

**GEOLOGICAL BRANCH
ASSESSMENT REPORT**

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Assessment Report for the
Sul Claim Group
Golden Mining Division
N.T.S. 82 K/8W
Latitude 50° 26'N, Longitude 116° 23'W

for

Gwen Resources Ltd.
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SUMMARY

The locations of previously unknown kimberlitic lamprophyric and kimberlitic dykes were recently disclosed in a scientific journal (Pope and Thirlwall 1992). At least thirteen separate occurrences were identified in a north-south belt closely associated with the "Windermere High", a local basement high. Independent whole rock and trace element analyses has confirmed the kimberlitic affinity of three of these exposures. The SUL Claim Group was acquired by staking in May, 1992 and consists of 6 2-post claim units which overlie at least one ultrapotassic occurrence. It has been tentatively assigned to Group A (kimberlitic lamprophyre) dykes on the basis of SiO₂, Al₂O₃ and TiO₂ content, mineralogy and texture.

Alkaline intrusive diatremes, dykes and sills have been examined and documented in the southern Rocky Mountain alkaline belt of British Columbia, extending from Fernie to northeast of Golden. Documented lithologies include olivine melilitites, alkaline to basaltic lamprophyres, kimberlites and diamond-bearing lamproites. The Cross kimberlite is a Permo-Triassic intrusive breccia located north of Elkford on the southeastern edge of a broadly defined north-trending belt of alkaline diatremes. The Joff Pipe is located on the western edge of this same belt east of Invermere and is considered to be either a kimberlite (Nassichuk et al. 1989) or an olivine melilite (Pell 1987). The northern portion of this belt is located in the area west of the Columbia Icefields and consists of diamond-bearing lamproites (Nassichuk et al. 1989). Of particular significance regarding this program is the presence of diamond-bearing lamproites in a mobile belt.

The objective of the 1993 program was to identify potential diamond-bearing properties in the Purcell Mountains and secure the ground by staking in order that a more in-depth examination may be carried out in the future. A cursory geological examination was made of the property, with some geological mapping, prospecting, and sampling carried out. Sampling included one 50kg sample of diatreme material for heavy mineral separation and several thin sections for description and lithological identification.

INTRODUCTION

The SUL claim group was staked on the basis of a recent scientific paper describing occurrences of kimberlite and kimberlitic lamprophyre dykes in an entirely new location. Pope and Thirlwall (1992) document thirteen occurrences with whole rock and trace element data for 9 of these localities. The described occurrences are comprised of ultrapotassic, silica-deficient, phlogopite-apatite ultrabasic dykes that compare favourably with world-wide kimberlitic data (Dawson 1980).

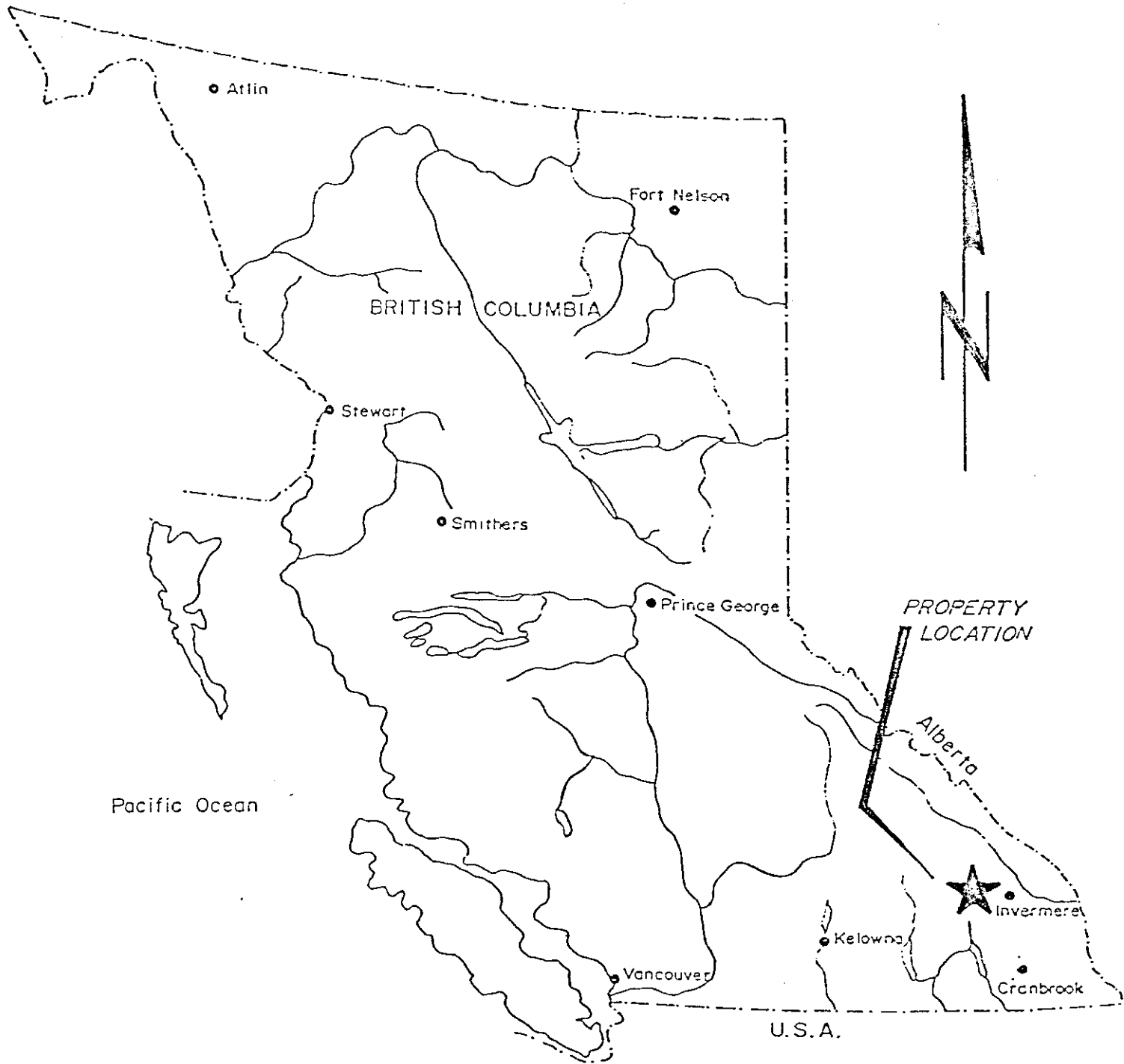
The dykes are distinct from other igneous lithologies in the area which include the Toby volcanics (altered submarine basalts), metadiabase dykes (interpreted to be coeval with Toby volcanics) and the Horsethief Creek quartz monzonite batholith. Sedimentary strata in the area consist of Proterozoic to Lower Paleozoic clastics and carbonates deposited on attenuated basement of the North American miogeocline. Recent mapping (Reesor 1973, Root 1983, Pope 1990) has identified a local basement high, the "Windermere" High", consisting of an inverted, dismembered high standing block of terraced basement which may have controlled emplacement of the ultrapotassic dykes (Pope and Thirlwall 1992).

The ultrapotassic dykes are kimberlitic in composition or of kimberlitic affinity. They have compositions that plot in rift associated and/or active orogen fields of ultrapotassic, major element discrimination diagrams (Pope and Thirlwall 1992). They are coeval with the Cross kimberlite in the southern Rocky Mountains and may be spatially and temporally related to diamond-bearing ultrapotassic lamproites northeast of Golden, B.C.

Dykes in the Toby - Horsethief Creek area (see Figure 1) represent a newly discovered series of ultrapotassic occurrences in an area previously thought devoid of such intrusives. They are exciting in terms of their kimberlitic composition or affinity and possible relationship with diamond-bearing lamproites in the southern Rocky Mountains.

Diamond Exploration Potential

Ultrapotassic/ultrabasic intrusive bodies have recently been identified and described in the Purcell Mountains west of Invermere, B.C. (see Figure 3). All of the occurrences are presently interpreted to be dykes (Pope and Thirlwall 1992) although the Law Creek occurrence (C12 claim group) may prove to be an elongated pipe. These intrusive bodies intrude host strata at high angles, contain abundant carbonate as a primary and/or replacement phase, have narrow chilled margins against host lithologies and show little or no assimilation of xenoliths (when present). Furthermore, they are reported to contain a minor component of ultramafic inclusions (pyroxenite nodules), foreign to host lithologies exposed in the surrounding area (Pope and Thirlwall 1992). These features are common for reported kimberlite intrusions around the world (Dawson 1980).



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SUL CLAIM GROUP

Location Map

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Scale: 1:8,000,000

Fig. No: 1

These occurrences have silica deficient compositions, are potassium enriched, and have high MgO and TiO₂ content (see Table 1). Mineralogically, they contain olivine (partially to completely altered to calcite or serpentine), euhedral to subhedral phlogopite phenocrysts, ilmenite and apatite. The mineralogy and composition (see Table 2) of these dykes compares favourably with kimberlites. Ubiquitous phlogopite as both a matrix and phenocryst phase in some occurrences described in the Toby - Horsethief Creek area is analogous to micaceous (Group II) kimberlites of South Africa.

They have been dated at 245 ± 2.4 Ma and are therefore coeval with the Cross kimberlite in the southern Rocky Mountains (241 ± 5 Ma and 249 ± 12 Ma, Smith et al. 1988; 244 Ma, Grieve 1982). The Cross kimberlite is near the southern end of a belt of alkaline diatremes documented in the Rocky Mountains (Pell 1986, Ijewliw 1986, Helmstaedt et al. 1987, Ijewliw and Schulze 1988, Nassichuk et al. 1989). This belt is approximately 40 kilometres wide and extends from Fernie, B.C. to the Bush Arm of McNaughton Lake (Reservoir), north of Golden, B.C. Alkaline lithologies described in the literature include alkaline to basaltic lamprophyres, olivine melilite, kimberlite and lamproite (Pell 1985, 1986, 1987; Ijewliw 1986; Ijewliw and Schulz 1988; Nassichuk et al. 1989).

Diatremes, dykes and sills of lamproitic composition have been described at the northern end of this belt, west of the Columbia Icefields and northeast of Golden, B.C. (Ijewliw and Schulze 1988). Several diamonds have reportedly been recovered from these lamproitic occurrences (JACK claims, 1 microdiamond and 1 macrodiamond recovered; MARK claims, 1 microdiamond recovered; Nassichuk et al. 1989) although not in economic quantities. However, diamonds have been recovered from ultrapotassic occurrences in the southern Rocky Mountains, a mobile belt.

The presence of dykes (and possible pipes) having kimberlitic composition in a previously unknown location together with documented diamond-bearing ultrapotassic intrusive bodies farther north is sufficient to warrant further exploration in the Toby - Horsethief Creek area to:

- 1) determine the diamond potential of these occurrences using an integrated exploration program including:
 - a) soil sample geochemistry,
 - b) whole rock geochemistry,
 - c) heavy mineral and xenolith inclusion suites, and
 - d) identification of kimberlitic indicator mineral suites.
- 2) identify additional occurrences of kimberlitic dykes through mapping, prospecting and remote sensing, and
- 3) determine the possible presence and location of hidden diatremes in subsurface or under cover using geophysics.

LOCATION AND ACCESS

The SUL claim group is located in the Purcell Mountains (latitude 50° 26'N, longitude 116° 23'W), approximately 35 kilometres west-southwest of the community of Invermere, B.C. on N.T.S. mapsheet 82 K/8W (see Figure 1). The claim group consists of 6 2-post claim units located over and adjacent to Sultana Creek, a south flowing tributary of Delphine Creek.

The claim group can be accessed by rough roads from Invermere, B.C. along Toby and Delphine Creeks. The road along Delphine Creek is passable for approximately 3 kilometres to a washout and then by foot or motorcycle for an additional 3 kilometres along road and trail. The condition of the road beyond the washout is adequate for vehicles to pass but the washout itself would require work and stabilization to re-establish should the need arise.

CLAIM STATUS

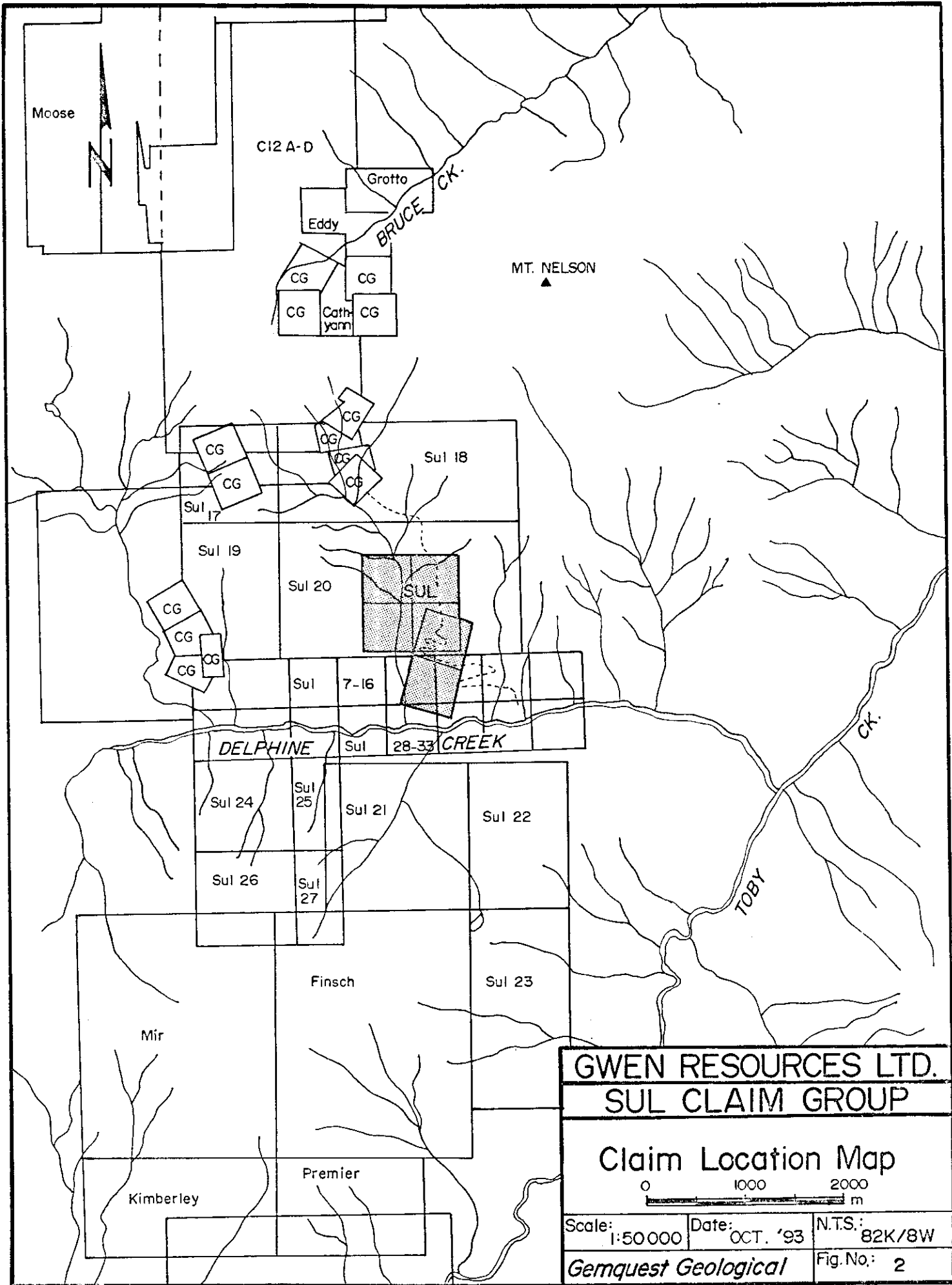
The property consists of 6 2-post claims (see Figure 2), staked in accordance with existing government claim location regulations. Significant claim data are summarized below:

| Claim Name | Units | Tenure # | Date of Record | Expiry Date* |
|------------|-------|----------|----------------|---------------|
| SUL 1 | 1 | 311872 | July 16, 1992 | July 16, 1995 |
| SUL 2 | 1 | 311873 | July 16, 1992 | July 16, 1995 |
| SUL 3 | 1 | 311874 | July 16, 1992 | July 16, 1995 |
| SUL 4 | 1 | 311875 | July 16, 1992 | July 16, 1995 |
| SUL 5 | 1 | 311876 | July 16, 1992 | July 16, 1995 |
| SUL 6 | 1 | 311877 | July 16, 1992 | July 16, 1995 |

Total: 6

Copies of claim application forms are provided in Appendix B.

*After 1993 assessment credit applied.



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SUL CLAIM GROUP

Claim Location Map

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PHYSIOGRAPHY AND CLIMATE

The SUL claim group is located within and adjacent to Sultana Creek. Relief in the area varies from 1433 metres (4700 feet) along Delphine Creek to more than 3307 metres (10,850 feet) on Mt. Nelson.

The claims are moderately well exposed along Sultana Creek and generally poorly exposed throughout the remainder of the claims. Vegetation in the area consists of a mixture of coniferous and deciduous trees with the undergrowth comprised largely of slide alder and/or Devil's Club.

The claims are located east of Kootenay and Trout Lakes in a regional topographic high and are therefore subject to heavier precipitation than further north and south. As a result, the region has many icefields and small glaciers developed in the area. The property is available for geological exploration from May to late October.

HISTORY

The area between Toby and Horsethief Creeks has over 80 documented mineral occurrences for commodities such as lead, silver and copper. These occurrences vary in size from small showings up to the 1.2 million tonne Mineral King deposit just north of Toby Creek and west of Jumbo Creek. Other mineral deposits include (from south to north): Silver Spray, Mineral King, Kootenay Queen, Hot Punch, Delphine, Nip and Tuck, Silver Queen, Ptarmigan and the Iron King. The majority of the mineral occurrences are fault and/or vein type deposits.

These mineral occurrences are coincident with a north-south belt of lamprophyric to kimberlitic dykes interpreted to be closely associated with the locus of the "Windermere High". Mineral deposits proximal to known dyke occurrences are briefly described below.

Mineral King Mine

The Mineral King mine is located north of Toby Creek and west of Jumbo Creek at an elevation between 1220 metres (4005 feet) and 1670 metres (5480 feet). The mine produced a total of 1,334,400 tons of ore at a reported grade of 8 percent zinc and 3 percent lead per ton before suspending operations in 1964 (Pope 1990).

The deposit is interpreted as a Stratabound Massive Replacement in a high angle fault panel within the Mineral King duplex system in the footwall of the Mount Forster fault. The orebody is hosted in brecciated dolomite of the Lower Gateway Formation and consists of galena, sphalerite, tetrahedrite (containing 6 to 7 percent silver), pyrite and barite with minor chalcopyrite and pyrrhotite (Pope 1990).

Kootenay Queen Group

The Kootenay Queen Mine is located in a cirque south of Delphine Creek at an elevation of 1980 metres (6495 feet). The orebody is hosted within the Mount Nelson Formation (white marker member), immediately below the interpreted position of the Windermere Unconformity. The orebody consists of galena, tetrahedrite and sphalerite with reported recovery of less than 100 tons having a grade of 2400 grams per tonne silver and 70 percent lead.

Silver Queen Mine

The Silver Queen mine is present at the base of a cliff on the west side of Mt. Slade at an elevation of 2900 metres. The mine consists of a system of small veins situated in the lower main dolomite of the Hadrynian Mount Nelson Formation and is associated with a green metadiabase dyke. The main workings were within a 20 centimetre wide vein hosting galena and sphalerite with minor chalcopyrite having a reported production of less than

100 tons with a grade of 2.35 kilograms per tonne silver and 59 percent lead.

Silver Spray Mine

The Silver Spray Mine is situated on the west side of Coppercrown Creek at an elevation of 2290 metres and is part of a group of claims which includes the Lady Bing, Gracie Fraction, Betsy and IOU properties. The workings are contained in the dolomite dominated upper portion of the Lower Gateway Formation immediately below the unconformably overlying Dutch Creek Formation. Up to 50 tons of ore were recovered consisting of galena, tetrahedrite and cerussite with minor sphalerite and copper carbonates in vertical and bedding parallel fractures.

Pretty Girl Group

The Pretty Girl Group is situated on the ridge crest between Law and Bruce Creeks at an elevation of 2720 metres (8925 feet). Stratigraphically the Pretty Girl Group is located within argillites of the Horsethief Creek Formation. The mineralization is reported to consist of tetrahedrite and chalcopyrite in a discontinuous quartz vein with less than 50 tons recovered at grades up to 188 grams per tonne silver and 27 per cent copper.

REGIONAL GEOLOGY

Stratigraphy

The stratigraphy of the Purcell Mountains (see Figure 3) consists of four separate and distinct, megascopic miogeoclinal sequences interpreted to have been deposited on passive North American continental crust. This Helikian to Lower and Upper Paleozoic package has undergone four major phases of deformation and local thermal metamorphism related to the Horsethief Creek Batholith. Igneous activity has episodically affected the sedimentary sequence and includes syn-depositional basaltic to andesitic flows and/or sills to post depositional intrusive dykes, sills and batholiths.

The sedimentary sequence exposed between Toby and Horsethief Creeks (the Toby Creek area) comprises the uppermost Helikian Belt-Purcell Supergroup, the Hadrynian Windermere Supergroup, and Lower Paleozoic strata to the Middle Devonian Starbird Formation. These strata are exposed in six separate panels bounded by thrust faults, and carried in the hanging wall of the northeast vergent Purcell Thrust (Pope 1990).

Proterozoic

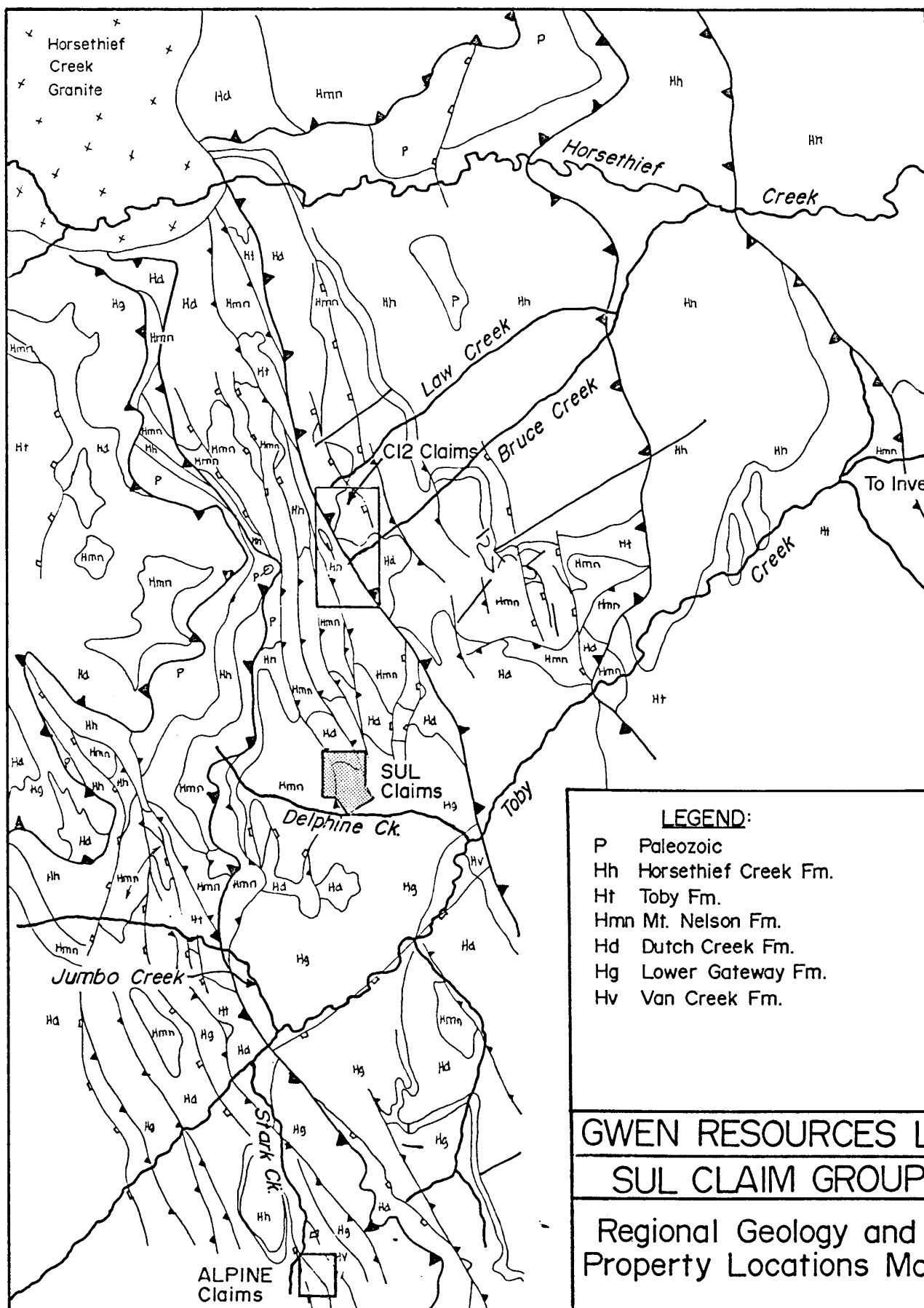
Belt-Purcell Supergroup

The Helikian Belt-Purcell Supergroup has an exposed thickness of 4300 metres (14,100 feet), from within the Van Creek Formation to the Mount Nelson Formation. The Belt-Purcell Supergroup is comprised predominantly of cliff-forming, buff weathering dolomitic lithologies with intercalated siliciclastic intervals.

The Van Creek Formation is the lowest formation exposed in the Toby Creek area. It consists of approximately 500 metres of medium- to coarse-grained, light grey to dark green quartzites, siltstones and silty argillites exposed in the core of an anticline. The Nicol Creek Formation is absent as the Van Creek quartzites apparently grade upward into over 1000 metres of pale green quartzites, silts and buff-weathering dolomitic silts of the Lower Gateway Formation.

The Lower Gateway Formation has been subdivided into two members, a basal transitional sequence and an upper dolomite dominated sequence. The transitional sequence is up to 100 metres thick. The base is identified as the first occurrence of carbonate above which are distinctive thin bedded, red spotted quartzites with interbedded green siltstone and buff weathering dolomitic siltstone and dolomite.

The Upper Gateway Formation is dominated by thin bedded dolomite which passes upward into a 90 metre thick, cream to buff weathering dolomitic unit. The dolomite has cryptalgal and stromatolitic laminations and cream coloured chert



LEGEND:

| | |
|-----|----------------------|
| P | Paleozoic |
| Hh | Horsethief Creek Fm. |
| Ht | Toby Fm. |
| Hmn | Mt. Nelson Fm. |
| Hd | Dutch Creek Fm. |
| Hg | Lower Gateway Fm. |
| Hv | Van Creek Fm. |

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SUL CLAIM GROUP

**Regional Geology and
 Property Locations Map**

| | | |
|----------------------------|----------------|----------------|
| Scale: 1:200,000 | Date: OCT. '93 | N.T.S.: 82K/8W |
| GEMQUEST GEOLOGICAL | | Fig. No.: 3 |

intercalations. The dolomite ranges from blue-grey micrite to light coloured coarse sucrose textured dolomite.

A sharp contact has been mapped separating the Upper Gateway Formation from the overlying Dutch Creek Formation. The contact is a narrow, rusty-weathering zone interpreted to represent a hiatus along a parallel unconformity. The Dutch Creek Formation varies from 300 to 1000 metres over less than 5 lateral kilometres and consists of dark coloured, fine-grained quartzite-argillite couplets.

The contact with the overlying Mount Nelson Formation is always very sharp with an abrupt change in facies and sedimentary characteristics evident across the contact, which is interpreted as a paraconformity. The Mount Nelson Formation is approximately 1300 metres thick, consisting of thick, well-bedded white orthoquartzite, buff weathering dolomites and purple weathering dolomites and argillites.

The Mount Nelson Formation has been subdivided into the:

- a) lower quartzite, a useful 50 to 150 metre thick marker horizon consisting of white, well-sorted, fine- to medium-grained pure quartz arenites,
- b) lower main dolomite - an approximately 400 metre thick sequence which conformably overlies and is gradational with the lower quartzite, comprised of cryptalgal to stromatolitic laminated, pale grey weathering dolomites with interbedded carbonaceous argillites capped by a cream-coloured stromatolitic, crystalline cherty-dolomite unit approximately 20 metres thick overlain in sharp contact by,
- c) the middle quartzite - an apple green coloured sequence consisting of massive, fine- to coarse-grained quartz arenites, impure sandstones and argillites having A-B to A-E Bouma sequences evident,
- d) orange dolomite sequence - approximately 180 metres thick consisting of varicoloured buff weathering dolomitic siltstones, argillites and impure sandstones underlying bright orange-buff weathering silty and sandy crystalline dolomites with abundant cryptalgal and stromatolitic laminations and intercalated chert.
- e) white markers conformably overlie the orange dolomite and are up to 70 metres thick. The white markers consist of cream, buff and silver-grey dolomites with purple, green and buff dolomitic mudstones and local interbeds of pure white magnesite up to 1 metre thick,
- f) purple sequence - gradationally overlies the white markers, consisting of purple weathering dolomitic sandstones and siltstones which grade upward into purple weathering

argillite. Mudchip breccias and monomict pebble conglomerates are interbedded with siltstones and argillites and the sequence is overlain by a pebble to boulder conglomerate with a purple weathering sandy argillitic matrix in sharp contact with the purple shales. The pebble to boulder conglomerate is the interpreted locus of an intraformational unconformity with a thickness between 2 and 10 metres thick,

- g) upper middle dolomite - approximately 80 metres thick and similar to the lower main dolomite. It is distinguished by abundant algal allochems which are typically replaced by black chert,
- h) upper quartzite - a distinctive cliff-forming unit consisting of white quartzites more than 260 metres thick (equivalent to the upper Mount Nelson Quartzite (Atkinson 1975)). The upper quartzite consists of well sorted medium- to coarse-grained, essentially pure arenites. They are distinguished from the lower quartzite on the basis of massive bedding and poorly preserved sedimentary structures.
- i) upper dolomite - the uppermost unit in the Belt-Purcell exposed below the Windermere unconformity. The upper dolomite is gradational with the underlying quartzite over 10 metres consisting of interbedded purple argillite, quartzite and dolomite. The upper dolomite is comprised of pale to dark grey dolomite interbedded with quartz and dolomite pebble conglomerates with dolomitic quartz sands.

Windermere Supergroup

The Windermere Supergroup varies in thickness in the Toby Creek area, from 80 metres to over 3 kilometres and is in sharp contact with the underlying Belt-Purcell Supergroup across an unconformity with considerable topography, interpreted as a result of a local basement high, the "Windermere High" (Reesor 1973). The Windermere Supergroup was deposited above this unconformity and consists of a basal conglomeratic unit, the Toby Formation, and the overlying argillite and pebble conglomerate dominated Horsethief Creek Formation.

The Toby Formation is the basal unit of the Windermere Supergroup and overlies different levels of the Belt-Purcell stratigraphy in the separate fault panels, interpreted to indicate active faulting during sedimentation (Pope 1990). Four distinct facies have been identified in the Toby Creek area but their stratigraphic position relative to one another is uncertain due to rapid lateral facies changes.

The Toby Formation consists of:

- a) a basal boulder breccia lithofacies consisting of monomict clast-supported boulder breccias.

- b) a diamictite lithofacies - the most commonly developed facies consisting of rounded quartzite and subangular dolomite boulders (derived from the immediately underlying Mount Nelson Formation) in a sandy argillite matrix.
- c) a sparse clast diamictite lithofacies consisting of graded fine- to coarse-grained, poorly sorted arenites and argillites with a minor component of rounded quartzite pebbles or cobbles.
- d) a siltstone-argillite lithofacies which comprises the bulk of, and is the dominant lithology in, the upper portion of the Toby Formation, consisting of well-sorted and graded fine quartz arenites and argillites which typically exhibit complete Bouma sequences.

The Toby volcanics are the oldest igneous rocks identified in the Toby Creek area and are believed to be altered submarine basalts related to regional Hadrynian extension. The flows are holocrystalline and glomeroporphyritic basaltic andesites, having plagioclase phenocrysts in a fine-grained plagioclase groundmass.

Green metadiabase dykes have also been identified and have been interpreted as the metamorphic equivalent to the Toby volcanics. They are the most common igneous rocks and are always intruded at a high angle to bedding. They are typically altered, consisting of anhedral masses of chlorite, anhedral to euhedral carbonate and sericite and skeletal opaques. Chlorite pseudomorphs after pyroxene and amphibole have been identified. Bulk mineralogical proportions indicate these dykes were most probably originally basaltic in composition and have been subsequently hydrated.

The Toby Formation is gradational into the overlying Horsethief Creek Formation, in which five lithofacies have been identified. These lithofacies define a rudimentary stratigraphy of facies within the Horsethief Creek Formation as individual lithological units are inconsistent due to rapid lateral thickness and facies variations.

The lithofacies identified in the Horsethief Creek Formation are as follows:

- a) siltstone-argillite - dominant in the lower half of the Horsethief Creek Formation and separate the remaining lithofacies throughout the formation. This lithofacies consists of thick sequences of thin bedded, graded siltstone and argillite and finely laminated black, green and grey argillite.
- b) black carbonate - an easily traced marker used to identify and map the base of the Horsethief Creek Formation consisting of thin bedded, dark grey to black limestone with variable

quartz sand and silt in a calcitic matrix and thin calcareous quartz-arenite beds.

- c) dolomite - buff weathering dolomite, up to 30 metres thick, dolomite pebble-conglomerate beds and dolomite supported quartzite occur throughout the Horsethief Creek Formation.
- d) quartz feldspar arenites and pebble conglomerates - consist of pebble conglomerates comprised of grain-supported crystalline quartz and quartz feldspar grains with variable red jasper, green to grey argillite, quartzite and dolomite clasts in a quartz, feldspar, carbonate, sericite and chlorite matrix. Clasts are generally 1 to 2 centimetres in diameter but may exceed 10 centimetres in length. Quartz feldspar arenite beds are similar to the pebble conglomerates but have a greater proportion of matrix and are generally poorly sorted.
- e) red and varicoloured argillites - are present at the top of the Horsethief Creek Formation and consist of variably coloured argillites with interbedded pink carbonate, and varicoloured impure arenites.

Lower Paleozoic

The Paleozoic succession is comprised of the Lower Cambrian Cranbrook Formation, Middle Cambrian Jubilee Formation, Ordovician-Silurian Beaverfoot Formation, Middle Devonian Mount Forster Formation and the Upper Devonian Starbird Formation. The Paleozoic stratigraphy neither hosts nor have Paleozoic clasts been identified in kimberlitic dykes and therefore will not be described at this point. The reader is referred to Pope (1989), Root (1985, 1983) and Reesor (1973), for a complete description of Paleozoic stratigraphy in the Invermere area of the Purcell Mountains.

Middle Cretaceous

The Horsethief Creek Batholith (see Figure 3) is a quartz monzonite intrusion present north of Horsethief Creek and therefore out of the Toby Creek area. However, granitic apophyses and aplitic dykes are present throughout the Toby Creek area and thermal metamorphism related to the batholith has affected the strata of the area.

Permo-Triassic

Ultrabasic to ultrapotassic dykes have been recently described (Pope and Thirlwall 1992) extending from Stark Creek northward to Law Creek (see Figure 3). These intrusive dykes define a broad north-south belt spatially associated with the intersection of the Nelson Creek Fault and the Bruce Creek Synform, possible basement control related to the locus of the Windermere High.

The dykes occur as rusty weathering, variably carbonatized intrusive bodies ranging from 50 centimetres to 10 metres thick. They have narrow chilled margins and show little or no evidence of contact metamorphism with either the host lithologies or xenoliths. This has been interpreted as evidence for rapid intrusion and quick cooling of the dyke material.

The dykes are xenolithic, ranging from approximately 5 percent to more than 50 percent by volume, and include: pyroxenite nodules, abundant Belt-Purcell quartzite and argillite lithologies and quartz-feldspar xenocrysts derived from either Hudsonian basement or Mesozoic-Cretaceous granitoid intrusives.

The occurrences have been subdivided into two suites based on petrology and chemistry (Table 2). Group A are considered to be lamprophyres having kimberlitic affinity whereas Group B dykes are considered to be true kimberlites (Pope and Thirlwall 1992).

Group A dykes are typically light green in colour with phlogopite phenocrysts absent to abundant (porphyritic) in a carbonate rich matrix with carbonate-iron oxide pseudomorphs after euhedral olivine. Olivine pseudomorphs are recognized by their crystal outline and relict serpentine fractures. The matrix consists of fine-grained carbonate, opaques and iron-stained, fine-grained laths, possibly an alteration product of mica or feldspar.

Xenoliths are abundant in Group A dykes, typically more than 50 percent xenoliths by volume, and are almost exclusively derived from the underlying Belt-Purcell Supergroup. However, granitic gneiss and ultrabasic xenoliths have also been reported (Pope 1990). Group A dykes correspond to the lamprophyre-kimberlite field (Pope and Thirlwall 1992), transitional between rift and active orogen type volcanics.

Group B dykes are typically dark green in colour. They are porphyritic with phlogopite phenocrysts up to 8 centimetres long and apatite phenocrysts up to 0.5 millimetres in diameter poikilitically enclosed by phlogopite. Phlogopite crystals are also intergrown with calcite, suggesting that calcite may also be primary.

Matrix phases identified include: carbonate, serpentine, chlorite, phlogopite, apatite and opaques. The presence of abundant matrix calcite distinguishes Group B dykes from lamproites (Pope and Thirlwall 1992, after Dawson 1987). The xenolith content of Group B dykes is typically less than 10 percent by volume.

Group B dykes are petrologically similar to kimberlites (Pope and Thirlwall 1992) and are classified as continental rift associated ultrapotassic rocks with kimberlitic affinities. They can be further defined as micaceous kimberlites, due to the high

proportion of phlogopite, matrix calcite and apatite (Pope and Thirlwall 1992).

Table 1: Whole rock chemistry for ultrapotassic - ultrapotassic dykes in the Toby - Horsethief Creek area (modified from Pope and Thirlwall 1992). The "A" series analyses are from Group A (kimberlitic lamprophyre) dykes, "B" series analyses are from Group B (true kimberlite) dykes while the last three are independent analyses of material from the Toby - Horsethief Creek area.

| | A1 | A2 | A3 | A4 | A5 | A6 | B1 | B2 | B3 | C-12 | Alpine | Suitana |
|------------------------------------|---------------|--------------|--------------|--------------|--------------|--------------|--------------|--------------|--------------|---------------|--------------|--------------|
| SiO ₂ | 41.49 | 39.89 | 38.08 | 40.44 | 40.16 | 37.25 | 34.12 | 35.45 | 33.35 | 43.69 | 37.68 | 37.65 |
| Al ₂ O ₃ | 12.71 | 11.57 | 12.00 | 11.09 | 12.02 | 12.77 | 7.78 | 6.92 | 7.81 | 11.44 | 5.00 | 5.04 |
| Fe ₂ O ₃ | 14.32 | 12.77 | 13.10 | 12.37 | 14.72 | 12.61 | 15.60 | 16.27 | 15.93 | 15.30 | 9.41 | 9.30 |
| MgO | 11.17 | 14.65 | 15.09 | 14.09 | 11.15 | 15.10 | 14.18 | 17.58 | 14.43 | 5.33 | 13.24 | 13.68 |
| CaO | 14.11 | 14.87 | 15.23 | 15.85 | 14.18 | 13.66 | 15.74 | 11.83 | 14.71 | 6.95 | 13.97 | 13.95 |
| Na ₂ O | 0.05 | 0.04 | 0.02 | 1.23 | 0.02 | 0.06 | 0.00 | 0.01 | 0.06 | 3.06 | 0.02 | 0.01 |
| K ₂ O | 2.67 | 1.60 | 1.73 | 0.54 | 2.38 | 4.07 | 5.33 | 4.99 | 5.56 | 0.06 | 2.14 | 1.85 |
| TiO ₂ | 3.31 | 2.80 | 2.97 | 2.54 | 3.15 | 3.11 | 2.64 | 2.68 | 2.68 | 2.58 | 0.98 | 0.98 |
| P ₂ O ₅ | 0.74 | 0.71 | 0.72 | 0.68 | 0.72 | 0.66 | 1.81 | 2.09 | 1.97 | 0.27 | 0.31 | 0.33 |
| MnO | 0.20 | 0.20 | 0.21 | 0.22 | 0.20 | 0.21 | 0.33 | 0.22 | 0.29 | 0.22 | 0.17 | 0.17 |
| LOI | 19.87 | 20.19 | 20.89 | 21.60 | 19.03 | 18.84 | 18.83 | 19.26 | 17.36 | 11.10 | 16.99 | 16.99 |
| Total | 100.80 | 99.10 | 99.14 | 99.07 | 98.65 | 99.49 | 97.51 | 98.04 | 96.99 | 100.00 | 99.91 | 99.85 |
| K ₂ O/Na ₂ O | 53.4 | 40 | 86.5 | 0.43 | 119 | 67 | 499 | 93 | 0.02 | 107 | 185 | 2.6 |
| Ni | 263 | 335 | 362 | 306 | 217 | 326 | 507 | 631 | 478 | 32 | 131 | 110 |
| Cr | 689 | 808 | 857 | 917 | 669 | 489 | 650 | 698 | 591 | 59 | 627 | 111 |
| V | 366 | 323 | 334 | 328 | 351 | 304 | 258 | 285 | 265 | 268 | 152 | 43 |
| Sr | 654 | 928 | 992 | 564 | 646 | 342 | 531 | 544 | 514 | 143 | 381 | 129 |
| Rb | 88 | 52 | 55 | 27 | 76 | 131 | 340 | 348 | 356 | --- | --- | --- |
| Ba | 309 | 238 | 280 | 116 | 262 | 601 | 511 | 623 | 588 | 50 | 60 | 270 |
| Zr | 213 | 186 | 200 | 170 | 206 | 173 | 230 | 228 | 235 | --- | --- | --- |
| Nb | 110 | 104 | 127 | 127 | 107 | 113 | 67 | 69 | 73 | --- | --- | --- |
| Y | 30 | 27 | 29 | 27 | 29 | 26 | 21 | 19 | 23 | 6 | 15 | 6 |
| La | 62 | 54 | 58 | 62 | 59 | 63 | 37 | 39 | 38 | 10 | 10 | 20 |
| Ce | 119 | 106 | 110 | 111 | 115 | 108 | 77 | 78 | 75 | --- | --- | --- |
| Nd | 50 | 45 | 46 | 44 | 49 | 42 | 35 | 34 | 33 | --- | --- | --- |

Structure

Four major phases of deformation have been identified in the Toby Creek area, Helikian-Devonian extension (D1), Jurassic-Paleocene contraction (D2-D3) and Eocene extension (D4).

The first phase of deformation resulted in unconformities at the base of the Dutch Creek and Mount Nelson Formations (D1a) and the unconformity at the base of the Windermere Supergroup (D1b). Thinning of Paleozoic strata onto the Windermere High is interpreted to reflect the effects of D1c deformation together with the development of small fault-bounded sub-basins.

Contraction during the Columbian (D2) and Laramide (D3) orogenies resulted in a series of northeast vergent thrust faults and the development of a regional foliation (S1). Three major thrust sheets are evident in the Toby Creek area with one, the Mount Nelson thrust sheet, comprised of four smaller fault panels. The three major thrust sheets represent out-of-sequence faults, having propagated toward the hinterland, carried in the hanging wall of the Purcell Thrust.

Contraction during D2 and D3 produced east-vergent imbricate thrust faults and west vergent backthrusts. Many of these faults were subsequently reactivated during the fourth phase (D4) of deformation. High angle brittle faults are also a result of D4.

LOCAL GEOLOGY

The SUL claim group contains two known ultramafic occurrences, one at approximately 1737 metres (5700 feet) in the Sultana Creek drainage and the other along a trail at approximately 1722 metres (5650 feet), south of Sultana Creek. The occurrence within Sultana Creek is in situ outcrop while the occurrence along the trail may be proximal float. One day was spent by the authors sampling and mapping along Sultana Creek and so the nature of the exposure along the trail is currently unresolved.

The occurrence along Sultana Creek is present on the southern bank and is exposed through glacio-fulvial cover. The outcrop extends approximately 10 metres above the level of the creek and approximately 5 meters along the creek. It is a medium green weathering rock with light to medium orange weathering interstitial carbonate. Phlogopite phenocrysts are present up to 0.75 cm in long dimension and imparts a porphyritic texture to the exposure. The phlogopite has a greenish colour and is interpreted to be partially altered to chlorite. The abundance and nature of the carbonate suggests that much of it is primary but additional secondary carbonate has been developed through alteration of primary olivine. Primary olivine is recognizable, in thin section, by relict crystal shape (outlined by opaques) and characteristic serpentine fractures.

The occurrence is located at or near the locus of three thrust faults and has intruded host lithologies of the Hadrynian Dutch Creek Formation (Pope 1990). Geological mapping by the authors is not sufficient at this time to determine the timing of intrusion relative to the thrust faults however it is interpreted to pre-date thrusting as it has been foliated. In addition, the nearby Toby Creek occurrence has been dated at 245 ± 2.4 Ma and therefore, as the intrusives are almost certainly coeval, was intruded during the earliest stages of the Columbian Orogeny.

1992-93 PROGRAM

A total of two man-days were spent on the property by the authors, in an attempt to determine the economic potential of the property. On June 10th, 1992, the authors hiked in to the property and collected a 50kg sample of the Sultana Creek diatreme material, as well as completed limited geologic mapping of the surrounding area (1:10,000 scale).

A total of 3 hand samples were taken from the property area, and were representative of diatreme lithologies encountered. Two were taken from float in Sultana Creek and the third from the occurrence along the trail at approximately 1722 metres. Thin sections were made of the selected material which was representative of the occurrences. Petrographic examinations and lithological identification were completed by R. Walker, and a single thin section was examined and described by S. Digel at the University of Calgary (see Appendix C).

One 50kg bulk sample was taken from diatreme float in Sultana Creek, representing various phases of the diatreme. Furthermore, the 50kg sample was split into two sub-equal portions. One split was kept in Cranbrook for sample duplication and the remainder shipped to Loring Labs in Calgary, where it was crushed to -6 mesh. A representative portion was split from the crushed and homogenized sample and analyzed for whole rock, gold and 47-element ICP at Activation Labs in Ontario (geochemical analyses have not been returned as of this writing and are still pending). The remaining crushed material (18.5 kg) was then placed in an acid bath for 24 hours to remove carbonates, washed, deslimed and further processed through heavy-liquid media.

The heavy liquid separation resulted in two sets of mineral separates: one containing minerals having specific gravities between 2.96 and 3.30 and a second having minerals greater than 3.30 S.G.. The two sets were further separated using a Frantz Isodynamic Magnetic Separator into four fractions: non-magnetic, weakly para-magnetic, para-magnetic and magnetic. The resulting heavy mineral concentrates were examined by R. Walker. Representative indicator minerals were picked under binocular microscope and have been sent away for electron microprobe analysis at the University of Calgary to confirm mineral composition.

Samples analyzed by ICP methods were crushed to -80 mesh and dissolved in aqua-regia solution.

RESULTS

One rock sample taken in October, 1992 from each of Law Creek, Sultana Creek and Stark Creek was sent to Eco-Tech Laboratories in Kamloops, B.C. for whole rock and ICP analysis. The results are tabulated in Table 1 and the complete results are included as Appendix C. A representative split from sample SUL93-2 was sent for geochemical analysis (whole rock and "Gold + 47" elements) but unfortunately no results have been returned.

It can be seen that the whole rock compositions obtained compare favourably with the results documented for the Toby - Horsethief Creek area by Pope and Thirlwall (1992). The results confirm ultrabasic (<45 wt.% SiO₂), potassium-enriched compositions having high K₂O/Na₂O ratios and high Mg values (>3 wt.%). Furthermore, the trace element data (see Table 2) is generally comparable to kimberlitic values established from worldwide occurrences as compared to the range for average ultramafic compositions. Deviation from kimberlitic compositions is expected as the samples submitted for geochemical analysis were not free of contamination by host lithologies. Xenoliths of host lithologies are present and include both carbonate- and silica-rich compositions (eg. Upper Gateway Formation and Mount Nelson Formation). Therefore, a characteristic kimberlitic composition will be diluted by inclusion of sedimentary inclusions (xenoliths). Such dilution is interpreted to be evident in the independent analyses from samples taken in October, 1992 (Columns C-E in Table 2) whereas samples taken by Pope and Thirlwall (1992) from the same occurrences have values more comparable to world-wide kimberlite values.

Sample SUL93-2 was a bulk sample taken of diatreme material found in Sultana Creek below the occurrence itself. It was processed at Loring Laboratories as described in "1992-93 Program" above. The weights of the fractions obtained through heavy liquid separation are tabulated below.

| Sample ID | Original Weight (kg) | Middlings (g) (2.96-3.3 S.G.) + 28 Mesh | | | | Heavies (>3.3 S.G.) - 28 Mesh | | |
|-----------|----------------------|--|--------------|-----------|---------------|----------------------------------|---------------------|--------------|
| | | Mag. (g) | Non-Mag. (g) | Magn. (g) | Non-Magn. (g) | Para. (g) | Weakly Non-Mag. (g) | Non-Mag. (g) |
| Sul93-2 | 18.5 | 0.19 | 18.34 | 0.48 | 0.78 | 0.09 | 0.29 | 1.23 |

Initial results have been returned from electron microprobe analysis at the University of Calgary and are tabulated in Table 3 to 5. Clinopyroxenes (Table 3) are subhedral to anhedral, light to medium translucent grass green in colour and are interpreted as diopsides. Microprobe analysis confirms a calcium rich and while chrome-bearing are not the high chrome diopsides considered an indicator mineral. However, Nassichuk et al. (1989) have derived a factor plot for low-chrome diopsides associated with eclogitic diamond inclusions from world-wide occurrences. The clinopyroxene data has not been compared with these data however.

Table 2: Representative trace element content for ultramafic and kimberlitic lithologies. Column A - average ultramafic, B - kimberlite, C - C12 (Law Creek), D - Alpine West, E - Sultana Creek, F - Group A kimberlitic lamprophyres and G - Group B (true kimberlites). Column A and B data from Dawson (1980); F and G from Pope and Thirlwall (1992). All values in ppm.

| Element | A | B | C | D | E | F | G |
|---------|------|----------|-----|-----|-----|---------|---------|
| B | 7 | 0-170 | 2 | 2 | 2 | | |
| Ba | 20 | 137-1970 | 50 | 60 | 270 | 116-601 | 511-623 |
| Ce | 1.93 | 45-522 | | | | 106-119 | 75-78 |
| Co | 110 | 35-130 | 43 | 35 | 18 | | |
| Cr | 3090 | 550-2900 | 59 | 627 | 111 | 489-917 | 591-698 |
| Cu | 47 | 10-300 | 160 | 2 | 68 | | |
| La | 0.92 | 26-200 | 10 | 10 | 20 | 54-63 | 37-39 |
| Nb | 1.3 | 32-450 | | | | 104-127 | 67-73 |
| Nd | 1.44 | 25-180 | | | | 42-50 | 33-35 |
| Ni | 1450 | 710-1600 | 32 | 131 | 110 | 217-362 | 478-631 |
| Sr | 22 | 40-1900 | 143 | 381 | 129 | 342-992 | 514-544 |
| V | 50 | 21-250 | 268 | 152 | 43 | 304-366 | 258-285 |
| Y | 2.9 | 4-75 | 6 | 15 | 8 | 26-30 | 19-23 |
| Zn | 56 | 15-287 | 78 | 23 | 46 | | |
| Zr | 16 | 84-700 | | | | 170-213 | 228-235 |

Table 3 - Electron Microprobe Analyses of clinopyroxenes

Clinopyroxene - Sultana Claim Group

| Formula Based on 6 O | | | | | | | | | | Number of ions in formula | | | | | | | | | | End Member Components | | |
|----------------------|------|-------|-------|------|-------|-------|------|-------|--------|---------------------------|------|------|------|------|------|------|------|------|------|-----------------------|-------|-------|
| SiO2 | TiO2 | Al2O3 | FeO* | MnO | MgO | CaO | Na2O | Cr2O3 | Total | Si | Ti | Al | Fe | Mg | Mn | Ca | Na | Cr | Sum | X Wo | X En | X Fs |
| 52.79 | 0.17 | 1.05 | 7.34 | 0.16 | 14.86 | 22.37 | 0.48 | 0.01 | 99.23 | 1.97 | 0.00 | 0.05 | 0.23 | 0.83 | 0.01 | 0.90 | 0.03 | 0.00 | 4.02 | 0.459 | 0.424 | 0.117 |
| 54.00 | 0.15 | 0.63 | 5.95 | 0.15 | 15.99 | 22.61 | 0.41 | 0.03 | 99.92 | 1.99 | 0.00 | 0.03 | 0.18 | 0.88 | 0.00 | 0.89 | 0.03 | 0.00 | 4.01 | 0.457 | 0.449 | 0.094 |
| 55.58 | 0.19 | 0.46 | 3.61 | 0.08 | 17.53 | 23.27 | 0.19 | 0.05 | 100.96 | 2.00 | 0.01 | 0.02 | 0.11 | 0.94 | 0.00 | 0.90 | 0.01 | 0.00 | 3.99 | 0.461 | 0.483 | 0.056 |
| 53.95 | 0.16 | 0.49 | 7.82 | 0.23 | 14.50 | 21.48 | 0.92 | 0.07 | 99.62 | 2.01 | 0.00 | 0.02 | 0.24 | 0.80 | 0.01 | 0.86 | 0.07 | 0.00 | 4.01 | 0.450 | 0.422 | 0.128 |
| 54.16 | 0.15 | 0.78 | 5.04 | 0.11 | 16.86 | 22.19 | 0.31 | 0.06 | 99.66 | 1.99 | 0.00 | 0.03 | 0.15 | 0.92 | 0.00 | 0.87 | 0.02 | 0.00 | 4.00 | 0.448 | 0.473 | 0.079 |
| 53.18 | 0.11 | 0.95 | 7.34 | 0.20 | 15.14 | 21.68 | 0.58 | 0.03 | 99.21 | 1.98 | 0.00 | 0.04 | 0.23 | 0.84 | 0.01 | 0.87 | 0.04 | 0.00 | 4.01 | 0.447 | 0.435 | 0.118 |
| 54.05 | 0.17 | 0.77 | 4.82 | 0.10 | 16.46 | 23.30 | 0.23 | 0.07 | 99.97 | 1.98 | 0.00 | 0.03 | 0.15 | 0.90 | 0.00 | 0.92 | 0.02 | 0.00 | 4.00 | 0.466 | 0.458 | 0.075 |
| 52.63 | 0.36 | 2.30 | 8.63 | 0.23 | 13.10 | 21.97 | 0.96 | 0.00 | 100.18 | 1.96 | 0.01 | 0.10 | 0.27 | 0.73 | 0.01 | 0.88 | 0.07 | 0.00 | 4.02 | 0.468 | 0.388 | 0.144 |
| 53.88 | 0.17 | 0.91 | 6.99 | 0.22 | 14.90 | 22.37 | 0.72 | 0.08 | 100.24 | 1.99 | 0.00 | 0.04 | 0.22 | 0.82 | 0.01 | 0.88 | 0.05 | 0.00 | 4.01 | 0.461 | 0.427 | 0.112 |
| 51.28 | 0.00 | 0.10 | 22.47 | 2.54 | 3.00 | 23.24 | 0.01 | 0.01 | 102.65 | 2.01 | 0.00 | 0.00 | 0.74 | 0.18 | 0.08 | 0.98 | 0.00 | 0.00 | 3.99 | 0.517 | 0.093 | 0.390 |
| 53.63 | 0.10 | 3.54 | 6.43 | 0.24 | 13.25 | 21.88 | 1.59 | 0.13 | 100.79 | 1.96 | 0.00 | 0.15 | 0.20 | 0.72 | 0.01 | 0.86 | 0.11 | 0.00 | 4.02 | 0.483 | 0.407 | 0.111 |
| 53.65 | 0.20 | 4.52 | 2.53 | 0.07 | 16.24 | 21.03 | 1.21 | 1.36 | 100.81 | 1.92 | 0.01 | 0.19 | 0.08 | 0.87 | 0.00 | 0.81 | 0.08 | 0.04 | 4.00 | 0.461 | 0.496 | 0.043 |
| 55.26 | 0.18 | 0.50 | 3.61 | 0.07 | 18.27 | 22.87 | 0.24 | 0.11 | 101.11 | 1.99 | 0.00 | 0.02 | 0.11 | 0.98 | 0.00 | 0.88 | 0.02 | 0.00 | 4.00 | 0.447 | 0.497 | 0.055 |
| 54.00 | 0.13 | 4.13 | 2.53 | 0.05 | 17.02 | 21.92 | 0.83 | 0.89 | 101.50 | 1.92 | 0.00 | 0.17 | 0.08 | 0.90 | 0.00 | 0.84 | 0.06 | 0.03 | 4.00 | 0.461 | 0.498 | 0.042 |

Table 4 - Electron Microprobe Analyses of garnets

Garnets - SUL Claim Group

Formula Based on 24 O

| Molecular Weight of Oxides | | | | | | | | | Number of ions in formula | | | | | | | | End-member components | | | | | |
|----------------------------|------------------|--------------------------------|--------|--------|--------|--------|--------------------------------|---------|---------------------------|-------|-------|-------|-------|-------|-------|-------|-----------------------|-------|-------|-------|-------|--|
| SiO ₂ | TiO ₂ | Al ₂ O ₃ | FeO* | MnO | MgO | CaO | Cr ₂ O ₃ | Total | Si | Ti | Al | Fe | Mg | Mn | Ca | Cr | Sum | Alm | Gro | Pyr | Spes | |
| 60.090 | 79.900 | 101.940 | 71.850 | 70.940 | 40.320 | 56.080 | 152.020 | | | | | | | | | | | | | | | |
| 37.060 | 0.000 | 21.230 | 32.830 | 0.490 | 4.110 | 4.210 | 0.060 | 99.990 | 5.920 | 0.000 | 3.998 | 4.386 | 0.978 | 0.066 | 0.721 | 0.008 | 16.077 | 0.713 | 0.117 | 0.159 | 0.011 | |
| 37.810 | 0.000 | 21.040 | 32.840 | 0.300 | 6.780 | 0.840 | 0.030 | 99.640 | 5.986 | 0.000 | 3.927 | 4.348 | 1.600 | 0.040 | 0.143 | 0.004 | 16.048 | 0.709 | 0.023 | 0.261 | 0.007 | |
| 37.520 | 0.000 | 20.830 | 32.290 | 0.350 | 5.810 | 2.360 | 0.000 | 99.160 | 5.989 | 0.000 | 3.920 | 4.310 | 1.382 | 0.047 | 0.404 | 0.000 | 16.052 | 0.702 | 0.066 | 0.225 | 0.008 | |
| 37.790 | 0.000 | 21.160 | 32.690 | 0.330 | 5.750 | 2.350 | 0.000 | 100.070 | 5.978 | 0.000 | 3.947 | 4.325 | 1.356 | 0.044 | 0.398 | 0.000 | 16.048 | 0.706 | 0.065 | 0.221 | 0.007 | |
| 36.790 | 0.000 | 20.420 | 34.850 | 0.370 | 4.350 | 1.780 | 0.020 | 98.580 | 5.984 | 0.000 | 3.915 | 4.740 | 1.054 | 0.051 | 0.310 | 0.003 | 16.057 | 0.770 | 0.050 | 0.171 | 0.008 | |
| 38.390 | 0.000 | 21.270 | 30.850 | 0.290 | 8.160 | 1.030 | 0.050 | 100.040 | 5.992 | 0.000 | 3.914 | 4.027 | 1.898 | 0.038 | 0.172 | 0.006 | 16.048 | 0.656 | 0.028 | 0.309 | 0.006 | |
| 37.460 | 0.000 | 20.880 | 32.530 | 0.490 | 6.050 | 1.760 | 0.000 | 99.170 | 5.979 | 0.000 | 3.929 | 4.342 | 1.439 | 0.066 | 0.301 | 0.000 | 16.057 | 0.706 | 0.049 | 0.234 | 0.011 | |
| 38.260 | 0.050 | 21.410 | 26.780 | 0.460 | 9.770 | 1.750 | 0.050 | 98.530 | 5.973 | 0.006 | 3.940 | 3.496 | 2.273 | 0.061 | 0.293 | 0.006 | 16.048 | 0.571 | 0.048 | 0.371 | 0.010 | |
| 38.450 | 0.020 | 21.940 | 28.450 | 0.220 | 9.430 | 0.920 | 0.070 | 99.500 | 5.960 | 0.002 | 4.009 | 3.688 | 2.178 | 0.029 | 0.153 | 0.009 | 16.029 | 0.610 | 0.025 | 0.360 | 0.005 | |
| 38.190 | 0.000 | 21.460 | 29.550 | 0.250 | 8.720 | 1.050 | 0.030 | 99.250 | 5.975 | 0.000 | 3.958 | 3.866 | 2.033 | 0.033 | 0.176 | 0.004 | 16.045 | 0.633 | 0.029 | 0.333 | 0.005 | |

Table 5 - Electron Microprobe Analyses of opaques

Opaques - SUL Claim Group

Formula Based on 32 O

| Mineral Identity | | | | | | | | | | Number of ions in formula | | | | | | | | |
|---------------------|------|------|-------|-------|------|-------|------|-------|--------|---------------------------|------|------|------|------|------|------|------|------|
| | SiO2 | TiO2 | Al2O3 | FeO* | MnO | MgO | CaO | Cr2O3 | Total | Si | Ti | Al | Fe | Mg | Mn | Ca | Na | Cr |
| Chromite | 0.03 | 1.24 | 26.96 | 21.65 | 0.20 | 14.18 | 0.02 | 35.42 | 99.70 | 0.01 | 0.23 | 7.76 | 4.42 | 5.16 | 0.04 | 0.01 | 0.00 | 6.84 |
| Chromite | 0.06 | 1.06 | 22.26 | 19.55 | 0.22 | 14.49 | 0.02 | 41.97 | 99.63 | 0.01 | 0.20 | 6.51 | 4.06 | 5.36 | 0.05 | 0.01 | 0.00 | 8.23 |
| Chromite | 0.04 | 1.09 | 24.28 | 19.19 | 0.20 | 14.90 | 0.02 | 39.68 | 99.40 | 0.01 | 0.20 | 7.04 | 3.95 | 5.46 | 0.04 | 0.01 | 0.00 | 7.71 |
| Chromite | 0.03 | 0.81 | 27.03 | 18.52 | 0.22 | 15.30 | 0.02 | 38.16 | 100.09 | 0.01 | 0.15 | 7.67 | 3.73 | 5.50 | 0.04 | 0.01 | 0.00 | 7.27 |
| Chromite | 0.10 | 1.06 | 25.34 | 17.40 | 0.20 | 15.82 | 0.02 | 40.36 | 100.30 | 0.02 | 0.19 | 7.20 | 3.51 | 5.68 | 0.04 | 0.01 | 0.00 | 7.69 |
| Chromite | 0.14 | 1.10 | 25.43 | 17.23 | 0.19 | 15.86 | 0.02 | 41.37 | 101.34 | 0.03 | 0.20 | 7.15 | 3.44 | 5.64 | 0.04 | 0.01 | 0.00 | 7.80 |
| Chromite | 0.04 | 1.02 | 24.74 | 17.16 | 0.23 | 15.33 | 0.01 | 41.22 | 99.75 | 0.01 | 0.19 | 7.09 | 3.49 | 5.56 | 0.05 | 0.00 | 0.00 | 7.92 |
| Chromite | 0.04 | 0.98 | 22.40 | 17.10 | 0.25 | 15.25 | 0.02 | 43.67 | 99.71 | 0.01 | 0.18 | 6.49 | 3.51 | 5.59 | 0.05 | 0.01 | 0.00 | 8.48 |
| Chromite | 0.05 | 0.98 | 23.48 | 16.88 | 0.20 | 15.41 | 0.01 | 42.53 | 99.54 | 0.01 | 0.18 | 6.77 | 3.45 | 5.62 | 0.04 | 0.00 | 0.00 | 8.23 |
| Chromite | 0.10 | 0.90 | 23.23 | 16.76 | 0.20 | 15.63 | 0.03 | 43.15 | 100.00 | 0.02 | 0.16 | 6.67 | 3.42 | 5.68 | 0.04 | 0.01 | 0.00 | 8.31 |
| Chromite | 0.07 | 1.02 | 24.29 | 16.73 | 0.23 | 15.86 | 0.01 | 42.19 | 100.40 | 0.02 | 0.19 | 6.91 | 3.38 | 5.71 | 0.05 | 0.00 | 0.00 | 8.06 |
| Chromite | 0.04 | 0.97 | 22.95 | 16.69 | 0.19 | 15.70 | 0.02 | 43.29 | 99.85 | 0.01 | 0.18 | 6.61 | 3.41 | 5.72 | 0.04 | 0.01 | 0.00 | 8.36 |
| Chromite | 0.07 | 0.78 | 24.48 | 16.59 | 0.19 | 15.71 | 0.02 | 41.95 | 99.79 | 0.02 | 0.14 | 7.00 | 3.37 | 5.69 | 0.04 | 0.01 | 0.00 | 8.05 |
| Chromite | 0.05 | 0.84 | 21.94 | 16.31 | 0.21 | 15.56 | 0.02 | 45.29 | 100.22 | 0.01 | 0.15 | 6.32 | 3.33 | 5.67 | 0.04 | 0.01 | 0.00 | 8.75 |
| Chromite | 0.12 | 0.87 | 23.90 | 16.25 | 0.20 | 16.07 | 0.01 | 42.65 | 100.07 | 0.03 | 0.16 | 6.82 | 3.29 | 5.80 | 0.04 | 0.00 | 0.00 | 8.17 |
| Chromite | 0.08 | 1.07 | 24.66 | 16.25 | 0.23 | 16.28 | 0.02 | 42.26 | 100.85 | 0.02 | 0.19 | 6.96 | 3.26 | 5.82 | 0.05 | 0.01 | 0.00 | 8.01 |
| Chromite | 0.08 | 0.82 | 24.90 | 16.21 | 0.22 | 16.19 | 0.01 | 41.25 | 99.68 | 0.02 | 0.15 | 7.10 | 3.28 | 5.84 | 0.05 | 0.00 | 0.00 | 7.89 |
| Chromite | 0.08 | 0.84 | 24.51 | 16.14 | 0.21 | 16.23 | 0.02 | 40.67 | 98.70 | 0.02 | 0.15 | 7.06 | 3.30 | 5.92 | 0.04 | 0.01 | 0.00 | 7.86 |
| Chromite | 0.10 | 0.72 | 24.51 | 16.10 | 0.18 | 15.78 | 0.02 | 42.17 | 99.58 | 0.02 | 0.13 | 7.01 | 3.27 | 5.71 | 0.04 | 0.01 | 0.00 | 8.10 |
| Chromite | 0.11 | 0.74 | 25.32 | 15.91 | 0.17 | 16.30 | 0.01 | 42.20 | 100.76 | 0.03 | 0.13 | 7.13 | 3.18 | 5.81 | 0.03 | 0.00 | 0.00 | 7.97 |
| Chromite | 0.07 | 0.74 | 23.92 | 15.86 | 0.20 | 15.59 | 0.02 | 43.54 | 99.94 | 0.02 | 0.14 | 6.84 | 3.22 | 5.64 | 0.04 | 0.01 | 0.00 | 8.35 |
| Chromite | 0.12 | 0.76 | 25.72 | 15.72 | 0.20 | 16.57 | 0.02 | 41.14 | 100.25 | 0.03 | 0.14 | 7.26 | 3.15 | 5.91 | 0.04 | 0.01 | 0.00 | 7.79 |
| Chromite | 0.12 | 0.81 | 24.02 | 15.37 | 0.22 | 16.19 | 0.02 | 42.31 | 99.06 | 0.03 | 0.15 | 6.90 | 3.13 | 5.88 | 0.05 | 0.01 | 0.00 | 8.15 |

Most of the garnets identified in the heavy mineral are light to medium pink in colour and are almost certainly not pyrope rich. However, some medium red garnets are also present and will be sent away for quantitative analysis. Garnets analyzed to date do not have high calcium-high chrome compositions consistent with G10 garnets associated with kimberlites. However, the data has not yet been evaluated relative to eclogitic diamond compositions. The colour of the garnets is not the deep orange associated with eclogitic garnets however. Analysis of the garnet compositions is not yet complete and more need to be submitted for quantitative analysis before any conclusions can be made.

Non-magnetic to weakly paramagnetic opaque compositions (Table 5) document a chromite composition for the majority of the opaques submitted for analysis. However, the chromite content of these samples is approximately 25-20 weight percent lower than chromite compositions associated with diamondiferous lithologies. However these represent preliminary results and more grains need to be picked and submitted for analysis before any conclusions can be reached.

Thin Section Descriptions

Sample BRWT93-14 - The sample consists of several large inclusions (clasts) up to 3 cm in long dimension oriented sub-parallel to parallel to a moderately well defined foliation. The foliation is comprised of sub-parallel to parallel subordinate chlorite in a groundmass consisting of 30-35% opaques, 40-50% fine-grained carbonates minor phlogopite and minor serpentine. The inclusions are rounded to sub-rounded and predominantly comprised of a granoblastic mosaic of carbonate, coarse at the core and finer grained at the rim. A skeletal biotite is present with individual laths separated by coarse granoblastic carbonate. The biotite is yellow-brown to red-brown pleochroic and has golden-yellow to black pleochroic (altered) rims. Portions of the biotite inclusion are corroded.

This sample is interpreted to represent Toby conglomerate.

Sample BRWT93-15 - The sample consists of a highly altered (hydrated), fine-grained, opaque-rich lithology with a poorly developed foliation defined by opaques and fine-grained serpentine. Three thin (<<1 mm thick) carbonate veinlets cross-cut the plane of thin section, consisting of a granoblastic mosaic of carbonate grains. A porphyritic texture is developed, characterized by the presence of completely altered, euhedral to subhedral phenocrysts of pseudomorphed olivine and pyroxene. The phenocrysts have been completely pseudomorphed by serpentine and carbonate. The original identity of the phenocrysts has been surmised based upon relict crystal shape (columnar to lath shaped - pyroxene) and serpentine alteration along fractures (typical of olivine). The phenocrysts have a glomeroporphyritic texture comprised of clusters of 2 - 6 phenocrysts. A hematite-

rich rind is present, developed as an irregular alteration front extending for up to 1 cm into the plane of the thin section and is interpreted to represent late stage alteration of the sample.

The sample is interpreted to be an essentially completed hydrated mafic to ultramafic igneous intrusive. Original composition is difficult to surmise due to the extent of alteration, however the presence of pyroxene and olivine and apparent lack of feldspar (no sericite alteration evident), together with the ubiquitous presence of serpentine suggests an ultramafic precursor, consistent with a kimberlitic composition.

CONCLUSIONS AND RECOMMENDATIONS

Dykes of kimberlitic affinity are known to occur in a north-south belt (Pope and Thirlwall 1992), within which the SUL claim group lies (see Figure 3). "(The) common petrogenesis and extensional, fault-controlled structural setting, coupled with essentially identical ages of emplacement, lead us to postulate that the Toby - Horsethief Creek dykes and the Cross kimberlite evolved in parallel, during a single phase of continental extension" (Pope and Thirlwall 1992).

Due to the backlog caused by intense diamond exploration in Canada this season, geochemical analysis (whole rock and "gold+47" element analysis) of the bulk sample sent in early July to Loring Laboratories in Calgary has not yet been returned, therefore it is impossible to comment on the geochemistry of the occurrence beyond that previously discussed for the sample obtained in October, 1992.

Preliminary geochemical results of samples taken from the occurrences does confirm composition of kimberlitic affinity. The Sultana Creek occurrence appears to be a dyke or diatreme having textural similarities to the Toby Creek occurrence, described by Pope and Thirlwall (1992) as a true kimberlite. However, geochemical data from a sample submitted in October, 1992 is more comparable to lamprophyres of kimberlitic affinity.

At present, it is not clear where these intrusives originated relative to the diamond stability field. However, regardless of their diamond potential, the fact that they originated in the lower levels of the crust or from within the mantle suggests potential for hosting Rare-Earth Elements (REEs), Platinum Group Elements and gold.

If the results of the program (some still outstanding), do confirm a kimberlitic affinity for the dykes in terms of composition, elemental suites, kimberlitic indicator minerals and/or deep seated xenoliths (peridotitic and/or eclogitic inclusions) a follow-up program of trenching and drilling is recommended. Trenching, possibly with blasting, would enable the collection of a large sample of pristine dyke material for a determination of diamond content. Drilling is recommended upon favourable geochemical results, in terms of a kimberlitic indicator (heavy mineral) and/or xenolith suite. A preliminary budget has been prepared for the above work and is presented on the following page.

The fact that kimberlites and lamprophyres of kimberlitic affinity have been documented in a new, previously unknown, terrane clearly indicates that more exploration is warranted in the Purcell Mountains. The close association postulated by Pope and Thirlwall (1992) between diatremes mapped in the Rocky Mountains in terms of alkaline composition, age of intrusion and volcanic association demands that these occurrences be thoroughly

evaluated for their potential as hosts for diamonds. The recovery of several diamonds in the Golden cluster of diatremes and the kimberlitic affinity of the Purcell Mountain occurrences suggests these occurrences may be possible diamond hosts.

PROPOSED BUDGET

| | |
|------------------------------------|-------------|
| Pre-Field | \$ 5,000.00 |
| Field Program | |
| Diamond Drilling..... | \$30,000.00 |
| Personnel | \$20,000.00 |
| Helicopter Support | \$ 5,000.00 |
| Analytical | \$ 5,000.00 |
| Camp Costs | \$ 2,000.00 |
| Food/Grocery | \$ 4,000.00 |
| Truck/Equipment Rentals | \$ 1,000.00 |
| Fuel | \$ 1,000.00 |
| Transportation | \$ 2,000.00 |
| Supplies | \$ 1,000.00 |
| Miscellaneous | \$ 1,000.00 |
| | <hr/> |
| | \$77,000.00 |
| Contingency on Field Program (10%) | \$ 7,700.00 |
| Post-Field | \$ 5,000.00 |
| | <hr/> |
| TOTAL: | \$89,700.00 |

REFERENCES

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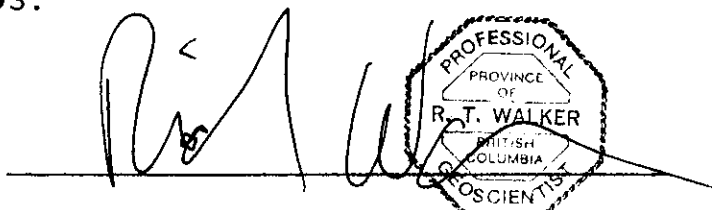
Appendix A
Statement of Qualifications

STATEMENT OF QUALIFICATIONS

I, Richard T. Walker, of 1916 - 5th Street South, Cranbrook, BC, hereby certify that:

- 1) I am a graduate of the University of Calgary of Calgary, Alberta, having obtained a Bachelors of Science in 1986.
- 2) I obtained a Masters of Geology at the University of Calgary of Calgary, Alberta in 1989.
- 3) I am a member in good standing with the Association of Professional Engineers, Geologists and Geophysicists of Alberta.
- 4) I am a member of good standing with the Association of Professional Engineers and Geoscientists of the Province of British Columbia.
- 5) I am a consulting geologist and Principal with the firm of Gemquest Geological with offices at 1916-5th St. S., Cranbrook, British Columbia.
- 6) I am the author of this report which is based on work I personally performed on the property between July 17th, 1992, and June 10th, 1993.
- 7) I was personally involved in the acquisition of the claims described herein.
- 8) I hold at present 21,500 common shares of Gwen Resources Ltd..

Dated at Cranbrook, British Columbia this 13th day of October, 1993.



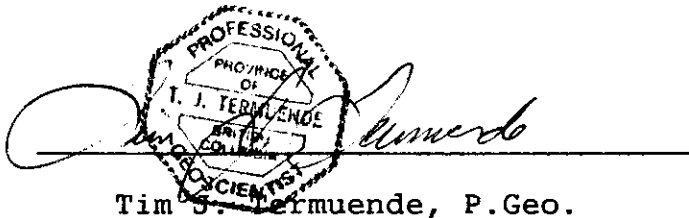
Richard T. Walker, P.GEO.

STATEMENT OF QUALIFICATIONS

I, Tim J. Termuende, of 2720-17th Street South, Cranbrook, BC, hereby certify that:

- 1) I am a graduate of the University of British Columbia, Vancouver, B.C. with a Geological Sciences Degree (B.Sc.) in Geology (1987).
- 2) I am a member in good standing of the Association of Professional Engineers and Geoscientists of the Province of British Columbia, member # 19201.
- 3) I am a consulting geologist and Principal with the firm of Gemquest Geological with offices at 1916-5th St. S., Cranbrook, British Columbia.
- 4) I personally carried out geologic work on the SUL claims between July 17, 1992, and June 10, 1993.
- 5) I was personally involved in the acquisition of the claims described herein.
- 6) I own at present 20,500 common shares of Gwen Resources Ltd..

Dated at Cranbrook, British Columbia this 13th day of October, 1993.


Tim J. Termuende, P. Geol.

Appendix B
Claim Records



RECORD OF 2 POST CLAIM - MINERAL TENURE ACT

SECTION 23

MAP NO. 82K8W

TENURE NO. 311872

MINING RECEIPT NO. 15 RECORDED AT NELSON BC DATE OF RECORD JULY 16 1992

DO NOT WRITE IN THIS SHADED AREA

[Handwritten Signature]
DEPT. COMMISSIONER

GOLDEN

MINING DIVISION

PLEASE PRINT CLEARLY

NAME TIM J. TERMUENDE AGENT FOR SELF

ADDRESS 2720-17th ST. E.

CRANBROOK, BC

TELEPHONE (604) 426-3112 POSTAL CODE V1C 4H4

CLIENT NUMBER 126599

APPLICATION TO RECORD A 2 POST CLAIM

hereby apply for a record of a 2 post claim for the location as outlined on the attached copy of mineral titles reference map No. 82K13W in the GOLDEN Mining Division

ACCESS: Describe how you gained access to the location, include references to roads, trails, topographic features, permanent landmarks, and a description of the post location.

ACCESS VIA TERRY CREEK ROAD TO DELPHIC CREEK, THEN CAPTURED FOR 2 KM, THEN BRIDGE NAME ROAD TO ELEVATION 2500'. INITIAL POST LOCATED (ELEVATION 2500' 1400M) BEARING 360°, FROM THE CONFLUENCE OF SULTANA AND DELPHIC CR.

I have securely affixed the portion of the metal identification tag embossed "INITIAL POST (NO. 1)" to the initial post and impressed this information on the tag.

TAG NUMBER 644036M
INITIAL POST (NO. 1)

CLAIM NAME SUL 1

LOCATOR T. TERMUENDE

AGENT FOR SELF

DATE COMMENCED JULY 16/92

TIME COMMENCED 3:00P

DIR TO NO 2 POST 344°

METRES TO RIGHT 500

METRES TO LEFT -

I have securely affixed the portion of the metal identification tag embossed "FINAL POST (NO. 2)" to the final post (or the witness post) and impressed this information on the tag.

TAG NUMBER 644036M
FINAL POST (NO. 2)

CLAIM NAME SUL 1

LOCATOR T. TERMUENDE

AGENT FOR SELF

DIST. FROM NO. 1 POST 500M

DATE COMPLETED JULY 16/92

TIME COMPLETED 3:50P

*If witness post placed for final post:
Bearing to true position of final post: 344°
distance 375M metres

I have complied with all the terms and conditions of the Mineral Tenure Act Regulation pertaining to the location of 2 post claims and have attached a plan of the location on which the positions of the initial and final posts (and witness and identification posts if applicable) are indicated.

[Handwritten Signature]
Signature of Locator

PAID
GOVERNMENT AGENT CRANBROOK

AUG 4 1992
10:00:15
AMT. 4.00
TRANS. # _____

RECORDING STAMP



RECORD OF 2 POST CLAIM - MINERAL TENURE ACT

SECTION 23

MAP NO. 82K8W

TENURE NO. 311873

MINING RECEIPT NO. 15 RECORDED AT NELSON B.C. DATE OF RECORD JULY 16 1992

DO NOT WRITE IN THIS SHADED AREA GOLDEN MINING DIVISION

[Handwritten Signature]
DEPUTY COMMISSIONER

PLEASE PRINT CLEARLY

1. TIM J. TERMUENDE AGENT FOR SELF
2720-17th ST. S.
CRAWBROOK BC
(604) 426-3112
126599

APPLICATION TO RECORD A 2 POST CLAIM

hereby apply for a record of a 2 post claim for the location as outlined on the attached copy of mineral titles reference map No. 82K/8W in the GOLDEN Mining Division

ACCESS Describe how you gained access to the location, include references to roads, trails, topographic features, permanent landmarks, and a description of the post location.

ACCESS VIA TOBY CR ROAD FROM INVERMERE, THEN ALONG DEWINE ROAD TO 3KM, THEN ALONG DIRT ROAD TO 5500' ELEVATION. INITIAL POST LOCATED 140M LEADING 340°/2 FROM THE CONFLUENCE OF SULTANA AND DEWINE CR.

I have securely affixed the portion of the metal identification tag embossed "INITIAL POST (NO. 1)" to the initial post and impressed this information on the tag:

TAG NUMBER 644037A1
INITIAL POST (NO. 1)

CLAIM NAME SUL 2
 LOCATOR T. TERMUENDE
 AGENT FOR SELF
 DATE COMMENCED JULY 16/92
 TIME COMMENCED 3:00P
 DIR. TO NO. 2 POST 344°
 METRES TO RIGHT —
 METRES TO LEFT 500

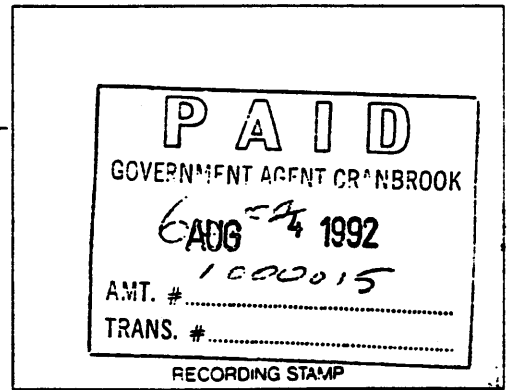
I have securely affixed the portion of the metal identification tag embossed "FINAL POST (NO. 2)" to the final post (or the witness post) and impressed this information on the tag:

TAG NUMBER 644037A1
FINAL POST (NO. 2)

CLAIM NAME SUL 2
 LOCATOR T. TERMUENDE
 AGENT FOR SELF
 DIST. FROM NO. 1 POST 500M
 DATE COMPLETED JULY 16/92
 TIME COMPLETED 3:55P
 "If witness post placed for final post:
 Bearing to true position of final post 344°
 distance 75 metres.

I have complied with all the terms and conditions of the Mineral Tenure Act Regulation pertaining to the location of 2 post claims and have attached a plan of the location on which the positions of the initial and final posts (and witness and identification posts if applicable) are indicated.

[Handwritten Signature]
Signature of Locator





MAP NO. 82KSW

TENURE NO. 311874

MINING RECEIPT NO 15 RECORDED AT NELSON BC DATE OF RECORD JULY 16 92

DO NOT WRITE IN THIS SHADED AREA

GOLDEN

[Handwritten Signature]
OFFICIAL REGISTRAR

PLEASE PRINT CLEARLY

NAME TIM J. TERMUENDE AGENT FOR SELF

ADDRESS 3720-17th ST. S

(VANCOUVER, BC)

TELEPHONE (604) 426-3112 POSTAL CODE V1C 4H4

CLIENT NUMBER 126599

APPLICATION TO RECORD A 2 POST CLAIM

hereby apply for a record of a 2 post claim for the location as outlined on the attached copy of mineral titles reference map No 82K15W in the GOLDEN Mining Division

ACCESS Describe how you gained access to the location, include references to roads, trails, topographic features, permanent landmarks, and a description of the post location

ACCESS VIA TONY CR ROAD FROM INTERMERE TRAIL ALONG DELHIANE CR 3 KM TO DALL ROAD, THEN NORTHWARD TO ELEVATION 5300'. INITIAL POST LOCATED 100M BEARING 300M FROM THE CONFLUENCE OF SULTANA AND DELHIANE CR.

I have securely affixed the portion of the metal identification tag embossed "INITIAL POST (NO. 1)" to the initial post and impressed this information on the tag

TAG NUMBER L44035M
INITIAL POST (NO. 1)

CLAIM NAME SUL 3

LOCATOR T. TERMUENDE

AGENT FOR SELF

DATE COMMENCED JULY 16/92

TIME COMMENCED 11:45 AM 3:00 PM

DIR TO NO 2 POST 164°

METRES TO RIGHT 500

METRES TO LEFT —

I have securely affixed the portion of the metal identification tag embossed "FINAL POST (NO. 2)" to the final post for the witness post and impressed this information on the tag

TAG NUMBER L44035M
FINAL POST (NO. 2)

CLAIM NAME SUL 3

LOCATOR T. TERMUENDE

AGENT FOR SELF

DIST. FROM NO. 1 POST 520M

DATE COMPLETED JULY 16/92

TIME COMPLETED 4:50 P

*If witness post placed for final post

Bearing to true position of final post: _____

distance _____ metres

I have complied with all the terms and conditions of the Mineral Tenure Act Regulation pertaining to the location of 2 post claims and have attached a plan of the location on which the positions of the initial and final posts (and witness and identification posts if applicable) are indicated.

[Handwritten Signature]
Signature of Locator

PAID
GOVERNMENT AGENT OR BROOK
AUG 4 1992
60.
AMT. # 160015
TRANS. # _____
RECORDING STAMP



RECORD OF 2 POST CLAIM - MINERAL TENURE ACT

SECTION 23

MAP NO. 82KSW

TENURE NO. 311875

MINING RECEIPT NO. 15 RECORDED AT NELSON BC DATE OF RECORD JULY 16 '92

DO NOT WRITE IN THIS SHADED AREA GOLDEN MINING DIVISION

[Handwritten Signature]
SOLD BY REGISTER

PLEASE PRINT CLEARLY

1. TIM J. TERMUENDE NAME AGENT FOR SELF NAME

2770-17th ST. S. ADDRESS ADDRESS

CRANBROOK, BC ADDRESS ADDRESS

(604) 476-3112 TELEPHONE VIC-4114 TELEPHONE TELEPHONE

CLIENT NUMBER 106599 CLIENT NUMBER

APPLICATION TO RECORD A 2 POST CLAIM

hereby apply for a record of a 2 post claim for the location as outlined on the attached copy of mineral titles reference map No. 82K105W in the GOLDEN Mining Division

ACCESS Describe how you gained access to the location, include references to roads, trails, topographic features, permanent landmarks, and a description of the post location.

ACCESS VIA TOBY CR. ROAD FROM INVERMERE TO DELMINE ROAD, THEN WESTWARD 2 KM TO DALL ROAD, THEN NORTHWARD TO ELEVATION 5500. INITIAL POST LOCATED FROM PLANNED 300M FROM THE CONFLUENCE OF SULTANA AND DELMINE CR.

I have securely affixed the portion of the metal identification tag embossed "INITIAL POST (NO. 1)" to the initial post and impressed this information on the tag

TAG NUMBER 644039M INITIAL POST (NO. 1)

CLAIM NAME SUL 4

LOCATOR T. TERMUENDE

AGENT FOR SELF

DATE COMMENCED JULY 16/92

TIME COMMENCED 3:00P

DIR TO NO 2 POST 169°

METRES TO RIGHT -

METRES TO LEFT 500

I have securely affixed the portion of the metal identification tag embossed "FINAL POST (NO. 2)" to the final post or the witness post and impressed this information on the tag

TAG NUMBER 644039M FINAL POST (NO. 2)

CLAIM NAME SUL 4

LOCATOR T. TERMUENDE

AGENT FOR SELF

DIST. FROM NO. 1 POST 500M

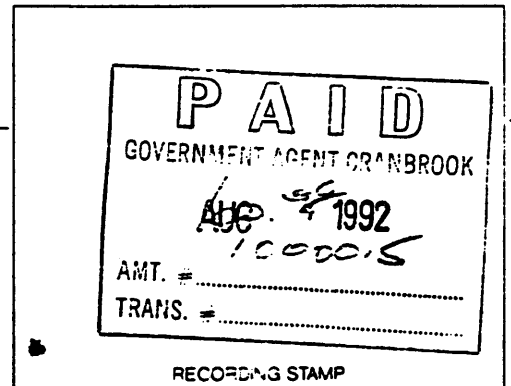
DATE COMPLETED JULY 16/92

TIME COMPLETED 4:55P

*If witness post placed for final post
Bearing to true position of final post: _____
distance _____ metres

I have complied with all the terms and conditions of the Mineral Tenure Act Regulation pertaining to the location of 2 post claims and have attached a plan of the location on which the positions of the initial and final posts (and witness and identification posts if applicable) are indicated.

[Handwritten Signature]
Signature of Locator





RECORD OF 2 POST CLAIM - MINERAL TENURE ACT

SECTION 23

MAP NO. 82K8W

TENURE NO. 311876

MINING RECEIPT NO. 15 RECORDED AT NELSON BC DATE OF RECORD JULY 16 19 92

DO NOT WRITE IN THIS SHADED AREA GOLDEN MINING DIVISION

[Handwritten signature]

PLEASE PRINT CLEARLY

NAME TIM T. TERMUENDE AGENT FOR SELF

ADDRESS 2720-17th ST. S.

CRAWFORD, BC

TELEPHONE (604) 476-7112 VIC -4114

CLIENT NUMBER 126579

APPLICATION TO RECORD A 2 POST CLAIM

hereby apply for a record of a 2 post claim for the location as outlined on the attached copy of mineral titles reference map No. 82K18W in the GOLDEN Mining Division

ACCESS Describe how you gained access to the location. include references to roads, trails, topographic features, permanent landmarks, and a description of the post location

ACCESS VIA TULLY CR ROAD FROM INTERSECTION WITH ROAD
DRIVING ROAD TO THE TRAIL HEAD. DRIVE ROAD TO CLEARING
500'. INITIAL POST LOCATED FROM CLEARING LOOKING FROM THE
CONJUNCTION OF SUSTA AND DRIVING CR.

I have securely affixed the portion of the metal identification tag embossed "INITIAL POST (NO. 1)" to the initial post and impressed this information on the tag:

TAG NUMBER 64404011
INITIAL POST (NO. 1)

CLAIM NAME SUL 5

LOCATOR T. TERMUENDE

AGENT FOR SELF

DATE COMMENCED JULY 1/92

TIME COMMENCED 11:35P

DIR. TO NO. 2 POST 1050

METRES TO RIGHT 500

METRES TO LEFT ---

I have securely affixed the portion of the metal identification tag embossed "FINAL POST (NO. 2)" to the final post (or the witness post) and impressed this information on the tag:

TAG NUMBER 64404011
FINAL POST (NO. 2)

CLAIM NAME SUL 5

LOCATOR T. TERMUENDE

AGENT FOR SELF

DIST FROM NO. 1 POST 500M

DATE COMPLETED JULY 16/92

TIME COMPLETED 5:15P

*If witness post placed for final post:

Bearing to true position of final post _____

distance _____ metres.

ACCESS

TAG INFORMATION

APPROVED FOR RECORDING

I have complied with all the terms and conditions of the Mineral Tenure Act Regulation pertaining to the location of 2 post claims and have attached a plan of the location on which the positions of the initial and final posts (and witness and identification posts if applicable) are indicated.

[Handwritten signature]
Signature of Locator

PAID
GOVERNMENT AGENT CRANBROOK
AUG 4 1992
AMT. # 60.00
TRANS. # 1570075

RECORDING STAMP



RECORD OF 2 POST CLAIM - MINERAL TENURE ACT

SECTION 23

MAP NO. 92K8W

TENURE NO. 311877

MINING RECEIPT NO. 15 RECORDED AT NELSON BC DATE OF RECORD JULY 16 19 92

DO NOT WRITE IN THIS SHADED AREA GOLDEN MINING DIVISION

PLEASE PRINT CLEARLY

APPLICATION TO RECORD A 2 POST CLAIM

NAME TIM T. TERMUENDE AGENT FOR SELF
ADDRESS 2770-17th ST. S.
CANBROOK, BC
TELEPHONE (604) 476-5112 VIC-4121
CLIENT NUMBER 120577

hereby apply for a record of a 2 post claim for the location as outlined on the attached copy of mineral titles reference map No. 82L/SW in the COX Mining Division

ACCESS Describe how you gained access to the location, include references to roads, trails, topographic features, permanent landmarks, and a description of the post location
ACCESS VIA TOST CR ROAD TO DEERING ROAD THEN WESTWARD
2 KM TO DALL ROAD, THEN NORTHWARD TO LOCATION
INITIAL POST LOCATED 90M NENEING 00842 FROM THE CONFLUENCE
OF SULTANA AND DEERING CR.

I have securely affixed the portion of the metal identification tag embossed "INITIAL POST (NO. 1)" to the initial post and impressed this information on the tag

TAG NUMBER 643991A1
INITIAL POST (NO 1)

CLAIM NAME SUL 6
LOCATOR T. TERMUENDE
AGENT FOR SELF
DATE COMMENCED JULY 16/92
TIME COMMENCED 4:35P
DIR. TO NO. 2 POST 105
METRES TO RIGHT —
METRES TO LEFT 700

I have securely affixed the portion of the metal identification tag embossed "FINAL POST (NO. 2)" to the final post (or the witness post) and impressed this information on the tag

TAG NUMBER 643991A1
FINAL POST (NO 2)

CLAIM NAME SUL 6
LOCATOR T. TERMUENDE
AGENT FOR SELF
DIST. FROM NO. 1 POST SDUM
DATE COMPLETED JULY 16/92
TIME COMPLETED 5:20P
*If witness post placed for final post:
Bearing to true position of final post _____
distance _____ metres.

I have complied with all the terms and conditions of the Mineral Tenure Act Regulation pertaining to the location of 2 post claims and have attached a plan of the location on which the positions of the initial and final posts (and witness and identification posts if applicable) are indicated.

Signature of Locator [Signature]

PAID
GOVERNMENT AGENT ORNBROOK
AUG 4 1992
AMT. # 60.
TRANS. # 1000015
RECORDING STAMP

Appendix C
Analytical Results

Sample: 6

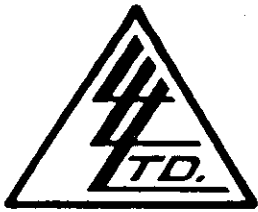
(Your sample ID: C)

Minerals

| Mineral | Mode | Special Characteristics |
|-----------|--------|--|
| Chlorite | 15% | Very fine-grained matrix phase is pleochroic colourless to green |
| Biotite | 10-15% | Phenocrysts up to 2 mm in size have extreme absorption and are pleochroic honey brown to black. Partial to near total alteration to colourless (Mg-rich?) chlorite especially well developed along cleavage planes |
| Quartz | 25% | Patches and lenses of polycrystalline quartz |
| Muscovite | 5% | Fine-grained matrix phase |
| Calcite | 45% | Equant 3 to 4 mm porphyroblasts have overgrown the fabric and mica phenocrysts. Dusty trails of tiny inclusions parallel to the external foliation can be traced through the calcite grains |
| Opaques | <5% | Idiomorphic, < 1mm in size. All have translucent orange/brown (hematite?) rims < 0.1 mm thick. |

Description

A well developed fabric is defined by the alignment of fine-grained chlorite and muscovite in the matrix. There is also compositional layering (calcite and chlorite-muscovite layers) parallel to the mica foliation and quartz-rich layers and lenses (< 0.5 mm thick). The fabric is somewhat irregular due to the high proportion of calcite porphyroblasts which deflect it.



LORING LABORATORIES LTD.

629 Beaverdam Rd. N.E.
Calgary, Alberta T2K 4W7

Tel: (403) 274-2777
Fax: (403) 275-0541

TO : GEMQUEST GEOLOGICAL
1916 - 5TH Street South
Cranbrook, B.C. V1C 1K4

FILE # 35928-D
DATE : SEPTEMBER 8, 1993
PAGE :

ATTN :

CERTIFICATE OF ANALYSIS

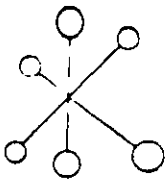
| SAMPLE ID | ORIGINAL WEIGHT (Kg) | MIDLINGS | | HEAVIES >3.3 SG | | | | |
|-----------|----------------------|--------------|---------------|-----------------|------------------------|--------------------|----------------------|------------------------|
| | | 2.9 - 3.3 SG | | MAGN. (Gm) | NON-MAG. +28 Mesh (Gm) | P.M. -28 Mesh (Gm) | W.P.M. -28 Mesh (Gm) | NON-MAG. -28 Mesh (Gm) |
| | | MAGN. (Gm) | NON-MAG. (Gm) | | | | | |
| | 27.1 | 0.17 | 6.67 | 0.58 | 0.65 | 0.35 | 0.98 | 0.69 |
| | 25.1 | 0.82 | 4.00 | 1.48 | 0.59 | 0.22 | 0.57 | 3.47 |
| | 22.0 | 0.14 | 8.00 | 3.17 | 1.57 | 2.46 | 3.70 | 0.61 |
| | 25.0 | 0.24 | 7.72 | 0.46 | 0.43 | 0.12 | 0.61 | 0.97 |
| | 29.5 | 0.17 | 33.53 | 0.66 | 44.40 | 0.14 | 0.83 | 65.18 |
| | 27.0 | 6.62 | 4.25 | 3.47 | 0.02 | 0.50 | 0.19 | 0.18 |
| | 20.0 | 0.13 | 4.35 | 2.59 | 3.85 | 0.51 | 2.09 | 4.77 |
| SUL-93-2 | 18.5 | 0.19 | 18.34 | 0.48 | 0.78 | 0.09 | 0.29 | 1.23 |
| | 31.5 | 0.76 | 20.74 | 4.49 | 0.91 | 1.52 | 6.62 | 0.66 |
| | | | | | | | | |

NOTE : P.M. = PARAMAGNETIC

W.P.M. = WEAK PARAMAGNETIC

I HEREBY CERTIFY that the above results are those assays made by me upon the herein described samples....


ASSAYER



ECO-TECH LABORATORIES LTD.

ASSAYING - ENVIRONMENTAL TESTING
10041 East Trans Canada Hwy., Kamloops, B.C. V2C 2J3 (604) 573-5700 Fax 573-4557

FEBRUARY 4, 1993

CERTIFICATE OF ANALYSIS ETK 93-20

RICK WALKER

1913-5th Street South
CRANBROOK, B.C.
VIC 1K4

SAMPLE IDENTIFICATION: 3 ROCK samples received JANUARY 25, 1993

| ET# | Description | BaO | P2O5 | SiO2 | MnO | Fe2O3 | MgO | Al2O3 | CaO | TiO2 | Na2O | K2O | L.O.I. |
|-----|---------------|-----|------|-------|-----|-------|-------|-------|-------|------|------|------|--------|
| 1- | | .02 | .27 | 43.69 | .22 | 15.30 | 5.33 | 11.44 | 6.95 | 2.58 | 3.06 | .06 | 11.10 |
| 2- | | .07 | .31 | 37.68 | .17 | 9.41 | 13.24 | 5.00 | 13.97 | .98 | .02 | 2.14 | 16.99 |
| 3- | SULTANG CREEK | .11 | .67 | 34.74 | .15 | 8.86 | 10.71 | 11.42 | 10.75 | 2.39 | .98 | 2.58 | 16.60 |

QC DATA

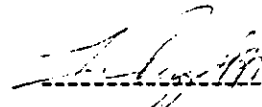
REPEAT #:

2- .02 .33 37.65 .17 9.30 13.68 5.04 13.95 .98 .01 1.85 16.99

STANDARD:

SY2 .07 .40 59.09 .32 6.39 2.87 12.11 8.06 .15 4.32 4.35 1.84
MRG-1 .03 .05 38.19 .17 17.77 13.80 8.30 14.73 3.67 .72 .19 2.40

NOTE: VALUES EXPRESSED IN PERCENT



ECO-TECH LABORATORIES LTD.
FRANK J. PEZZOTTI
B.C. CERTIFIED ASSAYER

SC93/KAMMISC#1

ECO-TECH LABORATORIES LTD.
 10041 EAST TRANS CANADA HWY.
 KAMLOOPS, B.C. V2C 2J3
 PHONE - 604-573-5700
 FAX - 604-573-4557

RICK WALKER ETK 93-20
 1916-5th Street South
 CRANBROOK, B.C.
 V1C 1K4

FEBRUARY 4, 1993

VALUES IN PPM UNLESS OTHERWISE REPORTED

3 ROCK SAMPLES RECEIVED JANUARY 25, 1993

| ET# | DESCRIPTION | AG | AL(%) | AS | B | BA | BI | CA(%) | CD | CO | CR | CU | FE(%) | K(%) | LA | MG(%) | MN | MO | NA(%) | NI | P | PB | SB | SN | SR | TI(%) | U | V | W | Y | ZN |
|-----|-----------------|----|-------|----|---|-----|----|-------|----|----|-----|-----|-------|------|----|-------|------|----|-------|-----|------|----|----|-----|-----|-------|-----|-----|-----|----|----|
| 1 | - | .4 | 2.91 | 30 | 2 | 50 | <5 | 4.54 | <1 | 43 | 59 | 160 | 8.99 | <.01 | 10 | 2.91 | 1486 | 1 | .01 | 32 | 850 | <2 | 5 | <20 | 143 | .01 | <10 | 268 | <10 | 6 | 78 |
| 2 | - | .2 | 2.18 | 10 | 2 | 60 | 5 | 8.69 | <1 | 35 | 627 | 2 | 3.79 | 1.62 | 10 | 5.57 | 1036 | <1 | <.01 | 131 | 1150 | <2 | <5 | <20 | 381 | .17 | <10 | 152 | <10 | 15 | 23 |
| 3 | - SULTANG CREEK | .2 | 2.08 | 10 | 2 | 270 | <5 | 6.97 | <1 | 18 | 111 | 68 | 4.33 | .15 | 20 | 5.34 | 991 | <1 | <.01 | 110 | 2630 | <2 | <5 | <20 | 129 | .01 | <10 | 43 | <10 | 8 | 46 |

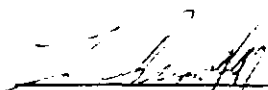
QC DATA

REPEAT #:

1 <.2 2.92 25 2 50 <5 4.51 <1 43 59 161 8.96 <.01 10 2.93 1483 1 .01 32 840 <2 <5 <20 144 .01 <10 270 <10 6 75

STANDARD 1991 - 1.8 1.76 70 4 125 <5 1.74 <1 19 63 82 3.68 .37 <10 1.00 703 <1 .01 20 630 10 5 <20 59 .10 <10 72 <10 13 81

NOTE: < = LESS THAN


 ECO-TECH LABORATORIES LTD.
 FRANK J. PEZZOTTI, A.Sc.T.
 B.C. Certified Assayer

SC93/KAMMISC#1

Appendix D
Statement of Expenditures

STATEMENT OF EXPENDITURES

The following expenses were incurred on the SUL claim group for the purpose of geological exploration within the period July 17, 1992 to June 10, 1993.

PERSONNEL

| | |
|---|-----------|
| R.T. Walker, P.Geo.: 1.0 days @ \$350.00/day | \$ 350.00 |
| T. J. Termuende, P.Geo. 1.0 days x \$350.00/day | \$ 350.00 |

EQUIPMENT RENTAL

| | |
|---|----------|
| 4WD Vehicle: 1.0 days x \$50.00/day | \$ 50.00 |
| Mileage: 322 km x \$.20/km | \$ 64.40 |
| Radios (2): 1.0 days x \$10.00/day | \$ 20.00 |

ANALYTICAL:

| | |
|--|-----------|
| 1 50kg. Bulk Samples | \$ 350.00 |
| Thin Section Preparation and Petrography | \$ 100.00 |

FIELD SUPPLY

| | |
|----------------------------------|----------|
| 2.0 man-days x \$20.00/day | \$ 40.00 |
|----------------------------------|----------|

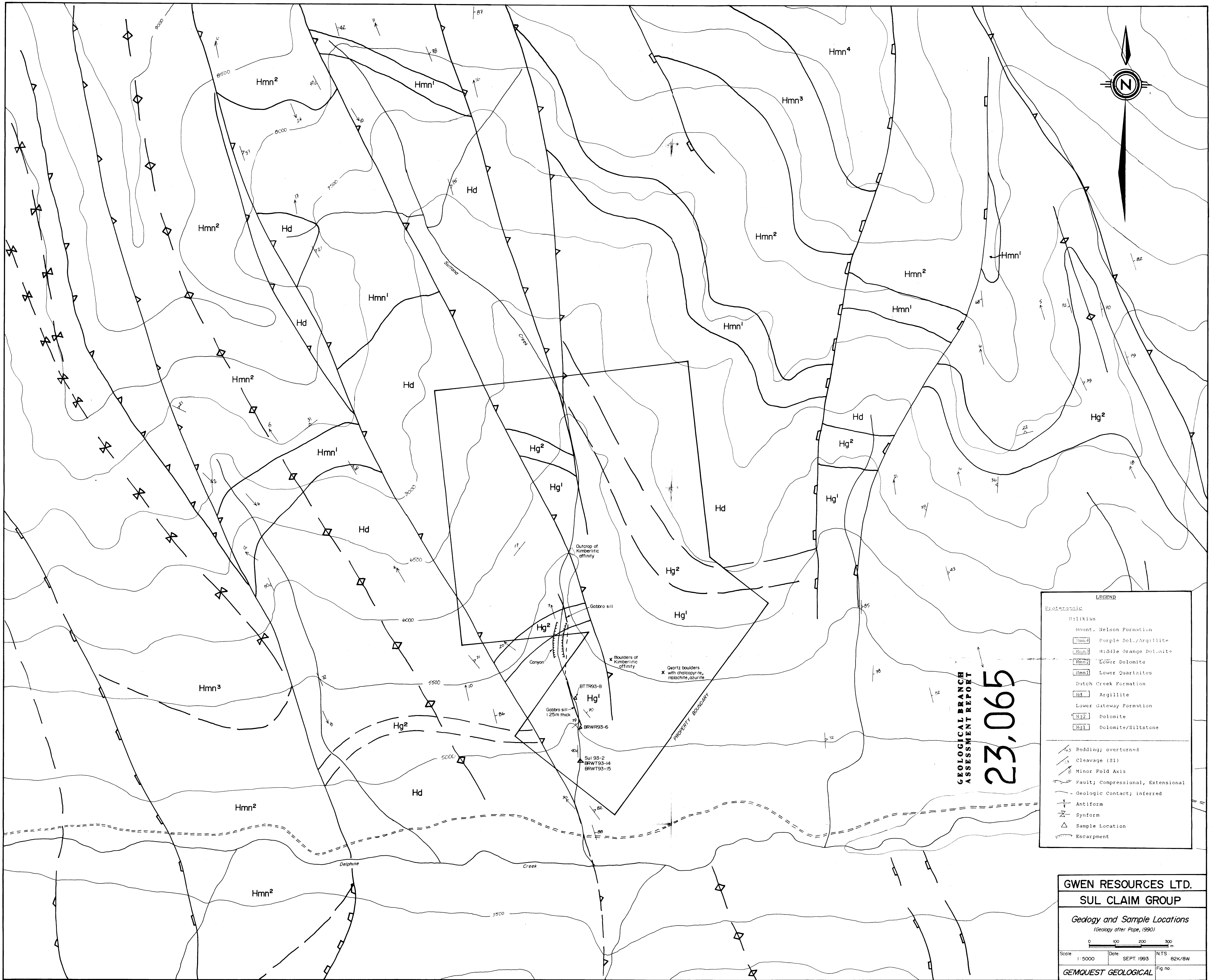
MISCELLANEOUS

| | |
|----------------|----------|
| Fuel | \$ 40.00 |
| Shipping | \$ 20.00 |

REPORT/REPRODUCTION

| | |
|--|-----------|
| R. T. Walker, P.Geo.: 1.0 days x \$350.00/day | \$ 350.00 |
| T. J. Termuende, P.Geo.: 1.0 days x \$350.00/day | \$ 350.00 |

TOTAL EXPENDITURES: \$2,084.40



LEGEND

Geotectonic

Halikian

Hout. Nelson Formation

Hmn⁴ Purple Dol./Argillite

Hmn³ Middle Orange Dolomite

Hmn² Lower Dolomite

Hmn¹ Lower Quartzites

Dutch Creek Formation

Hd Argillite

Lower Gateway Formation

Hg² Dolomite

Hg¹ Dolomite/Siltstone

Redding; overturned

Cleavage (S1)

Minor Fold Axis

Fault; Compressional, Extensional

Geologic Contact; inferred

Antiform

Synform

Sample Location

Escarpment

**GEOLOGICAL BRANCH
ASSESSMENT REPORT**
23,065

GWEN RESOURCES LTD.
SUL CLAIM GROUP

Geology and Sample Locations
(Geology after Pope, 1990)

Scale 1:5000 Date SEPT. 1993 NTS 82K/BW

GEMQUEST GEOLOGICAL Fig no.