably less than those in the surrounding lake.

In addition to the broad resistivity lows on the property there are several moderately strong, well-defined conductive zones of limited extent. The former have been attributed to conductive rock units while the latter are more likely due to discrete conductors.

## Electromagnetics

The EM anomalies resulting from this survey appear to fall within one of four general categories. The first type consists of discrete, well-defined anomalies which yield marked inflections on the difference channels. These anomalies are usually attributed to conductive sulphides or graphite and are generally given a "B", "T" or "D" interpretive symbol, denoting a bedrock source.

The second class of anomalies comprises moderately broad responses which exhibit the characteristics of a half space and do not yield well-defined inflections on the difference channels. Anomalies in this category are usually given an "S" or "H" interpretive symbol. The lack of a difference channel response usually implies a broad or flat-lying conductive source such as overburden. Some of these anomalies may reflect conductive rock units or zones of deep weathering.

The third class consists of moderately broad quadrature anomalies which are associated with negative inphase responses. The negative inphase is due to magnetite. The positive quadrature can be due to conductive material which overlies, or is contained within, the magnetite-rich rock unit. Most of these anomalies have been given an "S?" or "B?" interpretive symbol.

The fourth anomaly type comprises those which have been attributed to culture. These are prevalent in the southwestern portion of the Newman Peninsula, where most of the mining activity is concentrated.

In areas where EM responses are evident primarily on the quadrature components, zones of poor conductivity are indicated. Where these responses are coincident with magnetic anomalies, it is possible that the inphase component amplitudes have been suppressed by the effects of magnetite. Most of these poorly-conductive magnetic features give rise to resistivity anomalies which are only slightly below background. If it is expected that poorly-conductive economic mineralization may be associated with magnetite-rich units, most of these weakly anomalous features will be of interest. In areas where magnetite causes the inphase components to become negative, the apparent conductance and depth of EM anomalies may be unreliable.

The effects of conductive overburden are evident over portions of the survey area. Although the difference channels (DFI and DFQ) are extremely valuable in detecting bedrock conductors which are partially masked by conductive overburden, sharp undulations in the bedrock/overburden interface can yield anomalies in the difference channels which may be interpreted as possible bedrock conductors. anomalies usually fall into the "S?" or "B?" Such classification but may also be given an "E" interpretive symbol, denoting a resistivity contrast at the edge of a conductive unit.

As economic mineralization within the area may consist of massive to weakly disseminated sulphides, which may or may not be hosted by magnetite-rich rocks, it is difficult to assess the relative merits of EM anomalies on the basis of conductance. It is recommended that an attempt be made to compile a suite of geophysical "signatures" over areas of interest. Anomaly characteristics are clearly defined on the

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computer-processed geophysical data profiles which are supplied as one of the survey products.

A complete assessment and evaluation of the survey data should be carried out by one or more qualified professionals who have access to, and can provide a meaningful compilation of, all available geophysical, geological and geochemical data.

## CONDUCTORS IN THE SURVEY AREA

The electromagnetic anomaly maps show the anomaly locations with the interpreted conductor type, dip, conductance and depth being indicated by symbols. Direct magnetic correlation is also shown if it exists. The strike direction and length of the conductors are indicated when anomalies can be correlated from line to line. When studying the map sheets, consult the anomaly listings appended to this report.

<u>Sheet 1</u>

Conductors 10070D-10090E, 10090G-10110E,

#### 10140J-101500

Three weak conductors of possible bedrock origin are associated with islands in the southern portion of
the survey grid. Conductor 10070D-10090E is very close to a road on Sterrett Island. However, its significance is enhanced by its association with the contact of a small, but moderately strong, magnetic anomaly. There is no VLF coverage in this area.

Conductor 10090G-10110E extends from Sterrett Island into the lake. Unlike the previous conductor, this one gives rise to a strong resistivity low. It also appears to be related to the contact of a weak magnetic anomaly and is also situated in the area where no VLF coverage was obtained. This conductor is about 1 km south-southeast of the McDonald mine.

Conductor 10140J-101500 consists of two poorly defined responses located in Hawthorn Bay. They are non-magnetic, but the conductor is associated with a moderately strong magnetic unit, in the vicinity of a possible fold. The conductor is contained within a broad resistivity low near the intersection of NNW and NNE trending linears seen on the resistivity maps.

Conductors 10160C-10180B, 10220B-10230C, 10240E-10260G, 10280F-10300F, 10300E-10310C, 10460F-105000,

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10550D-10560E, 10560D-10570E, 10560C-10581D, 10600B-10620B, 10600C-10620C.

The conductors in this group are all located on the Newman Peninsula. Because of the numerous cultural features in this active mining area, nearly all of the anomalies have been given a 'B?' interpretive symbol. Some, or all of these, could be influenced by culture. Some can probably be eliminated by carrying out visual checks on the ground. Others will require careful follow-up with an appropriate system.

The significance of these conductors is obviously enhanced by their location. All are associated with resistivity lows yielding values of less than 100 ohm-m. Most yield partial weak magnetic correlation along strike although the entire peninsula is underlain by relatively non-magnetic rocks. All conductors show direct or flanking association with VLF anomalies.

In addition to the conductors listed above there are several other possible bedrock conductors which are evident on one line only. Most of these are also considered to be of interest as exploration targets. In

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many cases, the 'B?' anomalies occur within trends of 'S?'anomalies, suggesting the presence of bedrock conductors beneath conductive surficial material such as mine waste.

All of the anomalous responses on the Newman Peninsula are considered to be potential targets unless they can definitely be confirmed as being due to culture or overburden.

#### Conductors 10210K-10230I, 10220I-10230K,

#### 10230J-10260M

These three poorly-defined conductors are located in the northern part of Hawthorn Bay and are probably due mainly to conductive lake sediments. However, they are associated with the northern limb of a complex magnetic anomaly in an area of possible folding.

Amplitudes are moderately strong, with conductor 10210K-10230I yielding conductance grades of 4. Conductor 10230J-10260M also shows weak magnetic correlation at its south end.

# Conductors 19020C-10360F, 10380G-10390G,

10470G-10571H.

The first two conductors in this group are located near the western shore of Hagen Arm. They are contained within a strong, but fairly broad resistivity low which also correlates with a magnetic low. The southern conductor, 19020C-10360F does not correspond to a VLF anomaly, but this may be because of the apparent change in strike. This conductor exhibits a probable strike to the west-northwest, which coincides with a magnetic gradient in the same area. This conductor may be related to a faulted contact.

## Conductors 10340G-10350G, 10340H-10350H, 10340I-10380J, 10380I-10400K

The first two conductors in this group are of very limited strike length. The strike directions are dubious and may be towards the northwest, rather than north/south as shown on the EM map. Anomaly 10360I is possibly related to the strongest anomaly, 10350G. This conductor may be deep, but is likely due to a narrow source. Note the strong VLF correlation. Conductors 10340I-10380J and 10380I-10400K may reflect a single conductor. This very weak zone is of particular interest because of its apparent strike towards the west-northwest. There is no coincident VLF response but the conductors follow a weak resisitivity low which correlates with a break in the magnetic contour patterns.

Conductors 10380K-10421E, 10380L-10421F

These two conductors are parallel, striking north/south along the eastern face of a hill. They are separated by about 250 m, each exhibiting a probable strike length of approximately 750 m. Both reflect narrow, east-dipping bedrock sources of moderately strong conductance. Only one reponse yields a weak magnetic correlation of 7 nT. Both zones are nonmagnetic, but are situated on the flanks of a very weak magnetic response. These conductors give rise to a well-defined resistivity low and associated VLF responses.

This is considered to be a high priority target. The sharp resistivity and magnetic contrasts immediately west of 10380K - 10421E, suggest that these two conductors are likely in close proximity to a major contact. Additional work is recommended to check the source of these two conductors.

## Conductors 10360N-10380P, 103800-10421L, 10390P-

10410Q, 10421K, 10440M, 10440N, 10440O.

The first three conductors in this group reflect narrow bedrock sources which strike approximately north/south and dip to the west. This dip is opposite to that indicated by the two conductors to the west, suggesting the presence of a synclinal fold in the area.

The axis of conductor 103800-10421L shows partial correlation with a resistivity low, but appears to transect the local magnetic trends at a shallow angle. The VLF correlation is poor. The conductor axis follows a swampy area north of a small lake, and may therefore be partially masked by conductive overburden. Further work is recommended to check the source of these conductors as well as the isolated responses on line 10440. At least two of these conductors may extend to the north, beyond the survey limits.

## Conductors 10480M-10550J, 10490K-10583H,

10560J-10570N

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Conductor 10480M-10550J contains the strongest anomalies on the grid. Conductance grades of 7 are observed on line 10540 and 10591. The strong anomalies reflect a narrow, east-dipping, magnetic conductor which strikes about 10° east of north. The presence of two closely-spaced conductors on line 10583 indicates a second conductor or a folded portion of a single conductor.

The magnetic correlation suggests that pyrrhotite may be a contributing factor on at least two lines. The conductor follows a moderately strong magnetic anomaly which may be offset by a dextral fault trending east/west through anomaly 10490J. The strong resistivity low makes this one of the more attractive targets on the grid. Unfortunately, this conductor is also located on lines over which no VLF information was acquired.

Conductor 10490K-10583H to the east, and the very weak 10560J-10570N, to the north, are both located in relatively non-magnetic rock units. These may reflect

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structural breaks beneath swampy cover, and should also be subjected to further investigation.

Conductors 10790E-10820B, 10870C-10880D

These two weak conductors are located in the northwestern portion of sheet 1. Conductor 10790E-10820B is situated in the saddle of double-lobed magnetic anomaly. It is associated with a strong resistivity low but there is no VLF correlation.

The magnetic amplitudes are similar to those observed over the open pit operation of the Newman Peninsula to the south. Therefore, this conductor could be due to similar mineralization, and should be investigated.

The more northerly conductor, 10870C-10880D, is non-magnetic and poorly-defined, but is is also associated with a relatively strong resistivity low. This conductor is considered to be of relatively low priority.

#### BACKGROUND INFORMATION

This section provides background information on parameters which are available from the survey data. Those which have not been supplied as survey products may be generated later from raw data on the digital archive tape.

#### ELECTROMAGNETICS

DIGHEM electromagnetic responses fall into two general classes, discrete and broad. The discrete class consists of sharp, well-defined anomalies from discrete conductors such as sulfide lenses and steeply dipping sheets of graphite and sulfides. The broad class consists of wide anomalies from conductors having a large horizontal surface such as flatly dipping graphite or sulfide sheets, saline water-saturated sedimentary formations, conductive overburden and rock, and geothermal zones. A vertical conductive slab with a width of 200 m would straddle these two classes.

The vertical sheet (half plane) is the most common model used for the analysis of discrete conductors. All anomalies plotted on the electromagnetic map are analyzed according to this model. The following section entitled **Discrete Conductor Analysis** describes this model in detail, including the effect of using it on anomalies caused by broad conductors such as conductive overburden.

The conductive earth (half space) model is suitable for broad conductors. Resistivity contour maps result from the use of this model. A later section entitled **Resistivity Mapping** describes the method further, including the effect of using it on anomalies caused by discrete conductors such as sulfide bodies.

## Geometric interpretation

The geophysical interpreter attempts to determine the geometric shape and dip of the conductor. Figure 5-1 shows typical DIGHEM anomaly shapes which are used to guide the geometric interpretation.

#### Discrete conductor analysis

The EM anomalies appearing on the electromagnetic map are analyzed by computer to give the conductance (i.e., conductivity-thickness product) in siemens (mhos) of a vertical sheet model. This is done regardless of the interpreted geometric shape of the conductor. This is not an unreasonable procedure, because the computed conductance increases as the electrical quality of the conductor increases, regardless of its true shape. DIGHEM anomalies

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Conductor location	ł	ł	ł	ł	Ņ	ł	S <sub>1</sub> H E	ł
Channel CXI	$\bigwedge$	$\wedge$	$\wedge$	$\bigwedge$	介 -	$\sim$		
Channel CPI	$\sim$	$\mathcal{M}$	$\sim$	$\bigwedge$				
Channel DIFI	$\bigwedge$	$\bigwedge$	$\mathcal{N}$	$\bigvee$	$\sum$	$\sim$		
Conductor	<b>•</b>		$\setminus$		0			Lawrence and
	line	vertical thin dike	dipping thin dike	vertical or dipping thick dike	sphere; horizontal disk; metal roof; small fenced	wide horizontal ribbon; large fenced area	S = conductive overburden H = thick conductive cover or wide conductive rock unit E = edge effect from wide	flight line parallel to conductor
Patia at					yord	ureu	conductor	
amplitudes								
CXI / CPI	4/1	2/1	variable	variable	1/4	variable	1/2	< 1/4

Fig. 5-1 Typical DIGHEM anomaly shapes

are divided into seven grades of conductance, as shown in Table 5-1 below. The conductance in siemens (mhos) is the reciprocal of resistance in ohms.

Anomaly Grade	<u>siemens</u>		
7	> 100		
6	50 - 100		
5	20 - 50		
4	10 - 20		
3	5 - 10		
2	1 - 5		
1	< 1		

Table 5-1. EM Anomaly Grades

The conductance value is a geological parameter because it is a characteristic of the conductor alone. It generally is independent of frequency, flying height or depth of burial, apart from the averaging over a greater portion of the conductor as height increases. Small anomalies from deeply buried strong conductors are not confused with small anomalies from shallow weak conductors because the former will have larger conductance values.

-1

Conductive overburden generally produces broad EM responses which may not be shown as anomalies on the EM maps. However, patchy conductive overburden in otherwise resistive areas can yield discrete anomalies with a conductance grade (cf. Table 5-1) of 1, 2 or even 3 for conducting clays which have resistivities as low as 50 ohm-m. In areas where ground resistivities are below 10 ohm-m, anomalies caused by weathering variations and similar causes can have any conductance grade. The anomaly shapes from the multiple coils often allow such conductors to be recognized, and these are indicated by the letters S, H, and sometimes E on the electromagnetic anomaly map (see EM map legend).

For bedrock conductors, the higher anomaly grades indicate increasingly higher conductances. Examples: DIGHEM's New Insco copper discovery (Noranda, Canada) yielded a grade 5 anomaly, as did the neighbouring copper-zinc Magusi River ore body; Mattabi (copper-zinc, Sturgeon Lake, Canada) and Whistle (nickel, Sudbury, Canada) gave grade 6; and DIGHEM's Montcalm nickel-copper discovery (Timmins, Canada) yielded a grade 7 anomaly. Graphite and sulfides can span all grades but, in any particular survey area, field work may show that the different grades indicate different types of conductors.

Strong conductors (i.e., grades 6 and 7) are characteristic of massive sulfides or graphite. Moderate conductors (grades 4 and 5) typically reflect graphite or sulfides of a less massive character, while weak bedrock conductors (grades 1 to 3) can signify poorly connected graphite or heavily disseminated sulfides. Grades 1 and 2

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conductors may not respond to ground EM equipment using frequencies less than 2000 Hz.

The presence of sphalerite or gangue can result in ore deposits having weak to moderate conductances. As an example, the three million ton lead-zinc deposit of Restigouche Mining Corporation near Bathurst, Canada, yielded a well-defined grade 2 conductor. The 10 percent by volume of sphalerite occurs as a coating around the fine grained massive pyrite, thereby inhibiting electrical conduction.

Faults, fractures and shear zones may produce anomalies which typically have low conductances (e.g., grades 1 to 3). Conductive rock formations can yield anomalies of any conductance grade. The conductive materials in such rock formations can be salt water, weathered products such as clays, original depositional clays, and carbonaceous material.

On the interpreted electromagnetic map, a letter identifier and an interpretive symbol are plotted beside the EM grade symbol. The horizontal rows of dots, under the interpretive symbol, indicate the anomaly amplitude on the flight record. The vertical column of dots, under the anomaly letter, gives the estimated depth. In areas where anomalies are crowded, the letter identifiers, interpretive

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symbols and dots may be obliterated. The EM grade symbols, however, will always be discernible, and the obliterated information can be obtained from the anomaly listing appended to this report.

The purpose of indicating the anomaly amplitude by dots is to provide an estimate of the reliability of the conductance calculation. Thus, a conductance value obtained from a large ppm anomaly (3 or 4 dots) will tend to be accurate whereas one obtained from a small ppm anomaly (no dots) could be quite inaccurate. The absence of amplitude dots indicates that the anomaly from the coaxial coil-pair is 5 ppm or less on both the inphase and quadrature channels. Such small anomalies could reflect a weak conductor at the surface or a stronger conductor at depth. The conductance grade and depth estimate illustrates which of these possibilities fits the recorded data best.

Flight line deviations occasionally yield cases where two anomalies, having similar conductance values but dramatically different depth estimates, occur close together on the same conductor. Such examples illustrate the reliability of the conductance measurement while showing that the depth estimate can be unreliable. There are a number of factors which can produce an error in the depth estimate, including the averaging of topographic variations by the

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altimeter, overlying conductive overburden, and the location and attitude of the conductor relative to the flight line. Conductor location and attitude can provide an erroneous depth estimate because the stronger part of the conductor may be deeper or to one side of the flight line, or because it has a shallow dip. A heavy tree cover can also produce errors in depth estimates. This is because the depth estimate is computed as the distance of bird from conductor, minus the altimeter reading. The altimeter can lock onto the top of a dense forest canopy. This situation yields an erroneously large depth estimate but does not affect the conductance estimate.

Dip symbols are used to indicate the direction of dip of conductors. These symbols are used only when the anomaly shapes are unambiguous, which usually requires a fairly resistive environment.

A further interpretation is presented on the EM map by means of the line-to-line correlation of anomalies, which is based on a comparison of anomaly shapes on adjacent lines. This provides conductor axes which may define the geological structure over portions of the survey area. The absence of conductor axes in an area implies that anomalies could not be correlated from line to line with reasonable confidence.

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DIGHEM electromagnetic maps are designed to provide a correct impression of conductor quality by means of the conductance grade symbols. The symbols can stand alone with geology when planning a follow-up program. The actual conductance values are printed in the attached anomaly list for those who wish quantitative data. The anomaly ppm and depth are indicated by inconspicuous dots which should not distract from the conductor patterns, while being helpful to those who wish this information. The map provides an interpretation of conductors in terms of length, strike and dip, geometric shape, conductance, depth, and thickness. The accuracy is comparable to an interpretation from a high quality ground EM survey having the same line spacing.

The attached EM anomaly list provides a tabulation of anomalies in ppm, conductance, and depth for the vertical The EM anomaly list also shows the conductance sheet model. and depth for a thin horizontal sheet (whole plane) model, but only the vertical sheet parameters appear on the EM map. The horizontal sheet model is suitable for a flatly dipping thin bedrock conductor such as a sulfide sheet having a The list also shows the thickness less than 10 m. resistivity and depth for a conductive earth (half space) model, which is suitable for thicker slabs such as thick conductive overburden. In the EM anomaly list, a depth value of zero for the conductive earth model, in an area of thick cover, warns that the anomaly may be caused by conductive overburden.

Since discrete bodies normally are the targets of EM surveys, local base (or zero) levels are used to compute local anomaly amplitudes. This contrasts with the use of true zero levels which are used to compute true EM amplitudes. Local anomaly amplitudes are shown in the EM anomaly list and these are used to compute the vertical sheet parameters of conductance and depth. Not shown in the EM anomaly list are the true amplitudes which are used to compute the horizontal sheet and conductive earth parameters.

#### <u>Ouestionable Anomalies</u>

DIGHEM maps may contain EM responses which are displayed as asterisks (\*). These responses denote weak anomalies of indeterminate conductance, which may reflect one of the following: a weak conductor near the surface, a strong conductor at depth (e.g., 100 to 120 m below surface) or to one side of the flight line, or aerodynamic noise. Those responses that have the appearance of valid bedrock anomalies on the flight profiles are indicated by appropriate interpretive symbols (see EM map legend). The others probably do not warrant further investigation unless their locations are of considerable geological interest.

#### The thickness parameter

DIGHEM can provide an indication of the thickness of a steeply dipping conductor. The amplitude of the coplanar anomaly (e.g., CPI channel on the digital profile) increases relative to the coaxial anomaly (e.g., CXI) as the apparent thickness increases, i.e., the thickness in the horizontal plane. (The thickness is equal to the conductor width if the conductor dips at 90 degrees and strikes at right angles to the flight line.) This report refers to a conductor as thin when the thickness is likely to be less than 3 m, and thick when in excess of 10 m. Thick conductors are indicated on the EM map by parentheses "( )". For base metal exploration in steeply dipping geology, thick conductors can be high priority targets because many massive sulfide ore bodies are thick, whereas non-economic bedrock conductors are often thin. The system cannot sense the thickness when the strike of the conductor is subparallel to the flight line, when the conductor has a shallow dip, when the anomaly amplitudes are small, or when the resistivity of the environment is below 100 ohm-m.

### Resistivity mapping

Areas of widespread conductivity are commonly

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encountered during surveys. In such areas, anomalies can be generated by decreases of only 5 m in survey altitude as well as by increases in conductivity. The typical flight record in conductive areas is characterized by inphase and quadrature channels which are continuously active. Local EM peaks reflect either increases in conductivity of the earth or decreases in survey altitude. For such conductive areas, apparent resistivity profiles and contour maps are necessary for the correct interpretation of the airborne data. The advantage of the resistivity parameter is that anomalies caused by altitude changes are virtually eliminated, so the resistivity data reflect only those anomalies caused by The resistivity analysis also helps conductivity changes. the interpreter to differentiate between conductive trends in the bedrock and those patterns typical of conductive overburden. For example, discrete conductors will generally appear as narrow lows on the contour map and broad conductors (e.g., overburden) will appear as wide lows.

The resistivity profiles and the resistivity contour maps present the apparent resistivity using the so-called pseudo-layer (or buried) half space model defined by Fraser (1978)<sup>1</sup>. This model consists of a resistive layer overlying

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Resistivity mapping with an airborne multicoil electromagnetic system: Geophysics, v. 43, p.144-172

a conductive half space. The depth channels give the apparent depth below surface of the conductive material. The apparent depth is simply the apparent thickness of the overlying resistive layer. The apparent depth (or thickness) parameter will be positive when the upper layer is more resistive than the underlying material, in which case the apparent depth may be quite close to the true depth.

The apparent depth will be negative when the upper layer is more conductive than the underlying material, and will be zero when a homogeneous half space exists. The apparent depth parameter must be interpreted cautiously because it will contain any errors which may exist in the measured altitude of the EM bird (e.g., as caused by a dense tree cover). The inputs to the resistivity algorithm are the inphase and quadrature components of the coplanar coil-pair. The outputs are the apparent resistivity of the conductive half space (the source) and the sensor-source distance. The flying height is not an input variable, and the output resistivity and sensor-source distance are independent of the flying height. The apparent depth, discussed above, is simply the sensor-source distance minus the measured altitude or flying height. Consequently, errors in the measured altitude will affect the apparent depth parameter but not the apparent resistivity parameter.

The apparent depth parameter is a useful indicator of simple layering in areas lacking a heavy tree cover. The DIGHEM system has been flown for purposes of permafrost mapping, where positive apparent depths were used as a measure of permafrost thickness. However, little quantitative use has been made of negative apparent depths because the absolute value of the negative depth is not a measure of the thickness of the conductive upper layer and, therefore, is not meaningful physically. Qualitatively, a negative apparent depth estimate usually shows that the EM anomaly is caused by conductive overburden. Consequently, the apparent depth channel can be of significant help in distinguishing between overburden and bedrock conductors.

The resistivity map often yields more useful information on conductivity distributions than the EM map. In comparing the EM and resistivity maps, keep in mind the following:

- (a) The resistivity map portrays the absolute value of the earth's resistivity, where resistivity = 1/conductivity.
- (b) The EM map portrays anomalies in the earth's resistivity. An anomaly by definition is a change from the norm and so the EM map displays anomalies, (i) over narrow, conductive bodies

and (ii) over the boundary zone between two wide formations of differing conductivity.

The resistivity map might be likened to a total field map and the EM map to a horizontal gradient in the direction of flight<sup>2</sup>. Because gradient maps are usually more sensitive than total field maps, the EM map therefore is to be preferred in resistive areas. However, in conductive areas, the absolute character of the resistivity map usually causes it to be more useful than the EM map.

## Interpretation in conductive environments

Environments having background resistivities below 30 ohm-m cause all airborne EM systems to yield very large responses from the conductive ground. This usually prohibits the recognition of discrete bedrock conductors. However, DIGHEM data processing techniques produce three parameters which contribute significantly to the recognition of bedrock conductors. These are the inphase and quadrature difference channels (DIFI and DIFQ), and the resistivity and depth channels (RES and DP) for each coplanar frequency.

<sup>&</sup>lt;sup>2</sup> The gradient analogy is only valid with regard to the identification of anomalous locations.

The EM difference channels (DIFI and DIFQ) eliminate most of the responses from conductive ground, leaving responses from bedrock conductors, cultural features (e.g., telephone lines, fences, etc.) and edge effects. Edge effects often occur near the perimeter of broad conductive This can be a source of geologic noise. While edge zones. effects yield anomalies on the EM difference channels, they do not produce resistivity anomalies. Consequently, the resistivity channel aids in eliminating anomalies due to edge On the other hand, resistivity anomalies will effects. coincide with the most highly conductive sections of conductive ground, and this is another source of geologic recognition of a bedrock conductor in a noise. The conductive environment therefore is based on the anomalous responses of the two difference channels (DIFI and DIFQ) and resistivity channels (RES). The most favourable the situation is where anomalies coincide on all channels.

The DP channels, which give the apparent depth to the conductive material, also help to determine whether a conductive response arises from surficial material or from a conductive zone in the bedrock. When these channels ride above the zero level on the digital profiles (i.e., depth is negative), it implies that the EM and resistivity profiles are responding primarily to a conductive upper layer, i.e., conductive overburden. If the DP channels are below the zero level, it indicates that a resistive upper layer exists, and this usually implies the existence of a bedrock conductor. If the low frequency DP channel is below the zero level and the high frequency DP is above, this suggests that a bedrock conductor occurs beneath conductive cover.

The conductance channel CDT identifies discrete conductors which have been selected by computer for appraisal by the geophysicist. Some of these automatically selected anomalies on channel CDT are discarded by the geophysicist. The automatic selection algorithm is intentionally oversensitive to assure that no meaningful responses are missed. The interpreter then classifies the anomalies according to their source and eliminates those that are not substantiated by the data, such as those arising from geologic or aerodynamic noise.

#### <u>Reduction of geologic noise</u>

Geologic noise refers to unwanted geophysical responses. For purposes of airborne EM surveying, geologic noise refers to EM responses caused by conductive overburden and magnetic permeability. It was mentioned previously that the EM difference channels (i.e., channel DIFI for inphase and DIFQ for quadrature) tend to eliminate the response of conductive overburden. This marked a unique development in airborne EM technology, as DIGHEM is the only EM system which yields channels having an exceptionally high degree of immunity to conductive overburden.

Magnetite produces a form of geological noise on the inphase channels of all EM systems. Rocks containing less than 1% magnetite can yield negative inphase anomalies caused by magnetic permeability. When magnetite is widely distributed throughout a survey area, the inphase EM channels may continuously rise and fall, reflecting variations in the magnetite percentage, flying height, and overburden This can lead to difficulties in recognizing thickness. deeply buried bedrock conductors, particularly if conductive overburden also exists. However, the response of broadly distributed magnetite generally vanishes on the inphase difference channel DIFI. This feature can be a significant aid in the recognition of conductors which occur in rocks containing accessory magnetite.

## EM magnetite mapping

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The information content of DIGHEM data consists of a combination of conductive eddy current responses and magnetic permeability responses. The secondary field resulting from conductive eddy current flow is frequency-dependent and consists of both inphase and quadrature components, which are positive in sign. On the other hand, the secondary field resulting from magnetic permeability is independent of frequency and consists of only an inphase component which is negative in sign. When magnetic permeability manifests itself by decreasing the measured amount of positive inphase, its presence may be difficult to recognize. However, when it manifests itself by yielding a negative inphase anomaly (e.g., in the absence of eddy current flow), its presence is assured. In this latter case, the negative component can be used to estimate the percent magnetite content.

A magnetite mapping technique was developed for the coplanar coil-pair of DIGHEM. The technique yields a channel (designated FEO) which displays apparent weight percent magnetite according to a homogeneous half space model.<sup>3</sup> The method can be complementary to magnetometer mapping in Compared to magnetometry, it is far less certain cases. able to resolve closely spaced sensitive but is more magnetite zones, as well as providing an estimate of the amount of magnetite in the rock. The method is sensitive to 1/4% magnetite by weight when the EM sensor is at a height of 30 m above a magnetitic half space. It can individually resolve steep dipping narrow magnetite-rich bands which are

<sup>3</sup> Refer to Fraser, 1981, Magnetite mapping with a multi-coil airborne electromagnetic system: Geophysics, v. 46, p. 1579-1594. separated by 60 m. Unlike magnetometry, the EM magnetite method is unaffected by remanent magnetism or magnetic latitude.

The EM magnetite mapping technique provides estimates of magnetite content which are usually correct within a factor of 2 when the magnetite is fairly uniformly distributed. EM magnetite maps can be generated when magnetic permeability is evident as negative inphase responses on the data profiles.

Like magnetometry, the EM magnetite method maps only bedrock features, provided that the overburden is characterized by a general lack of magnetite. This contrasts with resistivity mapping which portrays the combined effect of bedrock and overburden.

#### Recognition of culture

Cultural responses include all EM anomalies caused by man-made metallic objects. Such anomalies may be caused by inductive coupling or current gathering. The concern of the interpreter is to recognize when an EM response is due to culture. Points of consideration used by the interpreter, when coaxial and coplanar coil-pairs are operated at a common frequency, are as follows:

- 1. Channel CPS monitors 60 Hz radiation. An anomaly on this channel shows that the conductor is radiating power. Such an indication is normally a guarantee that the conductor is cultural. However, care must be taken to ensure that the conductor is not a geologic body which strikes across a power line, carrying leakage currents.
- 2. A flight which crosses a "line" (e.g., fence, telephone line, etc.) yields a center-peaked coaxial anomaly and an m-shaped coplanar anomaly.<sup>4</sup> When the flight crosses the cultural line at a high angle of intersection, the amplitude ratio of coaxial/coplanar response is 4. Such an EM anomaly can only be caused by a line. The geologic body which yields anomalies most closely resembling a line is the vertically dipping thin dike. Such a body, however, yields an amplitude ratio of 2 rather than 4. Consequently, an m-shaped coplanar anomaly with a CXI/CPI amplitude ratio of 4 is virtually a guarantee that the source is a cultural line.
- 3. A flight which crosses a sphere or horizontal disk yields center-peaked coaxial and coplanar anomalies with a CXI/CPI amplitude ratio (i.e., coaxial/coplanar) of

<sup>4</sup> See Figure 5-1 presented earlier.

1/4. In the absence of geologic bodies of this geometry, the most likely conductor is a metal roof or small fenced yard.<sup>5</sup> Anomalies of this type are virtually certain to be cultural if they occur in an area of culture.

- 4. A flight which crosses a horizontal rectangular body or wide ribbon yields an m-shaped coaxial anomaly and a center-peaked coplanar anomaly. In the absence of geologic bodies of this geometry, the most likely conductor is a large fenced area.<sup>5</sup> Anomalies of this type are virtually certain to be cultural if they occur in an area of culture.
- 5. EM anomalies which coincide with culture, as seen on the camera film or video display, are usually caused by culture. However, care is taken with such coincidences because a geologic conductor could occur beneath a fence, for example. In this example, the fence would be expected to yield an m-shaped coplanar anomaly as in case #2 above. If, instead, a center-peaked coplanar anomaly occurred, there would be concern that a thick

<sup>&</sup>lt;sup>5</sup> It is a characteristic of EM that geometrically similar anomalies are obtained from: (1) a planar conductor, and (2) a wire which forms a loop having dimensions identical to the perimeter of the equivalent planar conductor.

geologic conductor coincided with the cultural line.

6. The above description of anomaly shapes is valid when is not conductively coupled to the the culture environment. In this case, the anomalies arise from inductive coupling to the EM transmitter. However, when the environment is quite conductive (e.g., less than 100 Hz), the cultural conductor may be ohm-m at 900 conductively coupled to the environment. In this latter case, the anomaly shapes tend to be governed by current gathering. Current gathering can completely distort the anomaly shapes, thereby complicating the identification of cultural anomalies. In such circumstances, the interpreter can only rely on the radiation channel CPS and on the camera film or video records.

#### MAGNETICS

The existence of a magnetic correlation with an EM anomaly is indicated directly on the EM map. In some geological environments, an EM anomaly with magnetic correlation has a greater likelihood of being produced by sulfides than one that is non-magnetic. However, sulfide ore bodies may be non-magnetic (e.g., the Kidd Creek deposit near Timmins, Canada) as well as magnetic (e.g., the Mattabi deposit near Sturgeon Lake, Canada).

The magnetometer data are digitally recorded in the aircraft to an accuracy of one nT (i.e., one gamma) for proton magnetometers, and 0.01 nT for cesium magnetometers. The digital tape is processed by computer to yield a total field magnetic contour map. When warranted, the magnetic data may also be treated mathematically to enhance the magnetic response of the near-surface geology, and an enhanced magnetic contour map is then produced. The response of the enhancement operator in the frequency domain is This figure shows that the illustrated in Figure 5-2. passband components of the airborne data are amplified 20 times by the enhancement operator. This means, for example, that a 100 nT anomaly on the enhanced map reflects a 5 nT anomaly for the passband components of the airborne data.

The enhanced map, which bears a resemblance to a downward continuation map, is produced by the digital bandpass filtering of the total field data. The enhancement is equivalent to continuing the field downward to a level (above the source) which is 1/20th of the actual sensorsource distance.

Because the enhanced magnetic map bears a resemblance to a ground magnetic map, it simplifies the recognition of trends in the rock strata and the interpretation of geological structure. It defines the near-surface local



CYCLES/METRE

Fig. 5-2

Frequency response of magnetic enhancement operator.

geology while de-emphasizing deep-seated regional features. It primarily has application when the magnetic rock units are steeply dipping and the earth's field dips in excess of 60 degrees.

Any of a number of filter operators may be applied to the magnetic data, to yield vertical derivatives, continuations, magnetic susceptibility, etc. These may be displayed in contour, colour or shadow.

#### VLF

VLF transmitters produce high frequency uniform electromagnetic fields. However, VLF anomalies are not EM anomalies in the conventional sense. EM anomalies primarily reflect eddy currents flowing in conductors which have been energized inductively by the primary field. In contrast, VLF anomalies primarily reflect current gathering, which is a non-inductive phenomenon. The primary field sets up currents which flow weakly in rock and overburden, and these tend to collect in low resistivity zones. Such zones may be due to massive sulfides, shears, river valleys and even unconformities. AMPLITUDE

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CYCLES / METRE



- 5 - 27 -

The VLF field is horizontal. Because of this, the method is quite sensitive to the angle of coupling between the conductor and the transmitted VLF field. Conductors which strike towards the VLF station will usually yield a stronger response than conductors which are nearly orthogonal to it.

The Herz Industries Ltd. Totem VLF-electromagnetometer measures the total field and vertical quadrature components. Both of these components are digitally recorded in the aircraft with a sensitivity of 0.1 percent. The total field yields peaks over VLF current concentrations whereas the quadrature component tends to yield crossovers. Both appear as traces on the profile records. The total field data are filtered digitally and displayed as contours to facilitate the recognition of trends in the rock strata and the interpretation of geologic structure.

The response of the VLF total field filter operator in the frequency domain (Figure 5-3) is basically similar to that used to produce the enhanced magnetic map (Figure 5-2). The two filters are identical along the abscissa but different along the ordinant. The VLF filter removes long wavelengths such as those which reflect regional and wave transmission variations. The filter sharpens short wavelength responses such as those which reflect local geological variations.

- 5-28 -
#### CONCLUSIONS AND RECOMMENDATIONS

This report provides a very brief description of the survey results and describes the equipment, procedures and logistics of the survey over the Babine Lake project.

Mining activity over much of the Newman Peninsula may have obscured some valid bedrock conductors in the southwestern quadrant of the survey block. However, several conductors, which are typical of massive sulphide responses, were identified in the survey area. The survey was also successful in locating many moderately weak or poorly defined conductors of limited strike extent which may warrant additional work. The various maps included with this report display the magnetic and conductive properties of the survey area. It is recommended that the survey results be reviewed in detail, in conjunction with all available geophysical, geological and geochemical information. Particular reference should be made to the computer generated data profiles which clearly define the characteristics of the individual anomalies.

Most anomalies in the Newman Peninsula have been given a 'B?' designation, and will require detailed follow-up. On the rest of the grid, anomalous responses are variable. Many have been attributed to conductive overburden or deep weathering, and several appear to be associated with magnetite-rich rock units. Others coincide with VLF anomalies, resistivity gradients, and/or magnetic gradients, which may reflect faults or shears. Such structural breaks are considered to be of particular interest as they may have influenced mineral deposition within the survey area. The fact that the current copper zones are located in a relatively non-magnetic rock unit, suggests that other lowintensity zones may be favourable areas for follow-up.

The interpreted bedrock conductors defined by the survey should be subjected to further investigation, using appropriate surface exploration techniques. Resistivity anomalies are also considered to be potential areas of interest. Anomalies which are currently considered to be of moderately low priority may require upgrading if follow-up results are favourable.

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It is also recommended that image processing of existing geophysical data be considered, in order to extract the maximum amount of information from the survey results. Current software and imaging techniques often provide valuable information on structure and lithology, which may not be clearly evident on the contour and colour maps. These techniques can yield images which define subtle, but significant, structural details. Respectfully submitted, DIGHEM SURVEYS & PROCESSING INC.

Paul A. Smith Geophysicist

PAS/sdp A1090ANOV.91R

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#### APPENDIX A

#### LIST OF PERSONNEL

The following personnel were involved in the acquisition, processing, interpretation and presentation of data, relating to a DIGHEM<sup>IV</sup> airborne geophysical survey carried out for Noranda Exploration Company, Limited, over the Babine Lake property, B.C.

Steve Kilty	Vice President, Operations
Dave Pritchard	Survey Operations Supervisor
Phil Miles	Senior Geophysical Operator
Dave Miles	Second Geophysical Operator
D.B. Wilton	Pilot (Questral Helicopters Ltd.)
Gordon Smith	Data Processing Supervisor
Paul A. Smith	Interpretation Geophysicist
Reinhard Zimmerman	Drafting Supervisor
Lyn Vanderstarren	Draftsperson (CAD)
Susan Pothiah	Word Processing Operator
Albina Tonello	Secretary/Expeditor

The survey consisted of 1797 km of coverage, flown from July 11 to July 22, 1990.

All personnel are employees of Dighem Surveys & Processing Inc., except for the pilot who is an employee of Questral Helicopters Ltd.

DIGHEM SURVEYS & PROCESSING INC.

Paul A. Smith Geophysicist

PAS/sdp

Ref: Report #1090A

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## APPENDIX B

## STATEMENT OF COST

Date: November 28, 1990

## IN ACCOUNT WITH DIGHEM SURVEYS & PROCESSING INC.

To: Dighem flying of Agreement dated May 22, 1990, pertaining to an Airborne Geophysical Survey of the Babine Lake Project, B.C.

Survey Charges

1768 km of flying @ \$75.00/km

\$132,600.00

## Allocation of Costs

-	Data	Acquisition	(60%)
_	Data	Processing	(20%)

- Interpretation, Report and Maps (20%)

DIGHEM SURVEYS & PROCESSING INC.

Paul A. Smith Geophysicist

PAS/sdp

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## A P P E N D I X C

# EM ANOMALY LIST

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COAXL 900		XIAL 0 HZ	COPI 90	ANAR 10 HZ	COPI 720	ANAR 10 HZ	. VERTI	ICAL Œ	. HORIZO	NTAL T	CONDUC	TIVE TH	MAG CORR	
ANOMALY	/ RE	AL	QUAD :	REAL	QUAD	REAL	QUAD	. COND I	)EPIH*	. COND D	EPTH	RESIS	DEPTH	
FID/INTE	RP P	PM	PPM	PPM	PPM	PPM	PPM	.SIEMEN	М	.SIEMEN	М	OHM-M	M	NT
LINE 111	60	(F	LIGHT	33				•		•				
E 3794	H	3	13	6	24	53	17	. 1.9	0	. 1	32	83	2	0
F 3816	S	2	4	3	4	16	12	. 1.0	0	. 1	34	76	15	0
LINE 111	.70	(F	LIGHT	33	)			•		•				
a 3578	S	3	10	5	17	42	41	. 2.2	0	. 1	43	82	9	0
в 3552	S	2	7	5	17	45	60	. 2.1	0	. 1	30	99	0	0
C 3524	S	4	5	8	9	27	47	. 5.4	14	• 1	24	146	U	0
D 3518	B?	1	2	1	2	2	4	• -	-	• -	-	100	-	0
E 3452	S	3	8	5 2	14	35	49	. 2.7	2	• 1	23	100	7	0
F 3428 G 3400	н S	2 3	5 11	5 6	10	- 51	26 54	. 1.0	0	. 1	23	104	7	0
T.TNE 111	80	/ ᠮ	т.т.снт		\			•		•				
	.00 C	۲ <u>ــــــــــــــــــــــــــــــــــــ</u>	6	55	, 10	22	18	. 4.1	10	. 1	41	81	9	0
B 3162	S	3	6	6	10	22	29	. 3.7	3	. 1	40	79	7	Ō
C 3181	н	5	5	8	11	28	32	. 6.1	4	. 1	33	84	0	0
D 3191	S	3	10	5	22	41	53	. 2.1	0	. 1	17	224	0	0
E 3224	S	2	5	0	9	27	50	. 1.1	0	. 1	22	715	0	110
F 3250	Н	5	8	7	8	36	15	. 5.1	15	. 1	44	195	5	0
G 3261	S?	4	20	10	39	98	107	. 2.1	0	. 1	20	134	0	0
н 3268	S?	1	2	1	2	2	4				-	-	-	0
I 3283	S	2	3	3	6	13	13	. 3.1	12	. 1	34	131	0	0
J 3305	S	3	14	5	24	34	32	. 1.6	0	. 1	26	138	0	0
LINE 111	.90	(F	LIGHI	33	)			•		•				
A 3069	S	ŝ	7	2	´ 11	31	37	. 2.0	0	. 1	41	112	5	0
в 3028	S?	11	27	21	47	79	86	. 4.5	0	. 1	31	57	5	0
C 2998	S?	0	1	0	1	2	4	. –	-		-	-	-	250
D 2976	S	2	4	2	8	25	26	. 1.8	0	. 1	6	528	0	0
E 2944	S	1	5	3	9	23	24	. 1.3	0	. 1	35	86	2	0
F 2930	S	4	11	5	21	52	76	. 2.2	0	. 1	22	135	0	0
G 2917	S	1	2	1	2	2	4		-	· -	-	-	-	U
LINE 112	200	(E	LIGHI	33	)			•		•	_			_
A 2651	S	3	8	4	14	37	38	. 2.6	0	. 1	34	95	1	0
в 2681	Н	2	3	2	6	14	15	. 2.4	15	. 1	58	111	19	0
C 2688	H	0	2	3	6	13	14	. 1.6	18	. 1	64	86	27	0
D 2699	H	7	19	14	22	57	61	. 4.4	0	. 1	32	63	4	0
E 2701	B?	1	2	1	2	2	4	• -		• -	-	-	-	0
F 2723	S?	1	2	1	2	2	4	• -	-	• <del>-</del>		- 120	-	0
G 2758	5	ځ	TT	/	20	τp	62	. 2.3	U	• 1	24	• • •	Ū	U
•*	ESTI	MA	TED DE	PIH :	MAY B	E UNR	ELIABI	E BECAU	SE THE	STRONG	ER PA	RT .		
•	OF 1	THE	CONDU	CIOR	MAY	BE DE	EPER (	NO ON HE	E SIDE	CH' THE	FLIG CINC	HI'.		
•	LINE	ś, (	JR BEC	AUSE	OF A	SHAL	LOW D	LP UR UV	LIKBURL	IGN EFFE	CT2.	•		

J-A

	CO7 90	XIAL )0 HZ	COPI 90	ANAR 00 HZ	COPI 720	ANAR 10 HZ	. VERT	ICAL KE	. HORIZO . SHEI	ONTAL ET	CONDUC	CTIVE IH	MAG CORR
ANOMALY/ FID/INTERP	REAL PPM	QUAD PPM	REAL PPM	QUAD PPM	REAL PPM	QUAD PPM	. COND : .SIEMEN	DEPIH* M	COND I SIEMEN	DEPIH M	RESIS OHM-M	DEPTH M	NT
LINE 11200 H 2783 S I 2803 S	(H 2 4	TLIGHT 3 5	33) 2 3	7 11	19 13	18 32	. 1.9 . 3.2	0 7	. 1 . 1	37 22	75 150	3 0	0 0
LINE 11210 A 2557 S B 2515 S C 2509 E D 2461 S E 2415 S F 2408 S	(1 4 7 2 2 0	FLIGHI 7 6 14 5 3 10	2 33) 5 9 13 1 3 3	14 11 25 12 17 20	21 47 67 11 10 48	33 49 57 28 15 60	. 3.0 . 5.6 . 4.3 . 0.9 . 1.6 . 0.6	0 10 0 0 0 0		42 31 26 24 19 14	88 75 137 157 156 195	7 1 0 0 0	0 0 0 0 0
LINE 11220 A 2187 S? B 2192 H C 2200 E D 2215 S E 2246 S F 2256 S G 2309 S	(] 1 4 6 1 2 0 1	FLIGH 2 16 11 7 11 15 5	33) 1 7 12 0 4 0 2	) 30 20 13 11 29 9	2 70 60 41 7 59 15	4 96 37 61 85 109 14	· 2.0 5.0 0.5 1.3 . 0.5 . 1.1	- 0 3 0 0 0 0 0		- 25 22 0 10 10 15	- 101 160 526 271 234 217	- 0 0 0 0 0 0	0 0 0 0 0 0
LINE 11230 A 2067 S? B 2054 S C 2043 H D 2021 S E 2014 S F 1996 S	(1 4 5 0 0	FLIGH 14 8 7 7 4 4	r 33 5 3 10 0 0	) 23 12 12 13 15 7	61 48 29 40 13 15	77 70 28 51 81 23	1.9 2.4 . 6.2 . 0.5 . 0.5	0 0 3 12 5 0 5 0 5 0	· · · 1 · 1 · 1 · 1 · 1	9 24 2 0 0 0	149 109 412 480 438 539	0 0 0 0 0	0 0 0 0 0
LINE 11240 A 1699 E B 1704 S C 1721 S D 1726 E E 1760 S F 1773 S	(1 4 5 0 0	FLIGH 2 12 6 8 4 2	r 33 1 6 7 0 0 0	) 21 11 22 9 2	2 45 29 55 23 2	4 70 59 62 32 4	· 2.3 · 2.3 · 5.5 · 0.5	- - - - - - - - - - - - - -	· · 1 · 1 · 1 · 1 · 1	- 18 0 3 -	137 440 515 505	0 0 0 0	0 0 0 0 0
LINE 11250 A 1626 S B 1568 S? C 1557 S D 1546 S	5 0 0 0	FLIGH 13 10 5 2	r 33 10 0 0	) 15 19 8 2	53 43 28 2	59 100 38 4	· · 3.9 · 0.5 · 0.5	) 3 5 0 5 0 -	· · 1 · 1 · 1 · -	27 3 4 -	105 470 668 -	0 0 0 0 -	0 0 0 0

		COA 90	XIAL 0 HZ	COPI 90	LANAR 00 HZ	COPI 720	ANAR 00 HZ	. VERT	ICAL KE	HORIZO	NTAL T	CONDUC	CTIVE IH	MAG CORR
ANOMA FID/IN	LY/ I FERP	REAL PPM	QUAD PPM	REAL PPM	QUAD PPM	REAL PPM	QUAD PPM	COND SIEMEN	DEPIH* M	. COND I .SIEMEN	DEPIH M	RESIS OHM-M	DEPIH M	NT
LINE 1 E 153 F 152	1250 1 S 2 S	(F 0 0	LIGHI 7 2	* 33) 0 0	11 2	27 2	62 4	. 0.5 . –	0 -	. 1 . –	13 -	588 -	0 -	0
G 149	1 S	0	2	U	2	2	4	•	-	• -	-	-	-	U
LINE 1 A 124 B 127 C 129	1260 9 S 9 S? 1 S	(F 2 0 0	LIGHT 7 10 5	33) 5 0 0	12 18 5	14 41 24	42 92 46	2.5 . 0.5 . 0.6	4 0 0	. 1 . 1 . 1	25 1 9	129 468 477	0 0 0	0 0 0
LINE 1 A 134	1261 9 S	(F 3	LIGHI 8	33) 1	12	-36	27	. 1.6	0	. 1	7	524	0	6
LINE 1 D 72 E 66	1270 0 S 0 S	(F 0 2	TLIGHI 3 7	33) 0 0	) 6 10	20 33	40 27	. 0.9 . 0.6	0 0	. 1 . 1	21 7	714 698	0 0	0 0
LINE 1 A 29 B 30 C 30 D 33 E 35 F 40 G 41	1280 6 E 0 S 8 S 1 E 1 S? 4 S 6 S	(F 1 5 0 0 0 3 1	TLIGHI 2 17 7 2 2 6 5	33) 1 8 4 0 0 0 0	) 33 15 2 2 11 8	2 66 35 2 2 30 5	4 109 48 4 4 18 21	· 2.2 · 2.2 · 0.9 · - · 0.7 · 0.5	- 0 - - 0 0	· - · · · · · · · · · · · · · · · · · ·	- 25 26 - 23 29	- 79 83 - 721 749	- 0 - - 0 0	0 0 550 550 0 0
LINE 1 A 461 B 460 C 459 D 458 E 456 F 454 G 453 H 453 I 447	1290 6 S? 1 S? 1 S 0 S 8 B? 5 S? 4 S? 0 S? 8 S?	(F 6 1 6 1 10 0 0 1	TLIGHI 7 20 18 33 4 2 2 2 2	33) 13 13 13 10 14 0 0 0	20 2 39 35 41 8 2 2 2	75 2 97 93 174 4 2 2 2	64 3 120 259 54 4 4	· 6.3 · - 3.0 · 1.4 · 3.2 · 0.5 · - ·	4 	· · · 1 · 1 · 1 · 1 · 1 · - · -	30 - 22 24 15 46 -	80 - 57 61 160 792 - - -	0 0 0 0 0	0 0 0 0 0 0 19
LINE 1 A 37 B 39 C 42 D 48	1300 0 S 6 S 5 S? 3 S * ES OF	(H 3 0 1 1 TIMAJ	TLIGHI 10 2 2 2 2 ED DE	26 4 0 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1	) 18 2 2 2 (AY B	40 2 2 2 E UNRI BE DEI	76 4 4 ELIABI	. 1.9    E BECAU	0 - - SE THE E SIDE	· 1 · - · - · - STRONGE	28 - - - ER PA	79 - - RT . HT .	0 - - -	0 0 0 0
-	т.т	NE. (	RREY	AUSE	OFA	SHAT		POROV	ERBURD	EN EFFFY	TIS.	-		

. HORIZONIAL CONDUCTIVE MAG COAXIAL COPLANAR COPLANAR . VERTICAL CORR 7200 HZ . DIKE SHEET EARTH 900 HZ 900 HZ • ANOMALY/ REAL QUAD REAL QUAD REAL QUAD . COND DEPTH\*. COND DEPTH RESIS DEPTH FID/INTERP PPM PPM PPM PPM PPM PPM .SIEMEN NT M .SIEMEN M OHM-M Μ (FLIGHT LINE 11310 26) 3. A 774 H 17. 0.8 B 679 S . (FLIGHT LINE 11320 26) A 977 E ---B 1001 B? ----\_ 4. C 1023 S? \_ -23. 0. 0.5 D 1068 S LINE 11330 (FLIGHT 26) -11 11 . 6.5 9. A 1386 H 18. 1.0 0. B 1338 S? 4. C 1314 S? -----D 1292 S? -E 1284 S \_ \_ (FLIGHT 26) LINE 11340 A 1501 H 0.5 26 . B 1554 S . C 1562 S? 4. \_ \_ -----. D 1606 S \_ ----E 1677 S \_ . LINE 11350 26) (FLIGHT A 1953 H ----B 1895 S \_ \_ \_ ----. . C 1891 S? \_ - . ----\_ \_ \_ . 18. 0.5 ο. D 1829 S 27) LINE 11361 (FLIGHT 2.3 0. A 694 S . Ο. 707 S 15. 2.6 в С 774 S? \_ \_ \_ -. . 0.8 795 S D . LINE 11370 (FLIGHT 27) 13. 3.4 15. A 1255 S 4. 1.6 B 1205 S 40. 96. 0.5 0. C 1196 S? D 1191 S? • -\_ 4. E 1149 S \* ESTIMATED DEPTH MAY BE UNRELIABLE BECAUSE THE STRONGER PART OF THE CONDUCTOR MAY BE DEEPER OR TO ONE SIDE OF THE FLIGHT .

LINE, OR BECAUSE OF A SHALLOW DIP OR OVERBURDEN EFFECTS.

ANOMALY/ REAL QUAD REAL QUAD REAL QUAD . COND DEPTH: COND DEPTH: RESIS DEPTH FID/INTERP PFM PFM PFM PFM PFM PFM PFM SIEMEN M SIEMEN M OHM-M M NT LINE 11380 (FLIGHT 27) A 1442 S 2 3 0 7 4 32 0.5 0 1 65 903 0 0 B 1448 S 2 4 0 5 4 32 0.5 0 1 1 65 903 0 0 D 1491 S7 2 4 0 4 12 22 0.5 0 1 1 24 440 0 0 LINE 11390 (FLIGHT 27) A 1912 S 2 3 2 6 14 23 2.8 20 1 40 83 8 0 B 1902 S 4 8 6 13 13 20 3.7 0 1 1 33 82 0 0 C 1887 S 5 7 3 3 12 13 0.9 0 1 1 25 92 6 0 D 1839 S 0 2 0 2 1 4 0 E 1822 S7 0 3 0 6 0 22 1.1 16 1 113 1035 0 0 F 1809 S 0 2 0 2 2 4 0 I 1716 S7 0 2 0 2 2 4 0 I 1716 S7 0 2 0 2 2 4 0 I 1716 S7 0 2 0 2 2 4 0 I 1716 S7 0 2 0 2 2 4 0 I 1716 S7 0 2 0 2 2 4 0 I 1716 S7 0 2 0 2 2 4 0 I 1716 S7 0 2 0 2 2 4 0 I 1716 S7 0 2 0 2 2 4 0 I 1716 S7 0 2 0 2 2 4 0 I 1716 S7 0 2 0 2 2 4 0 I 1716 S7 0 2 0 2 2 4 0 I 1716 S7 0 2 0 2 2 4 0 I 1716 S7 0 2 0 2 2 4 0 I 1716 S7 0 2 0 2 2 4 0 I 1716 S7 0 1 0 2 2 4 0 I 1716 S7 0 2 0 2 0 2 2 4 0 I 1716 S7 0 1 0 2 2 4 0 I 1716 S7 0 1 0 2 2 4 0 I 1716 S7 0 1 0 2 2 4 0 I 1716 1340 (FLIGHT 27) A 2534 S7 0 1 0 2 2 4		COA 90	XIAL 0 HZ	COPI 90	ANAR 00 Hz	COPI 720	LANAR )0 HZ	. VERTI	CAL E	. HORIZ	ONTAL ET	CONDUC	CTIVE EH	MAG CORR
Line 11380         (FLIGH 27)         .         .         .           A 1442 S         2         3         0         7         4         32         0.5         0         1         65         903         0         0           B 1448 S         2         3         0         7         4         32         0.5         0         1         67         846         0         350           C 1455 S         0         2         0         4         12         22         0.5         0         1         24         440         0         0           LINE 11390         (FLIGH 27)         .	ANOMALY/ F FID/INTERP	REAL PPM	QUAD PPM	REAL PPM	QUAD PPM	REAL PPM	QUAD PPM	. COND D .SIEMEN	EPIH* M	. COND .SIEMEN	DEPIH M	RESIS OHM-M	DEPTH M	NT
A 1442 S       2       3       0       7       4       32       0.5       0       1       65       903       0       0         B 1448 S7       2       4       0       5       4       32       0.5       0       1       65       903       0       0         D 1491 S7       2       4       0       4       12       22       0.5       0       1       24       440       0       0         LINE       11390       (FLIGHT 27)       -       0       0       125       92       6       0       0       0       125       92       6       0       0       0       126       92       12       12       1.1       16       1113       103       8       0       0       12       12	LINE 11380	(F	LIGHI	27				•		•				
B 1448 5? 2 4 0 5 4 32 . 0.5 0 . 1 67 846 0 350 C 1455 5 0 2 0 2 2 4 0 D 1491 5? 2 4 0 4 12 22 . 0.5 0 . 1 24 440 0 0 D 1491 5? 2 4 0 4 12 22 . 0.5 0 . 1 24 440 0 0 D 1491 5? 2 4 0 4 12 22 . 0.5 0 . 1 24 440 0 0 C 1897 5 2 3 2 6 14 23 . 2.8 20 . 1 40 83 8 0 B 1902 5 4 8 6 13 13 20 . 3.7 0 . 1 33 82 0 0 C 1897 5 5 7 3 3 12 13 . 0.9 0 . 1 25 92 6 0 D 1839 S 0 2 0 2 1 4 0 E 1822 S? 0 3 0 6 . 0 22 . 1.1 16 . 1 113 1035 0 0 F 1809 S 0 2 0 2 0 2 0 4 0 G 1801 S 0 3 0 5 26 25 . 1.0 0 . 1 33 300 8 0 H 1779 S 0 2 0 2 0 2 2 4 0 I 1716 S? 0 2 1 2 2 4 0 J 1706 S? 0 2 0 2 0 2 2 4 0 I 1716 S? 0 2 1 2 2 4 0 J 1716 S? 0 2 1 2 2 4 0 LINE 11410 (FLIGHT 27)	A 1442 S	2	3	0	<b>7</b>	4	32	. 0.5	0	. 1	65	903	0	0
C 1455 S 0 2 0 2 2 4 0 D 1491 S? 2 4 0 4 12 22 . 0.5 0 . 1 24 440 0 0 LINE 11390 (FLIGHT 27)	B 1448 S?	2	4	0	5	4	32	. 0.5	0	. 1	67	846	0	350
D 191 S? 2 4 0 4 12 22 0 0.5 0 1 2 24 440 0 0 0 LINE 11390 (FLIGHT 27)	C 1455 S	0	2	0	2	2	4	• -	-	• -	-	_	-	0
LINE 11390 (FLIGHT 27)	D 1491 S?	2	4	0	4	12	22	. 0.5	U	• 1	24	440	U	U
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$	LINE 11390	(F	LIGHI	27	l			•		•				
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$	A 1912 S	2	3	2	6	14	23	. 2.8	20	. 1	40	83	8	0
C 1887 S 5 7 3 3 12 13. 0.9 0. 1 25 92 6 0 D 1839 S 0 2 0 2 1 4 0 F 1809 S 0 2 0 2 0 4 0 G 1801 S 0 3 0 5 26 25. 1.0 0. 1 33 300 8 0 H 1779 S 0 2 0 2 2 4 0 I 1716 S? 0 2 1 2 2 4 0 I 1716 S? 0 2 0 2 2 4 0 I 1716 S? 0 2 0 2 2 4 0 I 1716 S? 0 2 0 2 2 4 0 I 1716 S? 0 2 0 2 2 4 0 I 1716 S? 0 2 0 2 2 4 0 I 1716 S? 0 2 0 2 2 4 0 I 1716 S? 0 2 1 2 0 2 2 4 0 I 1716 S? 0 2 1 2 0 2 2 4 0 I 1716 S? 0 2 1 2 0 2 2 4 0 I 1716 S? 0 2 1 2 0 2 2 4 0 I 1716 S? 0 2 1 2 0 2 2 4 0 IINE 11400 (FLIGHT 27)	B 1902 S	4	8	6	13	13	20	. 3.7	0	. 1	33	82	0	0
D 1839 S 0 2 0 2 1 4 0 F 1802 S 0 2 0 2 0 2 0 4 0 G 1801 S 0 3 0 5 26 25 . 1.0 0 . 1 33 300 8 0 H 1779 S 0 2 0 2 2 4 0 J 1708 S? 0 2 0 2 2 4 0 J 1708 S? 0 2 0 2 2 4 0 J 1708 S? 0 2 0 2 2 4 0 LINE 11400 (FLIGHT 27) A 1988 S 1 6 3 11 13 34 . 0.9 0 . 1 34 95 1 0 B 2125 S 1 2 0 2 2 4 0 J 1708 S? 0 5 0 6 21 34 . 0.5 0 . 1 53 219 7 0 B 2425 S? 0 5 0 6 21 34 . 0.5 0 . 1 67 828 0 0 C 2396 S 0 3 0 5 14 23 . 0.6 0 . 1 167 828 0 0 C 2396 S 0 3 0 5 14 23 . 0.6 0 . 1 163 396 0 0 D 2395 S? 0 1 0 2 2 4 0 LINE 11410 (FLIGHT 27) A 2514 S? 4 7 5 12 26 10 . 3.3 0 . 1 53 219 7 0 B 2425 S? 0 5 0 6 21 34 . 0.5 0 . 1 67 828 0 0 C 2396 S 0 3 0 5 14 23 . 0.6 0 . 1 167 828 0 0 C 2396 S 0 3 0 5 14 23 . 0.6 0 . 1 167 828 0 0 D 2339 S? 0 1 0 2 2 4 0 LINE 11420 (FLIGHT 27) A 3169 S 1 2 1 2 2 4 0 D 119 55 . 0.5 0 . 1 27 702 0 0 D 2030 S 1 2 1 2 2 4 0 D 3027 S 1 4 1 8 10 2 . 0.6 0 . 1 144 692 0 0 C 3060 S? 0 2 0 2 2 2 4 0 D 3027 S 1 4 1 8 10 2 . 0.6 0 . 1 144 692 0 0 C 3060 S? 0 2 0 2 2 4 0 J 307 S 1 2 0 2 2 4 0 J 3379 S? 1 2 0 2 2 4 0 D 3027 S 1 4 1 8 10 2 . 0.6 0 . 1 144 692 0 0 E 3003 S 1 2 0 2 2 4 0 J 102 I 14141 (FLIGHT 27) A 3379 S? 1 2 0 2 2 4 0 D 3027 S 1 4 1 8 10 2 . 0.6 0 . 1 144 692 0 0 E 3003 S 1 2 0 2 0 2 0 4 0 J 102 I 144 1 (FLIGHT 27) A 3379 S? 1 2 0 2 0 2 0 4 0 J 102 I 1442 (FLIGHT 27) A 3379 S? 1 2 0 2 0 2 0 4 0 J 102 I 1442 (FLIGHT 27)	C 1887 S	5	7	3	3	12	13	. 0.9	0	. 1	25	92	6	0
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$	D 1839 S	0	2	0	2	1	4	. –	16	• -	- 112	1025	-	0
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	E 1822 5: F 1809 5	0	2	0	2	0	22 4	• •••		• 1	- 112	1035	-	0
H 1779 S 0 2 0 2 2 4 0 I 1716 S? 0 2 1 2 2 4 0 J 1708 S? 0 2 0 2 2 4 0 LINE 11400 (FLIGHT 27)	G 1801 S	ŏ	3	Ő	5	26	25	. 1.0	0	. 1	33	300	8	0
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	H 1779 S	Ō	2	Ō	2	2	4	. –	_		_	-	_	Õ
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$	I 1716 S?	0	2	1	2	2	4	. –	-		-	-	-	0
LINE 11400 (FLIGHT 27) A 1988 S 1 6 3 11 13 34 0.9 0 1 34 95 1 0 B 2125 S 1 2 0 2 2 4 0 LINE 11410 (FLIGHT 27) A 2514 S? 4 7 5 12 26 10 3.3 0 1 53 219 7 0 B 2425 S? 0 5 0 6 21 34 0.5 0 1 67 828 0 0 C 2396 S 0 3 0 5 14 23 0.6 0 1 167 828 0 0 C 2396 S 0 3 0 5 14 23 0.6 0 1 163 96 0 0 D 2339 S? 0 1 0 2 2 4 0 LINE 11420 (FLIGHT 27) A 2754 S 2 3 0 3 15 14 1.0 0 1 18 476 0 0 B 2782 S 2 5 0 11 9 55 0.5 0 1 27 702 0 0 LINE 11430 (FLIGHT 27) A 3169 S 1 2 1 2 2 4 20 B 3150 S 5 6 6 10 24 32 4.9 0 1 33 102 0 0 C 3360 S? 0 2 0 2 2 4 0 LINE 11441 (FLIGHT 27) A 3769 S 1 2 0 2 2 4 0 LINE 11441 (FLIGHT 27) A 3769 S 1 2 0 2 2 4 0 J 3027 S 1 4 1 8 10 2 0.6 0 1 14 692 0 0 C 3060 S? 0 2 0 2 2 4 0 LINE 11441 (FLIGHT 27) A 3539 S 0 2 0 2 0 4 0 LINE 11441 (FLIGHT 27) A 3539 S 0 2 0 2 0 4 0 LINE 11441 (FLIGHT 27) A 3539 S 0 2 0 2 0 7 2 0 7 2 0 7 0 LINE 11441 (FLIGHT 27) A 3539 S 0 2 0 2 0 7 2 0 7 0 7 0 0 C 3000 S? 1 2 0 2 0 7 0 7 0 0 0 C 3000 S? 1 2 0 7 0 7 0 0 0 C 3000 S? 1 2 0 7 0 7 0 0 0 C 3000 S 1 2 0 7 0 7 0 0 0 0 C 3000 S 1 2 0 7 0 7 0 0 0 0 C 3000 S 1 2 0 7 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	J 1708 S?	0	2	0	2	2	4		-	• -	-	-	-	0
Inite 11400       (Initiality 2.1)       1       13       34       0.9       0       1       34       95       1       0         B 2125 S       1       2       0       2       2       4       -       -       -       -       -       -       0         LINE 11410       (FLIGHT 27)       . <td< td=""><td>T.TNE 11400</td><td><u>।</u></td><td>ग. र त्यमग</td><td>י 27</td><td>`</td><td></td><td></td><td>•</td><td></td><td>•</td><td></td><td></td><td></td><td></td></td<>	T.TNE 11400	<u>।</u>	ग. र त्यमग	י 27	`			•		•				
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	A 1988 S	1	6	. 27	' 11	13	34	. 0.9	0	. 1	34	95	1	0
LINE 11410 (FLIGHT 27)	B 2125 S	1	2	Ō	2	2	4	. –	_		_	-	_	Ō
$\begin{array}{cccccccccccccccccccccccccccccccccccc$								•		•				
A 2514 Sr       4       7       5       12       20       10       3.3       0       1       133       219       7       0         B 2425 S?       0       5       0       6       21       34       0.5       0       1       67       828       0       0         C 2396 S       0       3       0       5       14       23       0.6       0       1       16       396       0       0         D 2339 S?       0       1       0       2       2       4       -       -       -       -       -       0         LINE 11420       (FLIGHT 27)       .	LINE 11410	(1	LIGHI 7	: 27	10	26	10	• 22	0	•	52	210	7	0
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	A 2014 Dr B 0405 S?	4	/ 5	5	12	20	34	. 3.3	0	• 1	55	828	0	0
D 2339 S? 0 1 0 2 2 4	C 2396 S	Ő	3	Ő	5	14	23	. 0.6	ŏ	. 1	16	396	Ő	Ő
LINE 11420 (FLIGHT 27) A 2754 S 2 3 0 3 15 14 10 0 11 18 476 0 0 B 2782 S 2 5 0 11 9 55 0.5 0 1 27 702 0 0 LINE 11430 (FLIGHT 27) A 3169 S 1 2 1 2 2 4 20 B 3150 S 5 6 6 10 24 32 4.9 0 1 33 102 0 0 C 3060 S? 0 2 0 2 2 4 0 D 3027 S 1 4 1 8 10 2 0.6 0 1 14 692 0 0 E 3003 S 1 2 0 2 2 4 0 LINE 11441 (FLIGHT 27) A 3379 S? 1 2 0 2 2 4 0 LINE 11442 (FLIGHT 27) A 3539 S 0 2 0 2 0 4 690	D 2339 S?	Õ	1	Õ	2	2	4	. –	_		_	-	-	Õ
LINE 11420 (FLIGHT 27)		. –						•		•				
A 2734 S       2       3       0       3       13       14       1.0       0       1       18       476       0       0         B 2782 S       2       5       0       11       9       55       0.5       0       1       27       702       0       0         LINE 11430       (FLIGHT 27)       .	LINE 11420	(F	LIGHI	27	)	15	1/	• 10	٥	•	10	176	0	0
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	A 2724 5 B 2782 C	2	ט 5	0	د 11	0	14	. 1.0	0	• 1	27	4/0	0	0
LINE 11430 (FLIGHT 27)		2	5	v	**		55	. 0.5	Ŭ	• •	2.7	.02	Ŭ	0
A 3169 S 1 2 1 2 2 4 $-$ 20 B 3150 S 5 6 6 10 24 32 4.9 0 1 33 102 0 0 C 3060 S? 0 2 0 2 2 4 $-$ 0 D 3027 S 1 4 1 8 10 2 0.6 0 1 14 692 0 0 E 3003 S 1 2 0 2 2 4 $-$ 0 LINE 11441 (FLIGHT 27)	LINE 11430	(F	LIGHI	27	)			•		•				
B 3150 S 5 6 6 10 24 32 4.9 0 1 1 33 102 0 0 C 3060 S? 0 2 0 2 2 4 0 D 3027 S 1 4 1 8 10 2 0.6 0 1 14 692 0 0 E 3003 S 1 2 0 2 2 4 0 LINE 11441 (FLIGHT 27) 0 A 3379 S? 1 2 0 2 2 4 0 LINE 11442 (FLIGHT 27) 0 	A 3169 S	1	2	1	2	2	4	• -	-	• ~	-	-	-	20
C 3060 S? 0 2 0 2 2 4 0 D 3027 S 1 4 1 8 10 2 0.6 0 1 14 692 0 0 E 3003 S 1 2 0 2 2 4 0 LINE 11441 (FLIGHT 27) A 3379 S? 1 2 0 2 2 4 0 LINE 11442 (FLIGHT 27) A 3539 S 0 2 0 2 0 4 690	B 3150 S	5	6	6	10	24	32	. 4.9	0	. 1	55	102	0	0
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	C 3060 S?	1	Z A	1	2	10	4		-		-	- 602	-	0
LINE 11441 (FLIGHT 27) A 3379 S? 1 2 0 2 2 4 $       0$ LINE 11442 (FLIGHT 27) A 3539 S 0 2 0 2 0 4 $         -$	E 3003 S	1	2	Ō	2	2	4		-	• •	- 14		-	0
LINE 11441 (FLIGHT 27)		-	-	•	-	-	-	•		•				•
A 3379 S? 1 2 0 2 2 4 0 LINE 11442 (FLIGHT 27)	LINE 11441	(F	LIGHI	27	)			•		•				
LINE 11442 (FLIGHT 27) A 3539 S 0 2 0 2 0 4 690 .* ESTIMATED DEPTH MAY BE UNRELIABLE BECAUSE THE STRONGER PART . . OF THE CONDUCTOR MAY BE DEEPER OR TO ONE SIDE OF THE FLIGHT .	A 3379 S?	1	2	0	2	2	4	. –	-		-	-	-	0
A 3539 S 0 2 0 2 0 4 690 	LINE 11442	/ 1	ग राज्यमा	· 27				•		•				
	A 3539 S		2	0	2	0	4	• -	-	· –	_	_	-	690
.* ESTIMATED DEPTH MAY BE UNRELIABLE BECAUSE THE STRONGER PART . . OF THE CONDUCTOR MAY BE DEEPER OR TO ONE SIDE OF THE FLIGHT .		v	-		-		•	-		-				
. OF THE CONDUCTOR MAY BE DEEPER OR TO ONE SIDE OF THE FLIGHT .	.* ES.	<b>LIWEL</b>	ED DE	PIH 1	AY B	e unri	TLIABI	E BECAUS	E THE	STRONG	ER PA	RT .		
וויאד הסנציאובר הדא באזוומדומניה מה מדה הם העובים אבור ער ארבים ב	• OF	THE VE (	CONDU	CIOR	MAY	BE DEI	EPER C	NR TU ONE	SIDE	OF THE	FLIG	HI .		

	COA 90	XIAL 00 HZ	COPLANAR COPLANAR . VERTICAL 900 HZ 7200 HZ . DIKE		ICAL Æ	HORIZC	NTAL T	CONDUC	MAG CORR				
ANOMALY/ T FID/INTERP	REAL PPM	QUAD PPM	REAL PPM	QUAD PPM	REAL PPM	QUAD PPM	. COND I .SIEMEN	)EPTH* M	COND D SIEMEN	EPIH M	RESIS OHM-M	DEPIH M	NT
LINE 11442 B 3568 S C 3590 S	(F 1 1	TLIGHT 7 3	27) 0 1	12 4	25 11	38 18	. 0.5 . 0.6	0 0	. 1 . 1	26 18	703 349	0 0	0 0
LINE 11450 A 3965 H B 3869 S C 3846 S D 3822 S	(H 1 1 0 1	TLIGHT 2 5 2 2 2	27) 1 0 0 1	2 4 2 2	2 14 2 2	4 10 4 4	·	 	. – . 1 . –	- 27 - -	_ 338 _ _	- 0 -	0 0 0 0
LINE 11460 A 4089 H B 4107 S	(H 1 4	TLIGHI 2 4	27) 1 4	2 9	- 2 11	0 20	 . 4.4	_ 14	. <u>-</u> . 1	- 32	_ 449	- 0	0 0
LINE 11462 A 1031 S B 918 S C 904 S? D 873 S	(H 1 0 0 1	TLIGHI 2 4 2 3	28) 1 2 0 2	2 5 2 5	2 24 2 15	4 24 4 25	 . 1.0 	- 0 - 0	. – . 1 . –	- 22 - 39		  11	0 0 0 0
LINE 11471 A 638 S B 678 S C 686 S	(1 0 0 0	тысни 4 2 2	28) 1 0 1	) 6 2 2	21 2 2	29 4 4	. 0.5 . – . –	0  -	. 1 . <del>-</del> 	39  -	706  _	0 _ _	0 0 0
LINE 11481 A 1337 S B 1366 S	( ) 0 0	TLIGHI 2 2	28) 0 0	) 2 2	2 2	4 4	· · - · -	- -	. – . –		-	-	0 0
LINE 11490 A 1750 H B 1639 S C 1598 S D 1554 S	(I 2 0 1	FLIGHI 5 2 2 2 2	28) 3 1 1	) 2 2 2	14 2 2 2	22 4 4 4	2.2  	8 - -	· · 1 · - · -	57 - - -	119 - - -	18 _ _ _	0 0 0 0
LINE 11570 A 2538 H B 2474 S C 2451 S D 2395 S E 2348 S	3 0 1 1	FLIGHI 5 2 4 3 2	29) 5 0 2 3 1	) 5 5 7 2	24 16 12 22 2	22 11 10 14 4	. 3.2 . 1.0 . 0.9 . 1.7 . –	0 0 6 -	. 1 . 1 . 1 . 1 . 1	36 43 19 25 -	145 335 614 443 -	0 15 0 0	0 0 0 0
LINE 11580 A 448 S	) (] 1	FLIGHI 2 TED DE	י 30 1 ריידים	) 1 MAY B	2 EUNR	4 FIJTABI	· · -	- Se The	. –	- Er pa	- RT .	-	0
. OF . LI	THE NE. (	CONDU DR BEC	ICTOR AUSE	MAY OF A	BE DE SHAL	EPER (	TO ON POR OV	E SIDE ERBURD	OF THE EN EFFEX	FLIG	HT.		

	COA2 900		COPI 90	LANAR 00 HZ	COPI 720	ANAR 00 HZ	. VERT	ICAL KE	. HORIZ . SHE	ONTAL ET	CONDU EAR	CTIVE IH	VE MAG CORR			
ANOMALY/ : FID/INTERP	REAL PPM	QUAD PPM	REAL PPM	QUAD PPM	REAL PPM	QUAD PPM	. COND .SIEMEN	DEPIH* M	COND SIEMEN	DEPIH M	RESIS OHM-M	DEPTH M	NT			
LINE 11580 B 481 S	(F 0	LIGHT 3	' 30) 0	1	14	17	. 0.8	0	. 1	47	477	17	0			
C 530 S D 587 S E 644 S	0 5 2	0 6 7	0 4 3	2 62 10	2 26 22	4 34 42	 . 1.6 . 1.5	- 0 0	· – · 1 · 1	- 0 34	327 333	- 0 0	0 0 0			
LINE 11590 A 880 S	(F 0	LIGHT 5	' 30) 0	10	21	57	. 0.5	0	• • • 1	31	742	0	0			
LINE 11591 A 1372 H B 1401 B? C 1408 S	(F 5 5 0	LIGHT 11 8 2	30) 9 6 1	23 11 2	56 - 28 2	39 21 4	. 3.5 . 4.2 . –	2 8 -	. 1 . 1 . –	36 43 -	85 192 -	6 3 -	0 0 0			
LINE 11592 A 1171 S? B 1130 S C 1079 S? D 1069 S?	(F 0 1 3 1	LIGHT 1 3 11 2	30) 0 2 3 1	2 6 18 2	0 17 52 2	4 19 87 4	 . 1.9 . 1.5 	- 11 2 -	· · · 1 · 1 · -	- 41 20 -	- 438 462 -	- 0 0 -	330 0 0 0			
LINE 11600 A 1675 H B 1707 B? C 1713 S D 1825 S	(F 5 8 1 1	LIGHT 14 10 5 4	30) 7 7 5 2	21 14 8 7	49 30 26 16	62 20 20 21	· · · 3.0 · 5.4 · 2.0 · 1.4	0 5 0 0	· · · 1 · 1 · 1	33 38 33 17	109 148 724 441	1 1 0 0	0 0 0			
LINE 11620 A 474 S B 561 S C 597 S D 611 S E 631 S	(F 2 0 0 2 0	LIGHT 4 3 4 8 2	31) 1 0 3 0	5 5 6 16 2	17 19 24 15 2	12 22 20 54 4	· · · 1.8 · 1.0 · 0.6 · 1.4 · -	0 0 0 -	· · 1 · 1 · 1 · 1 · -	13 41 35 17 -	790 312 792 319 -	0 14 0 0 -	0 0 0 0			
LINE 11630 A 1060 S B 1042 S? C 1028 S D 931 S? E 877 S	(F 1 0 0 0 2	T.IGHT 6 2 6 5 5 5	31) 2 0 0 1 2	11 2 9 9 7	30 2 22 25 5	16 4 47 33 21	· · · 1.3 · - · 0.5 · 0.5 · 0.5 · 1.7	0 - 0 0 0	· · · 1 · - · 1 · 1 · 1	36  38 27 16	584 - 715 589 361	0 - 0 0 0	0 240 0 0 0			
LINE 11640 A 1195 S .* ES	H) O TIMAT	LIGHI 2 ED DE	31) 1	) AY BI	1 EUNRI	4 ELIABI	E BECAU	- ISE THE	STRONG	- ER PA	- RT .	-	0			
· OF	NE. (	NR BEC	AUSE	OFA	SHALI	LOW DI	POROV	ERBURD	EN EFFE	CTS.						

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. LINE, OR BECAUSE OF A SHALLOW DIP OR OVERBURDEN EFFECTS.

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	CO2 90	XIAL 00 HZ	COPI 90	LANAR DO HZ	COPI 720	LANAR DO HZ	. VERTICAL . . DIKE .		HORIZONTAL SHEET		CONDUCTIVE EARTH		MAG CORR
anomat.v/f	REAT.		REAT.	OLID	REAL.			EPTH*	. COND D	ертн	RESTS	DEPTH	
FID/INTERP	PPM	PPM	PPM	PPM	PPM	PPM	.SIEMEN	M	.SIEMEN	M	OHM-M	M	NT
		-					•		•				
D 1215 U	1)	, TIGUI	: 31) ?	) 	21	3	• 16	14	• 1	50	315	5	5
C 1293 S	1	4	2	6	15	29	. 1.1	14	. 1	47	380	0	2
D 1342 S	$\overline{2}$	6	2	4	15		. 1.0	Ő	. 1	24	162	3	Ő
E 1346 S	ĩ	2	1	2	2	4	. –	_	. –	_	-	_	Õ
F 1368 S	4	2	0	3	15	13	. 1.0	0	. 1	25	432	0	0
LINE 11650	í.	नाःसा	· 31				•		•				
A 1756 H	3	10		17	44	22	. 3.0	0	. 1	42	111	6	0
B 1737 S?	1	6	0	9	23	54	. 0.6	10	. 1	58	754	1	0
C 1726 S	0	5	0	6	15	50	. 0.5	0	. 1	49	773	0	9
D 1652 S	0	2	1	2	- 2	4	. –	-	. –	-	-	-	0
Е 1630 Н	1	4	2	7	21	25	. 1.5	11	. 1	39	586	0	0
F 1620 S	1	5	1	5	16	31	. 0.5	0	. 1	31	412	5	0
G 1602 S?	0	5	0	7	17	46	. 0.5	0	. 1	38	747	0	0
H 1563 S	1	3	1	4	16	15	. 1.0	0	. 1	20	303	0	0
LINE 11660	(1	LIGHI	: 31)	)			•		•				_
A 1840 S	3	6	7	11	5	22	. 3.4	0	. 1	44	120	6	0
B 1867 S	0	4	0	7	20	25	. 0.5	0	. 1	50	804	0	0
C 1884 S?	2	8	5	17	45	41	. 1.6	0	• 1	17	225	0	0
D 1899 S	0	2	1	12	2	4	. –	_	· -	- 17	400	-	51
E 1903 5 F 1003 C2	0	5	2	5 T2	44	25	. 0.5	0	• 1	27	400	0	0
G 2031 S	0	ר ז	1	6	18	31	. 0.5	0	· 1	40	553	0	0
	•	Ū	-	. •			•	•	•		220	•	•
LINE 11670	(1	TIGHI	31)				•	_	•		-	_	_
A 2427 H	3	12	10	22	51	62	. 2.7	0	. 1	43	105	10	0
B 2388 H	2	3	3	9	26	11	. 2.6	5	• 1	38	299	0	0
C 2309 S	2	5	<u>ک</u>	11	38	30	. 1./	10	• 1	42	276	2	U
D 2299 5:	1	2	1	2	13	20		0	• -	- 14	200	-	0
F 2254 S2	1	2	1	2	2	20	. 0.0	-	• •	-		-	0
G 2240 S	2	4	1	5	12	24	. 0.5	0	. 1	22	307	0	Ő
T THE 11600	/1	स. म. (भ का	1 21				•		•				
7 2222 H	r) 1	7	, TC :	) 15	10	5	• • • •	6	•	16	222	6	0
R 2535 H	ວ າ	5	4 2	0	30	25	. 2.3	0	• 1	40 20	222	0	0
C 2611 S	2	2	1	2	20	2.J 4	. 2.0		• -	20	205	-	0
D 2640 S	1	7	3	12	46	60	. 1.2	0	• -	22	367	0	0
	*	,	5	*4	-10	00	• •	v	•		507	0	0
LINE 11690	( I	I ICHI	: 31)	)			•		•				
a 3020 s	2	9	4	15	42	43	. 1.7	0	. 1	32	240	0	0
* ESTIMATED DEPTH MAY BE UNRELIABLE BECAUSE THE STRONGER PART													
. OF THE CONDUCTOR MAY BE DEEPER OR TO ONE SIDE OF THE FLIGHT											HT.		
. LI	NE, (	DR BEC	AUSE	OF A	SHALL	LOW DI	POROVI	ERBURD	EN EFFEC	TS.	•		

COAXIAL ( 900 HZ			COPI 90	ANAR 00 HZ	COPI 720	ANAR 00 HZ	. VERT	ICAL Œ	. HORIZO . SHEE	NTAL T	CONDUC	TIVE H	MAG CORR				
ANOMALY /	REAL	QUAD	REAL	QUAD	REAL	QUAD	. COND I	)EPIH*	. COND D	EPTH	RESIS	DEPTH					
FID/INTERP	PPM	PPM	PPM	PPM	PPM	PPM	.SIEMEN	M	.SIEMEN	М	OHMM	М	NT				
LINE 11690	(I	LIGHI	י 31)				•		•								
B 2985 S	3	15	6	30	99	83	. 1.7	0	. 1	16	201	0	0				
C 2980 S?	1	2	1	2	2	4	• -	10	• -	-	-	-	0				
D 2913 H F 2887 S	2	27	3 1	13	40 36	39 65	. 3.2	10	• 1	28 17	249 446	0	0				
F 2868 S	ō	2	1	2	2	4	. –	-		_	-	-	Ő				
G 2842 S	1	2	1	2	2	4		-		-	-	-	0				
LINE 11700	(1	TIGHI	32				•		•								
A 464 S	1	2	1	2	2	4	. –	-		-	-	-	0				
B 518 S	0	2	1	2	2	4	• -	_	• -	-	_	-	0				
C 526 S	1	3	3	7	, 23	21	. 1.4	4	. 1	26	307	0	0				
D 559 5	0	4	1	8 2	21	30 4	. 0.5	U _	• I	-	400	-	0				
	. U		*	2	~	•	•		•				Ŭ				
LINE 11710	(H	TICHI	r 32)	)	26	40	•	0	•	-	40.4	0	0				
A 924 S	0	7	2	12	36	42	. 0.8	0	. 1	29	494	0 21	0				
в 893 5 С 870 S	1	4	1	2	2	0 4		-	• 1	42	- 110	Z1 -	0				
D 845 S	ŏ	4	1	5	16	27	. 0.7	0	. 1	25	224	3	11				
E 836 S	1	3	2	8	20	31	. 1.1	9	. 1	42	453	0	0				
F 815 S	1	4	1	7	16	33	. 1.0	8	. 1	46	733	0	0				
G 798 S?	0	4	2	9	7	34	. 0.8	0	. 1	36	438	0	0				
LINE 11720	(I	LIGHI	32)	)			•		•								
A 1024 S?	Ò	5	2	9	25	22	. 0.5	0	. 1	19	472	0	0				
B 1065 S?	0	8	1	13	-30	70	. 0.5	0	. 1	25	550	0	0				
C 1076 S	1	5	1	5	2 10	38	. 0.5	U	• 1	46	/59	U	0				
E 1132 S?	0	2	1	2	2	4		_	· -	_	-	_	30				
	, –	_	_	_	-	-	•		•								
LINE 11730	(H	TICHI	נ 32	)	47	50	•	0	•	20	170	0	0				
A 1440 S	1	6 1	5	20	4/	59 24	. 1.0	0	• 1	29	568	0	0				
C 1385 S	Ō	2	0 0	2	2	4		-	•	_		-	Ő				
D 1339 S?	1	2	1	2	2	4	. –	-	. –	-	-	-	0				
E 1333 S?	1	2	0	2	2	4	• -	-		-	-	-	0				
T.TNE 11740	. (T	ग . १ ट्रामा	י כצ יו	•			•		•								
A 1559 S	2	9	5	, 15	33	44	. 1.7	0	. 1	32	161	0	0				
B 1611 H	3	10	2	14	48	45	. 1.5	0	. 1	25	403	0	0				
C 1629 S	0	2	1	2	2	4		-		-	-	-	0				
* FC	יעשריני		INTEL N	AY R	r UNRI	ELTARI	E BETAIL	SE THE	STRONG	R PA	RT .						
. OF	THE	CONDU	CTOR	MAY	BE DE	EPER C	R TO ON	E SIDE	OF THE	FLIG	HT.						
. LI	NE.	RBEX	AUSE	OF A	SHAL	LOW DI	POROV	ERBURD	EN EFFEC	TS.	•						

	COAXIAL 900 HZ			COPI 90	ANAR )0 HZ	COPI 720	ANAR . 00 HZ .	VERT DI	ICAL KE	HORIZ	ONTAL ET	CONDUC	MAG CORR	
ANOMAL FID/INT	Y/ I ERP	REAL PPM	QUAD PPM	REAL PPM	QUAD PPM	REAL PPM	QUAD . PPM .	COND SIEMEN	DEPIH* M	COND SIEMEN	DEPTH I M	RESIS OHM-M	DEPTH M	NT
LINE 11 D 1651	740 S	( I 0	TLIGHI 2	32) 0	7	14	13 .	0.5	0	. 1	36	623	0	12
LINE 11 A 2058 B 2038 C 1988 D 1955	750 S S H S	(F 1 2 2 0	тысни 3 7 9 5	2 32) 1 5 3 1	6 12 14 8	15 35 48 17	15 . 28 . 54 . 49 .	1.3 2.3 1.4 0.5	0 0 0 0	. 1 . 1 . 1	37 39 25 37	604 141 299 613	0 0 0 0	0 0 0 0
LINE 11 A 2120 B 2153 C 2193 D 2232 E 2244 F 2284	760 S S? H S B? S	(F 2 1 2 0 1 0	FLIGHT 8 2 8 2 2 2 2 2	2 32) 3 1 5 0 1 0	13 2 15 2 2 2	38 2 34 2 2 2 2	42 . 4 . 28 . 4 . 4 . 2 .	1.2  1.9  	0 12 - -	· 1 · - · 1 · -	36 _ 44 _ _ _	263 215 - -	0  8  	0 0 0 6 0
LINE 19 A 553 B 601 C 671 D 679 E 702 F 945 G 954	010 S S S S S? B?	(F 1 1 1 1 1 4	TLIGHI 2 2 2 5 8 10	36) 1 0 0 0 0 6	2 2 2 9 8 20	2 2 2 25 18 56	4 . 4 . 4 . 4 . 7 . 52 . 35 .	- - 0.5 0.5 2.9	- - 0 0	· · - · - · - · · - · · - · · · · ·	- - 24 16 0	- - 769 668 453	- - - 0 0	0 0 0 0 0 0
LINE 19 A 701 B 747 C 806 D 822 E 834 F 874 G 901 H 949 I 963 J 978 K 1081 L 1107 M 1233 N 1350 O 1455	020 S B? H S S S S S S S S S S S S S S S S S S S	(F 4 3 5 0 1 1 1 3 1 1 1	FLIGHI 9 1 7 4 7 6 7 6 5 2 6 5 2 3 4 2 4	35) 6 1 4 9 0 1 0 0 1 3 1 2	12 2 6 6 12 11 12 2 6 5 2 7 5 2 8	9 2 28 15 31 41 38 2 20 13 2 14 14 14 2 19	$\begin{array}{c} 4 \\ 1 \\ 24 \\ 28 \\ 29 \\ 4 \\ 28 \\ 20 \\ 4 \\ 13 \\ 10 \\ 4 \\ 22 \\ \end{array}$	3.4 2.9 3.8 5.5 0.5 0.5 0.6 0.7 - 4.8 2.2 - 1.0	0 - 0 7 0 0 - 0 0 - 0 0 - 5 0 - 0	· 1 · - · 1 · 1 · 1 · 1 · 1 · 1 · 1 · 1 · 2 · 1 · 1 · 2 · 1 · 1 · 1 · 1 · 1 · 2 · 1 · 1 · 2 · 1 · 1 · 2 · 1 · 1 · 1 · 2 · 1 · 1 · 1 · 1 · 1 · 1 · 1 · 1 · 1 · 1	27 - 36 49 27 18 6 - 35 55 - 44 42 - 45	90 - 77 46 88 668 616 - 913 - 46 56 - 370	0 - 2 18 0 0 0 - 0 0 - 13 10 - 0	
.*	ES OF	TIMAT THE	CONDU	PIH N ICTOR	íay Bi May I	E UNRI BE DEI	ELIABLE EPER OF	EBECAU R TO ON	SE THE E SIDE	STRON	ER PA	RT. HT.		

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. LINE, OR BECAUSE OF A SHALLOW DIP OR OVERBURDEN EFFECTS. .

COAXIAL COPLANAR					COPI	ANAR	•	VER	TICAL	•	HORIZ	ONTAL ETT	CONDU	CTIVE	MAG	
		5		50		720		•			•		101			CLAU
AN	OMALY/	REAL	QUAD	REAL	QUAD	REAL	QUAD	•	COND	DEPIH*	•	COND	DEPTH	RESIS	DEPTH	
FID	/INIERI	PPM	PPM	PPM	PPM	PPM	PPM	•5	SIEMEN	I M	•	SIEMEN	M	OHMM	M	NT
		-						•			•					
LIN	E 19030	) (1	LICH	r 35)	)			•			•					
Α	547 L	i	2	1	2	2	4	•		-		-	-	-		0
В	507 H	6	4	10	14	19	26	•	9.1	6	•	2	52	43	22	0
С	416 Ba	2 1	2	1	2	2	4	•		-	•	-	-	-		0
D	340 B3	2 1	2	0	2	2	4	•		-			-	-	-	0
Ε	305 L	14	3	6	11	21	8	•	22.8	30	•	1	78	113	31	0
F	293 L	7	6	2	4	2	2	•	7.3	3 0		1	125	440	11	0
G	282 г	7	5	10	7	14	6	•	13.2	2 0	•	2	51	35	18	0
н	262 S	1	3	2	5	9	19	•	1.7	7 0	•	1	26	507	0	0
Ι	220 S7	2 1	4	0	7	22	19	•	0.7	/ 0		1	50	932	0	0
J	202 S	1	- 5	1	6	13	17		1.4	0		1	22	422	0	0
K	168 H	1	2	1	2	. 2	4	•	-	-	•	-	-	-	-	7

.\* ESTIMATED DEPTH MAY BE UNRELIABLE BECAUSE THE STRONGER PART . OF THE CONDUCTOR MAY BE DEEPER OR TO ONE SIDE OF THE FLIGHT .

. LINE, OR BECAUSE OF A SHALLOW DIP OR OVERBURDEN EFFECTS. .







![](_page_98_Figure_0.jpeg)

![](_page_99_Figure_0.jpeg)

![](_page_100_Figure_0.jpeg)

![](_page_101_Figure_0.jpeg)

![](_page_102_Figure_0.jpeg)