District Geologist, Prince George

Off Confidential: 94.11.08

ASSESSMENT REPORT 23097 MINING DIVISION: Omineca

PROPERTY:	Brewster Lake
LOCATION:	LAT 53 22 00 LONG 124 31 00
	UTM 10 5913915 399080
	NTS 093F07E
CLAIM(S):	Brew 1-4
OPERATOR(S):	Cogema Resources Inc.
AUTHOR(S):	Fraser, D.;Pritchard, R.
REPORT YEAR:	1993, 75 Pages
COMMODITIES	
SEARCHED FOR:	Gold
KEYWORDS:	Jurassic,Hazelton Group,Volcanics,Cretaceous,Skeena Group
	Sediments
WORK	
DONE: Geo	physical
EMA	B 150.0 km;VLF

 $\begin{array}{rl} \text{Map(s)} & -5; \text{ Scale(s)} & -1:20 & 000 \\ \text{MAGA} & 150.0 & \text{km} \\ \text{Map(s)} & -1; \text{ Scale(s)} & -1:20 & 000 \end{array}$

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 Report 93-CND-78-04

DIGHEM^V SURVEY FOR COGEMA RESOURCES INC. BREWSTER LAKE PROPERTY BRITISH COLUMBIA

NTS 93F/7,8

GEOLOGICAL BRANCH ASSESSMENT REPORT

Dighem Surveys & Processing Inc. Mississauga, Ontario April 26, 1993

Ruth A. Pritchard Geophysicist

A1138APR.93R

British Coumbia Energy, Mines and Petroleum Resources	ASSESSMENT REPORT TITLE PAGE AND SUMMARY
TYPE OF REPORT/SURVEY(S) Geophysical	TOTAL COST \$12,800.00
AUTHORIS; Ruth A. Pritchard Supervised by Douglas C. Fraser	SNATUREISI
DATE STATEMENT OF EXPLORATION AND DEVELOPMENT FIL PROPERTY NAME(S)	ED
СОИMODITIES PRESENT	
B.C. MINERAL INVENTORY NUMBER(S), IF KNOWN MINING DIVISION Catifude 53 ⁰ 22'N LATITUDE	NTS
NAMES and NUMBERS of all mineral tenures in good standing (when we (12 units): PHOENIX (Lot 1706): Mineral Lease M 123: Mining c: Certified	ork was done) that form the property {Examples: TAX 1-4, FIRE 3 d Mining Lease ML 12 {claims involved]} :
Brew 1 to 4	· · · · · · · · · · · · · · · · · · ·
OWNER(S) (1) COGEMA Resources Inc. (1) 801-409 Granville St	SUB-RECORDER RECEIVED
MAILING ADDRESS 801-409 Granville St	M.R. #
Vancouver B.C. V6C 1T2	
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(1) COGEMA Resources Inc	ion, size, and entitude): Zolcaniclastic rócks and Greatceous Skeena
(1) COGEMA Resources Inc. (2) MAILING ADDRESS 801-409 Granville St, Vancouver B.C. V6C 1T2 SUMMARY GEOLOGY (lithology, egc. structure, alteration, mineralizati Jurassic Hazelton volcanics and v 	ion, size, and ettitude): Folcaniclastic, rócks, and. Creatceous Skeena

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SUMMARY

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This report describes the logistics and results of a DIGHEM^V airborne geophysical survey carried out for Cogema Resources Inc. over the Brewster Lake Property, British Columbia. Total coverage of the survey block amounted to 155 km, including tie lines. The survey was flown from March 8 to March 12, 1993.

The purpose of the survey was to detect zones of conductive mineralization and to provide information that could be used to map the geology and structure of the survey area. This was accomplished by using a DIGHEM^V multi-coil, multi-frequency electromagnetic system, supplemented by a high sensitivity Cesium magnetometer and a two-channel VLF receiver. The information from these sensors was processed to produce maps which display the magnetic and conductive properties of the survey area. An electronic navigation system, operating in the UHF band, ensured accurate positioning of the geophysical data with respect to the base map. Visual flight path recovery techniques were used in areas where transponder signals were blocked by topographic features.

The survey property contains several anomalous features, some of which may be of moderate to high priority as exploration targets. Most of the inferred bedrock conductors appear to warrant further investigation using appropriate surface exploration techniques. Areas of interest may be assigned priorities on the basis of supporting geophysical, geochemical and/or geological information. After initial investigations have been carried out, it may be necessary to re-evaluate the remaining anomalies based on information acquired from the follow-up program.

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LIST OF MAPS

DIGHEM EM anomaly map with interpretation Total field magnetic map Enhanced magnetic map Resistivity (900 Hz coplanar) map Resistivity (7200 Hz coplanar) map Resistivity (56,000 Hz coplanar) map Filtered total field VLF map A DIGHEM^V electromagnetic/resistivity/magnetic/VLF survey was flown for Cogema Resources Inc. from March 8 to March 12, 1993, over the Brewster Lake Property, British Columbia. The survey area can be located on NTS map sheets 93F/7,8 (see Figure 1).

The Brewster Lake property consists of 4 contiguous claims (66 units, 16.5 km²), located by COGEMA in the fall of 1992; the claims are listed on Table 1-1 and shown on Figure 2.

Table 1-1

List of Claims

NAME	RECORD NO	UNITS	STAKING DATE	EXPIRY DATE
BREW 1	314657	20	92/11/15	93/11/15
BREW 2	314658	20	92/11/15	93/11/15
BREW 3	314659	6	92/11/15	93/11/15
BREW 4	314660	20	92/11/15	93/11/15
	TOTAL	66		

Survey coverage consisted of approximately 155 line-km, including tie lines. Flight lines were flown in an azimuthal direction of $90^{\circ}/270^{\circ}$ with a line separation of 200 m.

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The survey employed the DIGHEM^V electromagnetic system. Ancillary equipment consisted of a magnetometer, radar altimeter, video camera, analog and digital recorders, a VLF receiver and an electronic navigation system. Details on the survey equipment are given in Section 2.

The instrumentation was installed in an Aerospatiale AS350B turbine helicopter (Registration C-GJIX) which was provided by Questral Helicopters Ltd. The helicopter flew at an average airspeed of 130 km/h with an EM bird height of approximately 30 m.

Section 2 also provides details on the data channels, their respective sensitivities, and the navigation/flight path recovery procedure. Noise levels of less than 2 ppm are generally maintained for wind speeds up to 35 km/h. Higher winds may cause the system to be grounded because excessive bird swinging produces difficulties in flying the helicopter. The swinging results from the 5 m² of area which is presented by the bird to broadside gusts.



FIGURE 1 COGEMA RESOURCES INC. BREWSTER LAKE PROPERTY - 1138

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Figure 2. Brewster Lake Claim Map.

SURVEY EQUIPMENT

This section provides a brief description of the geophysical instruments used to

acquire the survey data:

Electromagnetic System

Model: DIGHEMV

Type: Towed bird, symmetric dipole configuration operated at a nominal survey altitude of 30 metres. Coil separation is 8 metres for 900 Hz and 7200 Hz, and 6.3 metres for the 56,000 Hz coil-pair.

Coil orientations/frequencies:	coaxial	/ 900 Hz
-	coplanar	/ 900 Hz
	coaxial	/ 7,200 Hz
	coplanar	/ 7,200 Hz
	coplanar	/ 56,000 Hz
Channels recorded:	5 inphase c	hannels
	5 quadratur	e channels
	2 monitor o	channels
Sensitivity:	0.1 ppm at	900 Hz
• 17	0.2 ppm at	7,200 Hz
	0.5 ppm at	56,000 Hz
Sample rate:	10 per seco	ond

The electromagnetic system utilizes a multi-coil coaxial/coplanar technique to energize conductors in different directions. The coaxial coils are vertical with their axes

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in the flight direction. The coplanar coils are horizontal. The secondary fields are sensed simultaneously by means of receiver coils which are maximum coupled to their respective transmitter coils. The system yields an inphase and a quadrature channel from each transmitter-receiver coil-pair.

Magnetometer

Model:	Picodas 3340
Туре:	Optically pumped Cesium vapour
Sensitivity:	0.01 nT
Sample rate:	10 per second

The magnetometer sensor is towed in a bird 20 m below the helicopter.

Magnetic Base Station

Model: Scintrex MP-3

Type: Digital recording proton precession

Sensitivity: 0.10 nT

Sample rate: 0.2 per second

A digital recorder is operated in conjunction with the base station magnetometer to record the diurnal variations of the earth's magnetic field. The clock of the base station is synchronized with that of the airborne system to permit subsequent removal of diurnal drift.

VLF System

Manufacturer:	Herz Industries Ltd.		
Гуре:	Totem-2A		
Sensitivity:	0.1%		
Stations:	Seattle, Washington; Cutler, Maine; Lualualei, Hawaii;	NLK, NAA, NPM,	24.8 kHz 24.0 kHz 23.4 kHz

The VLF receiver measures the total field and vertical quadrature components of the secondary VLF field. Signals from two separate transmitters can be measured simultaneously. The VLF sensor is housed in the same bird as the magnetic sensor, and is towed 20 m below the helicopter.

Radar Altimeter

Manufacturer:	Honeywell/Sperr		
Туре:	AA 220		
Sensitivity:	1 ft		

The radar altimeter measures the vertical distance between the helicopter and the ground. This information is used in the processing algorithm which determines conductor depth.

Analog Recorder

Manufacturer:	RMS Instruments
Туре:	DGR33 dot-matrix graphics recorder
Resolution:	4x4 dots/mm
Speed:	1.5 mm/sec

The analog profiles are recorded on chart paper in the aircraft during the survey. Table 2-1 lists the geophysical data channels and the vertical scale of each profile.

Channel Name	Parameter	Scale units/mm	Designation on digital profile
1X9I 1X9Q 3P9I 3P9Q 2P7I 2P7Q 4X7I 4X7Q 5P5I 5P5Q ALTR CMGC CMGF VF1T VF1Q VF1Q VF2T VF2Q CXSP CPSP CXPL CPPL	coaxial inphase (900 Hz) coaxial quad (900 Hz) coplanar inphase (900 Hz) coplanar quad (900 Hz) coplanar quad (700 Hz) coplanar quad (7200 Hz) coaxial inphase (7200 Hz) coaxial quad (7200 Hz) coaxial quad (7200 Hz) coplanar inphase(56000 Hz) coplanar quad (56000 Hz) altimeter magnetics, coarse magnetics, fine VLF-total: primary stn. VLF-quad: primary stn. VLF-quad: primary stn. VLF-quad: secondary stn. VLF-quad: secondary stn. coaxial spherics monitor coplanar powerline monitor	2.5 ppm 2.5 ppm 2.5 ppm 5 ppm 5 ppm 5 ppm 10 ppm 10 ppm 3 m 20 nT 2.0 nT 2% 2% 2% 2%	CXI (900 Hz) CXQ (900 Hz) CPI (900 Hz) CPQ (900 Hz) CPQ (900 Hz) CPI (7200 Hz) CXI (7200 Hz) CXQ (7200 Hz) CYI (56 kHz) CPQ (56 kHz) ALT MAG VLF (primary total field) VLF (primary quadrature) VLF (secondary total field) VLF (secondary quadrature) CXS CPS CXP CPP

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Table 2-1. The Analog Profiles

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Table 2-2.	The Digital Profiles
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Channel <u>Name (Freq)</u>	Observed parameters	Scale <u>units/mm</u>
MAG ALT CXI (900 Hz) CXQ (900 Hz) CPI (900 Hz) CPQ (900 Hz) CPQ (900 Hz) CXI (7200 Hz) CXQ (7200 Hz) CPI (7200 Hz) CPQ (7200 Hz) CPQ (7200 Hz) CPQ (56 kHz) CPQ (56 kHz) VLF1T VLF1Q VLF2T VLF2Q CXS CXP CPS CPP	magnetics bird height vertical coaxial coil-pair inphase vertical coaxial coil-pair quadrature horizontal coplanar coil-pair inphase horizontal coplanar coil-pair quadrature vertical coaxial coil-pair quadrature horizontal coplanar coil-pair quadrature horizontal coplanar coil-pair inphase horizontal coplanar coil-pair quadrature horizontal coplanar coil-pair quadrature primary station total field primary station quadrature secondary station quadrature coaxial spherics monitor coplanar spherics monitor coplanar powerline monitor	20 nT 6 m 2 ppm 2 ppm 2 ppm 4 ppm 4 ppm 4 ppm 4 ppm 10 ppm 5% 5% 5%
	Computed Parameters	
MAG DFI (900 Hz) DFQ (900 Hz) RES (900 Hz) RES (7200 Hz) RES (56 kHz) DP (900 Hz) DP (7200 Hz) DP (56 kHz) CDT	enhanced magnetics difference function inphase from CXI and CPI difference function quadrature from CXQ and CPQ log resistivity log resistivity log resistivity apparent depth apparent depth conductance	200 nT 2 ppm 2 ppm .06 decade .06 decade .06 decade 6 m 6 m 1 grade

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Digital Data Acquisition System

Manufacturer:	RMS Instruments
Туре:	DGR 33
Tape Deck:	RMS TCR-12, 6400 bpi, tape cartridge recorder

The digital data are used to generate several computed parameters. Both measured and computed parameters are plotted as "multi-channel stacked profiles" during data processing. These parameters are shown in Table 2-2. In Table 2-2, the log resistivity scale of 0.06 decade/mm means that the resistivity changes by an order of magnitude in 16.6 mm. The resistivities at 0, 33 and 67 mm up from the bottom of the digital profile are respectively 1, 100 and 10,000 ohm-m.

Tracking Camera

Type: Panasonic Video

Model: AG 2400/WVCD132

Fiducial numbers are recorded continuously and are displayed on the margin of each image. This procedure ensures accurate correlation of analog and digital data with respect to visible features on the ground.

Navigation System

Model:Del Norte 547Type:UHF electronic positioning systemSensitivity:1 mSample rate:2 per second

The navigation system uses ground based transponder stations which transmit distance information back to the helicopter. The ground stations are set up well away from the survey area and are positioned such that the signals cross the survey block at an angle between 30° and 150°. The onboard Central Processing Unit takes any two transponder distances and determines the helicopter position relative to these two ground stations in cartesian coordinates.

The cartesian coordinates are transformed to UTM coordinates during data processing. This is accomplished by correlating a number of prominent topographical locations with the navigational data points. The use of numerous visual tie points serves two purposes: to accurately relate the navigation data to the map sheet and to minimize location errors which might result from distortions in uncontrolled photomosaic base maps.

Field Workstation

Manufacturer:	Dighem
Model:	FWS: V2.41
Туре:	80386 based P.C

A portable PC-based field workstation is used at the survey base to verify data quality and completeness. Flight tapes are dumped to a hard drive to permit the creation of a database. This process allows the field operators to display both the positional (flight path) and geophysical data on a screen or printer.

Global Positioning System

Manufacturer:	Trimble Navigation Ltd.
Туре:	Pathfinder (C/A code, dual channel)
Accuracy:	25 metres (5 metres in differential mode)
Update:	Once per second

The Pathfinder system uses signals broadcast by the NAVSTAR GPS satellites to provide positional readouts in latitude/longitude or UTM coordinates. The GPS unit is placed at each UHF transponder site to determine its exact location. The system can also be used aboard the helicopter to provide real-time navigation guidance. Recorded data can be downloaded to a field computer for immediate post-survey processing, or transmitted to the central processing facility for final plotting.

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PRODUCTS AND PROCESSING TECHNIQUES

The following products are available from the survey data. Those which are not part of the survey contract may be acquired later. Most parameters can be displayed as contours, profiles, or in colour.

Base Maps

A base map of the survey area has been produced from published topographic maps. These provide a relatively accurate, distortion-free base which facilitates correlation of the navigation data to the UTM grid. Photomosaics are useful for visual reference and for subsequent flight path recovery, but usually contain scale distortions. Orthophotos are ideal, but their cost and the time required to produce them, usually precludes their use as base maps.

Electromagnetic Anomalies

Anomalous electromagnetic responses are selected and analysed by computer to provide a preliminary electromagnetic anomaly map. This preliminary map is used, by the geophysicist, in conjunction with the computer-generated digital profiles, to produce the final interpreted EM anomaly map. This map includes bedrock surficial and cultural conductors. A map containing only bedrock conductors can be generated, if desired.

Resistivity

The apparent resistivity in ohm-m may be generated from the inphase and quadrature EM components for any of the frequencies, using a pseudo-layer halfspace model. A resistivity map portrays all the EM information for that frequency over the entire survey area. This contrasts with the electromagnetic anomaly map which provides information only over interpreted conductors. The large dynamic range makes the resistivity parameter an excellent mapping tool.

EM Magnetite

The apparent percent magnetite by weight is computed wherever magnetite produces a negative inphase EM response.

Total Field Magnetics

The aeromagnetic data are corrected for diurnal variation using the magnetic base station data. The regional IGRF can be removed from the data, if requested.

Enhanced Magnetics

The total field magnetic data are subjected to a processing algorithm. This algorithm enhances the response of magnetic bodies in the upper 500 m and attenuates the response of deeper bodies. The resulting enhanced magnetic map provides better definition and resolution of near-surface magnetic units. It also identifies weak magnetic features which may not be evident on the total field magnetic map. However, regional magnetic variations, and magnetic lows caused by remanence, are better defined on the total field magnetic map. The technique is described in more detail in Section 5.

Magnetic Derivatives

The total field magnetic data may be subjected to a variety of filtering techniques to yield maps of the following:

vertical gradient second vertical derivative magnetic susceptibility with reduction to the pole upward/downward continuations All of these filtering techniques improve the recognition of near-surface magnetic bodies, with the exception of upward continuation. Any of these parameters can be produced on request. Dighem's proprietary enhanced magnetic technique is designed to provide a general "all-purpose" map, combining the more useful features of the above parameters.

VLF

The VLF data are digitally filtered to remove long wavelengths such as those caused by variations in the transmitted field strength.

Multi-channel Stacked Profiles

Distance-based profiles of the digitally recorded geophysical data are generated and plotted by computer. These profiles also contain the calculated parameters which are used in the interpretation process. These are produced as worksheets prior to interpretation, and can also be presented in the final corrected form after interpretation. The profiles display electromagnetic anomalies with their respective interpretive symbols. The differences between the worksheets and the final corrected form occur only with respect to the EM anomaly identifier.

Contour, Colour and Shadow Map Displays

The geophysical data are interpolated onto a regular grid using a modified Akima spline technique. The resulting grid is suitable for generating contour maps of excellent quality.

Colour maps are produced by interpolating the grid down to the pixel size. The parameter is then incremented with respect to specific amplitude ranges to provide colour "contour" maps. Colour maps of the total magnetic field are particularly useful in defining the lithology of the survey area.

Monochromatic shadow maps are generated by employing an artificial sun to cast shadows on a surface defined by the geophysical grid. There are many variations in the shadowing technique. These techniques may be applied to total field or enhanced magnetic data, magnetic derivatives, VLF, resistivity, etc. Of the various magnetic products, the shadow of the enhanced magnetic parameter is particularly suited for defining geological structures with crisper images and improved resolution.

Resistivity Sengpiel Sections

The apparent resistivity and approximate thickness of two or more horizontal layers can be displayed simultaneously for all coplanar frequencies. An inversion algorithm has been developed by Dr. K.P. Sengpiel* of the B.G.R., which determines the generalized skin depth, or "centroid depth" of the inphase current concentration, as a function of frequency. The centroid depth is combined with the apparent resistivity over a broad frequency range to produce resistivity-thickness pseudo-sections. A coloured presentation yields a smoothed representation of the true resistivity-depth distribution within the limits of the model used.

Approximate Inversion of Airborne EM Data from Multilayered Ground: Sengpiel, K.P., Geophysical Prospecting 36, 446-459, 1988.

SURVEY RESULTS

GENERAL DISCUSSION

The survey results are presented on 1 map sheet for each parameter at a scale of 1:20,000. Table 4-1 summarizes the EM responses in the survey area, with respect to conductance grade and interpretation.

The anomalies shown on the electromagnetic anomaly map are based on a nearvertical, half plane model. This model best reflects "discrete" bedrock conductors. Wide bedrock conductors or flat-lying conductive units, whether from surficial or bedrock sources, may give rise to very broad anomalous responses on the EM profiles. These may not appear on the electromagnetic anomaly map if they have a regional character rather than a locally anomalous character. These broad conductors, which more closely approximate a half space model, will be maximum coupled to the horizontal (coplanar) coil-pair and should be more evident on the resistivity parameter. Resistivity maps, therefore, may be more valuable than the electromagnetic anomaly maps, in areas where broad or flat-lying conductors are considered to be of importance. Contoured resistivity maps, based on the 900 Hz, 7200 Hz and 56,000 Hz coplanar data are included with this report. [__]

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TABLE 4-1

EM ANOMALY STATISTICS

BREWSTER LAKE PROPERTY

CONDUCTOR GRADE	CONDUCTANCE RANGE SIEMENS (MHOS)	NUMBER OF RESPONSES
7	>100	0
6	50 ~ 100	1
5	20 - 50	3
4	10 - 20	6
3	5 - 10	21
2	1 - 5	62
1	<1	50
*	INDETERMINATE	52
TOTAL		195
CONDUCTOR MODEL	MOST LIKELY SOURCE	NUMBER OF RESPONSES
D	DISCRETE BEDROCK CONDUCTOR	19
В	DISCRETE BEDROCK CONDUCTOR	75

SCONDUCTIVE COVER61HROCK UNIT OR THICK COVER40TOTAL195

(SEE EM MAP LEGEND FOR EXPLANATIONS)

Excellent resolution and discrimination of conductors was accomplished by using a fast sampling rate of 0.1 sec and by employing common frequencies (900 Hz and 7200 Hz) on two orthogonal coil-pairs (coaxial and coplanar). The resulting "difference channel" parameters often permit differentiation of bedrock and surficial conductors, even though they may exhibit similar conductance values.

Anomalies which occur near the ends of the survey lines (i.e., outside the survey area), should be viewed with caution. Some of the weaker anomalies could be due to aerodynamic noise, i.e., bird bending, which is created by abnormal stresses to which the bird is subjected during the climb and turn of the aircraft between lines. Such aerodynamic noise is usually manifested by an anomaly on the coaxial inphase channel only, although severe stresses can affect the coplanar inphase channels as well.

Magnetics

A Scintrex proton precession magnetometer was operated at the survey base to record diurnal variations of the earth's magnetic field. The clock of the base station was synchronized with that of the airborne system to permit subsequent removal of diurnal drift. The background magnetic level has been adjusted to match the International Geomagnetic Reference Field (IGRF) for the survey area. The IGRF gradient across the survey block is left intact.

The total field magnetic data have been presented as contours on the base map using a contour interval of 5 nT where gradients permit. The map shows the magnetic properties of the rock units underlying the survey area.

The total field magnetic data have been subjected to a processing algorithm to produce an enhanced magnetic map. This procedure enhances near-surface magnetic units and suppresses regional gradients. It also provides better definition and resolution of magnetic units and displays weak magnetic features which may not be clearly evident on the total field map. A map of the first or second vertical magnetic derivative can also be prepared from existing survey data, if requested.

There is some evidence on the magnetic map which suggests that the survey area have been subjected to deformation and/or alteration. These structural complexities are evident on the contour map as variations in magnetic intensity, irregular patterns, and as offsets or changes in strike direction. Some of the more prominent linear features are also evident on the topographic base map. If a specific magnetic intensity can be assigned to the rock type which is believed to host the target mineralization, it may be possible to select areas of higher priority on the basis of the total field magnetic data. This is based on the assumption that the magnetite content of the host rocks will give rise to a limited range of contour values which will permit differentiation of various lithological units.

The magnetic results, in conjunction with the other geophysical parameters, should provide valuable information which can be used to effectively map the geology and structure in the survey area.

VLF

VLF results were obtained from the transmitting stations at Cutler, Maine (NAA - 24.0 kHz), Seattle, Washington (NLK - 24.8 kHz) and Lualualei, Hawaii (NPM - 23.4 kHz). Table 4-2 shows the VLF stations used for the contour map. Table 4-3 displays the VLF stations that were recorded as the primary and secondary stations and that are plotted as total field and quadrature on the digital profiles.

Table 4-2

<u>Area</u>

VLF Station(s)

Brewster Lake

Brewster Lake

Seattle (24.8 kHz - NLK) - all lines

Table 4-3

<u>Area</u>

Primary Station (VLF1) Seattle (24.8 kHz - NLK)

Cutler (24.0 kHz - NAA) - lines 30010-30250 39010-39020 Hawaii (23.4 kHz - NPM) - all other lines

Secondary Station (VLF2)

The VLF method is quite sensitive to the angle of coupling between the conductor and the propagated EM field. Consequently, conductors which strike towards the VLF station will usually yield a stronger response than conductors which are nearly orthogonal to it.

The VLF parameter does not normally provide the same degree of resolution available from the EM data. Closely-spaced conductors, conductors of short strike length or conductors which are poorly coupled to the VLF field, may escape detection with this method. Erratic signals from the VLF transmitters can also give rise to strong, isolated anomalies which should be viewed with caution. The filtered total field VLF contours are presented on the base map with a contour interval of one percent.

Resistivity

Resistivity maps, which display the conductive properties of the survey area, were produced from the 900 Hz, 7200 Hz and 56,000 Hz coplanar data. The maximum resistivity value, which is calculated for each frequency, is approximately 1.15 times the numerical value of the frequency. This cutoff eliminates the meaningless higher resistivities which would result from very small EM amplitudes. In general, the resistivity patterns show some agreement with the magnetic trends. This suggests that a few of the resistivity lows are probably related to bedrock features, rather than conductive overburden. There are some areas, however, where contour patterns appear to be strongly influenced by conductive surficial material.

Electromagnetics

The EM anomalies resulting from this survey appear to fall within one of two general categories. The first type consists of discrete, well-defined anomalies which yield marked inflections on the difference channels. These anomalies are usually attributed to conductive sulphides or graphite and are generally given a "B", "T" or "D" interpretive symbol, denoting a bedrock source. The second class of anomalies comprises moderately broad responses which exhibit the characteristics of a half space and do not yield well-defined inflections on the difference channels. Anomalies in this category are usually given an "S" or "H" interpretive symbol. The lack of a difference channel response usually implies a broad or flat-lying conductive source such as overburden. Some of these anomalies may reflect conductive rock units or zones of deep weathering.

The effects of conductive overburden are evident over portions of the survey areas. Although the difference channels (DFI and DFQ) are extremely valuable in detecting bedrock conductors which are partially masked by conductive overburden, sharp undulations in the bedrock/overburden interface can yield anomalies in the difference channels which may be interpreted as possible bedrock conductors. Such anomalies usually fall into the "S?" or "B?" classification but may also be given an "E" interpretive symbol, denoting a resistivity contrast at the edge of a conductive unit.

In areas where EM responses are evident primarily on the quadrature components, zones of poor conductivity are indicated. Where these responses are coincident with magnetic anomalies, it is possible that the inphase component amplitudes have been suppressed by the effects of magnetite. Most of these poorly-conductive magnetic features give rise to resistivity anomalies which are only slightly below background. If it is expected that poorly-conductive economic mineralization may be associated with magnetite-rich units, most of these weakly anomalous features will be of interest. In areas where magnetite causes the inphase components to become negative, the apparent conductance and depth of EM anomalies may be unreliable.

It is difficult to assess the relative merits of EM anomalies on the basis of conductance. It is recommended that an attempt be made to compile a suite of geophysical "signatures" over areas of interest. Anomaly characteristics are clearly defined on the computer-processed geophysical data profiles which are supplied as one of the survey products.

A complete assessment and evaluation of the survey data should be carried out by one or more qualified professionals who have access to, and can provide a meaningful compilation of, all available geophysical, geological and geochemical data.

The electromagnetic anomaly map shows the anomaly locations with the interpreted conductor type, dip, conductance and depth being indicated by symbols. Direct magnetic correlation is also shown if it exists. The strike direction and length of the conductors are indicated when anomalies can be correlated from line to line. When studying the map sheets, consult the anomaly listings appended to this report.
BREWSTER LAKE PROPERTY

- 4.10 -

The general magnetic strike in the survey block is approximately northwest/southeast. The strongest magnetic intensities are located in several magnetic features which trend northwest/southeast, the length of the survey area. Several possible north-northeast/south-southwest trending linear structural features intersect these magnetic trends.

Many possible bedrock anomalies are associated with several of these magnetic trends between lines 30160 and 30250. The magnetic features are truncated by possible structural breaks to the north and south of this group of anomalies. These anomalies are situated within a large resistivity low. Many of them are quite well-defined and are moderately strong. The possible structural feature which is located to the north of these conductors is also evident as a break in the resistivity contours. Another, small resistivity low is situated immediately north of this possible structural break over the west ends of lines 30090 through 30150. The anomalies within this resistivity low are generally moderately strong and well-defined. Several reflect thin dyke-like sources.

One other portion of the area seems to reflect a possible bedrock source. An elongate conductive zone extends south-southeast from fiducial 4318 on line 30010 to the east end of line 30250. The approximate limit of this zone is defined by the 100 ohmmetre contour on the 900 Hz resistivity map. Much of this zone is indicative of a

conductive half space at depth, although portions of the zone display higher resistivities on 56,000 Hz resistivity.

The zone appears quite continuous on the 900 Hz and 7200 Hz resistivity maps, whereas the 56,000 Hz map displays a less continuous zone intersected by more resistive features. The northern portion of this conductive zone is coincident with a north-northwest/south-southeast trending magnetic feature. This magnetic trend seems to be truncated by a possible structural feature in the vicinity of fiducial 3112 on line 30080. Several possible bedrock anomalies are associated with this magnetic trend. Several are coincident with the resistivity contrast at the edge of the conductive feature.

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BACKGROUND INFORMATION

This section provides background information on parameters which are available from the survey data. Those which have not been supplied as survey products may be generated later from raw data on the digital archive tape.

ELECTROMAGNETICS

DIGHEM electromagnetic responses fall into two general classes, discrete and broad. The discrete class consists of sharp, well-defined anomalies from discrete conductors such as sulfide lenses and steeply dipping sheets of graphite and sulfides. The broad class consists of wide anomalies from conductors having a large horizontal surface such as flatly dipping graphite or sulfide sheets, saline water-saturated sedimentary formations, conductive overburden and rock, and geothermal zones. A vertical conductive slab with a width of 200 m would straddle these two classes.

The vertical sheet (half plane) is the most common model used for the analysis of discrete conductors. All anomalies plotted on the electromagnetic map are analyzed according to this model. The following section entitled **Discrete Conductor Analysis** describes this model in detail, including the effect of using it on anomalies caused by broad conductors such as conductive overburden.

The conductive earth (half space) model is suitable for broad conductors. Resistivity contour maps result from the use of this model. A later section entitled **Resistivity Mapping** describes the method further, including the effect of using it on anomalies caused by discrete conductors such as sulfide bodies.

Geometric interpretation

The geophysical interpreter attempts to determine the geometric shape and dip of the conductor. Figure 5-1 shows typical DIGHEM anomaly shapes which are used to guide the geometric interpretation.

Discrete conductor analysis

The EM anomalies appearing on the electromagnetic map are analyzed by computer to give the conductance (i.e., conductivity-thickness product) in Siemens (mhos) of a vertical sheet model. This is done regardless of the interpreted geometric shape of the conductor. This is not an unreasonable procedure, because the computed conductance increases as the electrical quality of the conductor increases, regardless of its true shape. DIGHEM anomalies are divided into seven grades of conductance, as shown in Table 5-1 below. The conductance in Siemens (mhos) is the reciprocal of resistance in ohms.

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Fig. 5-1 Typical DIGHEM anomaly shapes

Anomaly Grade	<u>Siemens</u>
7	> 100
6	50 - 100
5	20 - 50
4	10 - 20
3	5 - 10
2	1 - 5
1	< 1

The conductance value is a geological parameter because it is a characteristic of the conductor alone. It generally is independent of frequency, flying height or depth of burial, apart from the averaging over a greater portion of the conductor as height increases. Small anomalies from deeply buried strong conductors are not confused with small anomalies from shallow weak conductors because the former will have larger conductance values.

Conductive overburden generally produces broad EM responses which may not be shown as anomalies on the EM maps. However, patchy conductive overburden in otherwise resistive areas can yield discrete anomalies with a conductance grade (cf. Table 5-1) of 1, 2 or even 3 for conducting clays which have resistivities as low as 50 ohm-m. In areas where ground resistivities are below 10 ohm-m, anomalies caused by weathering variations and similar causes can have any conductance grade. The anomaly shapes from the multiple coils often allow such conductors to be recognized, and these are indicated by the letters S, H, and sometimes E on the electromagnetic anomaly map (see EM map legend). For bedrock conductors, the higher anomaly grades indicate increasingly higher conductances. Examples: DIGHEM's New Insco copper discovery (Noranda, Canada) yielded a grade 5 anomaly, as did the neighbouring copper-zinc Magusi River ore body; Mattabi (copper-zinc, Sturgeon Lake, Canada) and Whistle (nickel, Sudbury, Canada) gave grade 6; and DIGHEM's Montcalm nickel-copper discovery (Timmins, Canada) yielded a grade 7 anomaly. Graphite and sulfides can span all grades but, in any particular survey area, field work may show that the different grades indicate different types of conductors.

Strong conductors (i.e., grades 6 and 7) are characteristic of massive sulfides or graphite. Moderate conductors (grades 4 and 5) typically reflect graphite or sulfides of a less massive character, while weak bedrock conductors (grades 1 to 3) can signify poorly connected graphite or heavily disseminated sulfides. Grades 1 and 2 conductors may not respond to ground EM equipment using frequencies less than 2000 Hz.

The presence of sphalerite or gangue can result in ore deposits having weak to moderate conductances. As an example, the three million ton lead-zinc deposit of Restigouche Mining Corporation near Bathurst, Canada, yielded a well-defined grade 2 conductor. The 10 percent by volume of sphalerite occurs as a coating around the fine grained massive pyrite, thereby inhibiting electrical conduction.

Faults, fractures and shear zones may produce anomalies which typically have low conductances (e.g., grades 1 to 3). Conductive rock formations can yield anomalies of any

conductance grade. The conductive materials in such rock formations can be salt water, weathered products such as clays, original depositional clays, and carbonaceous material.

On the interpreted electromagnetic map, a letter identifier and an interpretive symbol are plotted beside the EM grade symbol. The horizontal rows of dots, under the interpretive symbol, indicate the anomaly amplitude on the flight record. The vertical column of dots, under the anomaly letter, gives the estimated depth. In areas where anomalies are crowded, the letter identifiers, interpretive symbols and dots may be obliterated. The EM grade symbols, however, will always be discernible, and the obliterated information can be obtained from the anomaly listing appended to this report.

The purpose of indicating the anomaly amplitude by dots is to provide an estimate of the reliability of the conductance calculation. Thus, a conductance value obtained from a large ppm anomaly (3 or 4 dots) will tend to be accurate whereas one obtained from a small ppm anomaly (no dots) could be quite inaccurate. The absence of amplitude dots indicates that the anomaly from the coaxial coil-pair is 5 ppm or less on both the inphase and quadrature channels. Such small anomalies could reflect a weak conductor at the surface or a stronger conductor at depth. The conductance grade and depth estimate illustrates which of these possibilities fits the recorded data best.

Flight line deviations occasionally yield cases where two anomalies, having similar conductance values but dramatically different depth estimates, occur close together on the same

conductor. Such examples illustrate the reliability of the conductance measurement while showing that the depth estimate can be unreliable. There are a number of factors which can produce an error in the depth estimate, including the averaging of topographic variations by the altimeter, overlying conductive overburden, and the location and attitude of the conductor relative to the flight line. Conductor location and attitude can provide an erroneous depth estimate because the stronger part of the conductor may be deeper or to one side of the flight line, or because it has a shallow dip. A heavy tree cover can also produce errors in depth estimates. This is because the depth estimate is computed as the distance of bird from conductor, minus the altimeter reading. The altimeter can lock onto the top of a dense forest canopy. This situation yields an erroneously large depth estimate but does not affect the conductance estimate.

Dip symbols are used to indicate the direction of dip of conductors. These symbols are used only when the anomaly shapes are unambiguous, which usually requires a fairly resistive environment.

A further interpretation is presented on the EM map by means of the line-to-line correlation of anomalies, which is based on a comparison of anomaly shapes on adjacent lines. This provides conductor axes which may define the geological structure over portions of the survey area. The absence of conductor axes in an area implies that anomalies could not be correlated from line to line with reasonable confidence.

DIGHEM electromagnetic maps are designed to provide a correct impression of conductor quality by means of the conductance grade symbols. The symbols can stand alone with geology when planning a follow-up program. The actual conductance values are printed in the attached anomaly list for those who wish quantitative data. The anomaly ppm and depth are indicated by inconspicuous dots which should not distract from the conductor patterns, while being helpful to those who wish this information. The map provides an interpretation of conductors in terms of length, strike and dip, geometric shape, conductance, depth, and thickness. The accuracy is comparable to an interpretation from a high quality ground EM survey having the same line spacing.

The attached EM anomaly list provides a tabulation of anomalies in ppm, conductance, and depth for the vertical sheet model. The EM anomaly list also shows the conductance and depth for a thin horizontal sheet (whole plane) model, but only the vertical sheet parameters appear on the EM map. The horizontal sheet model is suitable for a flatly dipping thin bedrock conductor such as a sulfide sheet having a thickness less than 10 m. The list also shows the resistivity and depth for a conductive earth (half space) model, which is suitable for thicker slabs such as thick conductive overburden. In the EM anomaly list, a depth value of zero for the conductive earth model, in an area of thick cover, warns that the anomaly may be caused by conductive overburden.

Since discrete bodies normally are the targets of EM surveys, local base (or zero) levels are used to compute local anomaly amplitudes. This contrasts with the use of true zero levels which are used to compute true EM amplitudes. Local anomaly amplitudes are shown in the EM anomaly list and these are used to compute the vertical sheet parameters of conductance and depth. Not shown in the EM anomaly list are the true amplitudes which are used to compute the horizontal sheet and conductive earth parameters.

Questionable Anomalies

DIGHEM maps may contain EM responses which are displayed as asterisks (*). These responses denote weak anomalies of indeterminate conductance, which may reflect one of the following: a weak conductor near the surface, a strong conductor at depth (e.g., 100 to 120 m below surface) or to one side of the flight line, or aerodynamic noise. Those responses that have the appearance of valid bedrock anomalies on the flight profiles are indicated by appropriate interpretive symbols (see EM map legend). The others probably do not warrant further investigation unless their locations are of considerable geological interest.

The thickness parameter

DIGHEM can provide an indication of the thickness of a steeply dipping conductor. The amplitude of the coplanar anomaly (e.g., CPI channel on the digital profile) increases relative to the coaxial anomaly (e.g., CXI) as the apparent thickness increases, i.e., the thickness in the horizontal plane. (The thickness is equal to the conductor width if the conductor dips at 90

degrees and strikes at right angles to the flight line.) This report refers to a conductor as <u>thin</u> when the thickness is likely to be less than 3 m, and <u>thick</u> when in excess of 10 m. Thick conductors are indicated on the EM map by parentheses "()". For base metal exploration in steeply dipping geology, thick conductors can be high priority targets because many massive sulfide ore bodies are thick, whereas non-economic bedrock conductors are often thin. The system cannot sense the thickness when the strike of the conductor is subparallel to the flight line, when the conductor has a shallow dip, when the anomaly amplitudes are small, or when the resistivity of the environment is below 100 ohm-m.

Resistivity mapping

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Areas of widespread conductivity are commonly encountered during surveys. In such areas, anomalies can be generated by decreases of only 5 m in survey altitude as well as by increases in conductivity. The typical flight record in conductive areas is characterized by inphase and quadrature channels which are continuously active. Local EM peaks reflect either increases in conductivity of the earth or decreases in survey altitude. For such conductive areas, apparent resistivity profiles and contour maps are necessary for the correct interpretation of the airborne data. The advantage of the resistivity parameter is that anomalies caused by altitude changes are virtually eliminated, so the resistivity data reflect only those anomalies caused by conductivity changes. The resistivity analysis also helps the interpreter to differentiate between conductive trends in the bedrock and those patterns typical of conductive overburden. For

over narrow, conductive bodies and (ii) over the boundary zone between two wide formations of differing conductivity.

The resistivity map might be likened to a total field map and the EM map to a horizontal gradient in the direction of flight². Because gradient maps are usually more sensitive than total field maps, the EM map therefore is to be preferred in resistive areas. However, in conductive areas, the absolute character of the resistivity map usually causes it to be more useful than the EM map.

Interpretation in conductive environments

Environments having background resistivities below 30 ohm-m cause all airborne EM systems to yield very large responses from the conductive ground. This usually prohibits the recognition of discrete bedrock conductors. However, DIGHEM data processing techniques produce three parameters which contribute significantly to the recognition of bedrock conductors. These are the inphase and quadrature difference channels (DFI and DFQ), and the resistivity and depth channels (RES and DP) for each coplanar frequency.

² The gradient analogy is only valid with regard to the identification of anomalous locations.

The EM difference channels (DFI and DFQ) eliminate most of the responses from conductive ground, leaving responses from bedrock conductors, cultural features (e.g., telephone lines, fences, etc.) and edge effects. Edge effects often occur near the perimeter of broad conductive zones. This can be a source of geologic noise. While edge effects yield anomalies on the EM difference channels, they do not produce resistivity anomalies. Consequently, the resistivity channel aids in eliminating anomalies due to edge effects. On the other hand, resistivity anomalies will coincide with the most highly conductive sections of conductive ground, and this is another source of geologic noise. The recognition of a bedrock conductor in a conductive environment therefore is based on the anomalous responses of the two difference channels (DFI and DFQ) and the resistivity channels (RES). The most favourable situation is where anomalies coincide on all channels.

The DP channels, which give the apparent depth to the conductive material, also help to determine whether a conductive response arises from surficial material or from a conductive zone in the bedrock. When these channels ride above the zero level on the digital profiles (i.e., depth is negative), it implies that the EM and resistivity profiles are responding primarily to a conductive upper layer, i.e., conductive overburden. If the DP channels are below the zero level, it indicates that a resistive upper layer exists, and this usually implies the existence of a bedrock conductor. If the low frequency DP channel is below the zero level and the high frequency DP is above, this suggests that a bedrock conductor occurs beneath conductive cover.

The conductance channel CDT identifies discrete conductors which have been selected by computer for appraisal by the geophysicist. Some of these automatically selected anomalies on channel CDT are discarded by the geophysicist. The automatic selection algorithm is intentionally oversensitive to assure that no meaningful responses are missed. The interpreter then classifies the anomalies according to their source and eliminates those that are not substantiated by the data, such as those arising from geologic or aerodynamic noise.

Reduction of geologic noise

Geologic noise refers to unwanted geophysical responses. For purposes of airborne EM surveying, geologic noise refers to EM responses caused by conductive overburden and magnetic permeability. It was mentioned previously that the EM difference channels (i.e., channel DFI for inphase and DFQ for quadrature) tend to eliminate the response of conductive overburden. This marked a unique development in airborne EM technology, as DIGHEM is the only EM system which yields channels having an exceptionally high degree of immunity to conductive overburden.

Magnetite produces a form of geological noise on the inphase channels of all EM systems. Rocks containing less than 1% magnetite can yield negative inphase anomalies caused by magnetic permeability. When magnetite is widely distributed throughout a survey area, the inphase EM channels may continuously rise and fall, reflecting variations in the magnetite percentage, flying height, and overburden thickness. This can lead to difficulties in recognizing

deeply buried bedrock conductors, particularly if conductive overburden also exists. However, the response of broadly distributed magnetite generally vanishes on the inphase difference channel DFI. This feature can be a significant aid in the recognition of conductors which occur in rocks containing accessory magnetite.

EM magnetite mapping

The information content of DIGHEM data consists of a combination of conductive eddy current responses and magnetic permeability responses. The secondary field resulting from conductive eddy current flow is frequency-dependent and consists of both inphase and quadrature components, which are positive in sign. On the other hand, the secondary field resulting from magnetic permeability is independent of frequency and consists of only an inphase component which is negative in sign. When magnetic permeability manifests itself by decreasing the measured amount of positive inphase, its presence may be difficult to recognize. However, when it manifests itself by yielding a negative inphase anomaly (e.g., in the absence of eddy current flow), its presence is assured. In this latter case, the negative component can be used to estimate the percent magnetite content.

A magnetite mapping technique was developed for the coplanar coil-pair of DIGHEM. The technique yields a channel (designated FEO) which displays apparent weight percent magnetite according to a homogeneous half space model.³ The method can be complementary to magnetometer mapping in certain cases. Compared to magnetometry, it is far less sensitive but is more able to resolve closely spaced magnetite zones, as well as providing an estimate of the amount of magnetite in the rock. The method is sensitive to 1/4% magnetite by weight when the EM sensor is at a height of 30 m above a magnetitic half space. It can individually resolve steep dipping narrow magnetite-rich bands which are separated by 60 m. Unlike magnetometry, the EM magnetite method is unaffected by remanent magnetism or magnetic latitude.

The EM magnetite mapping technique provides estimates of magnetite content which are usually correct within a factor of 2 when the magnetite is fairly uniformly distributed. EM magnetite maps can be generated when magnetic permeability is evident as negative inphase responses on the data profiles.

Like magnetometry, the EM magnetite method maps only bedrock features, provided that the overburden is characterized by a general lack of magnetite. This contrasts with resistivity mapping which portrays the combined effect of bedrock and overburden.

³ Refer to Fraser, 1981, Magnetite mapping with a multi-coil airborne electromagnetic system: Geophysics, v. 46, p. 1579-1594.

Recognition of culture

Cultural responses include all EM anomalies caused by man-made metallic objects. Such anomalies may be caused by inductive coupling or current gathering. The concern of the interpreter is to recognize when an EM response is due to culture. Points of consideration used by the interpreter, when coaxial and coplanar coil-pairs are operated at a common frequency, are as follows:

- 1. Channels CXP and CPP monitor 60 Hz radiation. An anomaly on these channels shows that the conductor is radiating power. Such an indication is normally a guarantee that the conductor is cultural. However, care must be taken to ensure that the conductor is not a geologic body which strikes across a power line, carrying leakage currents.
- 2. A flight which crosses a "line" (e.g., fence, telephone line, etc.) yields a center-peaked coaxial anomaly and an m-shaped coplanar anomaly.⁴ When the flight crosses the cultural line at a high angle of intersection, the amplitude ratio of coaxial/coplanar response is 4. Such an EM anomaly can only be caused by a line. The geologic body which yields anomalies most closely resembling a line is the vertically dipping thin dike. Such a body, however, yields an amplitude ratio of 2 rather than 4. Consequently, an

⁴ See Figure 5-1 presented earlier.

m-shaped coplanar anomaly with a CXI/CPI amplitude ratio of 4 is virtually a guarantee that the source is a cultural line.

- 3. A flight which crosses a sphere or horizontal disk yields center-peaked coaxial and coplanar anomalies with a CXI/CPI amplitude ratio (i.e., coaxial/coplanar) of 1/4. In the absence of geologic bodies of this geometry, the most likely conductor is a metal roof or small fenced yard.⁵ Anomalies of this type are virtually certain to be cultural if they occur in an area of culture.
- 4. A flight which crosses a horizontal rectangular body or wide ribbon yields an m-shaped coaxial anomaly and a center-peaked coplanar anomaly. In the absence of geologic bodies of this geometry, the most likely conductor is a large fenced area.⁵ Anomalies of this type are virtually certain to be cultural if they occur in an area of culture.
- 5. EM anomalies which coincide with culture, as seen on the camera film or video display, are usually caused by culture. However, care is taken with such coincidences because a geologic conductor could occur beneath a fence, for example. In this example, the fence would be expected to yield an m-shaped coplanar anomaly as in case #2 above.

⁵ It is a characteristic of EM that geometrically similar anomalies are obtained from: (1) a planar conductor, and (2) a wire which forms a loop having dimensions identical to the perimeter of the equivalent planar conductor.

If, instead, a center-peaked coplanar anomaly occurred, there would be concern that a thick geologic conductor coincided with the cultural line.

6. The above description of anomaly shapes is valid when the culture is not conductively coupled to the environment. In this case, the anomalies arise from inductive coupling to the EM transmitter. However, when the environment is quite conductive (e.g., less than 100 ohm-m at 900 Hz), the cultural conductor may be conductively coupled to the environment. In this latter case, the anomaly shapes tend to be governed by current gathering. Current gathering can completely distort the anomaly shapes, thereby complicating the identification of cultural anomalies. In such circumstances, the interpreter can only rely on the radiation channels and on the camera film or video records.

MAGNETICS

The existence of a magnetic correlation with an EM anomaly is indicated directly on the EM map. In some geological environments, an EM anomaly with magnetic correlation has a greater likelihood of being produced by sulfides than one that is non-magnetic. However, sulfide ore bodies may be non-magnetic (e.g., the Kidd Creek deposit near Timmins, Canada) as well as magnetic (e.g., the Mattabi deposit near Sturgeon Lake, Canada).

The magnetometer data are digitally recorded in the aircraft to an accuracy of 0.01 nT for cesium magnetometers. The digital tape is processed by computer to yield a total field magnetic contour map. When warranted, the magnetic data may also be treated mathematically to enhance the magnetic response of the near-surface geology, and an enhanced magnetic contour map is then produced. The response of the enhancement operator in the frequency domain is illustrated in Figure 5-2. This figure shows that the passband components of the airborne data are amplified 20 times by the enhancement operator. This means, for example, that a 100 nT anomaly on the enhanced map reflects a 5 nT anomaly for the passband components of the airborne data.

The enhanced map, which bears a resemblance to a downward continuation map, is produced by the digital bandpass filtering of the total field data. The enhancement is equivalent to continuing the field downward to a level (above the source) which is 1/20th of the actual sensor-source distance.

Because the enhanced magnetic map bears a resemblance to a ground magnetic map, it simplifies the recognition of trends in the rock strata and the interpretation of geological structure. It defines the near-surface local geology while de-emphasizing deep-seated regional features. It primarily has application when the magnetic rock units are steeply dipping and the earth's field dips in excess of 60 degrees.





2 Frequency response of magnetic enhancement operator.

Any of a number of filter operators may be applied to the magnetic data, to yield vertical derivatives, continuations, magnetic susceptibility, etc. These may be displayed in contour, colour or shadow.

VLF

VLF transmitters produce high frequency uniform electromagnetic fields. However, VLF anomalies are not EM anomalies in the conventional sense. EM anomalies primarily reflect eddy currents flowing in conductors which have been energized inductively by the primary field. In contrast, VLF anomalies primarily reflect current gathering, which is a non-inductive phenomenon. The primary field sets up currents which flow weakly in rock and overburden, and these tend to collect in low resistivity zones. Such zones may be due to massive sulfides, shears, river valleys and even unconformities.

The VLF field is horizontal. Because of this, the method is quite sensitive to the angle of coupling between the conductor and the transmitted VLF field. Conductors which strike towards the VLF station will usually yield a stronger response than conductors which are nearly orthogonal to it. AMPLITUDE

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CYCLES / METRE

Fig. 5-3 Frequency response of VLF operator.

The Herz Industries Ltd. Totem VLF-electromagnetometer measures the total field and vertical quadrature components. Both of these components are digitally recorded in the aircraft with a sensitivity of 0.1 percent. The total field yields peaks over VLF current concentrations whereas the quadrature component tends to yield crossovers. Both appear as traces on the profile records. The total field data are filtered digitally and displayed as contours to facilitate the recognition of trends in the rock strata and the interpretation of geologic structure.

The response of the VLF total field filter operator in the frequency domain (Figure 5-3) is basically similar to that used to produce the enhanced magnetic map (Figure 5-2). The two filters are identical along the abscissa but different along the ordinant. The VLF filter removes long wavelengths such as those which reflect regional and wave transmission variations. The filter sharpens short wavelength responses such as those which reflect local geological variations.

CONCLUSIONS AND RECOMMENDATIONS

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This report provides a very brief description of the survey results and describes the equipment, procedures and logistics of the survey.

The survey was successful in locating several anomalous zones which may warrant additional work. The various maps included with this report display the magnetic and conductive properties of the survey area. It is recommended that the survey results be reviewed in detail, in conjunction with all available geophysical, geological and geochemical information. Particular reference should be made to the computer generated data profiles which clearly define the characteristics of the individual anomalies.

Most anomalies in the area are moderately weak and poorly-defined. Many have been attributed to conductive overburden or deep weathering, although a few appear to be associated with magnetite-rich rock units. Others coincide with possible structural features which may reflect faults or shears. Such structural breaks are considered to be of particular interest as they may have influenced mineral deposition within the survey area.

The interpreted bedrock conductors defined by the survey should be subjected to further investigation, using appropriate surface exploration techniques. Anomalies which are currently considered to be of moderately low priority may require upgrading if follow-up results are favourable.

It is also recommended that image processing of existing geophysical data be considered, in order to extract the maximum amount of information from the survey results. Current software and imaging techniques often provide valuable information on structure and lithology, which may not be clearly evident on the contour and colour maps. These techniques can yield images which define subtle, but significant, structural details.

Respectfully submitted,

DIGHEM SURVEYS & PROCESSING INC.

Patchard

Ruth A. Pritchard Geophysicist

RAP/sdp

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APPENDIX A

LIST OF PERSONNEL

The following personnel were involved in the acquisition, processing, interpretation and presentation of data, relating to a DIGHEM^v airborne geophysical survey carried out for Cogema Resources Inc. over the Brewster Lake Property, British Columbia.

Vice President, Operations Steve Kilty Robert Gordon Survey Operations Supervisor Dave Miles Senior Geophysical Operator Maurie Bergstrom Second Geophysical Operator Del Rokosh Pilot (Questral Helicopters Ltd.) Gordon Smith Data Processing Supervisor Ruth A. Pritchard Interpretation Geophysicist Lyn Vanderstarren Drafting Supervisor Steve Mast Draftsperson (CAD) Susan Pothiah Word Processing Operator Albina Tonello Secretary/Expeditor

The survey consisted of 155 km of coverage, including tie lines, flown from March 8 to March 12, 1993.

All personnel are employees of Dighem Surveys & Processing Inc., except for the pilot who is an employee of Questral Helicopters Ltd.

DIGHEM SURVEYS & PROCESSING INC.

Ruth A. Pritchard Geophysicist

RAP/sdp

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APPENDIX B

STATEMENT OF EXPENDITURE

Date: April 26, 1993

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IN ACCOUNT WITH DIGHEM SURVEYS & PROCESSING INC.

To: Dighem flying of Agreement dated December 15, 1992, pertaining to an Airborne Geophysical Survey over the Brewster Lake Property, British Columbia.

Survey Charges

150 km of traverse line flying @ \$77.00/km plus mobilization costs of \$1,250.00.

<u>\$12,800,00</u>

Allocation of Costs

- Data Acquisition	(60%)
- Data Processing	(20%)
- Interpretation, Report and Maps	(20%)

DIGHEM SURVEYS & PROCESSING INC.

RAchard

Ruth A. Pritchard Geophysicist

RAP/sdp

A1138APR.93R

APPENDIX C

STATEMENT OF QUALIFICATIONS

I, Ruth A. Pritchard of the City of Brampton, Province of Ontario, do hereby certify that:

- 1. I am a geophysicist, residing at 31 Barrington Crescent, Brampton, Ontario, L6Z 1N2.
- 2. I am a graduate of York University, Downsview, Ontario, with a Specialized Honours B.Sc. Earth Sciences Geophysics (1986).
- 3. I have been actively engaged in geophysical exploration since 1986.
- 4. The statements made in this report represent my best opinion and judgement.

I, Douglas C. Fraser of the City of Mississauga, Province of Ontario, do hereby certify that I am a Professional Engineer of Ontario, Registration Number 14934012, and that I have supervised the reporting herein.

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Ruth A. Pritchard Geophysicist

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Douglas C. Fraser Geophysicist

APPENDIX D

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EM ANOMALY LIST

	COAXIAL 1126 HZ		COPI 87	COPLANARCOPLANAR874HZ7219HZ			. VERTI	CAL E	HORIZONTAL SHEET		CONDUCTIVE EARIH		MAG CORR	
מאנ	OMALV / I	.	GIAD	DFAT.		DFAT.	OTIAD	• •	\FDTH★		भाष्यास	PFSTS	DEDIH	
FID	/INTERP	PPM	PPM	PPM	PPM	PPM	PPM	.SIEMEN	M III.	.SIEMEN	M	OHM-M	M	NT
 T TN	E 20010	/1	T T CLIT	1 1 2 1				•		•				
	42985	1	2	ر <u>عد</u>	2	2	4	•	-	• . –	_	_	_	0
B	4318H	4	1	3	12	2	2	. 72.0	70	. 2	65	51	33	ō
c	43385?	2	11	2	17	42	114	. 0.8	4	. 1	36	173	2	0
	E 30020	(Ŧ	य उत्स म	12)	1			•		•				
A	41355	1	8	0	12	38	52	. 0.5	0	. 1	32	690	0	0
В	41205	ō	4	1	5	18	35	. 0.4	0	. 1	19	525	0	0
С	4098H	1	2	1	2	2	4	. –	-		-	-	-	0
D	4095B?	8	12	13	20	31	41	. 3.9	15	. 2	43	48	15	0
Ε	4067S?	0	6	1	18	23	89	. 0.4	6	. 1	20	468	0	0
LIN	E 30030	(I	LIGHI	12))			•		•				
A	39205	Ò	1	0	1	2	4	. –	_	. –	-	-	-	0
в	3961S	0	8	0	14	17	50	. 0.4	1	. 1	40	692	0	0
C	3986S	1	2	Ō	2	2	4	. –	-		-	_		0
D	3994B?	1	2	1	2	2	4		-	. –	_		-	40
E	4007H	8	38	17	61	203	306	. 1.5	0	. 2	29	43	7	6
F	4014B?	1	1	1	2	2	4	• -	-	. –	-	-	-	0
LIN	E 30040	(1	TIGHT	12)	h			•		•				
A	3862S	1	2	0	2	2	4	. –	_	. –	-	_	_	0
В	3832S?	0	2	0	2	2	4	. –	-	. –	-	_	~~	210
С	3802S	0	6	0	9	19	78	. 0.4	0	. 1	26	592	0	0
D	3787B?	7	3	10	0	29	29	. 20.1	53	. 2	81	44	49	0
Е	3776H	9	19	13	46	46	113	. 3.0	9	. 2	36	39	12	0
F	3767B?	21	49	26	74	358	353	. 3.5	7	. 1	19	60	0	0
LIN	E 30050	(1	LIGHI	12)	ľ			•		•				
A	35455	ò	2	0	2	2	4	. –	_	. –	_	-	_	0
в	3620B?	10	10	17	19	10	48	. 6.2	32	. 1	47	100	14	0
С	3627H	9	10	10	9	12	31	. 5.7	26	. 2	57	33	31	0
D	3641S?	4	13	7	19	71	99	. 1.6	7	. 1	35	94	4	0
T.TN	E 30060	(1	त. इ.स.	י 121 י				•		•				
Δ	34705	0	7) 		12	71	. 0.4	4	. 1	35	646	0	0
B	34615	ŏ	2	ŏ	2	2	4			• -	-	-	_ _	ň
č	3421H	1	2	ĩ	2	2	4		-		_	_	_	30
D	3389H	1	1	1	2	2	4	. –	-	. –	-	-	-	0
 T TN		/1	ת דריד	1 1 2 1				•		•				
ע	2350C	(1)	an chi	. 12) 1	2	22	43	• 07	0	•	20	202	^	0
А	77709	U	4	T	J	22	41	. 0./	U	• 1	20	252	υ	U
	• * দণ্ড	יימאדיו	יים בייי	י שוסי		R TINDI	TART	E BECAUC	e ane	CTECAIC	וגרב כביה	лт		
	. OF	THE	CONDL		MAVI			R TO ONT		OF THE	ניזים את באד דק	ivit e Prr		
	ית ז	JE. C)R BFC	AUSE	OF A	SHALL		P OR OVE	RBURD	NATAT NA	TS_	···· •		
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	COAXIAL 1126 HZ		XTAL 26 HZ	COPLANAR C 874 HZ		COPI 723	ANAR 19 HZ	. VERT	. VERTICAL . . DIKE .		. HORIZONTAL . SHEET		CONDUCTIVE EARTH		
۵N	OMATV/	DENT.		PFAT.	OTAD	BEAT.	OTTAD		NFDTH*		neeru	RESTS	DEPTH		
FID	/INTERP	PPM	PPM	PPM	PPM	PPM	PPM	.SIEMEN	M N	SIEMEN	M	OHM-M	M	NT	
								•		•					
LIN	E 30070	(F	LIGHT	' 12)	-			•		•				•	
В	32695	0	2	0	2	2	4	•	-	• -	-	-		0	
C	3320B?	11	8	14	13	32	35	. 9.5	37	• 1	50	56	21	0	
D	3342H	0	2	T	2	2	4		-	• -	-	-	-	0	
LIN	E 30080	(F	TIGHT	12)	n			•		•					
	31805	1	2	1	2	2	4	. –	_	. –			-	0	
В	3130S	1	2	ī	2	2	4	. –	_		-	_	-	Ō	
Ċ	3114B?	1	2	1	2	2	4	-	-		_	_	_	Ō	
D	3109B?	2	17	13	25	70	73	0.7	2	2	51	43	25	0	
Ē	3090H	ĩ	2	1	2	2	4	. –	-		_	-	-	ō	
		_	-	_	_	-	_	•		•					
LIN	E 30090	(F	LIGHT	' 12))			•		•					
Α	2911B?	6	21	7	32	115	105	. 1.6	7	. 1	32	119	4	0	
В	2923H	0	7.	3	11	29	32	. 0.4	0	. 1	39	421	0	0	
С	2944S	0	5	0	6	6	27	. 0.4	0	. 1	47	769	0	0	
D	29615	0	8	0	11	23	78	. 0.4	0	. 1	26	620	0	0	
\mathbf{E}	29735	6	14	12	23	66	120	. 2.4	17	. 1	35	77	7	0	
F	3005H	1	2	1	2	2	4		-		-		-	0	
								•		•					
LIN	E 30100	(F	IIGHI	' 12)	Ì			•		•					
A	2870H	2	11	1	21	79	95	. 0.7	3	. 1	34	128	3	0	
В	2851B?	2	10	5	17	57	29	. 1.0	5	. 1	36	432	0	0	
С	2816S	0	3	1	9	22	0	. 0.4	0	. 1	36	631	0	0	
D	2802S	1	2	1	2	2	4		-			-	-	0	
E	2782H	1	2	1	2	2	4		-		-	-	-	0	
\mathbf{F}	2775H	4	34	19	83	143	483	. 0.6	0	. 1	24	63	1	0	
								•		•					
LIN	E 30110	(F	ITCHI	' 12))			•		•					
Α	2588B?	2	6	2	6	31	16	. 1.4	14	. 1	50	81	17	0	
В	2600D	8	12	15	14	37	34	. 3.9	15	. 1	43	148	6	0	
С	2643S	0	5	0	8	21	35	. 0.4	0	. 1	42	708	0	0	
D	2661S?	1	2	1	2	2	4		-		-	-	-	0	
\mathbf{E}	2681H	7	35	14	66	139	80	. 1.4	6	. 1	28	63	6	20	
\mathbf{F}	2700H	5	7	1	9	24	39	. 4.1	33	. 1	42	179	5	0	
								•		•					
LIN	E 30120	(1	LIGHT	12))	-		•		•				_	
A	2534B?	1	2	1	2	2	4	• -	-	• -	-	-	_	0	
В	2530D	22	28	23	40	130	90	. 6.3	1	. 1	31	73	2	0	
С	2511B?	0	2	0	2	2	4	. –	•••	• -		-	-	0	
Ð	2476S	1	11	1	14	37	78	. 0.4	0	. 1	25	611	0	0	
Έ	2474S?	1	2	0	2	2	4	• -	-	• -	-	-	-	120	
	•											•			
	•* <u>ES</u>		ED DE	PIH N	1AY BI		<u> LIABI</u>	E BECAU	SE THE	STRONG		RT.			
	. OF	THE	CONDU	CTOR	MAY I	BE DEI	PER C	R TO ON	E SIDE	OF THE	FLIG	HT .			
	. LI	NE, C	OR BEC	AUSE	OF A	SHALI	LOW DI	POROV	ERBURD	EN EFFE	CTS.	•			

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		004 112	XIAL 26 HZ	COPI 87	ANAR 4 HZ	COPI 721	ANAR 19 HZ	. VERTICAL . . DIKE .		. HORIZO	. HORIZONIAL . SHEET		CONDUCITIVE EARTH		
AN	OMALY/ 1	Y/ REAL QUAD REAL QUAD REAL QUAD . COND DEPIH)EPIH*	COND D	EPIH	RESIS	DEPIH							
FID	/INTERP	PPM	PPM	PPM	PPM	PPM	PPM	SIEMEN	М	SIEMEN	М	OHM-M	M	NT	
т тл		/1	תוכד זי	י יי				•		•					
E LTIN	50120 2429H	(1 4	15	. 12) 8	24	61	114	• 1.4	4	•	37	90	7	0	
		т	10	Ŭ	64	U1	TT4	•	-	• -	57	20		0	
LIN	E 30130	(I	LICHI	. 12)	•			•		•					
Α	2171D	14	8	21	6	23	21	. 14.2	15	. 2	41	37	14	0	
В	2176D	12	15	17	26	75	30	. 5.1	2	. 1	53	161	10	0	
С	2226S	0	7	2	7	31	41	. 0.4	0	. 1	41	532	0	0	
D	2270H	1	2	2	4	9	10	. 2.5	72	. 1	71	73	36	17	
 T TN	E 20140	/1	न मत्वम	1 1 2 1				•		•					
	2135B	1) 10	ТШСШ СПСП	. 12) 29	27	75	30	• 65	5	• 1	13	70	12	0	
а В	2133D	- 19	10	20	27	75	25	· 0.5	13	• 1	51 51	2/0	12	0	
Č	20036	0	10	0	5	1/	13	. 2.7	13	• 1	36	696	0	0	
n	20933 2077B?	1	4 Q	ט ר	13	4	69	. 0.4	0	• 1	21	513	0	0	
E	2070B?	Ō	5	3	7	27	31	0.4	Ő	. 1	22	433	Ő	60	
F	2050H	1	1	1	2	2	4		-	· · ·	_		_	0	
		_	_		_		-	•		•				-	
LIN	E 30150	(1	LICHI	12)	1			•		•					
Α	1843B	12	11	24	19	52	31	. 7.4	19	. 2	44	46	17	0	
в	1848B?	3	11	10	17	62	42	. 1.5	0	. 1	54	170	10	0	
С	1872S	1	2	1	2	2	4	• •	-	• •	-	-	-	20	
D	1897S?	0	15	2	21	73	77	. 0.4	0	• 1	15	303	0	40	
E	1903S?	0	2	3	20	42	106	. 0.4	0	. 1	15	300	0	20	
F	1922B?	1	2	1	2	2	4	• -		• -	-	-	-	0	
G	1926B?	4	5	6	21	47	37	. 3.8	47	. 1	34	157	2	0	
H	1946H	1	2	1	2	2	4	• -	-		-	-	-	0	
I	19615	3	9	4	14	34	73	. 1.5	16	. 1	39	123	7	0	
T.TN	E 30160	(1	य उद्धय	ר י	1			•		•					
Δ	1799B?	1	5	. 12,	5	25	15	• • 0.4	0	• _ 1	61	231	13	0	
B	1795B?	4	6	3	10	6		. 2.9	11	. 1	48	152		ŏ	
Ē	1791B?	4	5	4	11	6	35	. 3.1	24	. 1	41	153	2	Ō	
D	1787B?	Ō	9	3	12	15	38	. 0.4	0	. 1	41	198	0	Ō	
E	1760S	1	0	2	9	3	34	. 36.1	126	. 1	35	309	Ō	0	
F	1729H	1	1	1	2	2	4		-	. –	-			0	
G	1706H	3	9	5	12	33	40	. 1.7	20	. 1	49	77	17	0	
								•		•					
LIN	E 30170	(H	TIGHI	r 12))			•	_	•				_	
Α	1538B?	0	11	4	11	40	45	. 0.4	0	. 1	45	171	3	0	
В	1547D	2	15	7	19	62	77	. 0.6	0	. 1	35	125	2	0	
С	1551D	0	17	7	18	72	59	. 0.4	0	. 1	28	193	0	0	
D	1565B?	0	1	2	0	22	33	. 0.4	0	. 1	64	810	0	220	
	• • EXT		יירו רוסיי	-	ים ערא	ורוזאדז ק	7 T X D T	יזגריזם ש	יזה הנות	CUIDONICE	יארד כדי	•			
	•• ES	TTANT T			<u>ולם וראי</u> זי עראת	ישריו יקר יישריו יקר			асти ТПЕ Тор		ET TOT	кт . um			
	. Ur	NE (NR RFY	ALICE	OF 7	SHATI	OW DT	PORONI		EN FEFF	TS				
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		COAXIAL COPLANAR COPLANAR VER 1126 HZ 874 HZ 7219 HZ D		. VERT	VERITICAL HORIZONTAL DIKE SHEET		CONDU EAR	MAG CORR						
AN	OMALY/	REAL	QUAD	REAL	QUAD	REAL	QUAD	. COND I	DEPIH*	COND	DEPIH	RESIS	DEPTH	1.101
FID	/1NIERP	PPM	PPM	PPM	PPM	PPM	PPM	•SIEMEN	М	.SIEMEN	M	OHM-M	м	N.I.
LIN	E 30170	()	FLIGHI	12)				•		•				
Ε	1591S	ò	2	1	2	2	4		-			-	-	0
F	1598B?	3	9	6	12	45	41	. 1.3	19	. 1	47	288	7	0
G	1610H	2	4	1	4	16	17	. 2.1	49	. 1	62	143	24	0
H	1636H	12	3	4	54	28	204	. 35.1	46	. 1	30	58	5	0
LIN	E 30180	a	FLIGHI	12)				•		•				
A	1486B?	33	61	69	126	383	250	. 5.1	0	. 3	36	17	17	120
в	1485D	33	61	69	126	383	250	. 5.1	2	. 2	23	28	4	0
С	1482D	5	21	69	126	383	95	. 1.4	0	. 1	37	135	4	0
D	1471D	1	2	1	2	2	4		-			-	-	0
Έ	1468B	9	7	19	18	60	74	. 8.4	45	. 1	29	130	1	0
F	1463B?	8	7	19	18	159	166	. 6.4	36	. 1	26	97	0	0
G	1448B?	0	2	1	2	2	4	• -	-		-		-	0
H	1443B?	0	8	1	10	34	57	. 0.4	0	. 1	28	544	0	0
I	1434S	0	10	2	13	48	71	. 0.4	0	. 1	12	460	0	0
J	1422H	3	7	5	13	45	28	. 1.6	21	. 1	46	122	12	0
K	1393H	1	2	1	2	2	4	• -	-	• -	-	-	-	0
LIN	E 30190	()	FLIGHI	· 12)	}			•		•				
Α	1206B?	i	6	6	14	45	51	. 0.4	0	. 1	46	114	9	60
в	1228B	13	47	22	90	363	267	. 2.2	0	. 1	23	100	0	0
С	1229B	15	41	22	90	363	267	. 2.8	2	. 1	19	80	0	20
D	1234D	17	33	16	29	95	52	. 3.9	8	. 1	28	61	3	0
Е	1237B?	14	33	4	29	95	52	. 3.1	4	. 1	28	48	5	0
F	1243B?	22	15	35	69	253	172	. 12.9	17	. 1	20	47	0	50
G	1255S?	0	9	2	13	48	96	. 0.4	0	. 1	26	628	0	0
H	1258S?	0	11	2	13	48	96	. 0.4	0	. 1	20	587	0	0
I	1267S	0	20	6	30	137	127	. 0.4	0	. 1	16	211	0	0
ປ T	1284S	1	2	1	2	2	4	• -	-	• -	-	-	-	0
к ~	1314ff	Т	2	Ŧ	2	Z	4	• -	-	. –	-	-	-	U
LIN	E 30200	()	FLIGHI	12))			•		•				
A	1182B?	5	23	4	40	122	188	. 1.3	0	. 1	38	189	0	0
В	1179B?	5	23	6	40	122	188	. 1.3	5	. 1	22	217	0	30
С	1175B	26	26	89	103	239	104	. 9.2	13	. 3	30	19	11	0
D	1173B	26	26	89	103	239	37	. 9.2	16	. 3	36	20	16	0
Е	1169B?	15	23	58	78	222	116	. 4.5	13	. 1	28	59	3	0
F	1166B?	27	31	58	78	222	116	. 7.9	6	. 3	28	18	9	0
G	1161B?	31	32	58	54	231	331	. 9.3	20	. 2	32	32	12	0
H	1158D	31	47	60	85	274	331	. 6.0	3	• 3	21	16	4	0
I	1153D	30	26	60	85	274	98	. 11.0	15	• 4	40	11	23	0
	.* ES	TIMA	TED DE	ртн и	MAY BI	e unri	ELTABI	E BECAUS	SE THE	STRONG	FR PAI	rr .		
	• OF	THE	CONDU	CIOR	MAY	BE DE	EPER C	RTOON	E SIDE	OF THE	FLIG	ET .		
	. LE	NE, (OR BEC	AUSE	OF A	SHALL	LOW DI	P OR OV	ERBURD	EN EFFE	CTS.			

. LINE, OR BECAUSE OF A SHALLOW DIP OR OVERBURDEN EFFECTS.

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COAXIAL 1126 HZ		COPI 87	ANAR 4 HZ	COPLANAR 7219 HZ		. VERTICAL . . DIKE .		. HORIZONTAL . SHEET		CONDUCTIVE EARTH		MAG CORR		
AN	OMALY/	REAL	QUAD	REAL	QUAD	REAL	QUAD	. COND E)EPIH*	. COND D	EPIH	RESIS	DEPTH	
FID	/INTERP	PPM	PPM	PPM	PPM	PPM	PPM	.SIEMEN	М	SIEMEN	M	OHMM	М	NT
								•		•				
	E 30200	()	FLIGHI	: 12)			26	•	40	•	~ ~ ~	216	~	•
J V	112262	1 2	2	4 5	17	34	30	· 1.3	48	• I	34 34	210	0	0
T.	11169?	د ۱	14	2	20	71	134	. 1.0	- 0	• 1	24 14	482	0	0
м	1100H	1	2	1	20	2	154	-	~	• -			-	0 0
N	1070H	1	1	1	2	2	4	. –	-	. –	_	-	-	õ
								•		•				
LIN	E 30210	(1	FLIGHI	. 12)	I			•		•				
A	913B	18	28	32	64	192	68	. 4.9	2	. 2	28	37	4	0
в	915B	19	28	32	64	192	68	. 5.1	2	. 2	34	37	9	9
С	924B?	16	10	34	28	44	37	. 12.6	21	. 3	42	18	20	80
D	948B?	1	2	1	2	2	4	• -	-	• -	-	_	-	0
E	977S?	0	17	5	17	126	89	. 0.4	0	. 1	4	399	0	0
F	986H	5	18	10	28	101	76	. 1.6	5	. 1	24	106	0	0
6 	1012H	L.	2	T	2	2	4	• -	-	• -	-	-	***	0
LIN	E 30220	Ċ	FLIGHT	· 12)				•		•				
A	803D	32	63	73	133	396	227	. 4.7	0	. 2	24	21	6	30
В	792D	6	16	7	8	34	49	. 2.0	13	. 1	53	69	22	50
с	771D	7	11	8	16	19	31	. 3.3	23	. 1	53	67	22	0
D	749B?	4	11	4	13	55	68	. 1.7	17	. 1	43	159	8	120
Ε	737B?	2	7	4	8	51	25	. 0.8	4	. 1	20	596	0	0
F	705H	3	7	5	12	55	24	. 1.9	22	. 1	36	153	2	0
G	684H	2	27	9	43	146	207	. 0.5	0	. 1	20	90	0	0
								•		•				
<u>ر</u> ۱۳۳۸	405D3	' (. 1	പപ്പ	: 12) 1	`	2	,	•		•				•
A P	42001	E E	2	15	2	2	22	· -	10	• -	- 56	-		0
с С	434D 445B?	11	10	14	23	23 66	3Z 47	· 2.9	22	• 4	46	24 17	18	40
л П	460B?	1	2	1	23	2	47					-	-	60
Ē	4745?	ō	12	5	19	26	21	. 0.4	0	. 1	16	263	0	0
F	507S?	3	14	4	22	72	94	. 0.8	Ō	. 1	6	371	Ō	Ō
G	526B?	12	21	11	51	160	214	. 4.0	17	. 1	26	72	2	0
H	533B?	1	2	1	2	2	4		-	. –		-	-	0
								•		•				
LIN	E 30240	()	FLIGHI	12)	• • •		_	•	_	•				_
A	371D	13	19	3	32	67	2	. 4.9	6	. 1	42	55	13	0
В	3678?	1	2	1	2	2	4	• • •	_	• -		_	-	7
C N	302B1	LL ,	5	21 1	29	5	58 •	. 13.5	30	• 3	ρŢ	22	36	U
ע די	20152	1	17	1	2	2 0 A	107	• -	-	• •	-	202	~	0
ы С	3240: วอาบ	1 7	۲۱ ۲۱	0	20 71	04	<u>та</u> /	• U.4	C 2	• ⊥	0 2	200	0	100
Ľ	2720	2	5	U	эт	24	50	• 2.0	43	• ⊥	2	773	υ	120
	≥∃ ★.	ימאדידי	ארו כדאיד	елн и	IAY BI	รบงคว	TABL	E BETAILS	E THE	STRONGE		RT .		
	. OF	THE	CONDI	ICTOR	MAY	3E DFI	EPER O	R TO ONF	STDE	OF THE	FLIC	HT .		
	. LI	NE, (OR BEC	AUSE	OF A	SHALI	LOW DI	POROVE	RBURD	EN EFFEC	TS.	•		

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		CO2 112	AXIAL 26 HZ	COPI 87	ANAR 74 HZ	COPI 723	ANAR 19 HZ	. VERT	ICAL KE	. HORIZ	ZONTAL EET	CONDUC	CTIVE IH	MAG CORR
AN FID	OMALY/ /INTERP	REAL PPM	QUAD PPM	REAL PPM	QUAD PPM	REAL PPM	QUAD PPM	. COND : .SIFMEN	DEPIH* M	COND SIEME	DEPTH I M	RESIS OHMM	DEPIH M	NT
LIN	E 30250) (1	FLIGHI	12)	I			•		•				
Α	151B?	' i	2	1	2	2	4	. –			-	-	-	0
в	153B?	' 1	2	1	2	2	4		-		-	-	-	0
С	159H	3	8	9	13	50	49	. 1.8	26	. 1	52	69	22	210
D	176S?	4	16	9	24	105	138	. 1.5	11	. 1	26	146	0	0
E	182S?	1	2	1	2	2	4		-		-	-	-	0
F	187S?	1	9	3	9	6	37	. 0.5	5	. 1	8	419	0	0
G	2035?	0	10	0	7	72	52	. 0.4	1	. 1	15	472	0	0
H	221S	3	13	2	20	17	53	. 1.1	1	. 1	6	439	0	0
LIN	E 30260) (1	FLIGHI	r 11)	I			•		•				
A	1630H	Ó	5	2	8	24	34	. 0.4	0	. 1	58	183	17	0
В	1664S	0	4	1	5	3	7	. 0.4	0	. 1	45	663	0	0
L.TN	E 30270	1 /1	ल . रत्यमा	י דר י	1			•		•				
Δ	1540H	4	5	3	11	5	28	• . 3.2	32	. 1	61	288	13	20
В	15005	2	5	2	7	29	26	. 2.0	26	. 1	36	496	0	0
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ี มีนี้ย	E 30280) (I	FLICHI	' 11)	-	~		•		•				•
A	4055	Ť	2	T	2	2	4	• -	-	• -	-		-	0
в —	4265	. 2	5	4	9	24	.5	• 1.6	20	• 1	50	229	6	U
LIN	E 30290) (1	FLIGH	. 11)	l			•						
A	584S	1	6	3	10	35	27	. 0.4	0	. 1	55	228	8	0
В	542S	1	5	2	8	4	6	. 1.0	16	. 1	75	180	29	0
 T.TN				121	1			•		•				
Δ	47835	· (1 7	A	<u> </u>	7	12	1	• 19	23	•	71	176	24	0
B	47675	0		3	5	15	11	. 0.4	23	• 1	75	210	27	0
č	4734S	ĩ	3	5	5	7	16	. 0.4	Ő	. 1	60	117	39	0
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	E 39010	l) (]	ытені	' 12)		~		•		•				~
A D	4588H	15	2	10	0	166	4	• •		• -	-	-	-	0
D 	4499n 	. 15	20	10	43	100	104	• 4•0	2	• • •	32	24	o	U
LIN	LINE 39020 (FLIGHT		[12]	ł			•		•					
Α	4931B?	' 3	8	5	13	46	31	. 1.5	26	. 1	75	179	31	0
в	4945B	14	10	25	11	42	67	. 11.0	24	. 2	58	37	30	0
С	4960B?	' 1	2	1	2	2	4	. –			-	-	-	0
D	4961B?	6	9	8	11	29	29	. 3.2	26	. 2	55	31	29	50
\mathbf{E}	4966B?	5	5	0	4	16	43	. 5.7	40	. 2	54	41	26	70
F	4969B?	' 1	2	1	2	2	4		-		-	-	-	0
	• • •	יאורדווו	ירו רובונו	ע נאונוי	יס ערא	יכוואד ק	ינוגיד דק	יזגריזם ש		CITICAT	יגרד בודשי	•		
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H	4977B	17	39	44	56	168	122	•	3.4	+ 0	•	2	29	25	8	0
I	4981B	20	27	44	65	229	42		5.7	77		2	41	23	19	0
J	4984D	24	44	44	65	229	133		4.7	' 0		2	27	39	4	40
К	4993D	13	43	23	78	245	190		2.3	30		1	29	92	2	0
\mathbf{L}	4996D	9	43	23	78	245	190		1.6	50		1	27	86	1	30
М	5038B?	? 10	14	8	18	26	51	•	4.3	14	•	1	44	217	3	4
N	5048H	5	4	5	4	17	85	٠	5.4	56	•	1	55	169	18	0
0	5081S?	2 1	12	0	18	15	142	•	0.4	11	•	1	34	561	0	6

 $\boldsymbol{.*}$ ESTIMATED DEPTH MAY BE UNRELIABLE BECAUSE THE STRONGER PART .

. OF THE CONDUCTOR MAY BE DEEPER OR TO ONE SIDE OF THE FLIGHT .

. LINE, OR BECAUSE OF A SHALLOW DIP OR OVERBURDEN EFFECTS. .





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