ARIS SUMMARY SHEET

Off Confidential: 94.11.08 District Geologist, Prince George ASSESSMENT REPORT 23099 MINING DIVISION: Omineca **PROPERTY:** Yellow Moose 125 05 00 LOCATION: LAT 53 30 00 LONG 10 5929696 361811 UTM 093F06E 093G11E NTS Yel 1-9 CLAIM(S): OPERATOR(S): Cogema Resources Inc. Pritchard, R.; Fraser, D. AUTHOR(S): 1993, 84 Pages **REPORT YEAR:** COMMODITIES SEARCHED FOR: Gold Cretaceous - Oligocene, Volcanics, Alteration **KEYWORDS:** WORK Geophysical DONE: 360.0 km;VLF EMAB Map(s) - 5; Scale(s) - 1:20 000360.0 km MAGA Map(s) - 1; Scale(s) - 1:20 000

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DIGHEM^V SURVEY FOR COGEMA RESOURCES INC. YELLOW MOOSE PROPERTY BRITISH COLUMBIA

NTS 93F/6,11

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Dighem Surveys & Processing Inc. Mississauga, Ontario April 26, 1993 Ruth A. Pritchard Geophysicist

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PROPERTY NAME(S) Yellow Moose COMMODITIES PRESENT B.C. MINERAL INVENTORY NUMBER(S), IF KNOWN	
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SUMMARY

This report describes the logistics and results of a DIGHEM^V airborne geophysical survey carried out for Cogema Resources Inc. over the Yellow Moose property, British Columbia. Total coverage of the survey block amounted to 377 km, including tie lines. The survey was flown from March 8 to March 12, 1993.

The purpose of the survey was to detect zones of conductive mineralization and to provide information that could be used to map the geology and structure of the survey area. This was accomplished by using a DIGHEM^V multi-coil, multi-frequency electromagnetic system, supplemented by a high sensitivity Cesium magnetometer and a two-channel VLF receiver. The information from these sensors was processed to produce maps which display the magnetic and conductive properties of the survey area. An electronic navigation system, operating in the UHF band, ensured accurate positioning of the geophysical data with respect to the base map. Visual flight path recovery techniques were used in areas where transponder signals were blocked by topographic features.

The survey property contains several anomalous features, some of which may be of moderate to high priority as exploration targets. Most of the inferred bedrock conductors appear to warrant further investigation using appropriate surface exploration techniques. Areas of interest may be assigned priorities on the basis of supporting geophysical, geochemical and/or geological information. After initial investigations have

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INTRODUCTION

A DIGHEM^V electromagnetic/resistivity/magnetic/VLF survey was flown for Cogema Resources Inc. from March 8 to March 12, 1993, over the Yellow Moose Property, British Columbia. The survey area can be located on NTS map sheets 93F/6,11 (see Figure 1).

The Yellow Moose property consists of 9 contiguous claims (146 units, 36.5 km²), located by COGEMA in the fall of 1992; the claims are listed on Table 1 and shown on Figure 2.

Table 1-1

List of Claims

NAME	RECORD NO	UNITS	STAKING DATE	EXPIRY DATE
YEL 1	314661	20	92/11/11	93/11/11
YEL 2	314662	20	92/11/11	93/11/11
YEL 3	314663	20	92/11/11	93/11/11
YEL 4	314664	18	92/11/11	93/11/11
YEL 5	314665	4	92/11/10	93/11/10
YEL 6	314666	16	92/11/10	93/11/10
YEL 7	314667	16	92/11/11	93/11/11
YEL 8	314668	16	92/11/10	93/11/10
YEL 9	314669	16	92/11/10	93/11/10
		1.1.5		

TOTAL 146

Survey coverage consisted of approximately 377 line-km, including tie lines. Flight lines were flown in an azimuthal direction of $0^{\circ}/180^{\circ}$ with a line separation of 200 m.

The survey employed the DIGHEM^V electromagnetic system. Ancillary equipment consisted of a magnetometer, radar altimeter, video camera, analog and digital recorders, a VLF receiver and an electronic navigation system. Details on the survey equipment are given in Section 2.

The instrumentation was installed in an Aerospatiale AS350B turbine helicopter (Registration C-GJIX) which was provided by Questral Helicopters Ltd. The helicopter flew at an average airspeed of 130 km/h with an EM bird height of approximately 30 m.

Section 2 also provides details on the data channels, their respective sensitivities, and the navigation/flight path recovery procedure. Noise levels of less than 2 ppm are generally maintained for wind speeds up to 35 km/h. Higher winds may cause the system to be grounded because excessive bird swinging produces difficulties in flying the helicopter. The swinging results from the 5 m² of area which is presented by the bird to broadside gusts.

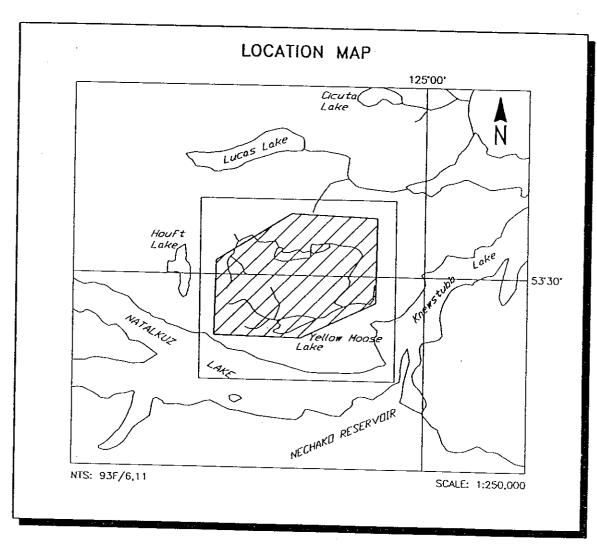


FIGURE 1 COGEMA RESOURCES INC. YELLOW MOOSE PROPERTY - 1138

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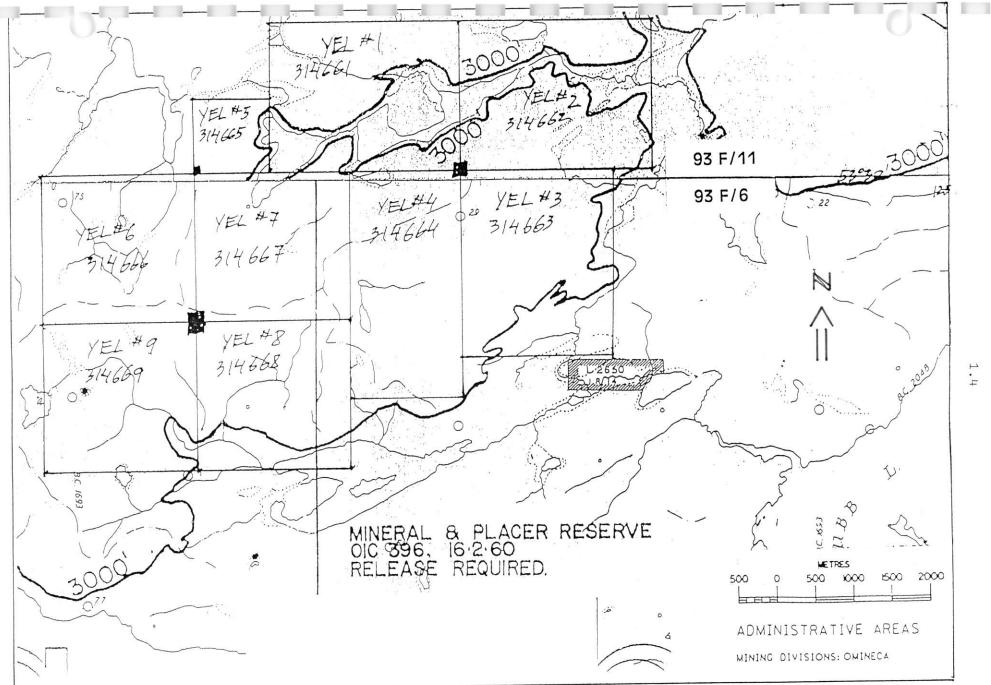


Figure 2. Yellow Moose Property Claim Map.

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This section provides a brief description of the geophysical instruments used to

acquire the survey data:

Electromagnetic System

Model: DIGHEM^V

Type: Towed bird, symmetric dipole configuration operated at a nominal survey altitude of 30 metres. Coil separation is 8 metres for 900 Hz and 7200 Hz, and 6.3 metres for the 56,000 Hz coil-pair.

Coil orientations/frequencies:	coaxial/900 Hzcoplanar/900 Hzcoaxial/7,200 Hzcoplanar/7,200 Hzcoplanar/56,000 Hz
Channels recorded:	5 inphase channels5 quadrature channels2 monitor channels
Sensitivity:	0.1 ppm at900 Hz0.2 ppm at7,200 Hz0.5 ppm at56,000 Hz
Sample rate:	10 per second

The electromagnetic system utilizes a multi-coil coaxial/coplanar technique to energize conductors in different directions. The coaxial coils are vertical with their axes

in the flight direction. The coplanar coils are horizontal. The secondary fields are sensed simultaneously by means of receiver coils which are maximum coupled to their respective transmitter coils. The system yields an inphase and a quadrature channel from each transmitter-receiver coil-pair.

Magnetometer

Model:	Picodas 3340
Туре:	Optically pumped Cesium vapour
Sensitivity:	0.01 nT
Sample rate:	10 per second

The magnetometer sensor is towed in a bird 20 m below the helicopter.

Magnetic Base Station

Model: Scintrex MP-3

Type: Digital recording proton precession

Sensitivity: 0.10 nT

Sample rate: 0.2 per second

A digital recorder is operated in conjunction with the base station magnetometer to record the diurnal variations of the earth's magnetic field. The clock of the base station is synchronized with that of the airborne system to permit subsequent removal of diurnal drift.

VLF System

Manufacturer:	Herz Industries Ltd.		
Туре:	Totem-2A		
Sensitivity:	0.1%		
Stations:	Seattle, Washington; Cutler, Maine; Lualualei, Hawaii;	NAA,	24.8 kHz 24.0 kHz 23.4 kHz

The VLF receiver measures the total field and vertical quadrature components of the secondary VLF field. Signals from two separate transmitters can be measured simultaneously. The VLF sensor is housed in the same bird as the magnetic sensor, and is towed 20 m below the helicopter.

Radar Altimeter

Manufacturer:	Honeywell/Sperry		
Туре:	AA 220		
Sensitivity:	1 ft		

The radar altimeter measures the vertical distance between the helicopter and the ground. This information is used in the processing algorithm which determines conductor depth.

Analog Recorder

Manufacturer:	RMS Instruments
Туре:	DGR33 dot-matrix graphics recorder
Resolution:	4x4 dots/mm
Speed:	1.5 mm/sec

The analog profiles are recorded on chart paper in the aircraft during the survey. Table 2-1 lists the geophysical data channels and the vertical scale of each profile.

Channel	Parameter	Scale	Designation on
Name		units/mm	digital profile
1X9I 1X9Q 3P9I 3P9Q 2P7I 2P7Q 4X7I 4X7Q 5P5I 5P5Q ALIR CMGC CMGF VF1T VF1Q VF2T VF2Q CXSP CPSP CXPL CPPL	coaxial inphase (900 Hz) coaxial quad (900 Hz) coplanar inphase (900 Hz) coplanar quad (900 Hz) coplanar quad (700 Hz) coplanar quad (7200 Hz) coaxial inphase (7200 Hz) coaxial quad (7200 Hz) coaxial quad (7200 Hz) coplanar inphase (56000 Hz) coplanar quad (56000 Hz) altimeter magnetics, coarse magnetics, fine VLF-total: primary stn. VLF-quad: primary stn. VLF-quad: primary stn. VLF-quad: secondary stn. Coaxial spherics monitor coplanar spherics monitor coplanar powerline monitor	2.5 ppm 2.5 ppm 2.5 ppm 5 ppm 5 ppm 5 ppm 10 ppm 10 ppm 10 ppm 20 nT 2.0 nT 2% 2% 2% 2%	CXI (900 Hz) CXQ (900 Hz) CPI (900 Hz) CPI (900 Hz) CPQ (900 Hz) CPQ (7200 Hz) CXI (7200 Hz) CXI (7200 Hz) CXQ (7200 Hz) CPI (56 kHz) CPQ (56 kHz) ALT MAG VLF (primary total field) VLF (primary quadrature) VLF (secondary total field) VLF (secondary total field) VLF (secondary quadrature) CXS CPS CXP CPP

 Table 2-1.
 The Analog Profiles

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Table 2-2.	The	Digital	Profiles
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		······
Channel		Scale
Name (Freq)	Observed parameters	<u>units/mm</u>
MAG ALT CXI (900 Hz) CXQ (900 Hz) CPI (900 Hz) CPQ (900 Hz)	magnetics bird height vertical coaxial coil-pair inphase vertical coaxial coil-pair quadrature horizontal coplanar coil-pair inphase horizontal coplanar coil-pair quadrature vertical coaxial coil-pair inphase vertical coaxial coil-pair inphase	20 nT 6 m 2 ppm 2 ppm 2 ppm 2 ppm 4 ppm 4 ppm 4 ppm 4 ppm 10 ppm 10 ppm 5% 5% 5%
	Computed Parameters	
DFQ (900 Hz) RES (900 Hz)	log resistivity log resistivity log resistivity apparent depth	200 nT 2 ppm 2 ppm .06 decade .06 decade .06 decade 6 m 6 m 1 grade

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Digital Data Acquisition System

Manufacturer:	RMS Instruments
Туре:	DGR 33
Tape Deck:	RMS TCR-12, 6400 bpi, tape cartridge recorder

The digital data are used to generate several computed parameters. Both measured and computed parameters are plotted as "multi-channel stacked profiles" during data processing. These parameters are shown in Table 2-2. In Table 2-2, the log resistivity scale of 0.06 decade/mm means that the resistivity changes by an order of magnitude in 16.6 mm. The resistivities at 0, 33 and 67 mm up from the bottom of the digital profile are respectively 1, 100 and 10,000 ohm-m.

Tracking Camera

Type: Panasonic Video

Model: AG 2400/WVCD132

Fiducial numbers are recorded continuously and are displayed on the margin of each image. This procedure ensures accurate correlation of analog and digital data with respect to visible features on the ground.

Navigation System

Model:Del Norte 547Type:UHF electronic positioning systemSensitivity:1 mSample rate:2 per second

The navigation system uses ground based transponder stations which transmit distance information back to the helicopter. The ground stations are set up well away from the survey area and are positioned such that the signals cross the survey block at an angle between 30° and 150°. The onboard Central Processing Unit takes any two transponder distances and determines the helicopter position relative to these two ground stations in cartesian coordinates.

The cartesian coordinates are transformed to UTM coordinates during data processing. This is accomplished by correlating a number of prominent topographical locations with the navigational data points. The use of numerous visual tie points serves two purposes: to accurately relate the navigation data to the map sheet and to minimize location errors which might result from distortions in uncontrolled photomosaic base maps.

Field Workstation

Manufacturer:	Dighem
Model:	FWS: V2.41
Туре:	80386 based P.C

A portable PC-based field workstation is used at the survey base to verify data quality and completeness. Flight tapes are dumped to a hard drive to permit the creation of a database. This process allows the field operators to display both the positional (flight path) and geophysical data on a screen or printer.

Global Positioning System

Manufacturer:	Trimble Navigation Ltd.
Туре:	Pathfinder (C/A code, dual channel)
Accuracy:	25 metres (5 metres in differential mode)
Update:	Once per second

The Pathfinder system uses signals broadcast by the NAVSTAR GPS satellites to provide positional readouts in latitude/longitude or UTM coordinates. The GPS unit is placed at each UHF transponder site to determine its exact location. The system can also

be used aboard the helicopter to provide real-time navigation guidance. Recorded data can be downloaded to a field computer for immediate post-survey processing, or transmitted to the central processing facility for final plotting. The following products are available from the survey data. Those which are not part of the survey contract may be acquired later. Most parameters can be displayed as contours, profiles, or in colour.

Base Maps

A base map of the survey area hav been produced from published topographic maps. These provide a relatively accurate, distortion-free base which facilitates correlation of the navigation data to the UTM grid. Photomosaics are useful for visual reference and for subsequent flight path recovery, but usually contain scale distortions. Orthophotos are ideal, but their cost and the time required to produce them, usually precludes their use as base maps.

Electromagnetic Anomalies

Anomalous electromagnetic responses are selected and analysed by computer to provide a preliminary electromagnetic anomaly map. This preliminary map is used, by the geophysicist, in conjunction with the computer-generated digital profiles, to produce the final interpreted EM anomaly map. This map includes bedrock surficial and cultural conductors. A map containing only bedrock conductors can be generated, if desired.

Resistivity

The apparent resistivity in ohm-m may be generated from the inphase and quadrature EM components for any of the frequencies, using a pseudo-layer halfspace model. A resistivity map portrays all the EM information for that frequency over the entire survey area. This contrasts with the electromagnetic anomaly map which provides information only over interpreted conductors. The large dynamic range makes the resistivity parameter an excellent mapping tool.

EM Magnetite

The apparent percent magnetite by weight is computed wherever magnetite produces a negative inphase EM response.

Total Field Magnetics

The aeromagnetic data are corrected for diurnal variation using the magnetic base station data. The regional IGRF can be removed from the data, if requested.

Enhanced Magnetics

The total field magnetic data are subjected to a processing algorithm. This algorithm enhances the response of magnetic bodies in the upper 500 m and attenuates the response of deeper bodies. The resulting enhanced magnetic map provides better definition and resolution of near-surface magnetic units. It also identifies weak magnetic features which may not be evident on the total field magnetic map. However, regional magnetic variations, and magnetic lows caused by remanence, are better defined on the total field magnetic map. The technique is described in more detail in Section 5.

Magnetic Derivatives

The total field magnetic data may be subjected to a variety of filtering techniques to yield maps of the following:

vertical gradient second vertical derivative magnetic susceptibility with reduction to the pole upward/downward continuations All of these filtering techniques improve the recognition of near-surface magnetic bodies, with the exception of upward continuation. Any of these parameters can be produced on request. Dighem's proprietary enhanced magnetic technique is designed to provide a general "all-purpose" map, combining the more useful features of the above parameters.

VLF

The VLF data are digitally filtered to remove long wavelengths such as those caused by variations in the transmitted field strength.

Multi-channel Stacked Profiles

Distance-based profiles of the digitally recorded geophysical data are generated and plotted by computer. These profiles also contain the calculated parameters which are used in the interpretation process. These are produced as worksheets prior to interpretation, and can also be presented in the final corrected form after interpretation. The profiles display electromagnetic anomalies with their respective interpretive symbols. The differences between the worksheets and the final corrected form occur only with respect to the EM anomaly identifier.

Contour, Colour and Shadow Map Displays

The geophysical data are interpolated onto a regular grid using a modified Akima spline technique. The resulting grid is suitable for generating contour maps of excellent quality.

Colour maps are produced by interpolating the grid down to the pixel size. The parameter is then incremented with respect to specific amplitude ranges to provide colour "contour" maps. Colour maps of the total magnetic field are particularly useful in defining the lithology of the survey area.

Monochromatic shadow maps are generated by employing an artificial sun to cast shadows on a surface defined by the geophysical grid. There are many variations in the shadowing technique. These techniques may be applied to total field or enhanced magnetic data, magnetic derivatives, VLF, resistivity, etc. Of the various magnetic products, the shadow of the enhanced magnetic parameter is particularly suited for defining geological structures with crisper images and improved resolution. - 3.6 -

Resistivity Sengpiel Sections

The apparent resistivity and approximate thickness of two or more horizontal layers can be displayed simultaneously for all coplanar frequencies. An inversion algorithm has been developed by Dr. K.P. Sengpiel[•] of the B.G.R., which determines the generalized skin depth, or "centroid depth" of the inphase current concentration, as a function of frequency. The centroid depth is combined with the apparent resistivity over a broad frequency range to produce resistivity-thickness pseudo-sections. A coloured presentation yields a smoothed representation of the true resistivity-depth distribution within the limits of the model used.

^{*} Approximate Inversion of Airborne EM Data from Multilayered Ground: Sengpiel, K.P., Geophysical Prospecting 36, 446-459, 1988.

SURVEY RESULTS

GENERAL DISCUSSION

The survey results are presented on 1 map sheet for each parameter at a scale of 1:20,000. Table 4-1 summarizes the EM responses in the survey area, with respect to conductance grade and interpretation.

The anomalies shown on the electromagnetic anomaly map are based on a nearvertical, half plane model. This model best reflects "discrete" bedrock conductors. Wide bedrock conductors or flat-lying conductive units, whether from surficial or bedrock sources, may give rise to very broad anomalous responses on the EM profiles. These may not appear on the electromagnetic anomaly map if they have a regional character rather than a locally anomalous character. These broad conductors, which more closely approximate a half space model, will be maximum coupled to the horizontal (coplanar) coil-pair and should be more evident on the resistivity parameter. Resistivity maps, therefore, may be more valuable than the electromagnetic anomaly maps, in areas where broad or flat-lying conductors are considered to be of importance. Contoured resistivity maps, based on the 900 Hz, 7200 Hz and 56,000 Hz coplanar data are included with this report.

TABLE 4-1

EM ANOMALY STATISTICS

YELLOW MOOSE PROPERTY

CONDUCTOR GRADE	CONDUCTANCE RANGE SIEMENS (MHOS)	NUMBER OF RESPONSES
7	>100	0 3
6	50 - 100	
5	20 - 50	4
4	10 - 20	23
3	5 - 10	62
2	1 - 5	195
1	<1	80
*	INDETERMINATE	159
TOTAL		526

CONDUCTOR MODEL	MOST LIKELY SOURCE	NUMBER OF RESPONSES
В	DISCRETE BEDROCK CONDUCTOR	269
S	CONDUCTIVE COVER	96
Н	ROCK UNIT OR THICK COVER	146
Ε	EDGE OF WIDE CONDUCTOR	15
TOTAL		526

(SEE EM MAP LEGEND FOR EXPLANATIONS)

Excellent resolution and discrimination of conductors was accomplished by using a fast sampling rate of 0.1 sec and by employing common frequencies (900 Hz and 7200 Hz) on two orthogonal coil-pairs (coaxial and coplanar). The resulting "difference channel" parameters often permit differentiation of bedrock and surficial conductors, even though they may exhibit similar conductance values.

Anomalies which occur near the ends of the survey lines (i.e., outside the survey area), should be viewed with caution. Some of the weaker anomalies could be due to aerodynamic noise, i.e., bird bending, which is created by abnormal stresses to which the bird is subjected during the climb and turn of the aircraft between lines. Such aerodynamic noise is usually manifested by an anomaly on the coaxial inphase channel only, although severe stresses can affect the coplanar inphase channels as well.

Magnetics

A Scintrex proton precession magnetometer was operated at the survey base to record diurnal variations of the earth's magnetic field. The clock of the base station was synchronized with that of the airborne system to permit subsequent removal of diurnal drift. The background magnetic level has been adjusted to match the International Geomagnetic Reference Field (IGRF) for the survey area. The IGRF gradient across the survey block is left intact.

The total field magnetic data have been presented as contours on the base map using a contour interval of 5 nT where gradients permit. The map shows the magnetic properties of the rock units underlying the survey area.

The total field magnetic data have been subjected to a processing algorithm to produce an enhanced magnetic map. This procedure enhances near-surface magnetic units and suppresses regional gradients. It also provides better definition and resolution of magnetic units and displays weak magnetic features which may not be clearly evident on the total field map. A map of the first or second vertical magnetic derivative can also be prepared from existing survey data, if requested.

There is some evidence on the magnetic map which suggests that the survey area have been subjected to deformation and/or alteration. These structural complexities are evident on the contour map as variations in magnetic intensity, irregular patterns, and as offsets or changes in strike direction. Some of the more prominent linear features are also evident on the topographic base map. If a specific magnetic intensity can be assigned to the rock type which is believed to host the target mineralization, it may be possible to select areas of higher priority on the basis of the total field magnetic data. This is based on the assumption that the magnetite content of the host rocks will give rise to a limited range of contour values which will permit differentiation of various lithological units.

The magnetic results, in conjunction with the other geophysical parameters, should provide valuable information which can be used to effectively map the geology and structure in the survey area.

VLF

VLF results were obtained from the transmitting stations at Cutler, Maine (NAA - 24.0 kHz), Seattle, Washington (NLK - 24.8 kHz) and Lualualei, Hawaii (NPM - 23.4 kHz). Table 4-2 shows the VLF stations used for the contour map. Table 4-3 displays the VLF stations that were recorded as the primary and secondary stations and that are plotted as total field and quadrature on the digital profiles.

Table 4-2

AreaVLF Station(s)Yellow MooseHawaii (23.4 kHz - NPM) - all linesTable 4-3

Area	Primary Station (VLF1)	Secondary Station (VLF2)
Yellow Moose	Seattle (24.8 kHz - NLK)	Hawaii (23,4 kHz - NPM)

The VLF method is quite sensitive to the angle of coupling between the conductor and the propagated EM field. Consequently, conductors which strike towards the VLF station will usually yield a stronger response than conductors which are nearly orthogonal to it.

The VLF parameter does not normally provide the same degree of resolution available from the EM data. Closely-spaced conductors, conductors of short strike length or conductors which are poorly coupled to the VLF field, may escape detection with this method. Erratic signals from the VLF transmitters can also give rise to strong, isolated anomalies which should be viewed with caution. The filtered total field VLF contours are presented on the base map with a contour interval of one percent.

- 4.6 -

Resistivity

Resistivity maps, which display the conductive properties of the survey area, were produced from the 900 Hz, 7200 Hz and 56,000 Hz coplanar data. The maximum resistivity value, which is calculated for each frequency, is approximately 1.15 times the numerical value of the frequency. This cutoff eliminates the meaningless higher resistivities which would result from very small EM amplitudes. In general, the resistivity patterns show some agreement with the magnetic trends. This suggests that a few of the resistivity lows are probably related to bedrock features, rather than conductive overburden. There are some areas, however, where contour patterns appear to be strongly influenced by conductive surficial material.

Electromagnetics

The EM anomalies resulting from this survey appear to fall within one of two general categories. The first type consists of discrete, well-defined anomalies which yield marked inflections on the difference channels. These anomalies are usually attributed to conductive sulphides or graphite and are generally given a "B", "T" or "D" interpretive symbol, denoting a bedrock source.

The second class of anomalies comprises moderately broad responses which exhibit the characteristics of a half space and do not yield well-defined inflections on the difference channels. Anomalies in this category are usually given an "S" or "H" interpretive symbol. The lack of a difference channel response usually implies a broad or flat-lying conductive source such as overburden. Some of these anomalies may reflect conductive rock units or zones of deep weathering.

The effects of conductive overburden are evident over portions of the survey areas. Although the difference channels (DFI and DFQ) are extremely valuable in detecting bedrock conductors which are partially masked by conductive overburden, sharp undulations in the bedrock/overburden interface can yield anomalies in the difference channels which may be interpreted as possible bedrock conductors. Such anomalies usually fall into the "S?" or "B?" classification but may also be given an "E" interpretive symbol, denoting a resistivity contrast at the edge of a conductive unit.

In areas where EM responses are evident primarily on the quadrature components, zones of poor conductivity are indicated. Where these responses are coincident with magnetic anomalies, it is possible that the inphase component amplitudes have been suppressed by the effects of magnetite. Most of these poorly-conductive magnetic features give rise to resistivity anomalies which are only slightly below background. If it is expected that poorly-conductive economic mineralization may be associated with magnetite-rich units, most of these weakly anomalous features will be of interest. In areas where magnetite causes the inphase components to become negative, the apparent conductance and depth of EM anomalies may be unreliable.

It is difficult to assess the relative merits of EM anomalies on the basis of conductance. It is recommended that an attempt be made to compile a suite of geophysical "signatures" over areas of interest. Anomaly characteristics are clearly defined on the computer-processed geophysical data profiles which are supplied as one of the survey products.

A complete assessment and evaluation of the survey data should be carried out by one or more qualified professionals who have access to, and can provide a meaningful compilation of, all available geophysical, geological and geochemical data.

The electromagnetic anomaly map shows the anomaly locations with the interpreted conductor type, dip, conductance and depth being indicated by symbols. Direct magnetic correlation is also shown if it exists. The strike direction and length of the conductors are indicated when anomalies can be correlated from line to line. When studying the map sheets, consult the anomaly listings appended to this report.

YELLOW MOOSE PROPERTY

The magnetic contour patterns within this block are, for the most part, quite complex. Magnetic strike directions vary greatly throughout the area, and many magnetic features do not extend for more than one or two lines. Many features exhibit evidence of deformation which is seen on the map as truncations or offsets in the magnetic contours.

The highest magnetic values are located within several highly magnetic units situated at the north edge of the survey block, over lines 20250 through 20370. These magnetic features are very complex. The magnetic unit which extends over lines 10340 to 20370 is coincident with a resistivity high, whereas other magnetic highs in this zone are associated with low resistivities.

A broad, relatively uniform magnetic feature is situated in the southeast corner of the survey block. It also contains relatively high magnetic values of over 57,450 nT. A few bedrock anomalies are associated with this zone, although most seem to be located in the periphery of this magnetic feature.

The resistivity patterns display limited agreement with the magnetic trends. In many areas of the survey block the lowest resistivities are seen on the 56,000 Hz resistivity map, suggesting that their sources are surficial. The larger lakes within the

- 4.10 -

block give rise to well-defined resistivity lows on the 56,000 Hz resistivity. There are, however, several zones of conductivity which may have deeper, bedrock sources. The eastern portion of the survey block is much more conductive than the west. The limit of the conductive eastern zone is approximately defined by the 150 ohm-metre contour on the 900 Hz resistivity map. Most of the bedrock anomalies interpreted from the survey data are situated within the highly conductive eastern zone. There are also many anomalies which seem to be indicative of half-spaces at depth.

A group of possible bedrock responses is situated at the western edge of the conductive zone, over lines 20230 through 20320 in the vicinity of tie line 29020. These anomalies are moderately well-defined. They are situated in the vicinity of a possible structural break inferred from the magnetic data.

BACKGROUND INFORMATION

This section provides background information on parameters which are available from the survey data. Those which have not been supplied as survey products may be generated later from raw data on the digital archive tape.

ELECTROMAGNETICS

DIGHEM electromagnetic responses fall into two general classes, discrete and broad. The discrete class consists of sharp, well-defined anomalies from discrete conductors such as sulfide lenses and steeply dipping sheets of graphite and sulfides. The broad class consists of wide anomalies from conductors having a large horizontal surface such as flatly dipping graphite or sulfide sheets, saline water-saturated sedimentary formations, conductive overburden and rock, and geothermal zones. A vertical conductive slab with a width of 200 m would straddle these two classes.

The vertical sheet (half plane) is the most common model used for the analysis of discrete conductors. All anomalies plotted on the electromagnetic map are analyzed according to this model. The following section entitled **Discrete Conductor Analysis** describes this model in detail, including the effect of using it on anomalies caused by broad conductors such as conductive overburden.

The conductive earth (half space) model is suitable for broad conductors. Resistivity contour maps result from the use of this model. A later section entitled **Resistivity Mapping** describes the method further, including the effect of using it on anomalies caused by discrete conductors such as sulfide bodies.

Geometric interpretation

The geophysical interpreter attempts to determine the geometric shape and dip of the conductor. Figure 5-1 shows typical DIGHEM anomaly shapes which are used to guide the geometric interpretation.

Discrete conductor analysis

The EM anomalies appearing on the electromagnetic map are analyzed by computer to give the conductance (i.e., conductivity-thickness product) in Siemens (mhos) of a vertical sheet model. This is done regardless of the interpreted geometric shape of the conductor. This is not an unreasonable procedure, because the computed conductance increases as the electrical quality of the conductor increases, regardless of its true shape. DIGHEM anomalies are divided into seven grades of conductance, as shown in Table 5-1 below. The conductance in Siemens (mhos) is the reciprocal of resistance in ohms.

Conductor location	ţ	ŧ	ţ	ŧ	Å	ŧ	S,Н Е ↓ ↓	ł
Channel CXI	\wedge	\wedge	\bigwedge	\wedge		\sim		
Channel CPI	\sim	\sim	\sim		\square	$\int \! \left(\right)$		\frown
Channel DIFI	\mathcal{N}	\bigwedge	\mathcal{N}	\bigvee	\mathcal{N}	\mathcal{N}	\frown	
Conductor	line	vertical	dipping	vertical or	sphere;	wide	S = conductive overburden	flight line
Ratio of		thin dike	thin dike	dipping thick dike	horizontal disk; metal roof; small fenced yard	horizontal ribbon; large fenced area	 H = thick conductive cover or wide conductive rock unit E = edge effect from wide conductor 	parallel to conductor
amplitudes CX1 / CPI	4/1	2/1	variable	variable	1/4	variable	1/2	< 1/4

-5+3-

Fig. 5-1 Typical DIGHEM anomaly shapes

Table 5-1. EM Anomaly Grades

Anomaly Grade	<u>Siemens</u>
7	> 100
6	50 - 100
5	20 - 50
4	10 - 20
3	5 - 10
2	1 - 5
1	< 1
· · · · · · · · · · · · · · · · · · ·	

The conductance value is a geological parameter because it is a characteristic of the conductor alone. It generally is independent of frequency, flying height or depth of burial, apart from the averaging over a greater portion of the conductor as height increases. Small anomalies from deeply buried strong conductors are not confused with small anomalies from shallow weak conductors because the former will have larger conductance values.

Conductive overburden generally produces broad EM responses which may not be shown as anomalies on the EM maps. However, patchy conductive overburden in otherwise resistive areas can yield discrete anomalies with a conductance grade (cf. Table 5-1) of 1, 2 or even 3 for conducting clays which have resistivities as low as 50 ohm-m. In areas where ground resistivities are below 10 ohm-m, anomalies caused by weathering variations and similar causes can have any conductance grade. The anomaly shapes from the multiple coils often allow such conductors to be recognized, and these are indicated by the letters S, H, and sometimes E on the electromagnetic anomaly map (see EM map legend). show that the different grades indicate different types of conductors. The presence of sphalerite or gangue can result in ore deposits having weak to moderate conductances. As an example, the three million ton lead-zinc deposit of Restigouche Mining Corporation near Bathurst, Canada, yielded a well-defined grade 2 conductor. The 10 percent

inhibiting electrical conduction.

For bedrock conductors, the higher anomaly grades indicate increasingly higher conductances. Examples: DIGHEM's New Insco copper discovery (Noranda, Canada) yielded a grade 5 anomaly, as did the neighbouring copper-zinc Magusi River ore body; Mattabi (copper-zinc, Sturgeon Lake, Canada) and Whistle (nickel, Sudbury, Canada) gave grade 6; and DIGHEM's Montcalm nickel-copper discovery (Timmins, Canada) yielded a grade 7 anomaly. Graphite and sulfides can span all grades but, in any particular survey area, field work may

Strong conductors (i.e., grades 6 and 7) are characteristic of massive sulfides or graphite. Moderate conductors (grades 4 and 5) typically reflect graphite or sulfides of a less massive character, while weak bedrock conductors (grades 1 to 3) can signify poorly connected graphite or heavily disseminated sulfides. Grades 1 and 2 conductors may not respond to ground EM equipment using frequencies less than 2000 Hz.

Faults, fractures and shear zones may produce anomalies which typically have low conductances (e.g., grades 1 to 3). Conductive rock formations can yield anomalies of any

by volume of sphalerite occurs as a coating around the fine grained massive pyrite, thereby

conductance grade. The conductive materials in such rock formations can be salt water, weathered products such as clays, original depositional clays, and carbonaceous material.

On the interpreted electromagnetic map, a letter identifier and an interpretive symbol are plotted beside the EM grade symbol. The horizontal rows of dots, under the interpretive symbol, indicate the anomaly amplitude on the flight record. The vertical column of dots, under the anomaly letter, gives the estimated depth. In areas where anomalies are crowded, the letter identifiers, interpretive symbols and dots may be obliterated. The EM grade symbols, however, will always be discernible, and the obliterated information can be obtained from the anomaly listing appended to this report.

The purpose of indicating the anomaly amplitude by dots is to provide an estimate of the reliability of the conductance calculation. Thus, a conductance value obtained from a large ppm anomaly (3 or 4 dots) will tend to be accurate whereas one obtained from a small ppm anomaly (no dots) could be quite inaccurate. The absence of amplitude dots indicates that the anomaly from the coaxial coil-pair is 5 ppm or less on both the inphase and quadrature channels. Such small anomalies could reflect a weak conductor at the surface or a stronger conductor at depth. The conductance grade and depth estimate illustrates which of these possibilities fits the recorded) data best.

Flight line deviations occasionally yield cases where two anomalies, having similar conductance values but dramatically different depth estimates, occur close together on the same

conductor. Such examples illustrate the reliability of the conductance measurement while showing that the depth estimate can be unreliable. There are a number of factors which can produce an error in the depth estimate, including the averaging of topographic variations by the altimeter, overlying conductive overburden, and the location and attitude of the conductor relative to the flight line. Conductor location and attitude can provide an erroneous depth estimate because the stronger part of the conductor may be deeper or to one side of the flight line, or because it has a shallow dip. A heavy tree cover can also produce errors in depth estimates. This is because the depth estimate is computed as the distance of bird from conductor, minus the altimeter reading. The altimeter can lock onto the top of a dense forest canopy. This situation yields an erroneously large depth estimate but does not affect the conductance estimate.

Dip symbols are used to indicate the direction of dip of conductors. These symbols are used only when the anomaly shapes are unambiguous, which usually requires a fairly resistive environment.

A further interpretation is presented on the EM map by means of the line-to-line correlation of anomalies, which is based on a comparison of anomaly shapes on adjacent lines. This provides conductor axes which may define the geological structure over portions of the survey area. The absence of conductor axes in an area implies that anomalies could not be correlated from line to line with reasonable confidence.

DIGHEM electromagnetic maps are designed to provide a correct impression of conductor quality by means of the conductance grade symbols. The symbols can stand alone with geology when planning a follow-up program. The actual conductance values are printed in the attached anomaly list for those who wish quantitative data. The anomaly ppm and depth are indicated by inconspicuous dots which should not distract from the conductor patterns, while being helpful to those who wish this information. The map provides an interpretation of conductors in terms of length, strike and dip, geometric shape, conductance, depth, and thickness. The accuracy is comparable to an interpretation from a high quality ground EM survey having the same line spacing.

The attached EM anomaly list provides a tabulation of anomalies in ppm, conductance, and depth for the vertical sheet model. The EM anomaly list also shows the conductance and depth for a thin horizontal sheet (whole plane) model, but only the vertical sheet parameters appear on the EM map. The horizontal sheet model is suitable for a flatly dipping thin bedrock conductor such as a sulfide sheet having a thickness less than 10 m. The list also shows the resistivity and depth for a conductive earth (half space) model, which is suitable for thicker slabs such as thick conductive overburden. In the EM anomaly list, a depth value of zero for the conductive earth model, in an area of thick cover, warns that the anomaly may be caused by conductive overburden.

Since discrete bodies normally are the targets of EM surveys, local base (or zero) levels are used to compute local anomaly amplitudes. This contrasts with the use of true zero levels which are used to compute true EM amplitudes. Local anomaly amplitudes are shown in the EM anomaly list and these are used to compute the vertical sheet parameters of conductance and depth. Not shown in the EM anomaly list are the true amplitudes which are used to compute the horizontal sheet and conductive earth parameters.

Questionable Anomalies

DIGHEM maps may contain EM responses which are displayed as asterisks (*). These responses denote weak anomalies of indeterminate conductance, which may reflect one of the following: a weak conductor near the surface, a strong conductor at depth (e.g., 100 to 120 m below surface) or to one side of the flight line, or aerodynamic noise. Those responses that have the appearance of valid bedrock anomalies on the flight profiles are indicated by appropriate interpretive symbols (see EM map legend). The others probably do not warrant further investigation unless their locations are of considerable geological interest.

The thickness parameter

DIGHEM can provide an indication of the thickness of a steeply dipping conductor. The amplitude of the coplanar anomaly (e.g., CPI channel on the digital profile) increases relative to the coaxial anomaly (e.g., CXI) as the apparent thickness increases, i.e., the thickness in the horizontal plane. (The thickness is equal to the conductor width if the conductor dips at 90

degrees and strikes at right angles to the flight line.) This report refers to a conductor as <u>thin</u> when the thickness is likely to be less than 3 m, and <u>thick</u> when in excess of 10 m. Thick conductors are indicated on the EM map by parentheses "()". For base metal exploration in steeply dipping geology, thick conductors can be high priority targets because many massive sulfide ore bodies are thick, whereas non-economic bedrock conductors are often thin. The system cannot sense the thickness when the strike of the conductor is subparallel to the flight line, when the conductor has a shallow dip, when the anomaly amplitudes are small, or when the resistivity of the environment is below 100 ohm-m.

Resistivity mapping

Areas of widespread conductivity are commonly encountered during surveys. In such areas, anomalies can be generated by decreases of only 5 m in survey altitude as well as by increases in conductivity. The typical flight record in conductive areas is characterized by inphase and quadrature channels which are continuously active. Local EM peaks reflect either increases in conductivity of the earth or decreases in survey altitude. For such conductive areas, apparent resistivity profiles and contour maps are necessary for the correct interpretation of the airborne data. The advantage of the resistivity parameter is that anomalies caused by altitude changes are virtually eliminated, so the resistivity data reflect only those anomalies caused by conductivity changes. The resistivity analysis also helps the interpreter to differentiate between conductive trends in the bedrock and those patterns typical of conductive overburden. For example, discrete conductors will generally appear as narrow lows on the contour map and broad conductors (e.g., overburden) will appear as wide lows.

The resistivity profiles and the resistivity contour maps present the apparent resistivity using the so-called pseudo-layer (or buried) half space model defined by Fraser (1978)¹. This model consists of a resistive layer overlying a conductive half space. The depth channels give the apparent depth below surface of the conductive material. The apparent depth is simply the apparent thickness of the overlying resistive layer. The apparent depth (or thickness) parameter will be positive when the upper layer is more resistive than the underlying material, in which case the apparent depth may be quite close to the true depth.

The apparent depth will be negative when the upper layer is more conductive than the underlying material, and will be zero when a homogeneous half space exists. The apparent depth parameter must be interpreted cautiously because it will contain any errors which may exist in the measured altitude of the EM bird (e.g., as caused by a dense tree cover). The inputs to the resistivity algorithm are the inphase and quadrature components of the coplanar coil-pair. The outputs are the apparent resistivity of the conductive half space (the source) and the sensor-source distance. The flying height is not an input variable, and the output resistivity and sensor-source distance are independent of the flying height. The apparent depth, discussed above, is

¹ Resistivity mapping with an airborne multicoil electromagnetic system: Geophysics, v. 43, p.144-172

simply the sensor-source distance minus the measured altitude or flying height. Consequently, errors in the measured altitude will affect the apparent depth parameter but not the apparent resistivity parameter.

The apparent depth parameter is a useful indicator of simple layering in areas lacking a heavy tree cover. The DIGHEM system has been flown for purposes of permafrost mapping, where positive apparent depths were used as a measure of permafrost thickness. However, little quantitative use has been made of negative apparent depths because the absolute value of the negative depth is not a measure of the thickness of the conductive upper layer and, therefore, is not meaningful physically. Qualitatively, a negative apparent depth estimate usually shows that the EM anomaly is caused by conductive overburden. Consequently, the apparent depth channel can be of significant help in distinguishing between overburden and bedrock conductors.

The resistivity map often yields more useful information on conductivity distributions than the EM map. In comparing the EM and resistivity maps, keep in mind the following:

- (a) The resistivity map portrays the apparent value of the earth's resistivity, where resistivity = 1/conductivity.
- (b) The EM map portrays anomalies in the earth's resistivity. An anomaly by definition is a change from the norm and so the EM map displays anomalies, (i)

over narrow, conductive bodies and (ii) over the boundary zone between two wide formations of differing conductivity.

The resistivity map might be likened to a total field map and the EM map to a horizontal gradient in the direction of flight². Because gradient maps are usually more sensitive than total field maps, the EM map therefore is to be preferred in resistive areas. However, in conductive areas, the absolute character of the resistivity map usually causes it to be more useful than the EM map.

Interpretation in conductive environments

Environments having background resistivities below 30 ohm-m cause all airborne EM systems to yield very large responses from the conductive ground. This usually prohibits the recognition of discrete bedrock conductors. However, DIGHEM data processing techniques produce three parameters which contribute significantly to the recognition of bedrock conductors. These are the inphase and quadrature difference channels (DFI and DFQ), and the resistivity and depth channels (RES and DP) for each coplanar frequency.

² The gradient analogy is only valid with regard to the identification of anomalous locations.

The EM difference channels (DFI and DFQ) eliminate most of the responses from conductive ground, leaving responses from bedrock conductors, cultural features (e.g., telephone lines, fences, etc.) and edge effects. Edge effects often occur near the perimeter of broad conductive zones. This can be a source of geologic noise. While edge effects yield anomalies on the EM difference channels, they do not produce resistivity anomalies. Consequently, the resistivity channel aids in eliminating anomalies due to edge effects. On the other hand, resistivity anomalies will coincide with the most highly conductive sections of conductive ground, and this is another source of geologic noise. The recognition of a bedrock conductor in a conductive environment therefore is based on the anomalous responses of the two difference channels (DFI and DFQ) and the resistivity channels (RES). The most favourable situation is where anomalies coincide on all channels.

The DP channels, which give the apparent depth to the conductive material, also help to determine whether a conductive response arises from surficial material or from a conductive zone in the bedrock. When these channels ride above the zero level on the digital profiles (i.e., depth is negative), it implies that the EM and resistivity profiles are responding primarily to a conductive upper layer, i.e., conductive overburden. If the DP channels are below the zero level, it indicates that a resistive upper layer exists, and this usually implies the existence of a bedrock conductor. If the low frequency DP channel is below the zero level and the high frequency DP is above, this suggests that a bedrock conductor occurs beneath conductive cover.

The conductance channel CDT identifies discrete conductors which have been selected by computer for appraisal by the geophysicist. Some of these automatically selected anomalies on channel CDT are discarded by the geophysicist. The automatic selection algorithm is intentionally oversensitive to assure that no meaningful responses are missed. The interpreter then classifies the anomalies according to their source and eliminates those that are not substantiated by the data, such as those arising from geologic or aerodynamic noise.

Reduction of geologic noise

Geologic noise refers to unwanted geophysical responses. For purposes of airborne EM surveying, geologic noise refers to EM responses caused by conductive overburden and magnetic permeability. It was mentioned previously that the EM difference channels (i.e., channel DFI for inphase and DFQ for quadrature) tend to eliminate the response of conductive overburden. This marked a unique development in airborne EM technology, as DIGHEM is the only EM system which yields channels having an exceptionally high degree of immunity to conductive overburden.

Magnetite produces a form of geological noise on the inphase channels of all EM systems. Rocks containing less than 1% magnetite can yield negative inphase anomalies caused by magnetic permeability. When magnetite is widely distributed throughout a survey area, the inphase EM channels may continuously rise and fall, reflecting variations in the magnetite percentage, flying height, and overburden thickness. This can lead to difficulties in recognizing

deeply buried bedrock conductors, particularly if conductive overburden also exists. However, the response of broadly distributed magnetite generally vanishes on the inphase difference channel DFI. This feature can be a significant aid in the recognition of conductors which occur in rocks containing accessory magnetite.

EM magnetite mapping

The information content of DIGHEM data consists of a combination of conductive eddy current responses and magnetic permeability responses. The secondary field resulting from conductive eddy current flow is frequency-dependent and consists of both inphase and quadrature components, which are positive in sign. On the other hand, the secondary field resulting from magnetic permeability is independent of frequency and consists of only an inphase component which is negative in sign. When magnetic permeability manifests itself by decreasing the measured amount of positive inphase, its presence may be difficult to recognize. However, when it manifests itself by yielding a negative inphase anomaly (e.g., in the absence of eddy current flow), its presence is assured. In this latter case, the negative component can be used to estimate the percent magnetic content.

A magnetite mapping technique was developed for the coplanar coil-pair of DIGHEM. The technique yields a channel (designated FEO) which displays apparent weight percent magnetite according to a homogeneous half space model.³ The method can be complementary to magnetometer mapping in certain cases. Compared to magnetometry, it is far less sensitive but is more able to resolve closely spaced magnetite zones, as well as providing an estimate of the amount of magnetite in the rock. The method is sensitive to 1/4% magnetite by weight when the EM sensor is at a height of 30 m above a magnetitic half space. It can individually resolve steep dipping narrow magnetite-rich bands which are separated by 60 m. Unlike magnetometry, the EM magnetite method is unaffected by remanent magnetism or magnetic latitude.

The EM magnetite mapping technique provides estimates of magnetite content which are usually correct within a factor of 2 when the magnetite is fairly uniformly distributed. EM magnetite maps can be generated when magnetic permeability is evident as negative inphase responses on the data profiles.

Like magnetometry, the EM magnetite method maps only bedrock features, provided that the overburden is characterized by a general lack of magnetite. This contrasts with resistivity mapping which portrays the combined effect of bedrock and overburden.

³ Refer to Fraser, 1981, Magnetite mapping with a multi-coil airborne electromagnetic system: Geophysics, v. 46, p. 1579-1594.

Recognition of culture

Cultural responses include all EM anomalies caused by man-made metallic objects. Such anomalies may be caused by inductive coupling or current gathering. The concern of the interpreter is to recognize when an EM response is due to culture. Points of consideration used by the interpreter, when coaxial and coplanar coil-pairs are operated at a common frequency, are as follows:

- 1. Channels CXP and CPP monitor 60 Hz radiation. An anomaly on these channels shows that the conductor is radiating power. Such an indication is normally a guarantee that the conductor is cultural. However, care must be taken to ensure that the conductor is not a geologic body which strikes across a power line, carrying leakage currents.
- 2. A flight which crosses a "line" (e.g., fence, telephone line, etc.) yields a center-peaked coaxial anomaly and an m-shaped coplanar anomaly.⁴ When the flight crosses the cultural line at a high angle of intersection, the amplitude ratio of coaxial/coplanar response is 4. Such an EM anomaly can only be caused by a line. The geologic body which yields anomalies most closely resembling a line is the vertically dipping thin dike. Such a body, however, yields an amplitude ratio of 2 rather than 4. Consequently, an

⁴ See Figure 5-1 presented earlier.

m-shaped coplanar anomaly with a CXI/CPI amplitude ratio of 4 is virtually a guarantee that the source is a cultural line.

- 3. A flight which crosses a sphere or horizontal disk yields center-peaked coaxial and coplanar anomalies with a CXI/CPI amplitude ratio (i.e., coaxial/coplanar) of 1/4. In the absence of geologic bodies of this geometry, the most likely conductor is a metal roof or small fenced yard.⁵ Anomalies of this type are virtually certain to be cultural if they occur in an area of culture.
- 4. A flight which crosses a horizontal rectangular body or wide ribbon yields an m-shaped coaxial anomaly and a center-peaked coplanar anomaly. In the absence of geologic bodies of this geometry, the most likely conductor is a large fenced area.⁵ Anomalies of this type are virtually certain to be cultural if they occur in an area of culture.
- 5. EM anomalies which coincide with culture, as seen on the camera film or video display, are usually caused by culture. However, care is taken with such coincidences because a geologic conductor could occur beneath a fence, for example. In this example, the fence would be expected to yield an m-shaped coplanar anomaly as in case #2 above.

⁵ It is a characteristic of EM that geometrically similar anomalies are obtained from: (1) a planar conductor, and (2) a wire which forms a loop having dimensions identical to the perimeter of the equivalent planar conductor.

If, instead, a center-peaked coplanar anomaly occurred, there would be concern that a thick geologic conductor coincided with the cultural line.

6. The above description of anomaly shapes is valid when the culture is not conductively coupled to the environment. In this case, the anomalies arise from inductive coupling to the EM transmitter. However, when the environment is quite conductive (e.g., less than 100 ohm-m at 900 Hz), the cultural conductor may be conductively coupled to the environment. In this latter case, the anomaly shapes tend to be governed by current gathering. Current gathering can completely distort the anomaly shapes, thereby complicating the identification of cultural anomalies. In such circumstances, the interpreter can only rely on the radiation channels and on the camera film or video records.

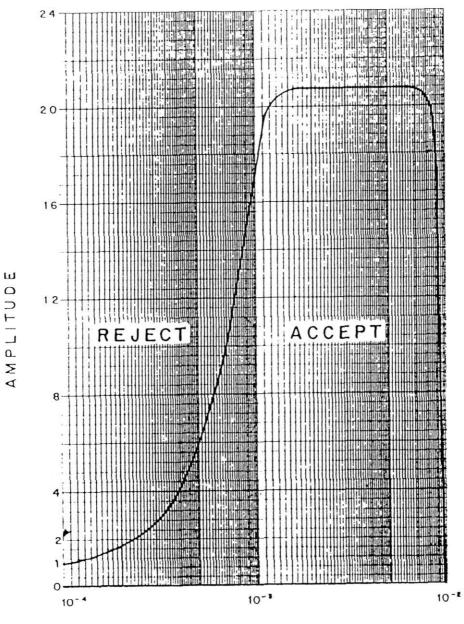
MAGNETICS

The existence of a magnetic correlation with an EM anomaly is indicated directly on the EM map. In some geological environments, an EM anomaly with magnetic correlation has a greater likelihood of being produced by sulfides than one that is non-magnetic. However, sulfide ore bodies may be non-magnetic (e.g., the Kidd Creek deposit near Timmins, Canada) as well as magnetic (e.g., the Mattabi deposit near Sturgeon Lake, Canada).

The magnetometer data are digitally recorded in the aircraft to an accuracy of 0.01 nT for cesium magnetometers. The digital tape is processed by computer to yield a total field magnetic contour map. When warranted, the magnetic data may also be treated mathematically to enhance the magnetic response of the near-surface geology, and an enhanced magnetic contour map is then produced. The response of the enhancement operator in the frequency domain is illustrated in Figure 5-2. This figure shows that the passband components of the airborne data are amplified 20 times by the enhancement operator. This means, for example, that a 100 nT anomaly on the enhanced map reflects a 5 nT anomaly for the passband components of the airborne data.

The enhanced map, which bears a resemblance to a downward continuation map, is produced by the digital bandpass filtering of the total field data. The enhancement is equivalent to continuing the field downward to a level (above the source) which is 1/20th of the actual sensor-source distance.

Because the enhanced magnetic map bears a resemblance to a ground magnetic map, it simplifies the recognition of trends in the rock strata and the interpretation of geological structure. It defines the near-surface local geology while de-emphasizing deep-seated regional features. It primarily has application when the magnetic rock units are steeply dipping and the earth's field dips in excess of 60 degrees.



CYCLES/METRE

Fig. 5-2

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Frequency response of magnetic enhancement operator.

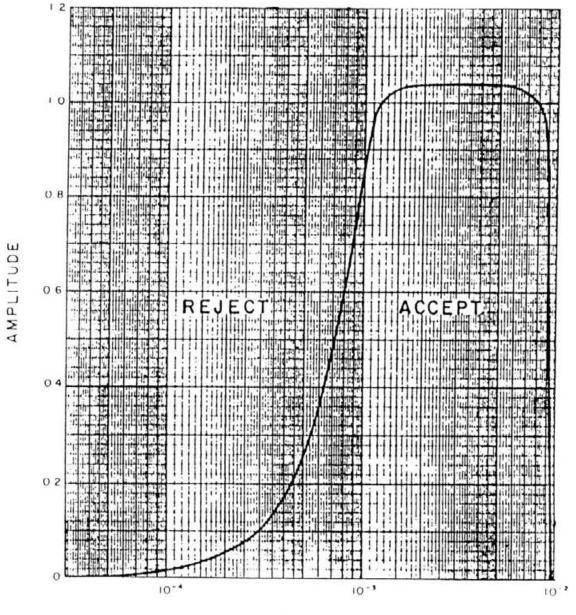
- 5.22 -

Any of a number of filter operators may be applied to the magnetic data, to yield vertical derivatives, continuations, magnetic susceptibility, etc. These may be displayed in contour, colour or shadow.

VLF

VLF transmitters produce high frequency uniform electromagnetic fields. However, VLF anomalies are not EM anomalies in the conventional sense. EM anomalies primarily reflect eddy currents flowing in conductors which have been energized inductively by the primary field. In contrast, VLF anomalies primarily reflect current gathering, which is a non-inductive phenomenon. The primary field sets up currents which flow weakly in rock and overburden, and these tend to collect in low resistivity zones. Such zones may be due to massive sulfides, shears, river valleys and even unconformities.

The VLF field is horizontal. Because of this, the method is quite sensitive to the angle of coupling between the conductor and the transmitted VLF field. Conductors which strike towards the VLF station will usually yield a stronger response than conductors which are nearly orthogonal to it.



CYCLES / METRE

Fig. 5-3 Frequency response of VLF operator.

- 5.24 -

1.00

The Herz Industries Ltd. Totem VLF-electromagnetometer measures the total field and vertical quadrature components. Both of these components are digitally recorded in the aircraft with a sensitivity of 0.1 percent. The total field yields peaks over VLF current concentrations whereas the quadrature component tends to yield crossovers. Both appear as traces on the profile records. The total field data are filtered digitally and displayed as contours to facilitate the recognition of trends in the rock strata and the interpretation of geologic structure.

The response of the VLF total field filter operator in the frequency domain (Figure 5-3) is basically similar to that used to produce the enhanced magnetic map (Figure 5-2). The two filters are identical along the abscissa but different along the ordinant. The VLF filter removes long wavelengths such as those which reflect regional and wave transmission variations. The filter sharpens short wavelength responses such as those which reflect local geological variations.

CONCLUSIONS AND RECOMMENDATIONS

This report provides a very brief description of the survey results and describes the equipment, procedures and logistics of the survey.

The survey was successful in locating several anomalous zones which may warrant additional work. The various maps included with this report display the magnetic and conductive properties of the survey area. It is recommended that the survey results be reviewed in detail, in conjunction with all available geophysical, geological and geochemical information. Particular reference should be made to the computer generated data profiles which clearly define the characteristics of the individual anomalies.

Most anomalies in the area are moderately weak and poorly-defined. Many have been attributed to conductive overburden or deep weathering, although a few appear to be associated with magnetite-rich rock units. Others coincide with possible structural features which may reflect faults or shears. Such structural breaks are considered to be of particular interest as they may have influenced mineral deposition within the survey area.

The interpreted bedrock conductors defined by the survey should be subjected to further investigation, using appropriate surface exploration techniques. Anomalies which

- 6.1 -

are currently considered to be of moderately low priority may require upgrading if follow-up results are favourable.

It is also recommended that image processing of existing geophysical data be considered, in order to extract the maximum amount of information from the survey results. Current software and imaging techniques often provide valuable information on structure and lithology, which may not be clearly evident on the contour and colour maps. These techniques can yield images which define subtle, but significant, structural details.

Respectfully submitted,

DIGHEM SURVEYS & PROCESSING INC.

itchard

Ruth A. Pritchard Geophysicist

RAP/sdp

A1138APR.93R

APPENDIX A

LIST OF PERSONNEL

The following personnel were involved in the acquisition, processing, interpretation and presentation of data, relating to a DIGHEM^v airborne geophysical survey carried out for Cogema Resources Inc. over the Yellow Moose Property, British Columbia.

Steve Kilty Robert Gordon Dave Miles Maurie Bergstrom Del Rokosh Gordon Smith Ruth A. Pritchard Lyn Vanderstarren Steve Mast Susan Pothiah Albina Tonello Vice President, Operations Survey Operations Supervisor Senior Geophysical Operator Second Geophysical Operator Pilot (Questral Helicopters Ltd.) Data Processing Supervisor Interpretation Geophysicist Drafting Supervisor Draftsperson (CAD) Word Processing Operator Secretary/Expeditor

The survey consisted of 377 km of coverage, including tie lines, flown from March 8 to March 12, 1993.

All personnel are employees of Dighem Surveys & Processing Inc., except for the pilot who is an employee of Questral Helicopters Ltd.

DIGHEM SURVEYS & PROCESSING INC.

2. thank

Ruth A. Pritchard Geophysicist

RAP/sdp

A1138APR.93R

APPENDIX B

STATEMENT OF EXPENDITURE

Date: April 26, 1993

IN ACCOUNT WITH DIGHEM SURVEYS & PROCESSING INC.

To: Dighem flying of Agreement dated December 15, 1992, pertaining to an Airborne Geophysical Survey over the Yellow Moose Property, British Columbia.

Survey Charges

350 km of traverse line flying @ \$77.00/km plus mobilization costs of \$1,250.00.

<u>\$28,200.00</u>

Allocation of Costs

- Data Acquisition	(60%)
- Data Processing	(20%)
- Interpretation, Report and Maps	(20%)

DIGHEM SURVEYS & PROCESSING INC.

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Ruth A. Pritchard Geophysicist

RAP/sdp

A1138APR.93R

APPENDIX C

STATEMENT OF QUALIFICATIONS

I, Ruth A. Pritchard of the City of Brampton, Province of Ontario, do hereby certify that:

- 1. I am a geophysicist, residing at 31 Barrington Crescent, Brampton, Ontario, L6Z 1N2.
- 2. I am a graduate of York University, Downsview, Ontario, with a Specialized Honours B.Sc. Earth Sciences Geophysics (1986).
- 3. I have been actively engaged in geophysical exploration since 1986.
- 4. The statements made in this report represent my best opinion and judgement.

I, Douglas C. Fraser of the City of Mississauga, Province of Ontario, do hereby certify that I am a Professional Engineer of Ontario, Registration Number 14934012, and that I have supervised the reporting herein.

RADARE

Ruth A. Pritchard Geophysicist

Trase

Douglas C. Fraser Geophysicist

APPENDIX D

EM ANOMALY LIST

			XIAL 26 HZ		ANAR 4 HZ		ANAR 9 HZ			. HORIZ . SHE		CONDUX EAR		MAG CORR
								. COND I .SIEMEN		COND		RESIS OHM-M	DEPIH M	NT
LIN	E 20010	(T	LIGHI	' 8)				•	•	•				
A	6813B?	ò	19	0	29	43	162	. 0.4	2	. 1	11	423	0	0
В	6810B?	0	17	0	28	43	162	. 0.4	7.	. 1	15	416	0	0
С	6783S	0	2	0	2	2	4	. –		. –	-	-	-	0
D	6709S	0	2	0	18	42	143	. 0.4	1.	. 1	33	649	0	0
E	6693S	1	8	3	14	42	107	. 0.7	7.	. 1	28	277	0	0
LIN	E 20020	(E	LIGHI	' 8)				•						
Α	6537S	ò	2	o	2	2	4				-	-	-	40
в	6577S	1	1	1	1	0	0	. –			-	-	-	0
С	6610S	0	2	0	2	2	4			. –	-	-	-	0
D	6628S?	0	2	0	2	2	4	. –			-	-	-	5
E	6634B?	Ö	13	2	19	16	127	. 0.4	0	. 1	23	572	0	8
LIN	E 20031	(I	LIGHI	' 8)				•						
A	6498S	ò	4	1	6	15	48	. 0.4	0	. 1	78	804	0	0
В	6413S	0	2	0	2	2	4	. –			-	-	-	0
С	6396S?	0	9	0	16	9	125	. 0.4	4	. 1	33	622	0	0
LIN	E 20040	(I	LIGHI	. 8)				• •						
A	6277S	ò	4	Ō	6	23	39	. 0.4	0	. 1	78	849	0	0
В	6307S?	2	5	1	6	22	35		28		57	234	13	0
T.TN	E 20050	(1	TIGHI	. 8)				•	•	•				
A	6098S?	ò	23	0	35	11	44	. 0.4	6	. 1	14	409	0	0
B	6071S	Õ	2	Ō	2	2	4		-		_	_	_	Ō
ē	6049S	ō	2	Ō	2	2	4		-		-	-	_	Ō
D	6032S	1	2	1	2	2	4	. –	-	. –	-	-	-	0
Έ	6010S?	1	2	ī	2	2	4		-		-	-	-	0
TIN	E 20060	(1	TIGHI	. 8)				•		•				
A	57985?	ò	6	1	9	37	18	. 0.4	0	. 1	47	667	0	0
В	58785?	-	4	ō					-		41		-	Ō
C	5921S	Ō	14	2			58		0		19		1	0
D	5938S	1	2	1	2			. –	-	. –	-	<u> </u>	-	0
Ε	5949S	5	3	2	5	16	27	. 7.7	53	. 1	43	403	0	0
LIN	E 20070	(1	TIGHI	. 8)				•		•				
	5740S?		12	Ō		34	187	. 0.4	8	. 1	25	499	0	30
В	5727B?		13	0							44		0	0
Ĉ	5696S	1	7	Ō	11		60		0	. 1	33		Ō	0
D	5681S?	0	2	0	2	2	4		-		-	-	-	20
	- -	TTMA	יית רוקויו	י נותסי	יס עג <i>ו</i>	יסואד ד	ים אד. דק	e becau	מה עובי	CULDUAT	ואנו סקי	דידכו		
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			XIAL 26 HZ		ANAR 4 HZ		ANAR 19 HZ			HORIZ		CONDUX EAR		MAG CORR
	OMALY/ 1 /INTERP									COND SIEMEN		RESIS OHM-M		NT
T.TN	E 20070	(1	TICHI	' 8)				•		•				
E	5668S	0	4	0	7	13	51	. 0.4	0	. 1	36	708	0	0
F	5659S?	ō	6	1	9	22	54				48	587	3	0
G	5631B?	0	10	0	10	55	71		9.	. 1	14	378	0	0
н	5608B?	1	2	1	2	2	4	. –	— .		-	-	-	0
I	5604B?	3	6	5	11	34	52	. 1.9	36	. 1	53	141	17	20
LIN	E 20080	(I	TICHI	' 8)				•		•				
A		0	5	0	8	8	11	. 0.4	1	. 1	85	869	2	0
в	5501S?	Õ	6	Ō	20	17	151				16	439	ō	150
c	5515S?	Ō	13	2	20	11	153				18	509	Ō	0
Ď	5537S?	Ō	23	3	34	40	243		8		9	347	Ō	0
Е	5553E	0	28	5	48	69	365		9.		5	293	0	0
F	5572H	2	13	8	24	76	71	. 0.5	0	. 1	53	74	21	0
T.TN	E 20090	(1	TICHI	. 8)				•		•				
	5313S	0	2	0	2	2	4	. –	-	. –	_	-	-	0
B	5274S	ō	2	Ō	2	2	4	-	-	-	_	-	. –	Ō
ē	5240S	Ō	2	Ō	2	2	4		_	-	-	_	-	Ō
D	5225S	0	2	0	2	2	4	•	-	. –	_	-	_	0
Е	5216S	0	2	0	2	2	4	. –	-	. –		-		13
F	5197S	0	5	2	9	26	58	. 0.4	0	. 1	16	464	0	0
G	5177B?	0	30	2	41	132	202	. 0.4	4	. 1	4	330	0	0
н	5157B?	1	2	1	2	2	4	. –	— .	. –	-	-	-	0
I	5153B?	3	16	17	16	45	51	. 1.0	4	. 1	27	80	1	0
J	5147B?	8	35	11	53	155	312	. 1.6	0	. 1	27	71	2	0
	E 20100	(1	FLIGHI	. 8)				•		•				
A	49745	ò	4	Ō	7	24	51	. 0.4	2	. 1	54	741	0	0
в	5010S	0	9	0	12	22	88			. 1	31	618	0	0
С	5022S	0	7	0	8	23	55	. 0.4	0	. 1	20	561	0	30
D	5033S	2	13	8	22	94	87		0	. 1	22	165	0	0
Е	5075B?	1	2	1	2	2	4	. –			-	-	-	0
F	5076B?	5	37	12	64	208	272	. 0.8	0	. 1	23	112	0	0
G	5093B?	13	39	31	62	219	179		0	. 1	35	54	7	0
H	5097B?	14	39	31	62	219	179	. 2.6	0	. 2	29	41	4	0
LIN	E 20110	(1	LICHI	. 8)				•		•				
	48735?	ò	11	0 0	16	34	123	. 0.4	0	. 1	21	587	0	0
В	4836S	Ō	2	Ō	2	2	4		-		-		-	Ō
c	_	1	2	1	2	2	4		-		-	-		Ō
D	4782H	1	10	5	14	40	70	. 0.4	0	. 1	44	154	6	Ō
	•			-								•		
								E BECAU						
								r to on				HT. •		
	للغال و	NE, C	JK BEL	AUSE	Or A	SHALL	TM NT	P OR OV	EKBUKD	EN EFFE	CIS.	•		•

			XIAL 6 HZ		ANAR 74 HZ		ANAR 9 HZ			HORIZ		CONDUC EAR		MAG CORR
								. COND I .SIEMEN		COND SIEMEN		RESIS OHM-M	DEPIH M	NT
								•		•				
LIN	E 20110		LICHI					•		•				
Ε	4755S?	1	2	1	2	2	4		-	. –	-		-	100
F	4740B?	4	6	2	9	30	46		37		52	124	16	0
G 	4722B?	15	55	11	85	305	301	. 2.2	0	. 2	27	33	6	0
LIN	E 20120	(F	LIGHI	: 8)	l			•		•				
Α	4507B?	Ó	4	0	5	15	30	. 0.4	0	. 1	46	750	0	0
В	4522S	0	2	1	3	23	12	. 1.0	0	. 1	47	202	23	0
С	4541S	0	5	1	7	18	32	. 0.4	0	. 1	55	603	0	0
D	4595S	0	2	1	2	2	4		-		-	-		0
E	4609S	0	6	3	9	27	47	. 0.4	0	. 1	27	241	0	0
F	4622B?	2	9	2	12	39	70	. 1.0	9	. 1	46	186	8	200
G	4633H	1	6	3	11	41	54	. 0.6	8	. 1	47	142	10	0
H	46485?	0	14	9	22	59	143		-	. 1	26	192	0	0
I	4681H	17	36	28	18	69	42		3		30	29	9	0
T TN	E 20130	(F	LIGHI	. 8)	1			•		•				
A	4460B?	5	12	. 14	18	15	37	. 2.1	21	. 1	52	128	17	0
В	4457B?	3	12	14	18	15	37		16		52	77	21	Ō
č	4450B?	ō	2	1	2	2	4				-	_	_	Ō
Ď	4441S	ŏ	õ	2	10	34	31		6	. 1	42	426	3	Ō
Ē	44135	ŏ	4	ō		12	29		ō		44	265	20	Ō
F	4358H	ŏ	4	3	6	25	26		ŏ		42	126	5	60
Ġ	4324H	4	17	4	25	68	126		ŏ	. 1	32	99	2	Ő
H	4296B?	1	2	1	23	2	4	· ····	Ľ	•	-	_	-	õ
I	4290D:	7	11	26	48	125		. 3.6	23	•	27		4	Ő
—— т тм	 E 20140	/1	LIGHI	n 01				•		•				
	4081B?	3	9	[8] 7	16	44	38	. 1.6	22	. 1	50	66	20	0
A B	4095B?	4	9	8	14	43	50		23	. 1	44		10	ŏ
C	4095B: 4114S	ů.	5	2	7	24	33		23	. 1	44	246	2	ŏ
D	41143 4133S	1	2	0	2	24	4		<u> </u>	• -	-	240	-	ŏ
E	41333 4172S	ō	2	1	2	2	4	• _	_	• _	_	_	_	Ő
F	41723 4197H	1	2	1	2	2	4	•	_	• _	_	_	-	ŏ
G	41971 4219B?		2	1	2	2	4	• _	_	• _	_	_	_	Ő
H	4219B: 4230H	8	13	11	20	68		. 3.8	7	. 2	30	32	6	ŏ
Ĩ	423011 4237B?		33	5	56	139	189		Ó	. 2	38		12	0
J	4237B?		33	15	56	139	189		0 0	. 1	28		2	0
		-						•		•				
	E 20150		LICHI					•		•				-
A		4	10	11	20	50	38		22	. 2	50	48	23	0
В	3990S	0	2	1	2	2	4	• -	-	• -	-	-	-	0
	. OF	THE	CONDU	JCTOR	MAY	BE DE	EPER C	E BECAUS R TO ONI P OR OVI	E SIDE	OF THE	E FLIG			

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			XTAL 26 HZ		ANAR 4 HZ		ANAR 19 HZ			. HORIZO . SHEE		CONDUC		MAG CORR
								. COND D. .SIEMEN		. COND D .SIEMEN		RESIS OHM-M	DEPIH M	NT
LIN	E 20150) (I	LICHL	8)				•		•				
c	39205	0	5	2	5	17	24	. 0.4	4	. 1	41	456	1	0
D	3900H	0	2	6	5	11	11	. 0.4	0	. 1	64	72	31	0
Е	3873B?	15	28	39	84	153	204	. 3.8	4	. 2	23	23	3	0
F	3837H	5	5	2	13	49	55	. 4.7	52	. 1	39	173	6	40
	E 20160) (I	LIGHT	8)				•		•				
A	3633H	1	2	1	2	2	4		-	. –	-	-	-	0
в	3646B?		2	0	2	2	4	. –	_	. –	-	-		60
С	3658S?		2	1	2	2	4		-		-	-	-	0
D	3681S?	' 1	2	1	2	2	4	• -	-		-	-	-	0
\mathbf{E}	3722B?	° 0	7	5	7	26	26	. 0.4	0	. 1	62	158	21	0
F	3742H	2	3	2	6	15	26	. 2.8	50	. 1	71	142	29	0
G	3781H	1	2	1	2	2	4		-	. –	-	-	-	0
H	3804H	2	9	2	13	46	71	. 0.9	13	. 1	37	217	3	0
	E 20170) (I	LIGHT	8)				•		•				
A	3522H	10	6	14	32	109	100	. 13.5	45	. 1	55	63	25	0
в	3516E	2	18	14	32	109	100	. 0.6	0	. 1	50	237	10	0
С	3501H	0	6	5	12	26	26	. 0.4	0	. 1	47	170	9	0
D	3467S	0	2	1	2	2	4		-		-	-	-	10
Ε	3434S?	2 3	6	2	6	21	13	. 2.5	32	. 1	58	227	14	0
F	3395H	5	6	7	14	43	37	. 3.4	30		52	79	19	0
G	3378B?		11	8	15	52	39	. 1.6	24	. 1	63	65	32	0
Η	3374B?		2	1	2	2	4		-	• -	-	-	-	0
Ι	3367B?		8	6	15	46		. 14.3	38		45	52	18	0
J	3365B?		8	6	15	46		. 12.1	36		43	45	18	0
K	3349B?		12	5	13	46	60	. 5.5	30	. 1	53	105	19	0
L	3345B?	? 1 -	2	1	0	2	4	• -	-	• -	-	-	-	30
LIN	E 20180) (1	LICHT	8)				•		•				
Α	3104H	1	2	1	2	2	4		-	. –	-	-	-	0
в	3134H	1	5	2	8	26	37		14		57	103	22	30
С	3145H?		9	5	12	33	46	. 0.5	5	. 1	54	100	20	0
D	3171B?		2	1	2	2	-	• -	_	• -	-	-	-	0
E	3174B?		7	1	8	37	40		29	. 1	39	632	0	0
F	3202S?		2	1	2	2	4			• •	-	-	-	0
G	3239B?		12	5	16	45	68		21	. 1	50	102	18	30
H	3252B?		2	1	2	2		• -	_	. 1	-			0
I	3262B?		10	3 5	19	60 65	51		9		42	100	7 8	0
J 	3278B?	? 7	14	C	20	65	92	. 2.6	19	• 1	39	123	Ö	U
	E 20190	•	FLIGHT	•				•		•				
Α	3049H	1	2	1	2	2	4	• -	-	• -	-	-	-	0
	• =	יעידעראנ	יירו רושויו	n Ling	ים עגו	יכתאדד ים	יסגדוק	E BECAUS	n nur		יגרד כדי	• •		
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								P OR OVE						÷
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			XTAL 6 HZ		ANAR 4 HZ		ANAR 19 HZ			. HORIZO		CONDUC		MAG CORR
AN	OMALY/ I	REAL	OUAD	REAL	OUAD	REAL	QUAD	. COND I)EPIH*	. COND D	EPIH	RESIS	DEPTH	
								.SIEMEN		SIEMEN		OHM-M	М	NT
 T TNT		(1						•		•				
B	E 20190 3022H	(r 4	LIGHI 7	'8) 6	11	2	9	. 3.2	28	. 1	57	67	24	ο
В С	2991S?	40	8	1	11	32	- 65		20	. 1	25	543	24	280
	29915: 29715?	0	8	3	12	35	66		_	. 1	30	328	ŏ	170
D	29713: 2925H	2	° 7	1 1	7	28	35		26		52	140	16	0
E				1 4							50		10	0
F	2906H	4	7		9	27	25		35	• 1	50		14	
G	2882H	1	2	1	2	2	4	• -	-		-	-	-	0
LIN	E 20200	(F	LIGHI	8)				•		•				
A	2628B?	12	64	36	114	286	374	. 1.6	0	. 1	23	52	1	0
В	2631B?	18	65	36	115	286	374		-	. 1	21	48	0	0
Ĉ	2656H	1	2	1	2	2		. –	-			-	-	0
D	2670H	3	5	3	9	29		. 2.2	36	. 1	53	136	16	0
Ē	2684S?	1	2	1	2	2			-	. –	<u> </u>	-	-	0
F	2717S	0	8	2	11	26	58	. 0.4	0	. 1	33	584	0	0
Ğ	2732S?	0	7	2	7	28	41		0		41	438	0	190
H	2753B?	1	2	1	2	2	4		-	. –	-	-	_	0
I	2757B?	1	2	1	2	2		. –	_		-	-	-	0
Ĵ	2760B?	14	30	12	49	158	157	. 3.5	1	. 1	21	88	0	0
ĸ	2778H	1	2	1	2	2	4	. –	_			-	-	50
L	27885?	1	1	1	2	ō	4		-	. –	-	-	-	7
M	2805B?	9	8	11	6	5		. 7.5	36	. 2	62	45	33	20
N	2818H	3	7	8	12	36	~~	. 2.2	31		57	36	31	0
								•		•				
LIN	E 20210	(I	LIGHI	•				•		•				
A	2576B ?	4	26	3	26	94	176		0	. 1	14	278	0	0
В	2567S?	0	10	2	17	56	89		. 0	. 1	13	414	0	0
С	2548H	4	3	7	12	17		. 6.4	36	. 1	40		3	0
D	2535H	1	4	2	5	24		. 0.8	16	. 1	57	131	18	0
Ε	2523S?	3	8	4	9	31		. 1.7	6	. 1	35		0	0
F	2484S	0	6	1	9	8	29	. 0.4	0	. 1	42	637	0	0
G	2473S	1	2	1	2	2	4		-		-	-	. –	210
н	2436B?	13	38	23	60	230	255		0	. 1	18	67	0	6
I	2433B?	9	38	23	60	230	255		0		28	57	2	0
J	2424B?	14	20	17	53	127	112		7		20		0	0
K	2408B?	6	10	14	37	58	11	. 2.8	18	. 1	23	163	0	6
\mathbf{L}	2389H	1	2	1	2	2	4		-	. –	-	-	-	0
М	2378H	6	25	15	40	111	145	. 1.5	0	. 2	34	33	11	0
 T T T T T T		/1	LIGHI	1 01				•		•				
	E 20220	•				100	100	• 7 9	~	•	17	252	0	•
	2115B?	5	25	6			190		0		17		0	0 0
В	21435?	4	10	5	15	33	68	. 1.9	11	. 1	33	168	0	U
	* ES	PIMAT	TED DE	PIH N	MAY BI	e unri	TTABL	E BECAUS	SE THE	STRONG	TR PA	RT:		

* ESTIMATED DEPTH MAY BE UNRELIABLE BECAUSE THE STRONGER PART .

•

. OF THE CONDUCTOR MAY BE DEEPER OR TO ONE SIDE OF THE FLIGHT .

. LINE, OR BECAUSE OF A SHALLOW DIP OR OVERBURDEN EFFECTS.

			XIAL 26 HZ		ANAR 4 HZ		ANAR 19 HZ			. HORIZ		CONDUC EARI		MAG CORR
ΔN	OMATVI	DFAL.	OTIAD	DFAT.	OTIAD	DENT.		. COND I	1 PDTU+		NEDTH	DECTO	הניסיות	
								.SIEMEN		.SIEMEN		OHM-M	M	NT
		•						•		•				
	E 20220	•	LICHI					•		•				
С	2165H	4	5	5	8	26	31	. 3.5	31	. 1	48	87	13	0
D	2224S	1	2	1	2	2	4	-	-		-	-	-	0
E	2254H	13	20	6	15	106		. 4.3	1		28	36	4	0
\mathbf{F}	2257B?		23	6	38	87		. 4.0	0		27	45	1	0
G	2277S?		16	5	24	79		. 2.8	2		25	135	0	0
н	2282H	8	6	7	10	41	57	. 7.5	38	. 1	47	79	15	0
I	2291B?		2	1	2	2		. –	-	. –	-	-	-	0
J	2299B?	11	9	10	25	70	81	. 9.1	19	. 2	35	23	13	0
Κ	2305B?	3	15	10	32	79	105	. 0.8	0	. 2	30	32	8	120
\mathbf{L}	2309B?	6	19	12	36	53	88	. 1.9	2	. 2	33	37	9	0
								•		•				
	E 20230	•	TICHT	•		_		•		•				•
A	2009B?		2	1	2	2	4	• -	-	• -	_	-	-	0
В	1997B?		74	22	118	394	543		0	. 1	15	95	0	0
С	1972B?		54	12	82	132	472	. 1.6	0	. 1	19	116	0	0
D	1924S?		2	1	2	2	4	•	_	• -	— `	-		0
E	1874S	2	6	6	25	56		. 1.4	25		23	210	0	0
F	1845H	1	2	1	2	2	4	• -	-	• -	_	_	-	30
G	1824B	42	127	64	215	742	755	. 3.7	0	. 2	18	28	2	0
H	1809B?		1	1	2	2	4	• •	<u> </u>		-		-	0
I	1803B?		2	1	2	2	4		-	• -	-	-	-	0
J	1782E	23	62	62	160	353	514		0	. 2	38	39	13	0
Κ	1775B?		65	62	160	353	514		1	. 2	30	27	10	0
\mathbf{L}	1769B?		27	23	68	147	232	. 3.9	15		33	33	12	0
М	1751B?	11	30	22	49	131	132	. 2.6	0	. 2	35	22	14	0
								•		•				
	E 20240	•	LIGHT	•		100	1 67	•	•	•	10	105	•	20
A	1450S?		28	6	47	180	167		0	. 1	18	105	0	20
B	1472H	9	9	12	32	19	86		21		32	66	3	0
C	1593S?	-	19	6	29	57	84		7	. 1	26	176	0	0
D	1638H	1	2	1	2	2	4	-	-	• -	-	_	-	0
E	1661H	5	11	5	19	60	76		21		42	91	11	0
F	1674E	26	24	38	120		151		11		39	42	14	0
G	1677H	26	24	38	120	370	151		21	. 2	24	32	5	0
H	1698B?		2	1	2	2	4		_	• -	-		-	0
I	1700B?	11	40	30	63	178	181	. 2.1	0	. 2	33	25	12	0
T.TN	E 20250	(1	LIGHT	8)				•		•				
A	1393S?	•	2	1	2	2	4	•		•	_	_		0
B	1390S?		16	27	112		510	-	21	. 1	15	76	-	0
C	13903. 1385E	15	65	27	112				31		15	76	0	0
C	TOOL	T.	05	41	112	495	510	. 2.0	1	. 1	5	287	0	15
	* 70	ጥፐለአጣ	יידרו רוביי	י גדרס	ים עמ	יכואד ק	T.TADT	E BECAUS	aun ar	CULOUIS	יארד סק	• ידיכ		
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			XTAL 26 HZ		ANAR 74 HZ		ANAR 19 HZ			HORIZ		CONDUX EAR		MAG CORR
								. COND I					DEPTH	
FID,	/INTERP	PPM	PPM	PPM	PPM	PPM	PPM	.SIEMEN	М	SIEMEN	I M	OHM-M	M	NT
								•		•				
	E 20250	-	TIGHI			~~~	0.1	• • •	2	•	25	40	10	
D	1369H	7	20	18	12	32	21		3		35 63	42 59	10 26	0 0
E	1357B?		4	6	13	37 2	0 4		15	• 2	-	- 59	20	0
F	1342B?		2	0	2 6		4 18			•	- 49	- 98	11	0
G	1331H	2	3	3 1		23 2	18 4		22	•	49	-	-	110
H	1312S?		2 2	1	2 2	2	4 4		_	• _	_	_	_	0
I	1232H	1 1	2	1	2	2	4 4		_	• _	_	_	_	ŏ
J	1191E		2	1	2	2	4 4	• _	_	• _	_		_	ŏ
K L	1188B? 1181B?		2	1	2	2	4 4	• _	_	• _	_	_	_	Ö
10 	TTOTE:	±.	4	т.	2	2	4	• -		•				Ŭ
T.TN	E 20260	(1	TIGH	r 8)				•		•				
A	882S	1	2	1	2	2	4		-		_	_	_	0
В	893S	3	19	5	29	105	142	. 0.8	0	. 1	20	164	0	0
č	912B?		2	1	2	2	4	. –	_		_	_	_	Ō
D	916B?		2	1	2	2	4	. –	-		-	-	-	40
Ē	935H	3	8	6	12	41	48	. 2.0	19	. 1	51	62	20	0
F	993B?		2	1	2	2	4		-	. –	-	-	-	0
Ģ	1009B?		8	5	13	41	37	. 4.1	3	. 1	52	69	16	0
Ĥ	1037B?		6	13	16	49		. 11.0	44	. 2	58	29	33	0
Ι	1049B?		2	1	2	2	4		-	. –	-	-	-	0
J	1066H	4	7	4	13	45	57	. 2.9	36	. 1	50	70	20	14
K	1084E	8	20	11	32	107	141	. 2.3	3	. 2	32	33	9	0
\mathbf{L}	1100H	15	38	31	66	212	195	. 3.0	1	. 2	32	24	11	0
		•						•		•				
LIN	E 20270	(I	TIGH	r 8))			•		•				
Α	784S	5	12	4	18	66	79		28	. 1	32	177	3	0
В	764S	3	16	4	33	107	194		0	. 1	15		0	0
С	755B?		16	27	47	119	122		28	. 1	52	76	22	0
D	739B?		12	24	67	26	132		32	. 2	33	20	14	0
Έ	729B?	19	15	15	24	23	36		25	. 3	52	20	30	0
F	693H	4	21	4	36	94	189		5	. 1	25		0	0
G	640B?		26	14			113		0	. 2	27	45	2	0
H	621B?		2	1	2	2	4			• -		-		0
I	591B?		18	44	90		72		20		26		7	0
J	587B?		60	28	93				0		25			20
K	560H	2	4	1			15		34		45			15
L	547H	7	17	14				. 2.5	11	• 2	31	32	9	0
M	528B?		2	1	2	2	4	• -	-	• -	-	-	-	0
N	518H	1	2	1	2	2	4	• -	-	• -	-	-	-	0
								•		•				
	E 20280	•	FLIGH			107	040	•		•	20	40	-	C 0
Α	5645E	15	40	24	62	187	248	. 2.9	4	. 1	30	48	7	60
	• _L 77-	*****	ייד רובוו	י דשורת:	ירד קראתו	ירת אדן יק	ניד דאנזי	יזגראזכו קו		CULTON	יגרו נויסי	•		
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			XTAL 26 HZ		LANAR 74 HZ		ANAR 19 HZ			. HORIZ . SHE		CONDUC		MA) COR
	OMALY/ : /INTERP							. COND I .SIEMEN		COND SIEMEN		RESIS OHM-M	DEPTH M	N
								•						
LIN	E 20280	•	TIGHI	r 9))			•						
в	5656B?	1	20	22	32	11	16		0		36	17	16	
С	5673B?	13	18	30	52	68	162		20		53	23	30	
D	5676B?	13	35	30	51	178	154		0		35	25	14	5
Е	5693B?	14	44	18	37	33	265		13 .		4	220	0	
F	5699B	17	31	19	27	99	192		21		28	92	5	
G	5736H	7	13	8	21	44	67		23		52	91	19	
Η	5767B?	10	17	4	15	32	43		15 .		54	62	23	1
I	5776B?	12	13	7	9	39	20		21		50	55	21	
J	5779B?	8	6	7	9	39	21		43		59	61	28	4
K	5790H	13	19	14	30	90	104		8		31	44	6	
L	5807B?	32	58	51	131	472	228	. 5.0	7		22	24	5	
М	5833B?	1	2	0	2	2	4					-		7
N	5850B?	17	53	30	87	259	317		0		29	30	9	
0	5853B?	20	44	30	87	259	317		8.		33	27	13	
P	5867B?	8	14	10	21	75	50	. 3.5	10 .	. 2	39	37	13	
Q	5874B?	1	2	1	2	. 2	4	• -		. –	-	-	-	
	E 20290	(1	TIGHI	. 9)	h			•	•	•				
A	5594S?	1	2	1	2	2	4	• _	_	· _	-	-	_	
B	5577B?	9	19	12	27	91	113	. 2.8	7	. 1	32	70	4	
č	5572B?	1	2	1	2	2	4	- 2.0	-		-	-	-,	
D	5545B?	14	5	16	36	25	160	. 28.3	42	-	45	45	19	5
Ē	5540H	1	2	1	2	2	4	. –	-	. –	-	-	_	-
F	5534B?	11	15	7	76	208	377	. 4.6	25	-	32	63	7	
G	5505S?	5	7	6	10	32	25		7		46	87	.9	
H	5470H	11	11	8	14	16	~	. 6.7	34		61	48	32	
I	5462B?		9	5	11	11	53		22		50	47	21	
J	5448H	4	16	9	23	78	~ ~	. 1.3	0	2	36	44	10	
ĸ	5428B?	5	13	16	24	70	57		14		47	25	24	
L	5406B?	15	27	25	49	111	165		9	2	35	25	13	
								•		•				
LIN	E 20300	•	LIGH	: 9)	ŀ			•						
А	5149B?	7	19	7	26	96	103	. 2.3	3.	. 1	32	85	3	
В	5183H	1	2	1	2	2	4				-	-		
С	5218H	7	18	13	30	99	52		12 .	. 2	52	39	26	
D	5244B?	1	2	1	2	2	4	–		. –	-	-	-	
Ε	5262B?	1	2	1	2	2	4			. –	-	-	-	
\mathbf{F}	5267B?	5	14	10	24	78	90		15 .	. 1	46	56	19	
G	5281B?	7	31	13	53	161	236		0		33	62	6	
Η	5291B?	3	11	5	23	63	99		26 .		49	65	22	
T	5306B?	15	14	18	29	92	76	. 8.2	24 .	. 2	39	24	18	

. LINE, OR BECAUSE OF A SHALLOW DIP OR OVERBURDEN EFFECTS.

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			AXIAL 26 HZ		ANAR 4 HZ		ANAR 9 HZ			. HORIZA . SHE		CONDUC		MAC CORE
								. COND I .SIEMEN		. COND I .SIEMEN		RESIS OHMM	DEPIH M	M
LIN	E 20300	(1	LIGHT	9)				•		•				
J	5313B?		10	28	31	94	61	. 1.3	25	. 3	53	19	33	10
Κ	5336B?	26	31	33	54	177	70		7		35	25	14	18
\mathbf{L}	5340B?	13	8	33	54	177	133	. 11.9	38	. 2	39	25	18	C
М	5353B?	23	31	39	1	165	32	. 6.3	16	. 3	40	20	20	C
LIN	E 20310	(1	LIGHT	9)				•		•				
Α	5045H	ì	2	1	2	2	4	. –	-		-	_	. <u> </u>	C
В	5027B?		24	11	39	125	172	. 2.1	3	. 1	32	83	3	(
С	4990H	1	2	1	2	2					_	-	-	(
D	4969H	5	14	10	21	68	73	. 1.7	0	. 2	42	40	15	(
Ε	4924H	5	18	4	28	80	159	. 1.5	5	. 1	27	109	0	30
F	4893B?	1	2	1	2	2	4	. –	-	. –	-	-	-	1:
G	4886B?	8	12	16	39	64	107	. 3.8	18	. 3	45	20	23	(
н	4879B?	1	2	1	2	2	4		-	. –	-		_	(
Ι	4876B?	16	15	21	28	90	54	. 8.3	24	. 3	52	22	29	(
J	4850B?	5	8	15	20	34	23	. 3.0	26	. 2	56	25	31	(
T.TN	E 20320	/τ	LIGHT	٥١				•		•				
	4616B?	9	25	9) 11	38	121	158	. 2.2		• • 1	32	76	5	
	4610B; 4634B?	15	15	16	24	75	158 59		4 26		-3∠ 53	76 50	5 25	1
C	4654D. 4660H	1	2	10	24	2	4		<u> </u>		- 55	- 50	29 	(
D	4681H	7	16	8	26	93	90		8	•	37	_ 50	10	(
E	4690H	5	9	9	20	36	11		0		41	30 37	28	
F	4715S	3	16	3	22	72	146		0		20	231	28	19
Ġ	4727S	4	21	1	31	88	185		3		11	327	0	ш. (
	4745E	23	49	36	82	252	225		0	. 2	33	38	9	, i
ï	4752B?	1	2	1	2	252	225 4		_	• •	-			
	4765B?	8	19	18	37	99	82		18	. 2	47	26	25	Ċ
ĸ	4784B?	1	2	1	2	2	4		-	. –	-	-	-	Ċ
——- Т.Т.М	E 20330	75	LICHT	9)				•		•				
	4546H	1	10	2	15	43	100	. 0.4	0	. 1	36	145	2	(
	4516H	î	2	1	2		100 4			• •		140	<u> </u>	
	4492B?	13	13	12	55	192	250		29	. 1	31	- 50	7	
	4486B?	5	40	12	56	146	290		0		34	50 55	10	
Ē	4472H	1	2	1	2	2	200 4			• -	-		-	
F	4458S?	4	12	2	16	46	100		11	. 1	35	- 246	0	, (
	4443S?	Ō	12	2	18	55	95		0		25	240 265	0 0	, (
	4426H?	17	41	24	83	249	272		2		20	265 51	0	, (
Ï	4410H	4	4 4	24	5	16	26		55		43	73	14	, (
Ĵ	4393H	9	17	13	27	84	72		12			26	29	
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OF THE CONDUCTOR MAY BE DEEPER OR TO ONE SIDE OF THE FLIGHT . LINE, OR BECAUSE OF A SHALLOW DIP OR OVERBURDEN EFFECTS. . •

			XIAL 26 HZ		ANAR 4 HZ		ANAR 19 HZ			. HORIZ . SHE		CONDUC		MAG CORR
	omaly/ : /INTERP									COND SIEMEN		RESIS OHM-M		NT
	E 20330	(1	LIGHT	<u>م</u>				•		•				
	4376H	12 12	8	' 9) 27	23	33	19	. 10.9	26	. 3	46	14	26	0
LIN	E 20340	(F	LIGHT	' 9)				•		•				
A	4128B?	7	5	4	8	30	32	. 8.2	41	. 1	41	181	4	0
В	4171H	7	13	8	22	55	86	. 3.2	13	. 2	39	42	13	0
С	4177B?	9	16	12	34	62	80	. 3.4	14	. 2	34	35	10	0
D	4180H	9	5	12	42	35	143				34	32	11	0
Е	4212H	9	24	10	36	85	151				35	53	9	30
F	4225B?	10	2	4	5	16	43				65	64	32	G
G	4242B?	6	7	5	13	53	60				44	132	10	0
H	4266S?	10	33	11	52	197	231				19	91	0	20
Ι	4278B?	9	6	7	23	68	86				43	102	11	0
J	4288B?	8	12	4	16	40	43				49	28	24	0
K	4296E	11	15	20	24	66	79		17	. 2	58	29	33	0
L	4319H	1	2	1	2	2	4		-		-	-	-	0
LIN	E 20350	(F	LIGHT	9)				•		•				
Α	4050H	. 8	11	14	41	121	198	. 4.0	27	. 2	33	38	10	C
В	4023H	1	2	1	2	2	4			. –	-	—		C
С	3988S?	5	10	2	14	53	74	. 2.4	28	. 1	36	294	1	C
D	3970B?	1	2	1	2	2	4		-	. –	-	-	-	13
Ε	3968B?	8	51	11	75	190	458	. 1.1	0	. 1	19	99	0	0
F	3963B?	1	2	1	2	2	4	. –	-		-	-	-	40
G	3949B?	13	8	16	25	10	38	. 13.0	43	. 3	58	21	36	13
H	3941B?	10	2	24	24	66	33	. 55.5	43	• 2	63	28	37	0
LIN	E 20360	(F	LICHT	9)				•		-				
Α	3748H	ì	2	1	2	2	4	. –		. –	-	-		c
в	3776H	3	7	3	10	33	34		8	. 2	45	38	18	Ō
С	3799H	7	10	8	17	56	27				54	77	22	Ċ
D	3834S?	7	26	11	39	148	169	. 1.7			23	79	0	C
Е	3847B?	10	12	9	24	62	50	. 5.8	27	. 3	54	18	33	C
F	3855B?	17	14	28	17	73	28	. 9.9	31	. 2	66	27	41	C
G	3883B?	9	29	24	50	145	114	. 2.0	6	. 2	38	44	14	C
LIN	E 20370	(F	LIGHT	9)				•		•				
	3574S	1	7	4	11	30	53	. 0.4	6	. 1	27	368	0	C
в	3557B?	3	16	8	23	73	125				41	115	11	C
č	3553B?	1	2	ī	2	2	4		_	- -			_	C
D	3545H	1	2	1	2	2	4		-		-	-	_	C
Ē	3529H	14	28	22	24	35	86		4	. 2	29	32	7	C
	• OF	THE	CONDU	CIOR	May i	BE DEF	PER O	e becau r to on p or ov	E SIDE	OF THE	FLIG			

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			XIAL 26 HZ		ANAR 4 HZ		ANAR 19 HZ		VERTI DIK		. HORIZA . SHE		CONDUC EAR		MZ COF
	OMALY/ : /INTERP				_						. COND :		RESIS OHM-M		ľ
		/-						•			•				
	E 20370	•	TICHL	•			40	•	~ ~	24	•		01	10	
F	3501H	6	9	10	27	53	42		3.9	34	. 1	46	81 77	16 7	
G	3480B?	10	36	15	54	156	194		2.0	5	. 1	34			
Н	3477H	10	36	15	54	156	194		2.0	8	. 2	56	34	31	
I	3461H	1	2	1	2	2	4		_				-	-	4
J	3453B?		8	8	12	48	57		9.4	44		53	79	22	
K	3444B?		11	0	18	53	93		2.7	.35		60			
L	3434B?		50	8	76	219	403		1.1	0	. 1	21	86	0	
M	3418B?			53	93	299	256		5.4	5	. 3	32	13	15	
N 	3411B?	1	2	0	2	2	4	•	-	-	• -	-	-	_	
	E 20380		LIGHT	-				•			•				
Α	3170B?		7	14	14	46	28		8.3	31	. 1	35	75	5	
В	3174B?		2	1	2	2	4		-	-	• -		-	_	
С	3196B?		7	13	15	51	14		8.6	35		57	40	29	
D	3204B?		22	19	46	151	173		4.2	15		37	29	15	
Е	3250E	14	32	24	55	186	178	•	3.1	1	. 1	46	107	12	
F	3259B?		1	1	0	2		•	-	-		-	-	-	
G	3263H	9	6	14	20	42	13	•	9.8	27	. 2	33	24	11	
H	3276H	1	2	1	2	2		•	-	-		-		_	
I	3284H	7	18	13	28	88	81		2.1		. 1	48	54	21	
J	3302H	19	27	36	46	98	81		5.6	4	. 3	36	19	15	
К	3309B?		46	57	86	270	176		6.2	3	. 3	28	13	10	
\mathbf{L}	3312B?		51	57	86	270	176		1.6	0	. 4	27	11	11	
М	3326B?		2	1	2	2	1		-	-	• -	-	. –	-	
N	3339B?	8	20	7	31	111	69		2.3	6	. 2	30	30	8	
0	3342B?		19	7	23	99	111	•	2.6	3	. 2	27	30	5	
P	3348B?	1	2	1	2	2	4	•	-	-	•	-	-	-	
LIN	E 20390	(I	LIGHI	9)				•			•				
Α	3067S?	2	19	11		87	136	•	0.5	0	. 1	29	108	0	
в	3053B?	11	3	16	27	81	4	•	41.6	42	. 2	43	44	16	
С	3049H	1	2	1	2	2	4	•	-	-			-	-	
D	3025H	24	22	44	98	314	269	•	9.3	16	. 3	24	18	6	
\mathbf{E}	2982H	1	2	1	2	2	4	•	-	-		-	-	-	
F	2941H	8	5	15	13	41	43	•	10.1	48	. 3	53	19	32	
G	2924B?	2	14	29	22	65	23	•	0.6	3	. 4	50	8	33	
Н	2919B?	1	2	1	2	2	4		-	-		-	-	-	
	2903B?		26	12	40	119	208	•	1.7	3	. 3	43	16	23	
J	2898B?	1	2	1	2	2	4		-	-	. –	-	-	-	
LTN	E 20400	ርጉ	LIGHT	9)				•			•				
	2685B?			1	2	2	4		· _	_		_		-	
	.* ES . OF	FIMAI THE	TED DE CONDU	PIH M CIOR	1ay be May e	E UNRE BE DEF	LIABI PER C	R	TO ONE	SIDE	STRONG OF THE EN EFFE	FLIG			

			AXIAL 26 HZ		LANAR 74 HZ		LANAR 19 HZ			HORIZA SHE		CONDUC		MAG CORR
								CONDI						
FID	/INTERP	PPM	PPM	PPM	PPM	PPM	PPM	.SIEMEN	м.	SIEMEN	M	OHM-M	M	NT
		•						•	•	•				
	E 20400	•	FLIGHI	•		200	272	•	~ `	, 			0	•
B C	2690B? 2698B?		60 17	38	92 28	288 83	373 42		0.		28 49	24	9	0
				18					5.			20	26	0
D	2701B? 2716H		17	18	28	83	42		9.		41	17	21	0
E		2 5	15	39	25	46	142		7.		31	15	14	0
F	2747H		4	10	14	120	100		55 .		36	40	12	0
G	2768H	7		18	48	167	165		0.		35	19	15	20
H	2780B?		35	32	60	67	146		0.		28	15	10	0
Ĩ	2786B?		7	32	46	119	51		40 .		41	24	20	0
J	2790B?		15	5	22	129	60		14		50	24	28	0
K	2804B?			20	29	105	60		19 .		46	14	27	0
L	2810B?			1	2	2	4				-	-		0
М	2848B?	14	5	32	21	58	26	. 28.2	33 .	. 4	35	10	18	0
								•	•	•				
	E 20410		FLIGHI			~ • • •		•					-	
A	2619B?			53	108	346	252		11 .		26	17	9	0
B	2614B?		53	26	33	58	251		0.		31	16	13	0
С	2594H	2	24	34	45	125	59		0.		36	16	18	0
D	2562B?			3	79	207	262		41 .		48	34	23	0
E	2556B?			35	135	427	447		4.		28	20	11	0
F	2538H	1	18	13	26	55	81		Ο.	. 4	38	11	21	0
G	2512B?		2	1	2	2	4			. —	-	-	_	0
H	2470B?	14	11	43	49	147	253	. 8.8	25 .	4	35	12	18	0
		•						•	•	ı				
	E 20420	•	FLIGHI	•				•		· _			_	
A	2284B?		38	29	70	222	186		0.		26	13	9	0
В	2297H	21	56	50	98	330	218		3.		29	15	12	0
C	2328H	8	16	15	42	117	88		19 .		35	34	13	40
D	2341H	12	18	23	30	82	97		19 .		38	19	18	0
E	2353H	15	46	50	77	173	162		2.	4	32	11	16	0
F	2388B?	2	18	11	30	91	43	. 0.4	3.	. 4	53	10	35	0
			त्वर भरत्य का					•	٠					
_	E 20430		FLIGHI	•		60	10	• • •	10		50	~ 1		•
A				3		62	16		18.		50	21	29	0
B	2158H		12	26		53			25		31	16	12	0
C	2148B?		35	27	60	146	173		10 .	. 3	37	20	18	13
D	2132H	1		1	2	2	4	-		. –	-	-	-	0
E	2122H	1	2	1	2	2	4	-		_	-	-	-	0
F	2107B?			1			4	-		-	-	-	-	0
G	2093B?			1			4	-			-	-	-	7
H	2089B?			1		2	4			. –	-	-	-	0
I	2071B?	1	2	1	2	2	4			. –	-	-	-	6
	•											•		
								E BECAUS						
								R TO ONE				Π.		
	• 11	NE, (JK BEC	AUSE	of A	SHALI	TO NOT	P OR OVE	RBURDI	IN LIFFEX	JIS.	•		

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			XIAL 26 HZ		ANAR 4 HZ		LANAR L9 HZ			. HORIZO . SHEE		EARI		MZ COI
								. COND : .SIEMEN		. COND D .SIEMEN		RESIS OHM-M	DEPTH M	1
	E 20430	(F	LIGHT	9}				•		•				
J	2064H	14	37	22	59	176	154	. 2.9	12	. 3	41	13	24	
ĸ	2053B	53	84	98	146	426	259		3		25		12	
	2045B?	21	18	65	50	150	39		29		59	12	40	
M		4	27	6	45	129	268		1		35	76	9	1
LINI	E 20440	(H	LICHI	' 9)				•		•				
Α	1809H	1	2	1	2	2	4		-	. –	-	-		
В	1839B?	6	10	6	17	55	66		32	. 1	49	71	20	
С	1851B?	5	14	5	21	85	73	. 2.0	10	. 1	54	126	17	
D	1864E	14	19	68	148	484	438	. 5.1	12	. 1	43	· 5 9	14	
Ε	1877H	17	31	42	56	14	85	. 4.1	5	. 4	29	10	13	
F	1902B?	1	2	1	2	2	4	. –	-	. –	-	-	-	
G	1917B	78	134	127	219	610	572	. 7.3	4	. 4	26	8	13	
Н	1961H	4	14	15	26	84	72	. 1.4	8	. 2	50	23	27	
LIN	E 20450	(I	TIGHT	' 9)				•		•				
Α	1671H	9	2	43	3	153	10		0		19	14	10	
в	1649E	13	23	23	33	110	67	. 3.8	22		55	44	28	
С	1637H	1	4	2	9	31	54	. 1.1	38		63	140	26	
D	1612B?	7	21	32	11	40	58	. 1.9	8	. 2	46	28	23	
Е	1608B?	1	2	1	2	2	4			• •	-	-	-	
F	1598B	42	76	82	122	369	248	. 5.6	2	. 4	30	10	15	
G	1595B	50	76	82	122	369	248	. 7.1	2	. 4	27	9	13	
Н	1583B?	16	12	23	27	72	70	. 10.1	30	. 4	50	11	32	
Ι	1576B?	11	7	20	15	64	73	. 11.6	43	. 3	56	18	35	
J	1567B?	1	2	1	2	2	4			. –	-	-	-	
Κ	1555H	28	30	49	50	163	75	. 8.6	10	. 4	36	9	20	
L	1528B?	1	2	1	2	2	4	. –	-	. –			-	
М	1515B?	1	2	1	2	2	4		-	• -	-	-	-	
LIN	E 20460	(E	TICHT	' 9)				•		•				
Α	1322H	30	62	55	105		295		0	. 3	25	13	8	
	1338H	1	2	1	2	2	4		-		-	-	-	
	1362B?	1	2	1	2	2	0		-	. –		-	-	
	1395B?	6	11	36	47	134	30		21		32	14	14	
	1415B?	9	21	11	30	110	172		12		50	36	25	
	1426B?	14	33	28	55	195	199		4	. 3	33	19	14	
	1442B?	40	49	70	91	257	94	. 8.2	18	. 4	41	10	26	
	1448B?	36	13	62	70	211	45	. 38.4	18	. 4	33	11	16	
1 	1467B?	11	8	18	12	39	32	. 9.3	43	. 1	- 59	64	2 9	
	E 20470	•	LIGHT					•		•				
Α	1280H	30	32	74	146	488	506	. 8.6	19	. 3	24	14	8	
										STRONGE				
	OF	THE	CONDU	CTOR	MAY F	BE DEF	PFR C		E STDE	OF THE	ED L FI	ייז		

			AXIAL 26 HZ		ANAR 74 HZ		ANAR 19 HZ			. HORIZC . SHEE		CONDUC EARI		MAG CORR
								. COND I .SIEMEN		COND D SIEMEN		RESIS OHM-M	DEPTH M	NT
								•	•	•				
	E 20470	•	FLIGHI	-		~~		•					96	~
B		12	19	8	8	33	52		22 .		53	50	26	0
С	1247B?		20	14	27	87	63		18 .		55	56	26	0
D	1233H	1	2	1	1	2	4			• –	-	-	-	0
Е	1222B?		2	1	2	2	4			-	-	-		0
F	1217B?			1	2	2	4			, –	-	-		40
G	1202B?		9	6	11	32	48		31 .	. 2	61	29	35	0
н	1191B?			1	2	2	4			. –	-	-		0
Ι	1172B?	11	28	15	43	88	205		9.		37	39	14	0
J	1156B?	27	52	45	88	253	209	. 4.6	0.	, 3	30	20	11	0
К	1131B?	1	2	1	2	2	4				-	-	-	0
	E 20480	1 (1	FLIGHI	[9]				•	•	•				
A	916H	6		29	4 0	69	71	. 1.6	0	3	37	14	18	0
В	930H	6		16	15	39	34		36		60	36	33	Ō
C	945B?		2	1	2	2		. –	-		_	-	~	Ő
D	943D. 948H	1	2	1	2	2	-		-	_	_	_		Ő
E	940H	12	12	18	21	64	38	•	21	•		11	26	0
				18							44 43	10	20 25	0
F	973H	12	9		14	45	22		23	. 4		-	25 	
G	999B?			1	2	2	4				-	_		0
H	1006B?			1	2	2		• -		, –	-	-		10
I	1016B?			1	2	2		• -		. –	-	-	- 74	0
J	1032H	16		22	21	14	30		19		45	14	24	0
K	1048B?			1	2	2		• -		-	-	-	-	0
L	1052B?			22	73	226	207		0		25	91	0	0
М	1054B?	25	45	41	73	226	207	. 4.9	5	. 2	28	33	7	0
	E 20490) (]	FLIGH	r 9))			•	•	•				
A	878H	16		26	27	90	39	. 7.2	14	4	35	12	17	0
В	869H	15	13	23	24	68	16		27		52	19	30	0
ē	850H	2	11	21	32	44	62		15		51	23	29	0
D	833H	8		19	28	82	45		16		41	12	23	C
Ε	825B?			21	23	60	37		37		58	13	39	C
F	813H	6		16	18	52	59		18		48	18	27	C
G	801B?		32	17	51	179	213		3		29	36	6	Ō
H	791H	1		1	2	2	4		-		_	-	_	Ċ
ï	778B?			37	38	116	97		20	-	45	26	22	c
Ĵ	757B?			46	24	72	31		20		45	15	25	c
K	747B?			21	41	117	156		12		29	46	2.5 4	C
L				21	⁴¹ 2	2	150 4			• 2	- 29	40	4	C
M	743B? 741B?			11	2 33	112	130		- 16	. 2	- 62	- 34	35	0
		•						•		• –				-
LIN	E 20500) (:	FLIGH	Г 9))			•		•		•	•	
A	574H	4	16	13	21	69	53	. 1.3	14	. 3	52	19	31	0
	•		וכו בהחד		URD 17 101	וריונות ה	דרד גד דכ			CUDONA		•		
								E BECAUS						
								R TO ONI				ar .		
	T.T	NE. (UR BEC	CAUSE	OF A	SHAT	OWDI	F OK OM	RRURD	en effex	JIS.	•		

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			XTAL 26 HZ		ANAR 14 HZ		ANAR 9 HZ			. HORIZ . SHE		CONDUC		MAG CORR
	•							. COND I .SIEMEN		. COND : .SIEMEN		RESIS OHM-M	DEPIH M	NI
LINE	20500	(H	LIGHT	9))			•		•				
В	608H	14	14	15	19	53	25	. 7.4	18		45	12	26	30
С	627B?	9	27	18	41	131	152		0	. 2	36	36	12	(
D	637B?	7	22	18	34	107	141		0		56	44	27	(
	657B?		29	53	91	173	184		9		27	12	11	(
F	672B?		2	1	2	2			- 0	· - 2	- 43	- 36	- 18	(
G H	680B? 684B?	5 7	16 16	11 11	23 23	83 83	115 115				43 45	39	10	Č
п I	692H	1	10	1	23	2	115 4	· 2.4	-	· -	-	-	-	(
LINE	20510	(I	LIGHT	9))			•		•				
A	503H	i	2	1	2	2	4		-		-	-	-	(
В	461H	0	9	6	30	110		. 0.4	3	. 2	33	37	10	(
С	442B?	3	9	4	13	39		. 1.3	13	. 2	48	34	23	(
D	419H	1	2	1	2	2	4	• -	_	• -	-	-		(
E	402B?	23	24	45	49	144		. 8.0			38	16	20	(
F G	393B?	14 14	14 14	18	68 68	217 217	97 07		30		32 37	51 48	8 12	(
H	391B? 380B?	14	14 2	18 1	2	217	97 4	· 7.4	25 -	· 2 · -	-	40 -	12 -	
		_						•		•				
	20520	•	LIGHT			_		•		•				
A	200H	11	16	26	26	:7	23				41	10	24	(
B	223H	13	11	29	19	62	11				39	7	22	•
C	237B? 248B?		27 9	19 11	42 14	129 44	54 52	. 2.0	0 19	. 3 . 2	31 32	16 30	11 11	
D E	246D: 253B?		2	1	14 2	44 2	-	· 1.0	-	• 4	- -			
F	294H	5	2 4	62	40	49	108	•		•	25	13	8	1
G	307B?		17	23	33	40	10				43	18	22	4
H	325B?	16	30	26	48	177	197		4	. 2	34	30	11	l
LINE	29010	(H	LICHT	10))			•		•				
Α	672H	1	2	1	2	0	4		-			-	-	(
в	650B?		2	. 1	2	2	4	• -	-		-	-	-	1
С	644B?		18	17	_8	15	39		0	. 1	29	136	0	ł
D	631B?	10	19	20	77	15	189	. 3.3	4	. 1	20	135	0	I
E	569H	1	2	1	2	2	4	• -	-	• -	-	-	-	1
F	560H	1 1	2	1	2	2 2	4	• _	_	• -	-	-	-	I
G H	530H 511H	17	2 12	1 27	2 45	2 15	4 20	. 11.5	- 	• – • 5	- 41	-	- 25	1
н I	497B?		2	27	45 2	15 2	28 4		28		41 -	_′	25	(
J	497B: 493B?		2	1	2	2	* 1	• -	_	•	_	_	-	, I
K	495D: 455H	1	2	1	2	2	4		_		-	-		Ì
	* ES	TIMAT	FED DE	PIH N	ay Bi	e unri	LIABI	e becau	SE THE	STRONG	ER PA	RT.		
								r to on						
	• LI	NE, C	OR BEC	AUSE	OF A	SHALI	OW DI	P OR OV	ERBURD	EN EFFE	CTS.	•		

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COAXIAL 1126 HZ				COPLANAR 874 HZ		COPLANAR 7219 HZ				•	. HORIZONTAL . SHEET		CONDUCTIVE EARTH		MAG CORR	
	OMALY/ /INTERP		· •				-					COND D SIEMEN		RESIS OHM-M	DEPIH M	NT
LIN	E 29020) (I	LICH	r 10))											
A	1023E	ĺ	2	1	2	2	4		-	-	•	-	-		-	40
в	1042B?	34	35	2	76	3	27		9.4	26	•	1	17	78	0	80
С	1049B?	' 1	2	1	2	2	4		-	-		-	-	-	-	0
D	1055B?	28	26	57	71	3	231	•	9.8	28	•	1	24	45	5	30
Е	1061B?	' 1	2	1	2	2	4	•	-	-		-	-	-	-	0
F	1072B?	11	2	12	23	3	154		55.4	49		1	30	69	4	0
G	1106B?	4	54	5	95	4	571	•	0.6	i 0	٠	1	0	288	0	0

* ESTIMATED DEPTH MAY BE UNRELIABLE BECAUSE THE STRONGER PART .

. OF THE CONDUCTOR MAY BE DEEPER OR TO ONE SIDE OF THE FLIGHT .

. LINE, OR BECAUSE OF A SHALLOW DIP OR OVERBURDEN EFFECTS.

