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Corey Project Assessment Report

for the

1994 Exploration Program

Claims;

(Dwayne 1, Carl J, JoJo M, Corey 1-8, 10-12, 14-16, 18-45, Unuk 20 Tine 1, Ginger 1-2, Candy 1 Fr., DEL 1-2, Sul 1-2, Nica 1-2, Cumberland, Silver Pine, Middlesex, Ziphis, Ougma, Dee 1-6)

> Skeena Mining Division NTS 104B/7,8,9,10 Lat. 56 27N, Long. 130 25W

> > for

Kenrich Mining Corp. #1500 - 789 West Pender Street Vancouver, B.C.

Prepared by

Coast Mountain Geological Ltd. 1680 - 650 West Georgia Street Vancouver, B.C.

FILMED

GEOLOGICAL BRANCH ASSESSMENT REPORT

25.15

J. Chapman, P. Geo. January 19, 1995

ARIS SUMMARY SHEET

Off Confidential: 96.02.10 Regional Geologist, Smithers ASSESSMENT REPORT 23757 MINING DIVISION: Skeena **PROPERTY:** Corey LOCATION: LAT 56 30 00 LONG 130 29 00 UTM 09 6262503 408686 NTS 104B08W 104B09W CLAIM(S): Corey OPERATOR(S): Kenrich Min. Chapman, J.; Visser, Syd AUTHOR(S):**REPORT YEAR:** 1995, 60 Pages COMMODITIES SEARCHED FOR: Gold, Silver, Copper, Lead, Zinc KEYWORDS: Triassic-Jurassic, Stuhini Group, Hazelton Group, Bowser Lake Group Tertiary, Intrusives, Anticlinorium, Faults, Metamorphics WORK DONE: Geophysical, Geochemical, Physical 7.4 km;VLF EMGR IPOL 5.2 km LINE 12.0 km Map(s) - 1; Scale(s) - 1:5000MAGG 7.4 km ROCK 8 sample(s) ;ME SILT 6 sample(s) ;ME SOIL 20 sample(s) ;ME

SUMMARY

The 1994 exploration program on the Corey property was originally designed to include cut line grid establishment, geologic mapping, geochemical sampling, geophysical surveys and trenching on the Bench, Battlement and TV zones, to define drill targets. This program was curtailled due to the late season startup and heavy snowfall which occured shortly after mobilization of the crews. Four days of geological mapping were carried out and a total of 34 samples collected before deep snow restricted the work to the geophysical surveys. All work was done on the Bench Zone.

Twelve kilometres of grid lines were established on the Bench Zone to provide control for the IP, magnetometer and VLF-EM surveys, however 3.5kms of these are only flagged as the snow depth would have made cutting the lines of no value.

The limited amount of mapping completed on the Bench Zone confirmed a portion of the '93 mapping by Van Damme and Mosher. A core zone of Mt. Dilworth Formation occurs on the central portion of the grid flanked by Salmon River Formation. Observed were variously altered rhyolite flows, breccias and lapilli to ash tuffs in contact with well laminated black graphitic argillite. "Eskay" type transition breccia was observed on the western portion of the grid, occuring as angular white fragments within a black siliceous matrix.

Geochemical sampling consisted of 20 soil, 6 silt and 8 rock samples, all from the Bench grid, which were sent to Chemex Labs in North Vancouver for 30 element ICP and gold analyses. Weakly anomalous values in gold, silver, antimony, moly and zinc were recorded.

SJ Geophysics Ltd. carried out an I.P., magnetometer and VLF-EM survey, over part of the grid prior to being halted due to dangerous conditions brought on by heavy snowfalls. These surveys produced encouraging results which are detailed in a separate report included as Appendix A.

The program, although abbreviated, did produce significant geophysical anomalies which appear to correlate well with the postulated trend of the Eskay facies rocks, and the 1993 geochemical anomalies. These results warrant a significant exploration and drilling program.

A program of detailed geologic mapping, over an expanded grid, to be followed by additional and closer spaced geophysical surveys is strongly recommended for the Bench Zone. A five hole, 1000m, diamond drill program should be based on the targets derived by this work. This work program should be initiated as early in the season as conditions will allow.

The TV and Battlement Zones require the establishment of a cut grid, detailed mapping and trenching followed by geophysical surveys as noted in the recommendations of the 1993 report by Van Damme and Mosher. On the TV zone this work will provide drill targets for a 1,200 metre diamond drill program which should commence as soon as the targets are defined.

Budgets for the proposed work program are included in this report.

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INTRODUCTION

This report has been prepared at the request of Kenrich Mining Corp. to document the work carried out by Coast Mountain Geological Ltd. on the Corey Property during October/November 1994. An abbreviated grid cutting, mapping, sampling and geophysical program was implemented between Oct. 27 and Nov 14, 1994, on the Bench Zone. This program was terminated prematurely due to an unseasonably early and heavy snowfall. A total of \$227,285.64 was expended during the 1994 work program, which was higher than anticipated due to the adverse conditions.

The major portion of the work consisted of the geophysical surveys carried out by S.J. Geophysics Ltd., consisting of 6 kms. of I.P., magnetometer and VLF-EM over 7 lines covering the centre of the Bench grid. The geophysical report and maps form Appendix A of this report.

Geologic mapping was carried out for 4 days during which time it was possible to confirm part of the work of Van Damme and Mosher (1993). Geochemical sampling consisted of one line of soil samples, 8 rock and 6 silt samples. A program of grid establishment, additional mapping, sampling, and geophysical surveys is recommended for the Bench, Battlement and TV zones during the 1995 field season.

LOCATION AND ACCESS

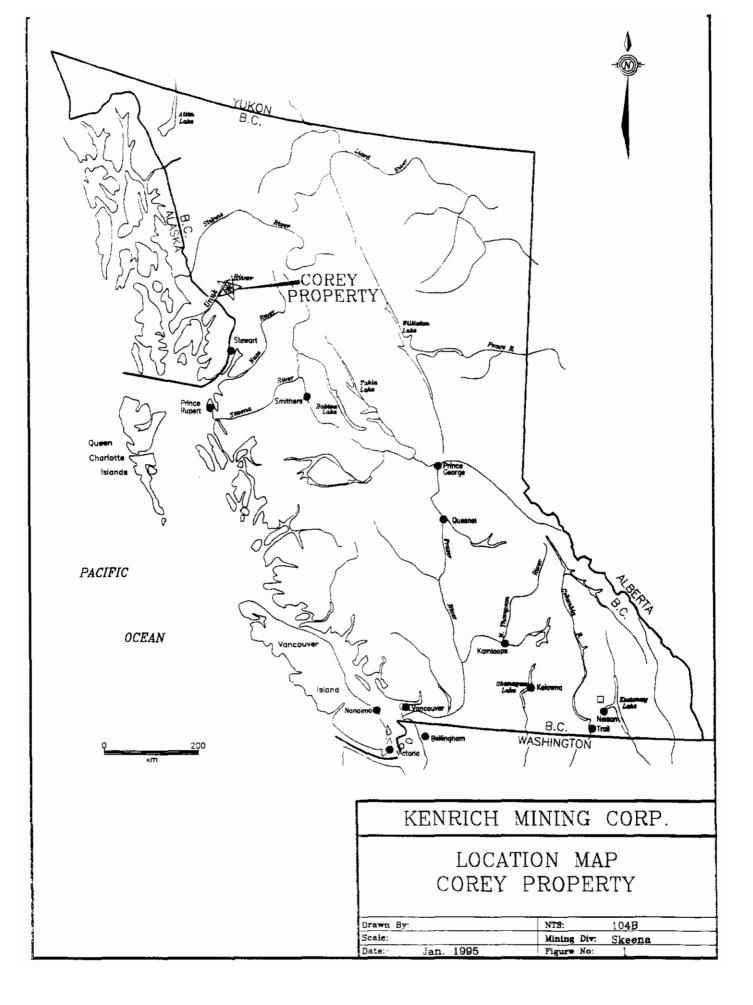
The Corey Property is located in northwestern British Columbia approximately 300kms. northwest of Smithers and 38kms southwest of Bell II on Highway 37 (Fig. 1). A recently completed private road to the Eskay Creek minesite passes within 12kms of the northern boundary of the property. National Topographic Sheets 104B/7,8,9, and 10 encompass the project area which lies within the Unuk River watershed. All of the claims are within the Skeena mining division and are centered at latitude 56 degrees 27 minutes north and longitude 130 degrees 25 minutes west.

Access to the property is by helicopter only, either from Bell II, the Bronson Creek airstrip 35kms to the northwest, the Bob Quinn airstrip 56kms to the north on Highway 37 or from the Eskay Creek road with prior permission. Regular scheduled flights connect Vancouver with Smithers.

PHYSIOGRAPHY AND CLIMATE

The project area is rugged and mountainous, typical of this region of northwestern British Columbia. Elevations range from 250m at the Unuk River to 2,256m at John Peaks, with the intervening slopes varying from moderate to very steep. Glaciers and ice fields occupy the upper reaches of many of the higher valleys, and semi permanent ice and snow may be found on north facing slopes.

Precipitation levels are very high, both as rain in the summer months and snow in the winter. This fosters a dense growth of mature stands of fir, sitka spruce, cedar, hemlock, aspen and alder in the valley bottoms, generally with a thick understory of ferns, salmonberry, huckleberry and devils club.



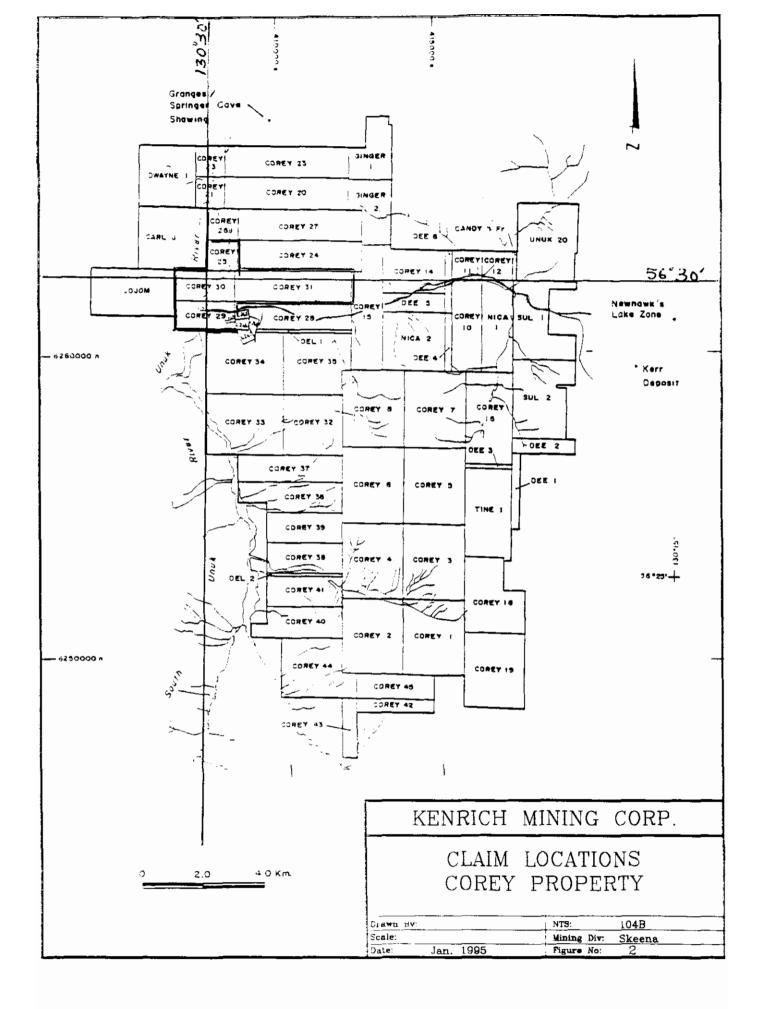
Treeline occurs at around 1,200m above which slide alder and related shrubbery gives way to grasses, mosses and lichen in the alpine areas.

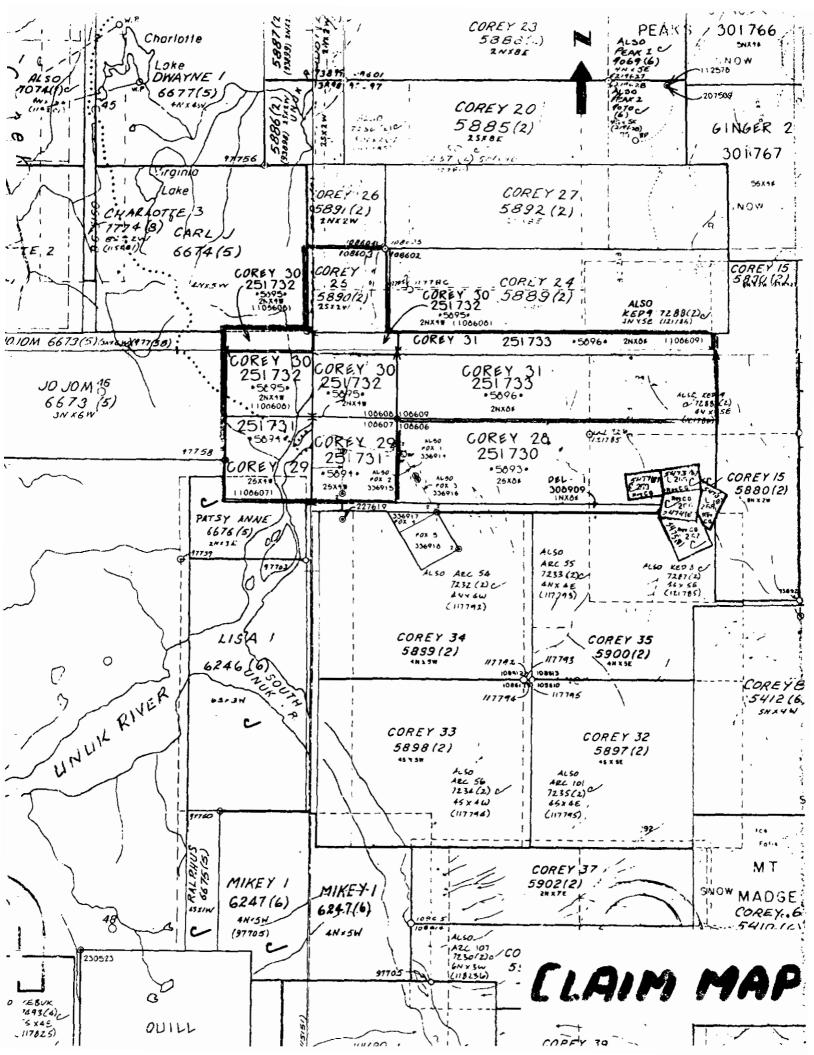
PROPERTY DESCRIPTION

The Corey property consists of 67 contiguous mineral claims and 5 reverted crown grants, totaling 80,000 acres, located within the Skeena Mining Division (Fig. 2). Kenrich Mining Corporation is the registered owner of the property and they maintain offices at #1500 - 789 West Pender Street, Vancouver, B.C.. A complete description of the claim data and status is presented in Table 1 from the 1993 Exploration Report by Van Damme and Mosher.

TABLE 1

NAME	NUMBER	UNITS	RECORD DATE	EXPIRY DATE
Dwayne 1	97756	16	May 13, 1988	May 13, 1995
Carl J	97757	20	May 13, 1988	May 13, 1995
Jo Jo	97758	18	May 13, 1988	May 13, 1995
Corey 1	251446	20	June 25, 1986	June 25, 1996
Corey 2	251447	20	June 25, 1986	June 25, 1996
Corey 3	251448	20	June 25, 1986	June 25, 1996
Corey 4	251449	20	June 25, 1986	June 25, 1996
Corey 5	251450	20	June 25, 1986	June 25, 1996
Corey 6	251451	20	June 25, 1986	June 25, 1996
Corey 7	251452	20	June 25, 1986	June 25, 1996
Corey 8	251453	20	June 25, 1986	June 25, 1996
Cory 10	251714	12	February 11, 1987	February 11, 1997
Corey 11	251715	4	February 11, 1987	February 11, 1997
Corey 12	251716	4	February 11, 1987	February 11, 1997
Corey 14	251717	12	February 11, 1987	February 11, 1997
Corey 15	251718	16	February 11, 1987	February 11, 1997
Corey 16	251719	18	February 11, 1987	February 11, 1997
Corey 18	251720	20	February 11, 1987	February 11, 1996
Corey 19	251721	20	February 11, 1987	February 11, 1996
Corey 20	251722	16	February 11, 1987	February 11, 1996
Corey 21	251723	4	February 11, 1987	February 11, 1996
Corey 22	251724	4	February 11, 1987	February 11, 1996
Corey 23	251725	16	February 11, 1987	February 11, 1996
Corey 24	251726	16	February 11, 1987	February 11, 1996
Corey 25	251727	4	February 11, 1987	February 11, 1996
Corey 26	251728	4	February 11, 1987	February 11, 1996
Corey 27	251729	16	February 11, 1987	February 11, 1996
Corey 28	251730	16	February 11, 1987	February 11, 1996
Corey 29	251731	8	February 11, 1987	February 11, 1996
Corey 30	251732	8	February 11, 1987	February 11, 1996





NAME	NUMBER	UNITS	RECORD DATE	EXPIRY DATE
Corey 31	251733	16	February 11, 1987	February 11, 1996
Corey 32	251734	20	February 11, 1987	February 11, 1996
Corey 33	251735	20	February 11, 1987	February 11, 1996
Corey 34	251736	20	February 11, 1987	February 11, 1996
Corey 35	251737	20	February 11, 1987	February 11, 1996
Corey 36	251738	14	February 11, 1987	February 11, 1996
Corey 37	251739	14	February 11, 1987	February 11, 1996
Corey 38	251740	12	February 11, 1987	February 11, 1996
Corey 39	251741	12	February 11, 1987	February 11, 1996
Corey 40	251742	12	February 11, 1987	February 11, 1996
Corey 41	251743	12	February 11, 1987	February 11, 1996
Corey 42	251744	5	February 11, 1987	February 11, 1996
Corey 43	251745	4	February 11, 1987	February 11, 1996
Corey 44	251746	20	February 11, 1987	February 11, 1996
Corey 45	251747	10	February 11, 1987	February 11, 1996
Tine 1	252211	18	February 10, 1989	February 10, 1997
Ginger 1	301766	20	June 26, 1991	June 26, 1996
Ginger 2	301767	20	June 26, 1991	June 26, 1996
Candy 1 Fr	303817	1	Sept. 10, 1991	Sept. 10, 1997
DEL 1	308909	8	April 16, 1992	April 16, 1996
DEL 2	308910	5	April 16, 1992	April 16, 1996
Cumberland (L265)	251492	1	August 01, 1985	August 01, 1997
Silver Pine (L266)	251493	1	August 01, 1986	August 01, 1997
Middlesex (L267)	251494	1	August 01, 1986	August 01, 1997
Ziphis (L268)	251495	1	August 01, 1986	August 01, 1997
Ougma (L269)	251496	1	August 01, 1986	August 01, 1997
Sul 1	251348	20	February 28, 1986	February 28, 1997
Sul 2	251349	20	February 28, 1986	February 28, 1997
Unuk 20	251377	20	February 28, 1988	February 28, 1997
Nica 1	252209	12	Sept. 10, 1988	Sept. 10, 1997
Nica 2	252210	16	Sept. 10, 1988	Sept. 10, 1997
Dee 1	253609	5	February 18, 1990	February 18, 1997
Dee 2	253610	4	February 18, 1990	February 18, 1997
Dee 3	253611	3	February 18, 1990	February 18, 1997
Dee 4	253612	4	February 18, 1990	February 18, 1997
Dee 5	253613	8	February 18, 1990	February 18, 1997
Dee 6	253614	4	February 18, 1990	February 18, 1997

PREVIOUS EXPLORATION

The earliest work conducted on what is now the Corey Property was the staking and excavation of two adits on the Cumberland group of claims between 1898 and 1903. A shipment of hand-cobbled ore is reported to have been made during the 1930's.

Only limited exploration was carried out within the area until the 1960's when a regional survey was conducted by Newmont during which time the Ox and Fox Claim Groups were staked, surrounding the earlier Cumberland crown grants. Up to 1983, the area south of Sulphurets Creek saw a series of small exploration programs conducted by E and B Explorations, Nor-Con Explorations and Dupont Canada.

In 1986 Catear Resources Ltd. staked the Corey 1-8 claims and conducted a program of rock and silt geochemistry and prospecting. At the same time Skelly Resources Ltd. staked the Sul 1-2 and Unuk 20 claims.

Bighorn Development Corp. optioned the Cory property in 1987 and subsequently staked an additional 516 claim units, Corey 10-45. A property wide program of silt, soil and rock geochemistry, prospecting and detailed evaluation was completed. Detailed work consisted of geological mapping, 40 metres of trenching and 590 metres of diamond drilling in six holes at the Cumberland prospect. During this period Bel Pac Industries Ltd. acquired the Sul 1-2 and Unuk 20 claims.

In 1988 Bighorn carried out a follow up program and completed 647 metres of diamond drilling n six holes on the C-10 prospect. At this time Kenrich Mining Corp., formerly Farquest Energy Corp., optioned the Sul 1-2 and Unuk 20 claims. Also Ambergate Explorations Inc., formerly Nica Ventures Inc., acquired the Nica 1 claim.

1989 saw Kenrich and Ambergate conduct geological and geophysical surveys on the combined claims.

During 1990, Ambergate drilled two holes totalling 86 metres on the Nica 1 and Kenrich drilled seven diamond drill holes totalling 476.4 metres on the Unuk 20 claim. The latter part of the '90 saw Kenrich-Ambergate augment their property holdings with the acquisition of the Corey 1-8 and Corey 10-45 claims.

In 1991 Placer Dome optioned the Sul 1-2, Nica 1 and Unuk 20 claims from Kenrich-Ambergate. An exploration program of geological mapping, geochemical sampling and ground geophysics was completed. Placer also evaluated the Cumberland and C-10 prospects at this time.

In 1992 Placer Dome carried out an extended program of geochemical, geophysical and diamond drilling on the option. The rest of the property underwent varying degrees of exploration or review by Kennecott Canada Inc., Inco Exploration and Technical Services Inc., and Homestake Canada Ltd. This work consisted primarily of reconnaissance geochemical and geological surveys.

In 1993, with the completion of an extensive geological, geochemical and limited geophysical and trenching program Kenrich and Ambergate further expanded the property's limits by purchasing the Swayne 1, Carl, and Jo Jo claims.

REGIONAL GEOLOGY

The Corey Property is located within supracrustal rock units of Stikinia. Stikinia is a terrane block of the Northwestern Cordillera, which has four tectonostratigraphic assemblages bounded by unconformities. These include the Palaeozoic aged Stikine Assemblage, several Triassic to Jurassic volcanic-plutonic arc complexes, the middle to late Jurassic Bowser overlap assemblage and the Tertiary Coast plutonic complex.

Stikine Assemblage rocks are exposed southwest of the property. They consist of Mississippian or older mudstone, chert, wacke, and feldspar-phyric volcanic flows and pyroclastic rocks.

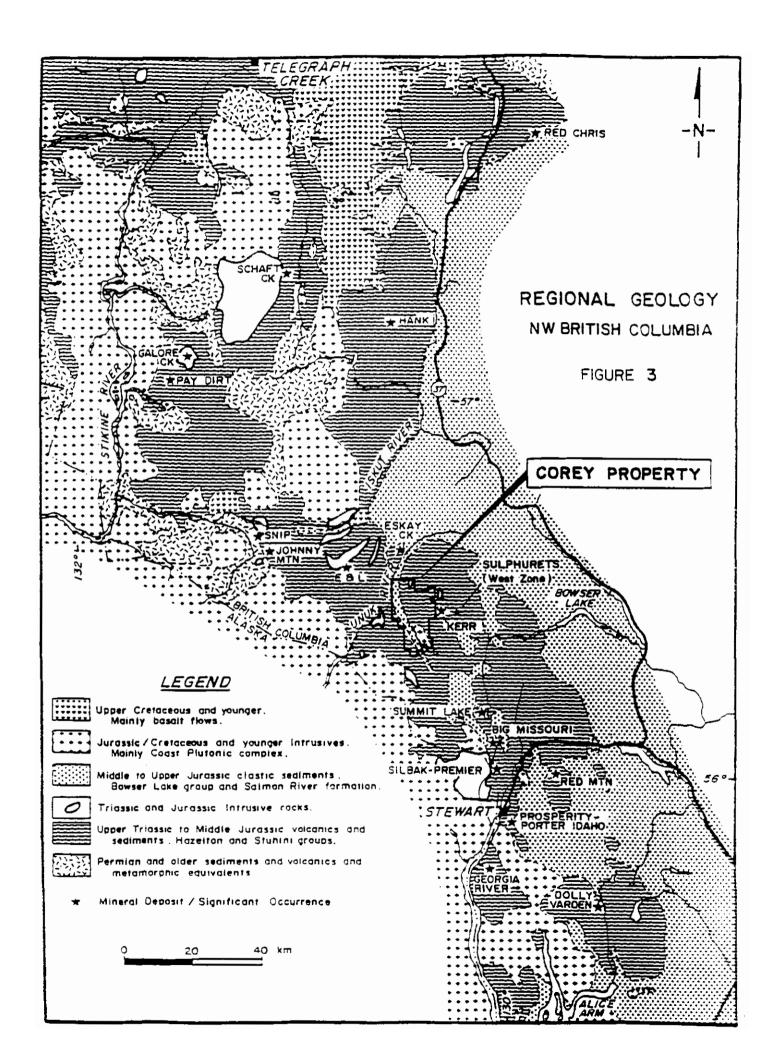
Five lithostratigraphic packages are recognized within the Triassic to middle Jurassic arc complex. Stratigraphic nomenclature (subject to revision) includes the older Stuhini Group and a younger Hazelton Group. The Hazelton is in turn subdivided into four units assigned formation status, including the Unuk River, Betty Creek, Mount Dilworth, and Salmon River Formations. Hazelton Group rocks host the Eskay Creek deposits. Aggregate package thickness may exceed 5,000 metres.

Stuhini Group units underlie the eastern portion of the property and extend north to Treaty Creek. Thick-bedded lithic wacke, limestone, pyroxene-phyric volcanic breccia and locally occurring granite cobble conglomerate are dominant rock types. Assigned ages are constrained by the rare occurrence of Carnian and Norian index fossils and radiometric dating of intrusive rocks. In the field these rocks are generally weakly to strongly foliated and of lower to middle greenschist metamorphic rank, marked by the presence of chlorite and epidote.

The oldest member of the Hazelton Group, the Unuk River Formation is composed of thickbedded, monotonous sequences of andesite pyroclastic and flow units interbedded with minor tuffaceous wacke and conglomerate. A distinguishing attribute is the presence of orthoclase megacrystic units, frequently referred to as "Premier Porphyry".

Betty Creek Formation units are composed of variably maroon to green-hued andesitic to dacitic pyroclastic breccia, tuff and flows, minor pillowed andesite flow and flow breccia, and relatively common interflow units of volcaniclastic grit, wackestone and siliceous mudstone. The unit is moderately well constrained by index fossils. The presence of welded-appearing ash tuff within the Betty Creek, along with the commonly reddish-hued aspect of rock members, has been cited as evidence supporting a subaerial depositional environment. However, in the immediate Eskay Creek area numerous occurrences of marine fossils have been discovered, suggesting a subaqueous environment existed at least locally.

The Mount Dilworth formation is an extensive, thin unit of grey to white-weathering dacite and rare rhyolite ash and lapilli tuff and breccia, plus feldspar-phyric flows and intrusive dykes. Interflow sedimentary rocks are uncommon, and the unit is remarkably consistent in appearance. It has been dated by radiometric techniques at approximately 190 million years, and is well constrained by Toarcian-aged (Middle Jurassic) fossils found in bounding rock units.



At Eskay Creek, felsic volcanic units thought to be equivalent to the Mount Dilworth occur deep in the deposit footwall, and locally are important mineralized host-rocks.

The Salmon River Formation is a dominantly epiclastic sequence composed of grey to browncoloured wacke, mudstone and conglomerate overlying a thin, but persistent calcareous wacke which is remarkably rich in fossil debris (particularly bellemnites and bivalves). At Eskay Creek, pillowed andesite flows and breccia and massive flow-banded rhyolite and autoclastic breccia are abundant in the lower Salmon River. The unit is well-constrained by fossils, assigned to a Toarcian to Bajocian age. The rhyolite is dated at approximately 170 million years. Mapping by the Geological Survey of Canada has further subdivided the Salmon River Formation into three north-trending facies belts: the eastern Troy Ridge facies, the central Eskay Creek facies, and the western Snippaker Mountain facies. The Troy Ridge facies is distinguished by the presence of rhythmically alternating thin-bedded black chert and light-coloured tuff. The Eskay Creek facies comprises calcareous and siliceous mudstone and basaltic to andesitic pillow lava. The relative abundance and thickness of volcanic material increases north of Eskay Creek. The Snippaker Mountain facies comprises plagioclase feldspar and hornblende-phyric flows and breccia interbedded with impure limestone, calcareous conglomerate and sandstone.

The Bowser overlap assemblage comprises a thick, regionally extensive epiclastic sequence of Middle and Upper Jurassic age. It is composed of thick and thin-bedded wacke, mudstone and conglomerate. Ammonite fossils of Bathonian age have been found in the Eskay Creek area, suggesting unit equivalency to the Bowser Lake Group/Ashman Formation. Bowser Lake units are remarkable monotonous and have been little-studied in the past. Separation from the underlying Salmon River Formation epiclastic units is best effected on the basis of fossils, as they are similar appearing.

The Tertiary Coast plutonic complex, composed principally of large bodies of Eocene-aged granitic rocks, outcrops to the south.

The Mesozoic volcanic and sedimentary rocks of the Stuhini and Hazelton Groups are considered to represent an arc-building sequence, accreted onto the adjacent Cache Creek terrane (to the east) by Middle Jurassic time. The geological environment present in the Eskay Creek area at the time of ore formation was a back-arc basin, marked by a north-trending graben.

Other features of regional interest, not necessarily incorporated into the Stikinia regional framework, include Late Triassic quartz diorite stocks and Jurassic diorite stocks and sills. Pleistocene to Recent basalt flows and tephra fields are relatively common in the region, particularly in valley bottoms. These have been dated as young as 2,600 year B.P.

Two structural elements dominate the region. To the west is a pronounced lineament, the Harrymel Fault Zone, which separates relatively undisturbed Mesozoic units in the west from highly folded and faulted identical rocks to the east. The Harrymel Fault displays evidence of increasing displacement to the south. North and east of the Corey Property is the McTagg Anticlinorium, a major north-plunging fold structure. Rocks east of the main anticline axis are thrust and folded to the east. West of the axis, including rocks of the Corey property, units are folded and thrust to the west. It

appears that both the Harrymel Fault Zone and the McTagg Anticlinorium are contemporaneous, having been formed during the Cretaceous Skeena Orogeny.

Deformation and metamorphic rank are variable. Stikine Assemblage rocks, ranging from schists to gneisses, exceed the biotite isograd. Stuhini Group units are variably schistose, have sausseritized feldspars and chlorite-epidote stable assemblages, suggesting at least lower greenschist rank metamorphism. Hazelton and Bowser Lake Group rocks are variably folded, and appear fresh and unmetamorphosed. Prehnite is found in the Salmon River Formation units immediately overlying the Eskay Creek deposits, and these rocks may be of prehnite - pumpellyite rank. In general metamorphic rank appears to reflect both the age of any particular unit and its proximity to an intrusive body, while deformation is most notable along zones of faulting.

Eskay Creek Deposit Summary

The Eskay Creek Deposit, owned by Prime Resources Inc., is the most important mineral deposit in the region and is located ten kilometers north of the Corey Property. Stated mineable reserves at Eskay are 1,200,000 tons grading 1.91 ounces per ton gold and 85..5 ounces per ton silver, with significant by-product zinc, lead and copper. The project has recently commenced production, and when in full production, it will yield an annual average of 240,000 ounces of gold and 10.5 million ounces of silver over an eight year mine life. Eskay will be the third largest silver producer in the world.

The mineralization at Eskay Creek occurs in Jurassic Salmon River Formation volcanic and sedimentary rocks, within a zone of complex faulting and folding. Footwall to the ore zone is the "Eskay Rhyolite Unit", up to 80 metres thick, which is massive, flow-banded, locally autobrecciaed, extensively altered and weakly mineralized (Sherlock et al, 1994). Overlying the rhyolite is up to 4 metres of "Transition Breccia", a polymictic fragmental unit with rhyolite, mudstone and massive sulphide clasts. Overlying the breccia is the "Contact Unit Mudstone", the main ore host, a massive-appearing, medium to thick-bedded black tuffaceous mudstone which is from 2 to 20 metres thick. Hanging-wall to the mudstone unit is a thick sequence of pillowed basalt, flow breccia and numerous sills. Basalt flows are intercalated with mudstone beds. Mudstone is the dominant rock type both upsection and laterally from the deposit area. Deeper in the footwall, units equivalent to the Mount Dilworth are present, including dacite flows, flow breccia, lapilli tuff and minor mudstone.

The economic orebody is hosted entirely by the Contact Mudstone. It is 600 metres long, 60 to 120 metres in dip extent, and from 2 to 12 metres thick. The ore minerals include sphalerite, tetrahedrite and electrum, with lesser chalcopyrite, boulangerite, bournonite, galena, pyrite and stibnite. Mineralization occurs as bedding-parallel layers and lenses, characterized by clastic, graded and slumped bedforms. The ore zone is considered to have formed through syngenetic processes, and can be classed as a volcanogenic massive sulphide deposit.

PROPERTY GEOLOGY

The 1993 report by Van Damme and Mosher details the geology of the Corey property. As no significant mapping was accomplished in 1994 the reader is referred to that report for a full treatment of

the property geology. That portion of the property geology which relates to the Bench, Battlement and TV zones, the north half, is reproduced here.

Argillite and siltstone crop out along the east bank of the Unuk River in the vicinity of the Bench and Battlement zones. These rocks lack diagnostic strata which would allow their assignment to either the Bowser Group or the Salmon River Formation, but they considered to belong to the Salmon River Formation.

To the east the geology is dominated by a thick package of mafic volcanic rocks which can be assigned to the Salmon River Formation. In their western exposures these rocks are dominated by massive chloritized and saussuritized flows, and yield eastward to pillowed flows carrying pyrite/pyrrhotite and calcite/chlorite amygdules. Pillows are frequently blocky and doubly cusped. Tops are up and east facing. Intercalated sediments, typical of the Salmon river Formation are notably absent in these rocks. Outcrop exposure may be limited by the recessive weathering nature of the sediments.

Rocks of the Unuk River Formation occur east of, and in fault contact with the Salmon River Formation. These rocks form a thick homogenoeous assemblage of andesitic ash, lapilli, and crystal tuff regularly interbedded with argillite and siltstone. In the central part of the area the Unuk River Formation units appear to interfinger the Salmon River Formation units. This may suggest a complex interaction of faults in juxtaposing the two formations.

Rocks of the Betty Creek Formation occur further east of, and upsection of the Unuk River Formation. At the base is highly contorted and drag-folded andesitic tuff, argillite, siltstone, and sandstone. Overlying these are heterolithic dacitic tuff, consisting of massive white to green angular lapilli, a high (>15%) lithic component, and brown weathering fiamme.

Mount Dilworth Formation rhyolites overlie Betty Creek Formation rocks to the east, forming a thin but continuous band from the western edge of John Peaks to Sulphurets Creek. It consists of two principal lithologies: a breccia with 10 cm to 3 m sized blocks and a welded ash flow tuff. Both would be assigned to a medial facies of the Mount Dilworth Formation elsewhere in the belt.

The Mount Dilworth Formation is succeeded upwards by a discontinuous calcareous, fossiliferous basal mudstone of the Salmon River Formation. Deformed fossil-rich units were observed. Also overlying the rhyolites are discontinuous, massive, mafic flows which show chilled margins in contact with interbedded thin calcareous wacke beds.

Rocks of the Betty Creek Formation overlie the Salmon River Formation further to the east. This repetition is likely due to folding. Rock units consist of bright green felsic siltstone alternating with red siltstone and argillite. Sedimentary textures are abundant, and include scour and fill channels, graded bedding, and climbing ripple marks.

A single noted occurrence of Jack Formation, here a granitoid pebble conglomerate, occurs at the extreme eastern limit of mapping. This occurrence is thought to be separated by a reverse fault from the Salmon River Formation.

The detailed geology of the Bench and Battlement Zones is shown on Map 4. Both Salmon River Formation rhyolite and basalt are present. Here, Salmon River rhyolites are distinctly different from Mount Dilworth Formation units of the John Peaks area, reflecting a more proximal facies depositional environment. In the central portions of the rhyolite it is flow-banded, brecciated, and flow folded. Autobrecciation appears to increase outward, as does the amount of silicification and presence of a perlitic texture. To the west, the rhyolite upper contact is marked by a distinctive "Eskay" type transition breccia in which massive, angular, white fragments are contained within a black siliceous matrix. Overlying the transitional breccia is a distinctive, thin but consistent, sedimentary rock comprised of alternating beds of cherty argillite and tuffaceous siltstone. Both the transition breccia and argillite are mineralized with arsenopyrite, sphalerite, galena, and pyrite as clasts, veins, and matrix in-fillings.

These rocks are in turn overlain by massive and pillowed mafic volcanic rocks with lesser intercalated and recessively weathering argillite, which yield abruptly to the north to sedimentary rocks. It is considered that the abruptness reflects the boundary between the lower flow dominated, and upper sediment-dominated members of the Salmon River Formation.

The detailed geology of the Battlement Zone consists of Salmon River rhyolites, lapilli tuff and tuff breccia with rare lithic fragments. These are overlain by sedimentary rocks that grade from cobble and boulder conglomerates to graphitic, tuffaceous argillite. At the western limit of mapping the rhyolites are in contact with massive, pillowed, and pillow breccia flows of the Salmon River Formation. In contrast, rhyolites at the northern and eastern limits of mapping are in contact with massive argillite of the Salmon River Formation.

At the Cumberland showing recently identified rocks include rhyolite breccia with black matrix which had previously been mapped as conglomerate, and a previously mapped sandstone was found to be well-bedded ash tuff containing small felsic fragments. These rocks are very similar in appearance to Salmon River Formation rhyolites found on the Bench Zone.

Intrusive Rocks

Apart from minor basaltic to andesitic dykes and sills present in the Salmon River Formation rocks, the only major intrusive bodies observed were the extensive John Peaks diorite, and the Lee Brant stock. A possible intrusive unit has also been identified in the Betty Creek Formation.

The John Peaks intrusive occurs along the northeast Corey Property boundary. It is a hornblende diorite which separates Triassic sediments to the east from upper Hazelton Group rocks to the west, and is tentatively correlative with rocks of the Texas Creek plutonic suite.

The Lee Brant stock has been mapped as a hornblende, biotite, microcline-bearing porphyritic quartz monzonite. It is of Tertiary age and forms a part of the Coast Plutonic Complex.

Of uncertain origin are possible subvolcanic dykes occurring at the base of the Betty Creek Formation. These rocks, apart from rare remnant feldspar phenocrysts, are massive and uniform in appearance. The protolith is obscured by intense silica alteration and pyrite. There is a strong similarity between these rocks and Mount Dilworth Formation subvolcanic dacite sills at Eskay Creek.

Structural Geology

Graded bedding preserved in sediments, welding in ash flow tuff, and top determinations in pillowed basalts allow the generalization that the stratigraphy in the area trends north to northwesterly and dips moderately to the east. Facing directions are to the east, except in those major drainages interpreted as reverse-normal, to high-angle thrust faults. In these areas "drag" folds are common. At the eastern limit of mapping intrafolial folding has been interpreted in the Jack Formation. In the Bench Zone area, a north-plunging syncline has been documented, possibly being a synclinal crest on a much larger anticlinal arch.

Predominant linear structures present are vertical to steeply-dipping faults that strike 010° to 020°. Abnormal apparent unit thicknesses may be due to structural repetition. Several crosscutting fault sets were noted, one is northwesterly-trending with moderate northeast to vertical dips. A possible conjugate fault set strikes northeast and dips moderately to the southeast. The two are responsible for the mapped horizontal offsets of the north-trending stratigraphy.

East striking, vertical to steeply north dipping faults are marked by watercourses such as Sulphurets Creek.

The dominant terrain-forming structures are high angle thrust or reverse normal faults. These generally trend north-northwest and dip up to 60 degrees. They commonly coincide with, or possibly ramp along sedimentary units. At least three section repetitions occur on the Corey Property because of thrusting.

1994 EXPLORATION PROGRAM

Crews were mobilized to the Corey property on October 27, 1994 to commence a program designed to define targets for a diamond drill program in early 1995. This work was to consist of establishment of a cut grid, detailed geological mapping, sampling, geophysical surveys and trenching on the TV, Bench and Battlement Zones. These targets had been located during the 1993 work program. Due to the late start up date the TV Zone, at approximately 800m, was already covered by over a foot of snow and could not be safely or efficiently worked.

Both the Bench and Battlement zones, at river level, were still free of snow at the time of mobilization. The Bench Zone was chosen as the higher priority target of the two and grid cutting commenced immediately. By Nov. 3 deep snow had terminated the mapping and sampling portions of the program, however the grid cutting was continued to provide control for the geophysical surveys. A total of 6kms. of combined mag, VLF-EM and IP were completed, over 7 lines, covering the central portion of the Bench grid. These surveys were ended on Nov. 14, due to excessive snow cover and dangerous conditions on the steep slopes.

On the Battlement Zone only the start point for a continuation of the Bench grid was located, no further work could be completed due to the adverse conditions.

DISCUSSION

Correlating the 1993 anomalous soil geochemical results, on the Bench and Battlement zones, with the mapped geology and the 1994 geophysics reveals a number of relationships as shown on Figure 4. Arsenic and antimony values commonly show a close spatial relationship both to each other and to north to northwesterly trending fault zones. This is demonstrated quite distinctly on the eastern side of the Bench grid and the western side of the Battlement grid, where the anomalies cluster along the trace of fault zones. Both of these elements are good indicators of potential Eskay type mineralization possibly leaking up a fracture system.

The most significant geochemical anomaly occurs on the Bench Zone where a linear 600m long zinc high, locally up to 150m wide, trends northeasterly across the central portion of the grid, from line 46+00N/49+00E through line 52+00N/50+00E. This coincides with a geologically and geophysically complex zone on the west limb of what is thought to be a northerly plunging syncline. Smaller lead, arsenic and antimony anomalies are associated with this zinc anomaly. An IP survey line at the south end of this zone indicates a section of rapidly alternating electrical response in the underlying rocks. This would be consistent with a sequence of interbeded argillites, tuffs and possibly rhyolites which have been mapped on lines 44+00N and 45+00N, at 48+00E. A zone of very high resistivity and low chargeability flanks the zinc anomaly to the west, which changes rapidly to a region of high chargeability and low resistivity to the east of the anomaly. The VLF-EM survey also recorded a significant response across this trend.

On the eastern side of the Bench grid, between 52+00E and 53+00E on lines 45+00N through 49+00N, both the VLF-EM and the IP surveys display anomalous responses over the northerly trending contact between the Mt. Dillworth and the Salmon River formations. Geochemically this zone is also anomalous with high values in lead, zinc, silver, arsenic and antimony, showing a close spatial relationship to the contact. This same pattern is repeated on the Battlement grid where almost all of the anomalous geochemical results are located along the same contact. The only exception on the Battlement zone is a cluster of copper, arsenic, antimony and silver anomalies which occur along the northerly trending fault on the west side of the grid.

The highest arsenic and antimony values are associated with north to northwesterly trending fault zones on the eastern portion of the Bench grid. In the vicinity of line 48+00N at 52+00E an EM anomaly correlates with a combined arsenic, antimony, lead and silver anomaly, and a mapped fault zone. A very similar association is evident on the Battlement grid where anomalous copper, arsenic, antimony and silver zones are associated with a north trending fault on the west side of the grid.

VLF-EM data from the southwestern corner of the Bench grid shows an anomaly very similar to that over the Mt. Dillworth - Salmon River formation in the northeast. This occurs at 46+00E on lines 42+00N and 43+00N, and may be offset to 45+25E on line 44+00N. The anomaly appears to continues to the south past line 42+00N, which was the limit of the survey. Mapping to date has not explained this feature. The IP survey extended only to the edge of this zone but the last reading on line 44+00N did indicate that a change was taking place.

The magnetic response over the Bench grid is generally flat for the area surveyed, with the exception of line 42+00N between 52+00E and 54+00E where a sharp spike is observed. This is likely due to a buried intrusive or dyke like body, however this area has not yet been mapped.

CONCLUSIONS

The 1994 exploration program on the Corey property examined only the Bench Zone, which was located during the 1993 work program along with several other high priority targets including the TV, Battlement, MM and GFJ showings. All information and recommendations regarding these other targets is excerpted from the 1993 Exploration Report on the Corey Property, by Van Damme and Mosher.

The Bench Zone contains a number of features which indicate that it may be condusive to hosting a significant mineral deposit. Highly prospective Eskay type stratigraphy underlies the grid area and appears to extend north on to the Battlement zone and south to the Cumberland showing, which has reported assays up to 0.27ozs./ton gold and 9.8% zinc. Structural complexity, necessary for both ground preparation and fluid movement, is evident at outcrop scale and on a more regional level related to the Unuk River fault system. From the Cumberland to Battlement zone constitutes a 4.5 km long belt of highly prospective but largely unexplored geology.

Soil geochemistry from the 1993 survey on the Bench zone has revealed a suite of anomalous elements including arsenic (1513ppm), antimony (956ppm), copper (211ppm), lead (171ppm), zinc (2134ppm) and silver (13.1ppm). This suite is very similar to those that occur at the Eskay deposit. Rock sampling during the 1993 field program also located low grade but significant values in silver (up to 63.6ppm), zinc (up to 7970ppm), arsenic (up to 1915ppm) copper and antimony.

Comparing the 1993 soil geochemistry over the Bench Zone with the geology reveals a strong correlation between anomalous arsenic and antimony values and structural features interpreted from air photos.

A 600m long zinc anomaly with related arsenic, lead and antimony highs, is located approximately parallel to the axis of the northerly plunging syncline which occupies the west central portion of the Bench grid. This geochemical anomaly correlates well with IP and VLF-EM conductors which trace a zone of complex geology on the west limb of the syncline. The VLF-EM anomaly located at the baseline on line 50+00N correlates well with the northern end of the elongate zinc anomaly. The EM anomaly at the western end of lines 45+00N through 42+00N may possibly be an extension of this zone as it continues the same trend.

The geology of the western limb of the syncline may be more complex than presently shown. The IP results indicate a zone of rapidly alternating electrical response in the rocks of this area which extends from line 44+00N through line 48+00N, between 46+00E at the southern end and 50+00E at the northern end, and possible further. A sequence of interbedded rhyolites, argillites and tuffs, as exists at Eskay, could produce this type of response. Significantly this correlates in part with the elongate zinc anomaly. Both the VLF-EM and the IP surveys targeted the Mt. Dillworth formation eastern contact as an area of interest.

RECOMMENDATIONS

Based on the results obtained from the 1994 and 1993 field work the Bench Zone warrants and requires further work. This should take the form of completion of the grid cutting and continuation of the 1:2000 scale mapping which was begun in 1994. Upon completion of the mapping program the remainder of the grid area should be covered by the IP, EM and magnetometer surveys. Following this, expansion of the grid to the south and north will be necessary to tie in the Cumberland showing and the Battlement zone. Results of this work will provide the data necessary to locate drill targets. As this zone is located at low elevation it will be the first area accessible in the spring, therefore the 1995 exploration program should commence as soon as the snow clears from the Bench grid.

On the TV zone a minimun of 6kms. of cut grid are required to provide control for detailed 1:2000 scale mapping, IP, EM, and magnetometer surveys. Additional soil sampling is required to infill and expand the current data base. Trenching, where feasible, and washing of existing outcrop will provide a greater understanding of the geology and more geochemical sampling opportunities. The results of this work will supply the data required to accurately locate a minimum of 8 diamond drill holestotaling 1200metres.

The Battlement Zone should be considered as an extension of the Bench grid, and tied in to the same surveys. This would include mapping at 1:2000 scale, IP, EM and magnetic surveys, soil sampling and trenching. Diamond drill testing of anomalies on this grid should be planned pending positive results of the Bench drilling.

Similarly, the Cumberland Showing to the south of the Bench Zone, should be tied in to the Bench grid and the 1:2000 scale mapping expanded to cover the showing and vicinity. Reconnaissance mapping to determine the extent of the favourable units, and expansion of the soil sampling coverage are reccomended.

The GFJ and MM showings will require detailled mapping, prospecting and contour soil sampling to be followed by hand trenching. This will necessarily be conducted in late summer due to the high elevation of the showings.

A program of property wide reconnaissance mapping and prospecting should be continued, with emphasis on the region to the south of Sulphurets Creek and to the west of the Bench and Battlement zones.

Prior to the 1995 field season the existing database for the Corey project should be digitized. This would allow for a much faster and simpler integration of current and future information. All of the 1995 geophysical data for the Bench Zone is presently on disk and has been integrated with a simplified geologic map.

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Consideration should be given to establishing a base camp on the Corey property, as opposed to renting the Granges camp. A suitable site exists at the Bench zone which would have the advantage of allowing foot access to at least one of the primary targets.

Costs for the above mentioned work program, if completed in its entirety including a Phase II drill program, would be approximately \$1,300,000.00

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STATEMENT OF EXPENDITURES

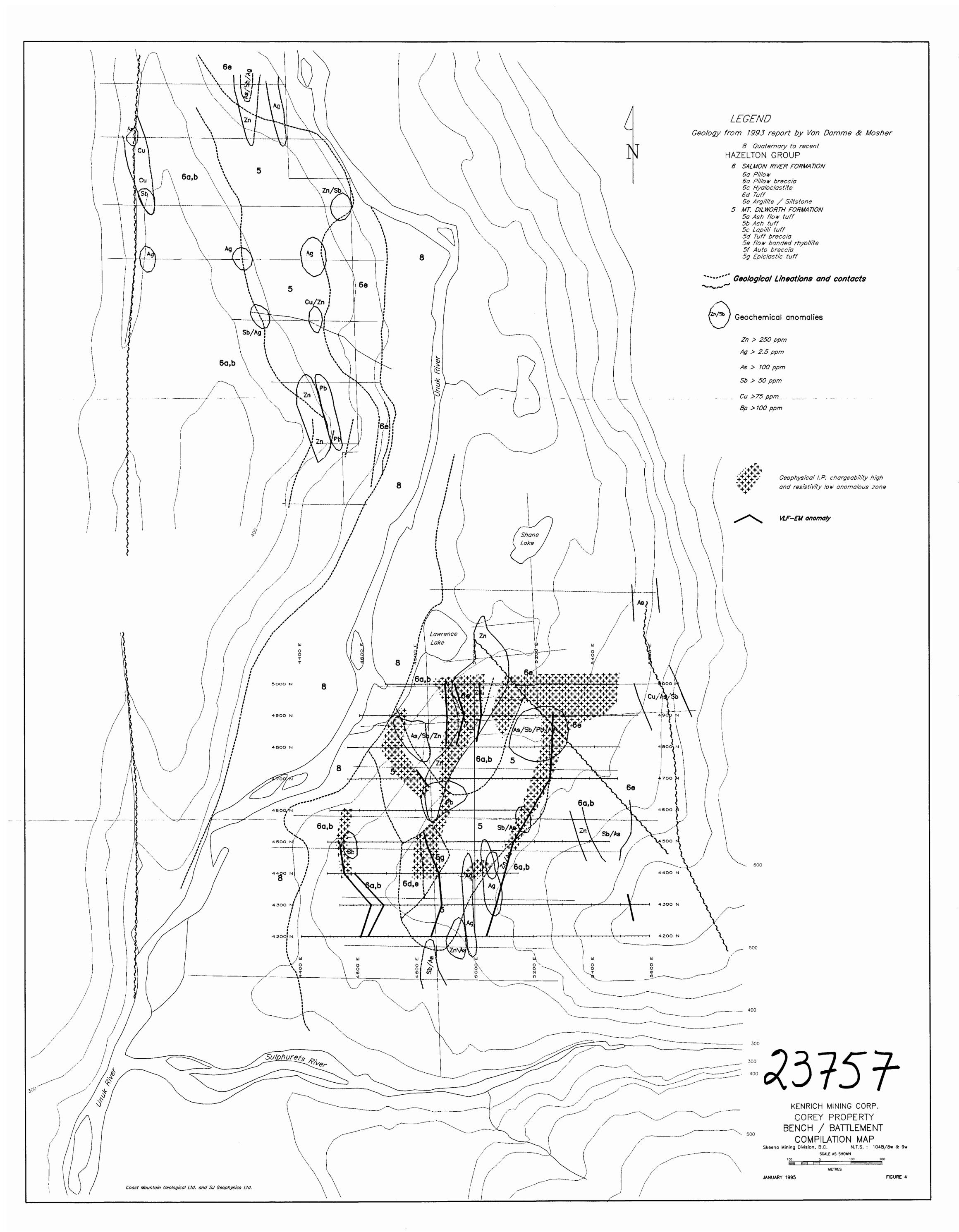
Personnel	Ŀ
	-

Personnel:			
D. Penner	Geologist	5.0 days @ \$350/day	\$ 1,750.00
		27.0 days @ \$400/day	10,800.00
W. Kushner	Geologist	24.8 days @ \$250/day	6,200.00
		21.0 days @ \$275/day	5,775.00
		5.0 days @ \$350/day	1,750.00
S. J. Visser	Geophysicist	7.0 days @ \$450/day	3,150.00
R. Arnold	Geologist	5.0 days @ \$350/day	1,750.00
S. Dudka	Geologist	5.0 days @ \$350/day	1,750.00
R. Schellenberg	Field Assistant	7.0 days @ \$225/day	1,575.00
N. Simpson	Field Assistant	7.0 days @ \$225/day	1,575.00
A. Smallwood	Geotechnician	31.5 days @ \$275/day	8,662.50
J. McConville	Cook	31.0 days @ \$250/day	7,750.00
C. Cutforth	Cook	2.0 days @ \$250/day	500.00
J. Chapman	Geologist	32.75 days @ \$350/day	11,462.50
J. Fraser	Geologist	1.00 days @ \$400/day	400.00
H. Page	Field Assistant	26.0 days @ \$275/day	7,150.00
T. Archibald	Geotechnician	16.0 days @ \$275/day	4,400.00
R. Hill	Geophysicist	4.75 days @ \$350/day	1,662.50
T. Hutchings	Geotechnician	2.5 days @ \$200/day	500.00
			\$78,562.50
Expenditures			
Accomodation			\$1,055.01
Meals			540.52
Groceries			6,933.24
Mob/Demob cos	sts - Flights, etc		5,974.18
Geophysical Mo	b/Demob Expenses:		2,296.25
	Flights		550.94
Ac	comodation in Smithers		464.00
Helicopter G	reat Slave	66.2 hrs @ \$570 /hr	37,734.00
N	orthern Mountain	22.74 hrs @ \$765 /hr	17,396.10
	CMG Truck rental:		
	Days used	20 days @ \$75 /day	1,500.00
	Stand-by days	12 days @ \$25 /day	300.00
	Kilometers travelled	10,063 kms@ \$0.25 /km	2,515.75
Field Expendibl	es: flagging, oil,		1,139.80
batteries etc.			
	Fuel:		
	Jet B	32 bbls @ \$197.94/	6,334.08
	Gasoline	1 bbls @ \$112.24/	112.24
	Propane refills	9 tanks @ \$45/	405.00
Linecutters		54 days @ \$250/day	13,500.00

Linecutter Mob/Demob costs		\$748.65
Handheld radios - 6 radios	156 days @ \$10 /day	1,560.00
Base station radios	26 days @ \$15 /day	390.00
Computer and Software	26 days @ \$15 /day	390.00
GPS units - 2 units	36 days @ \$10 /day	360.00
Field gear	148 mandays@ \$5 /manday	740.00
Camp Equipment	1000.00 /month	1000.00
Expediting Costs, Smithers Expediting		6,533.04
Travel Expenses		1,906.34
Communication		
Satellite Phone		832.99
B. C. Telephone		4,847.81
Freight		2,288.18
Maps and Reproduction		1,206.01
Assay Costs		660.13
Geophysical Surveys		
IP - In use	10 days @ \$1,250 /day	12,500.00
IP - Standby (Mob/Demob)	6 days @ \$938 /day	5,628.00
Mag/VLF - in use	5 days @ \$650 /day	3,250.00
Mag/VLF - Standby (Mob/Demob)	5 days @ \$488 /day	2,440.00
Base and 2 Field Units Standby	5 days @ \$200 /day	1,000.00
Management Fee		<u>1,690.88</u>
		\$148,723.14

TOTAL EXPENDITURES

\$227,285.64



STATEMENT OF QUALIFICATIONS

I, Jim Chapman, of 1455 Upland Trail, Bowen Island, British Columbia hereby certify:

- 1. I am a graduate of the University of British Columbia (1976) and hold a B.Sc. degree in geology.
- 2. I am presently employed as a consulting geologist by Coast Mountain Geological Ltd. of #1680 650 W. Georgia Street, Vancouver, British Columbia.
- 3. I have been employed in my profession by various mining companies for the past 20 years.
- 4. I am registered as a Professional Geoscientist with the Professional Engineers and Geoscientists Association of British Columbia
- 5. The information contained in this report was obtained through personal knowledge of the regional geology, unpublished reports from personnel on the property, publications listed in the bibliography and supervision of the field work.
- 6. I have no interest, direct or indirect, in the property described herein or in the securities of Kenrich Mining Corp.
- 7. This report may be used by Kenrich Mining Corp. for any and all corporate purposes, including, any public financing.

(Jim Chapman, P. Geo. Consulting Geologist

Dated at Vancouver, British Columbia, this 10th day of January 1995

APPENDIX A

KENRICH MINING CORP.

INDUCED POLARIZATION VLF-EM and MAGNETIC GROUND SURVEY ASSESMENT REPORT.

COREY PROJECT BENCH GRID

SKEENA MINING DISTRICT, BRITISH COLUMBIA NTS 104B/8W & 9W

> SJ GEOPHYSICS LTD. and COAST MOUNTAIN GEOLOGICAL LTD

JANUARY 1995

Report By Syd Visser

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INTRODUCTION

Total field magnetic, very low frequency electromagnetic (VLF-EM) and induced polarization (I.P.) surveys were completed on the Corey project for Kenrich Mining Corp. by SJ Geophysics Ltd. and Coast Mountain Geological Ltd. during the period of November 2 to November 17, 1994. The Corey Project claim group is located, 38 Km southwest of Bell II which is 275 Km to the north of Smitters on highway 37, in the Skeena mining district of B.C., NTS 104B/8W & 9W.

The surveys were performed on the Bench grid. The surveys were not completed due to an exceptional large and early snow fall.

This report is meant to be an addendum to a more complete property report written by Coast Mountain Geological Ltd., therefore, property location maps, claim names, local geology, sampling results and previous work done will not be shown or discussed in this report.

EQUIPMENT

The Induced Polarization (I.P.) equipment used was a Phoenix time domain transmitter, with a cycle time of 2 second on and 2 second off, and a Androtex, 6 dipole, time domain receiver. The delay time of the receiver was set at 80 millisecond with 10 integrating windows with widths of 80,80,80,80,160,160,160,320,320 and 320 millisecond each. The apparent resistivity was calculated using the recorded transmitter current and the nominal dipole spacing (50m) at each measurement location.

The magnetic and VLF-EM equipment used consisted of a EDA Omni Plus combined Proton Precession and VLF-EM field unit and a Omni-IV proton precession base magnetometer. The base station was located near the property to provide local corrections for diurnal drift. The magnetic data was corrected for diurnal at the end of each survey day.

All the data was transferred to a computer and plotted using Geopak Systems software.

FIELD WORK

The I.P. Geophysical crew, consisting of John Ashenhurst (technologist/supervisor) Zoran Dujakovic (Geophysicist) Ted Archibald (technician) Rod Hill and Neil Visser all with SJ Geophysics Ltd., and equipment mobilized from Vancouver on November 2/94. Inclement weather conditions slowed the mobilization into the Camp from Bell II which was completed on November 5/94 Accommodations were provided by Coast mountain Geological Ltd. near the property in the old Granges camp. A Helicopter was used daily to access the grid from the camp. The survey was stopped on November 14/94 due to dangerous working conditions because of excessive snow. The crew demobilized from camp on November 15 and drove to Vancouver on November 16 and 17, 1994. Poor road conditions forced a 2 day drive from Smitters.

The Grid was surveyed by I.P. using a Pole-Dipole array with an 'a' spacing of 50m and N=1 to 6. The survey direction was from west to east on all the surveyed lines. The slope was recorded at each station to produce a topographic map to aid in the interpretation and correct the station spacing (horizontal distances) for the plan maps.

The VLF-EM and magnetic survey was completed by Coast mountain personal under the supervision of SJ Geophysics Ltd. The VLF-EM transmitter stations used for the survey was Seattle (NLK., 24.5 Khz) and Annapolis (NSS., 21.4 Khz). The direction to Annapolis is approximately east and therefore was recorded only to check for possible cross structures if required.

DATA PRESENTATION

The data was plotted using Geopak presentation software. The I.P. speudosections station spacing is not corrected for horizontal distances. All of the grids on the plan maps are corrected for horizontal distances. The geology shown on some of the plan maps was digitized from a geology map supplied by Kenrich Mining Corp. The following is a list of the maps accompanying this report:

Magnetic survey

Plate G1a	Total Field Magnetometer Profiles
Plate G1b	Total Field Magnetic Contours

VLF-EM survey

Plate G2a	VLF-EM Survey Profiles
	Dip Angle, Quadrature & Topography
Plate G2b	VLF-EM Fraser Filtered Dip Angle
	Contours

1.1 . 301 PEY	
Plate G3a	Total Chargeability (N=1)
	Contours & Geology
Plate G3b	Total Chargeability (N=3)
	Contours & Geology
Plate G4a	Apparent Resistivity (N=1)
	Contours & Geology
Plate G4b	Apparent Resistivity (N=3)
	Contours & Geology
Line	I.P. Pseudosections
4400N-5000N	

I.P. survey

DISCUSSION

MAGNETICS

The magnetic response indicates a long wavelength magnetic response with an amplitude of approximately 200nT and local short wavelength responses of up to 1100nT. The amplitude of the long wavelength response which is either due to regional or deep seated effects increase from the northeast end of the grid to the southwest end.

With the exception of two locations on line 4700N, at the base line and 5450E, the short wavelength or near surface responses are located on the west side and south end of the grid. The high amplitude response between 5200E and 5400E on lines 4200N and 4300N suggest that the source may be an intrusive or volcanic rock unit in this area. The magnetic response near 4600E on lines 4300N to 4700N suggest that there is a possible contact in this region. This is correlated with the VLF data.

As can be seen on plate G1A and G1B there is no clear correlation between the known geology and the magnetic response.

VLF-EM

The VLF-EM indicates a number of VLF-EM anomalies as shown on Plates G2A to G2B. The main anomalies are located on the eastern side, the northern end and the southwest corner of the grid.

The anomaly on the eastern side of the grid appears to follows the eastern contact of the Mt. Dilworth Formation from line 4800N south to line 4500N and then appears to drift of to the east as shown on Plate G2B. This anomaly is most prominent on line 4700N, at approximately 5250E, where it correlates well with an I.P. resistivity low and a chargeability high

The second anomaly or anomalous zone is located on the northern end of the grid between 4900E and 5050E on line 5000N and continues down to line 4800N where it is weak. The western part of this anomaly may continue to the southern end of the grid but it is very weak and difficult to trace. The anomalous zone consist of a number of parallel conductors or contact zones. The anomaly correlates well with an I.P. resitivity low (directly east of a very high resistivity zone) and chargeability high zone.

The third VLF-EM anomaly on the southwestern end of the grid strikes from approximately 4750E on line 4500N to 4600E on line 4200N. This anomaly appears to be either two parallel anomalies or more likely a wide (25m) conductive bed. The I.P. was not completed over most of this part of the grid. There is a resitivity low and chargeability high on the westernmost end of lines 4400N and 4500N, which appears to correlate well with this anomaly.

The very weak VLF anomaly located near 4850E on the southern part of the grid may be a contact zone and is associated with a chargeability high and resitivity low on line 4400N where it appears to be associated with the eastern contact of the Mt. Dilworth Formation.

I.P. SURVEY

The Bench grid zone was covered with 7 lines, varying in length from 400 to 900m, with a pole-dipole I.P. survey using an 'a' spacing of 50m with N=1 to 6. The data was plotted as pseudosections and labeled with the respective line number. Plan maps showing the total chargeability data from N=1 and N=3 and the apparent resistivity data from N=1 and N=3 are also displayed on plates G3A, G3B, G4A, G4B. The geology, from Van Damme and Mosher 1993, was digitized and included on the plan maps.

There appears to be no direct relationship between the known geology and the I.P. response with the exception of the Eastern contact of the Mt. Dilworth Formation and the argillites. The Eastern contact of the Mt. Dilworth Formation is a low resitivity zone and a high chargeability region. The highest chargeability is located on line 4800N at approximately 5200E.

The rocks mapped as argillite/siltstone (Salmon River Formation) on the north central and northeastern part of the grid have a low resitivity and high chargeability as expected from argillites. This zone in the north central part of the grid appears to be more extensive than shown by the mapping. The rock unit mapped as pillow and pillow breccia to the south and west of the argillite have very high resitivity areas and generally low chargeabilities.

Line 4400N

This whole line has a very high chargeability anomaly with the exception of a near surface zone between 4600E and 4750E and a narrow zone between 4900E and 4950E. The high chargeability region a low correlates well with a resistivity area. The low chargeability correlates well with a high resitivity zone.

Line 4500N

The near surface low chargeability and high resistivity zone located between 4600E and 4750E on line 4400N extends across to line 4500N. The area between 4750E and 5050E appears to be interbedded high resitivity and low resistivity zones. The high resitivity zones also appear to have low chargeability. The area between 5050E and 5200E is a low resistivity and high chargeability zone..

The highest chargeability appears to be at depth near 4950E. It is not clear if this is a cumulative effect from the eastern and western near surface high chargeability regions or if it is a high chargeability zone at depth. The highest chargeability on this line appears to be at 4500E, the most westerly station. This zone appears to be very similar to the near surface high chargeability at 4500E, the most westerly station on this line.

Line 4600N

The anomalous chargeability at the western end of line 4500N extends to the western end of line 4600N. The high resitivity and low chargeability zone between 4550E and 4800E also continues and appears to have more depth extent on this line. A chargeability high and resitivity low region is located between 4800E and 5000E and possibly extends further east at depth. A near surface high resitivity and chargeability low extends from 5000E to the eastern end of the line.

Line 4700N

Although it appears from the section that there is a drastic change from line 4600N to 4700N they are very similar with the exception that the high chargeability zones have a higher amplitude and are more extensive. The resitivity high on the western end of the line is smaller but continues between 4600E and 4725E. There is also a near surface high resitivity zone which extends from 4800E to 5200E. This zone does not appear to have much depth extent except between 5100E and 5150E which correlates with a low

chargeability zone. In general the whole line with the exception of some near surface region has a very high chargeability and low resitivity.

Line 4800N

There is a drastic change between lines 4700N and 4800N. There is an extremely high resitivity zone between 4700E and 4900E on this line. It was very difficult to obtain any reliable data in this region because of the high resitivity but it does appear that the chargeabilities are very low.

Between 4900E and 5150E the resistivities and chargeabilities are very inhomogeneous likely due to rapid changes in the electrical character of the rocks and therefore very difficult to interpret. East of 5150E the resistivities are low and chargeabilities high especially in the region centered at 5250E.

Line 4900N

Line 4900N has low resistivity and high chargeability with the exception of a near surface high resistivity and low chargeability zone between 4800E and 4900E. There is an exceptional low resistivity zone between 4950E and 5050E which correlates with a very high chargeability. The resistivity may increase and chargeability decrease with depth in this region.

Line 5000N

Line 5000N is very similar to line 4900N with a high resistivity zone at 4850E and a very low resistivity and high chargeability region to the east. As with the previous line the resistivity may increase and chargeability decrease with depth in the central part of this line.

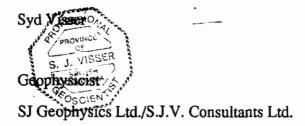
SUMMARY AND RECOMMENDATIONS

The magnetic data does not indicate a close correlation to known geology with the exception of possibly at the southern end of the grid. A denser line spacing and more complete survey would be required to find any correlation with geology.

The VLF-EM anomalies correlate with the eastern contact of the Mt. Dilworth formation, with argillites on the northern part of the survey area and indicates that there is another contact on the southwestern part of the grid. The only significant anomaly is indicated on line 4700N and possibly on the north central part of the grid. This survey should be completed along with denser line spacing. A local source EM system such as Max-Min HLEM, for depth of up to 50M, or a large loop system such as UTEM, for greater depth, should be considered for the search of massive sulphides.

The I.P. indicates that the argillites on the northern and northeastern part of the survey area have a low resistivity and high chargeability. Areas of interest may be the eastern Mt. Dilworth contact, the central part of lines 4400N and 4500N and the southeastern edge of the survey area which all have resitivity lows and chargeability highs.

The survey should be extended in all directions where possible. The resultant data should be closely correlated to known geology and geochemical information to determine further follow-up work and the type of instrumentation best suited for this work.



APPENDIX 1

STATEMENT OF QUALIFICATIONS

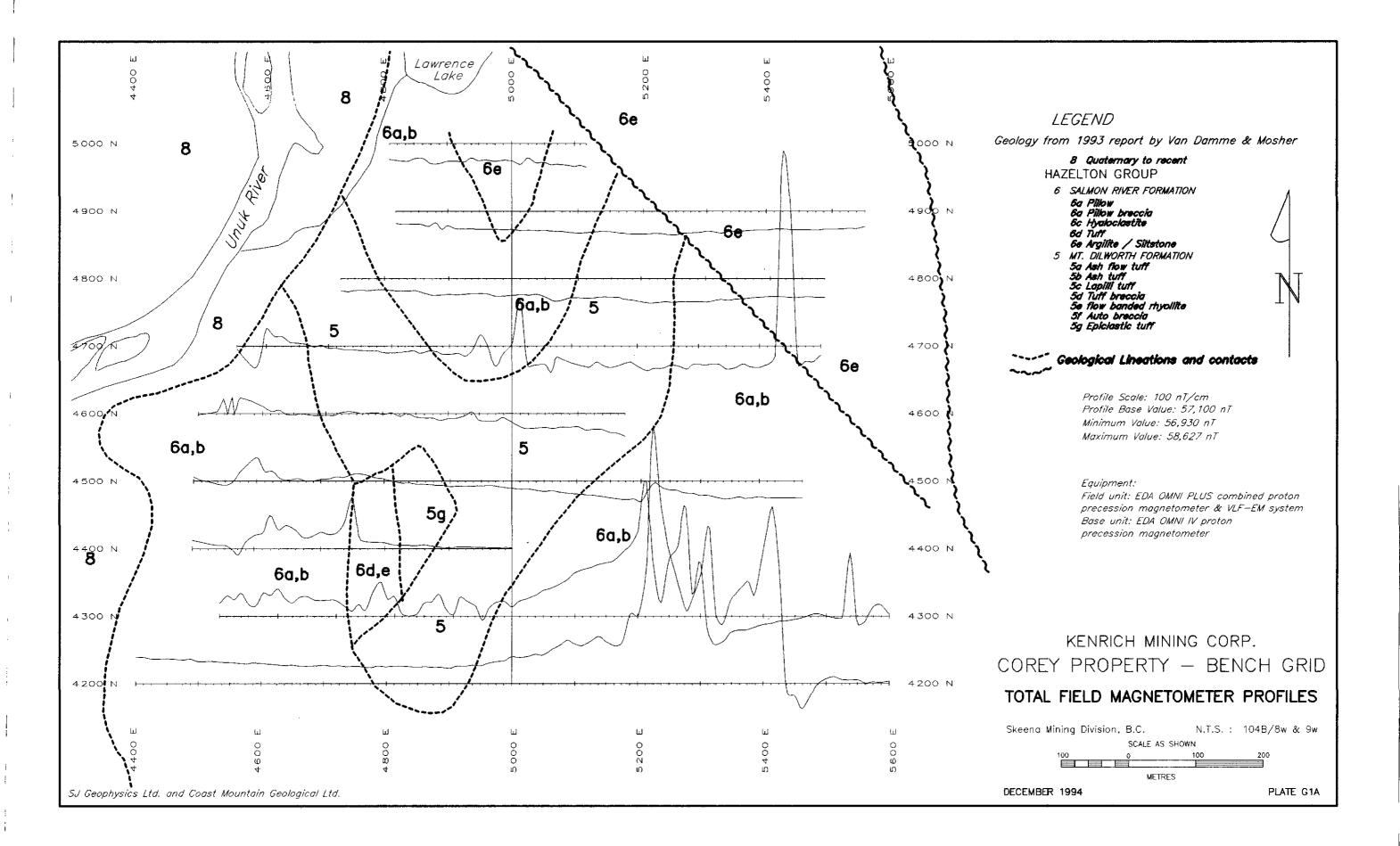
I, Syd J. Visser, of 11762 - 94th Avenue, Delta, British Columbia, hereby certify that,

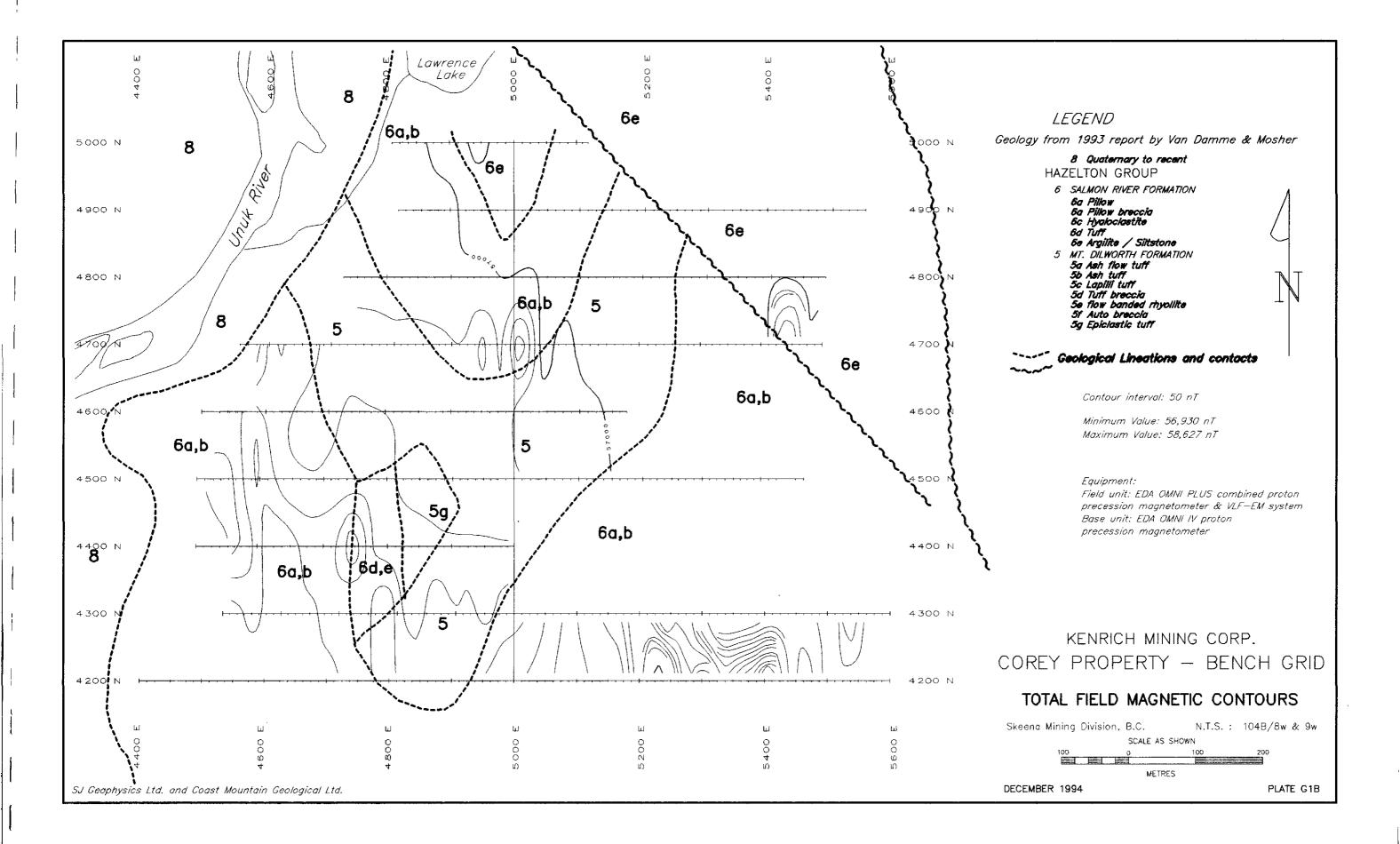
- 1) I am a graduate from the University of British Columbia, 1981, where I obtained a B.Sc. (Hon.) Degree in Geology and Geophysics.
- 2) I am a graduate from Haileybury School of Mines, 1971.
- 3) I have been engaged in mining exploration since 1968.
- 4) I am a professional Geoscientist registered in British Columbia.

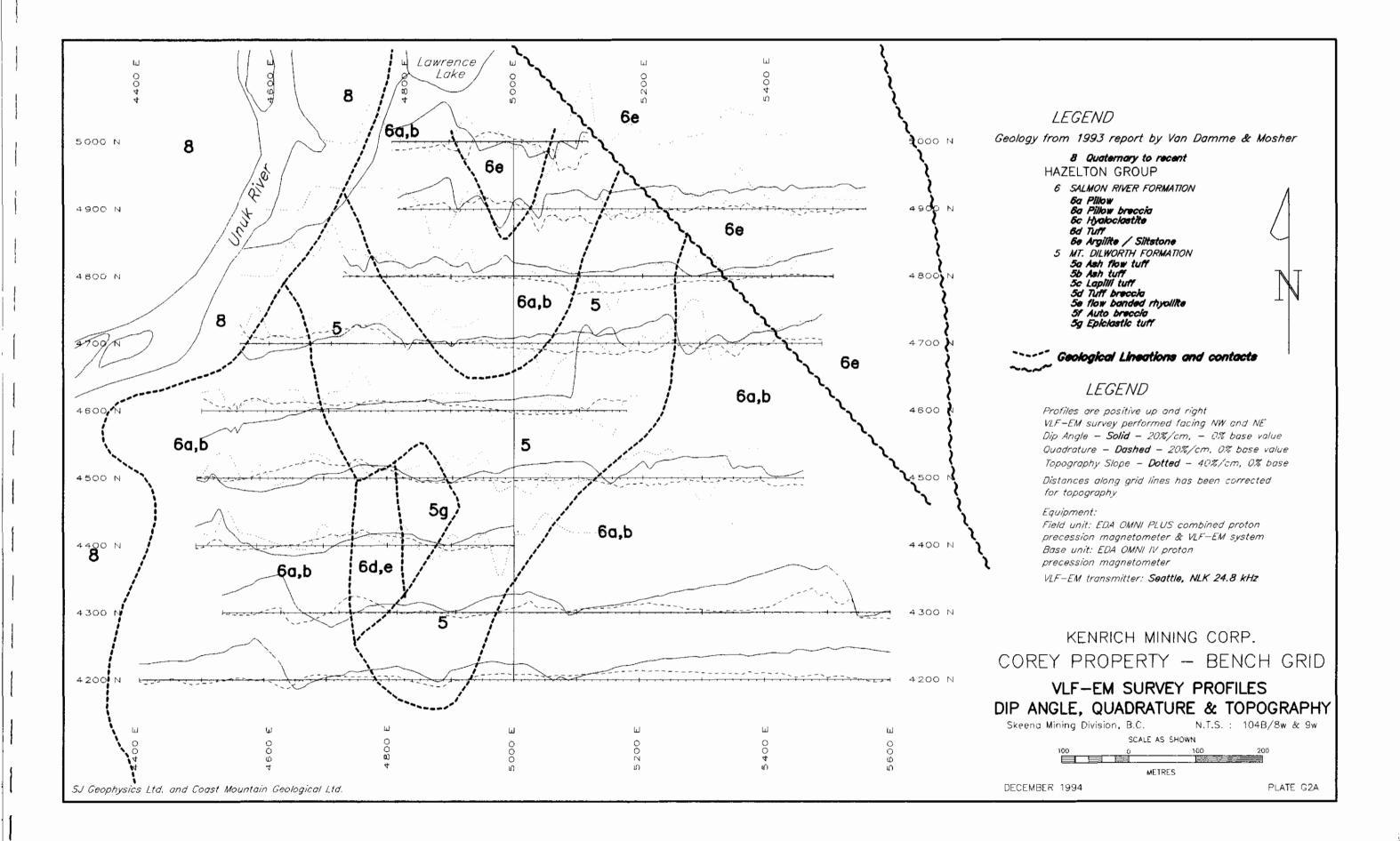
-VIS Syd J./Visser, B.So. P.Geo

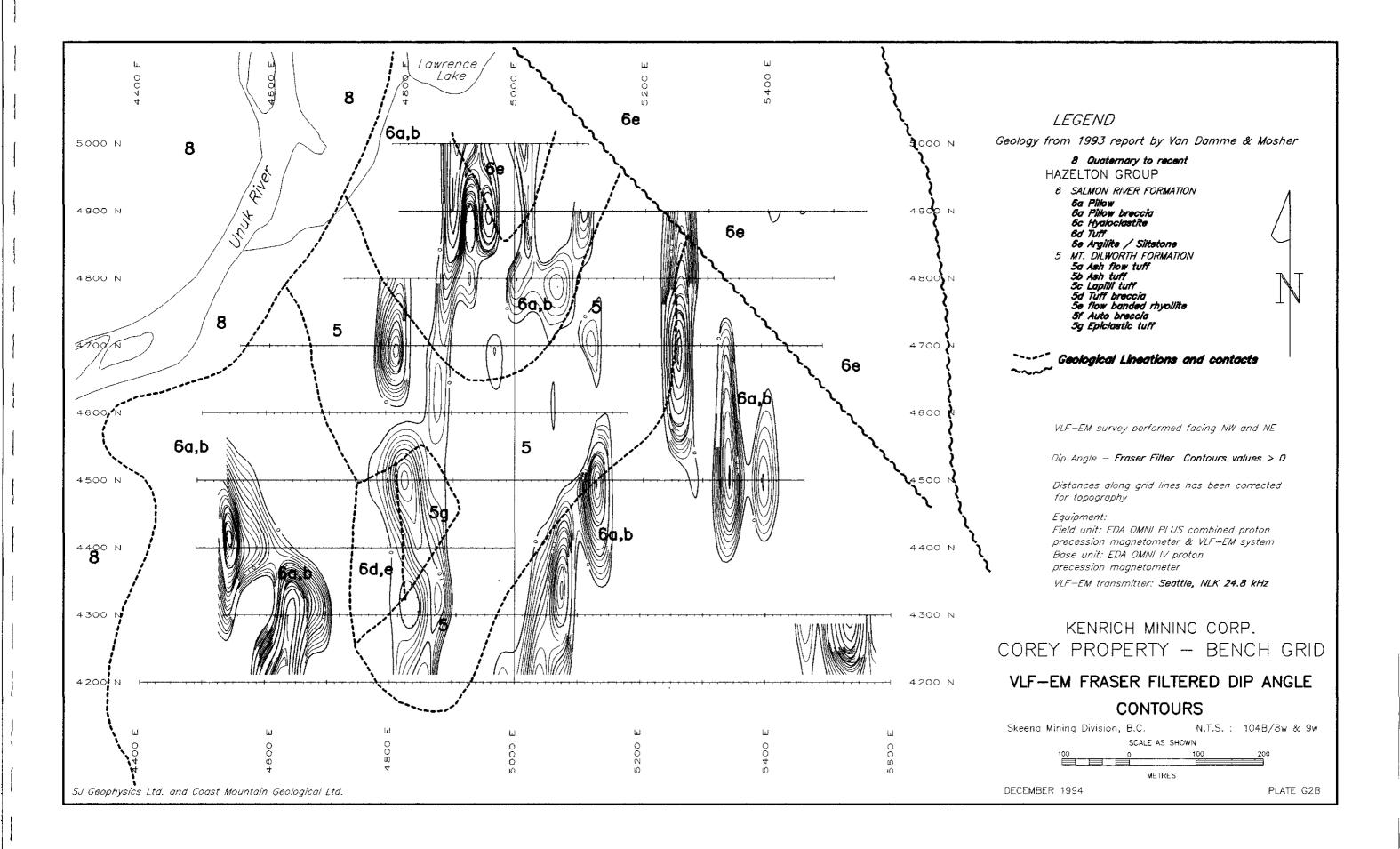
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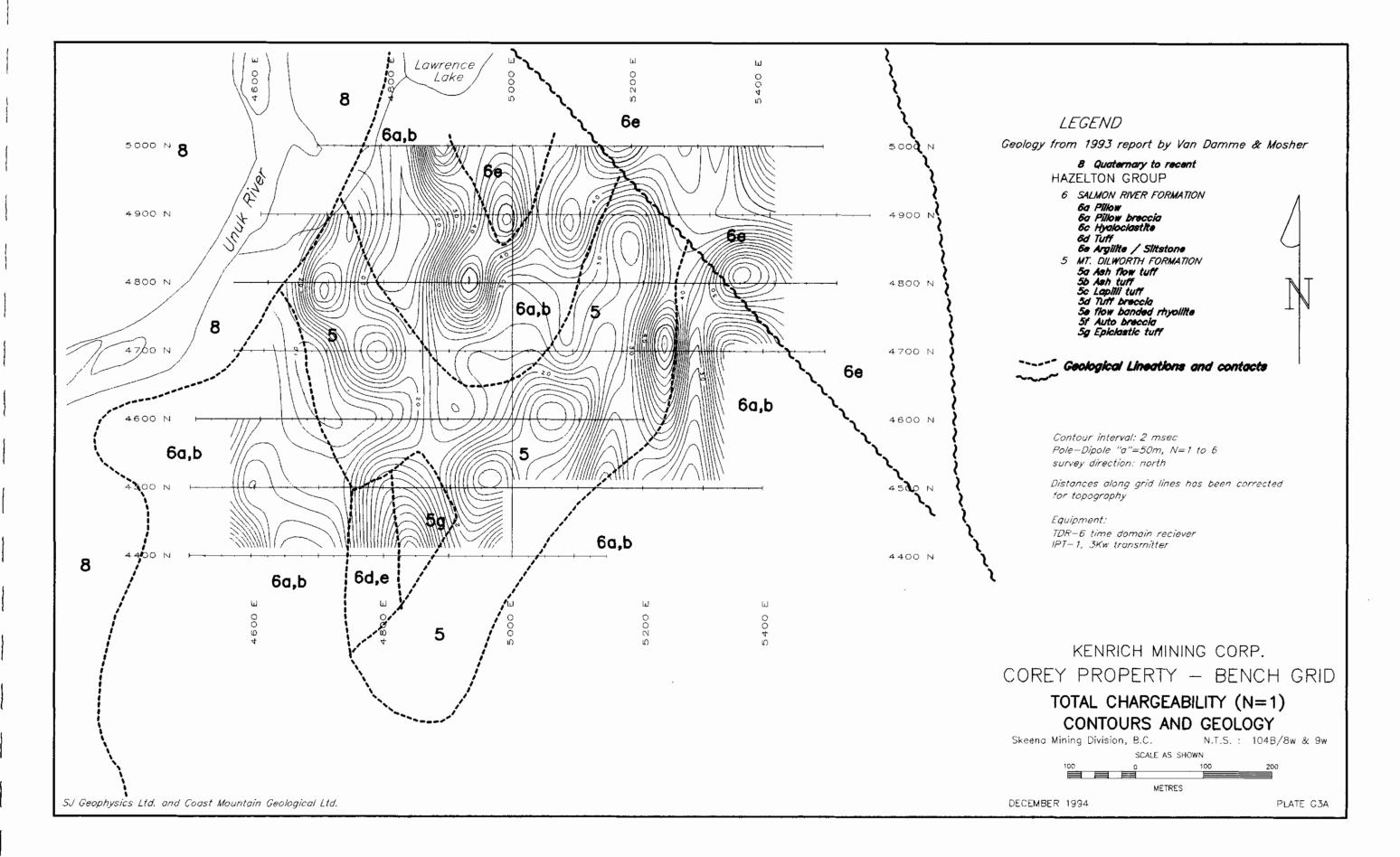
Geophysicist

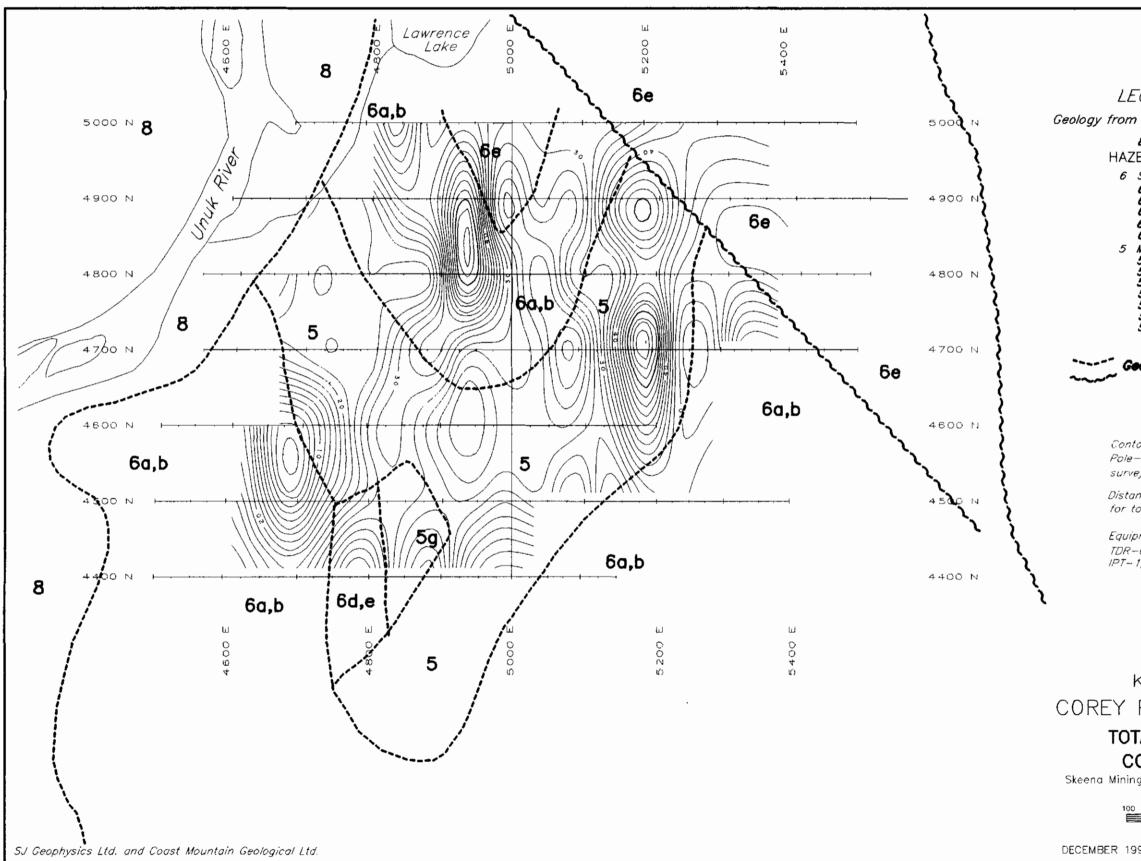






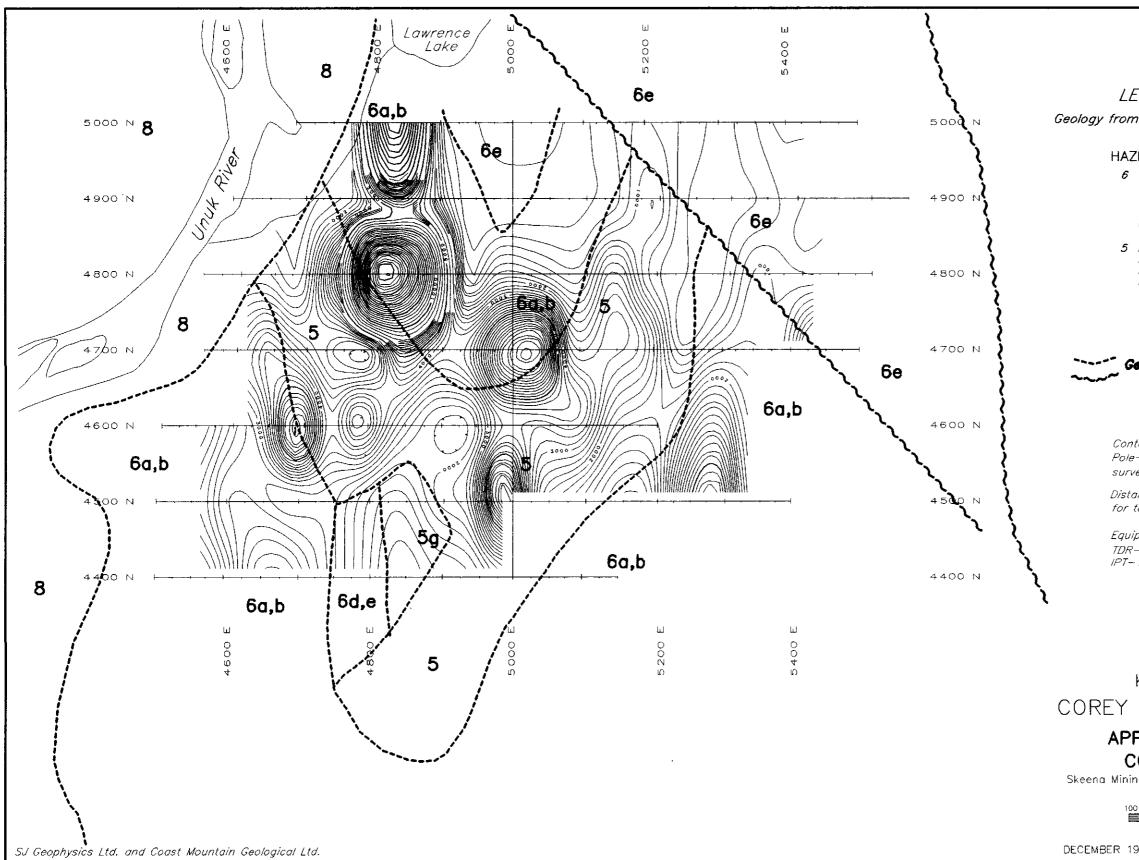




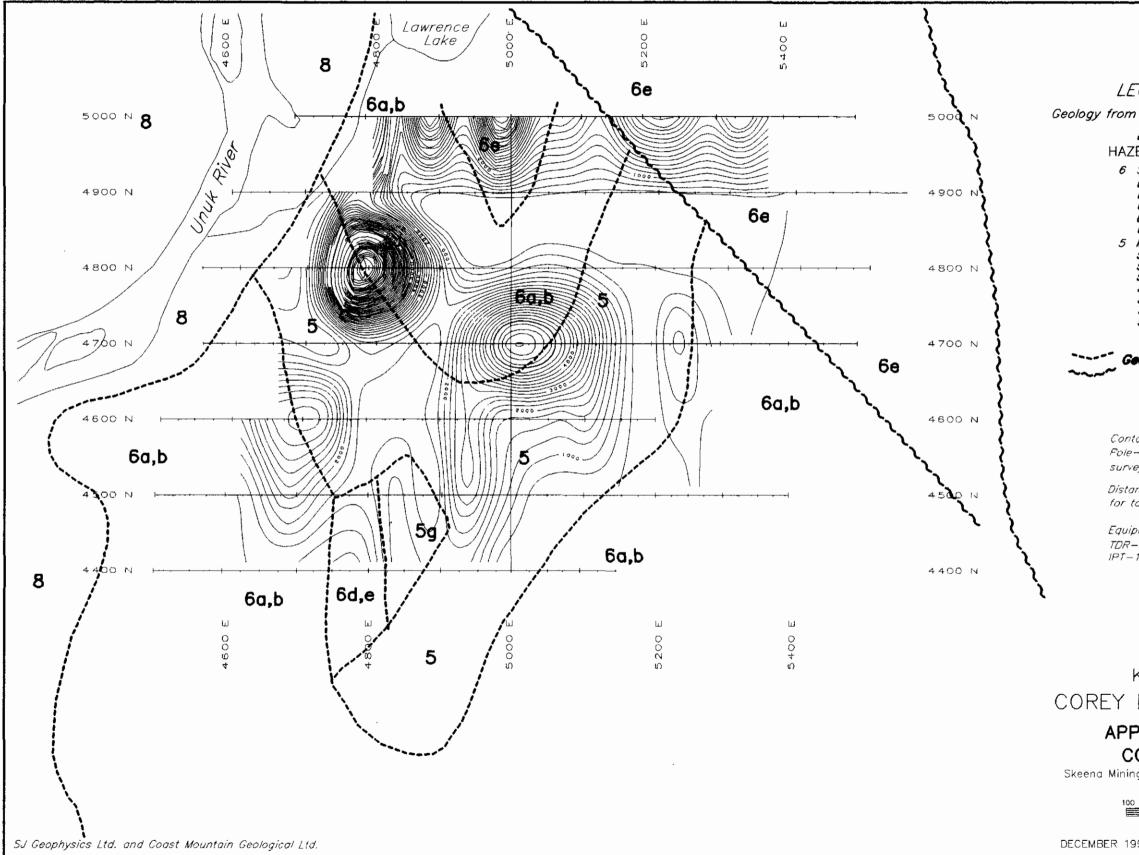


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6d Tuff 6e Argiilte / Siltstone
MT. DILWORTH FORMATION 5g Ash flow tuff
5b Ash tuff 5c Lapilli tuff 5d Tuff breccia
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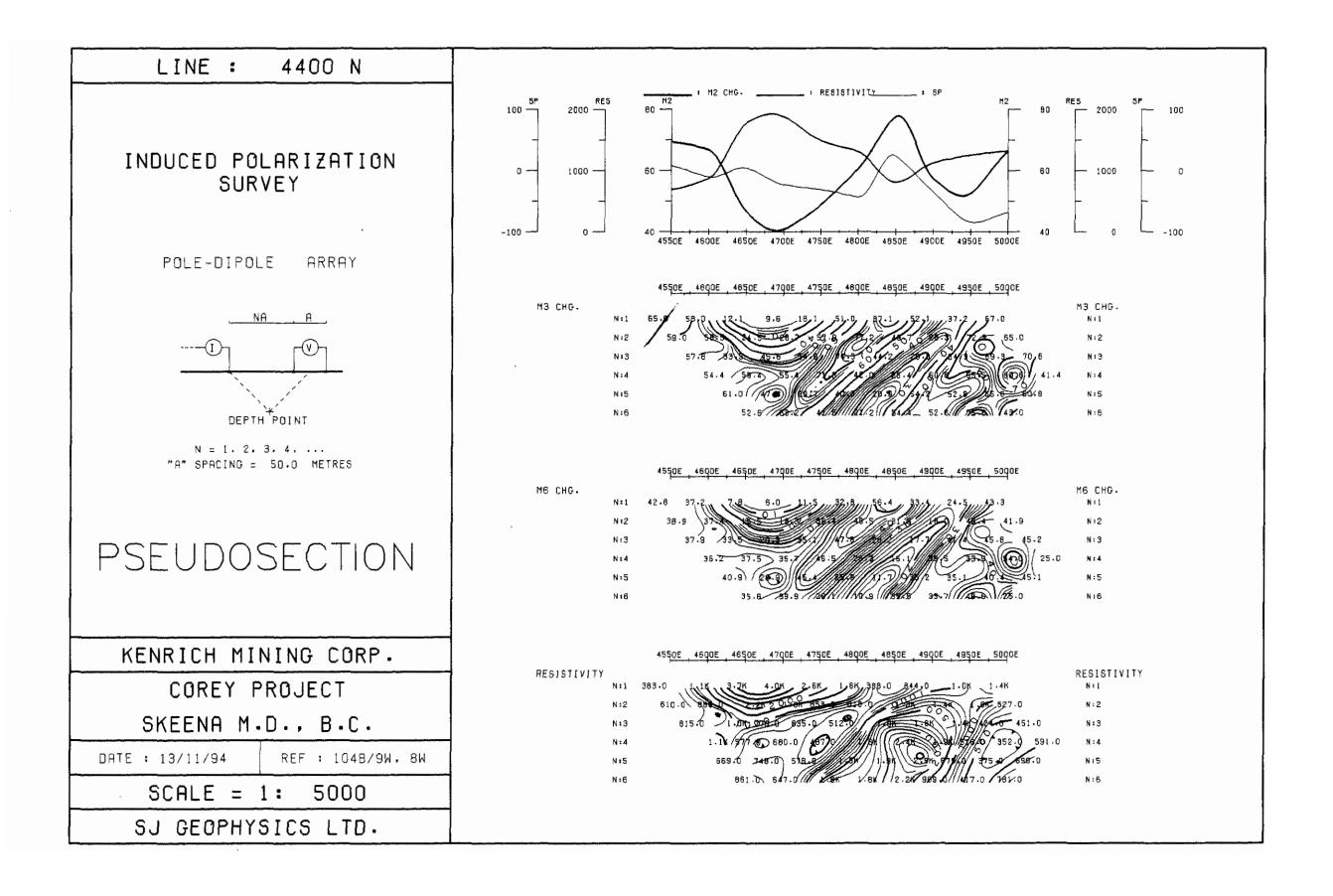


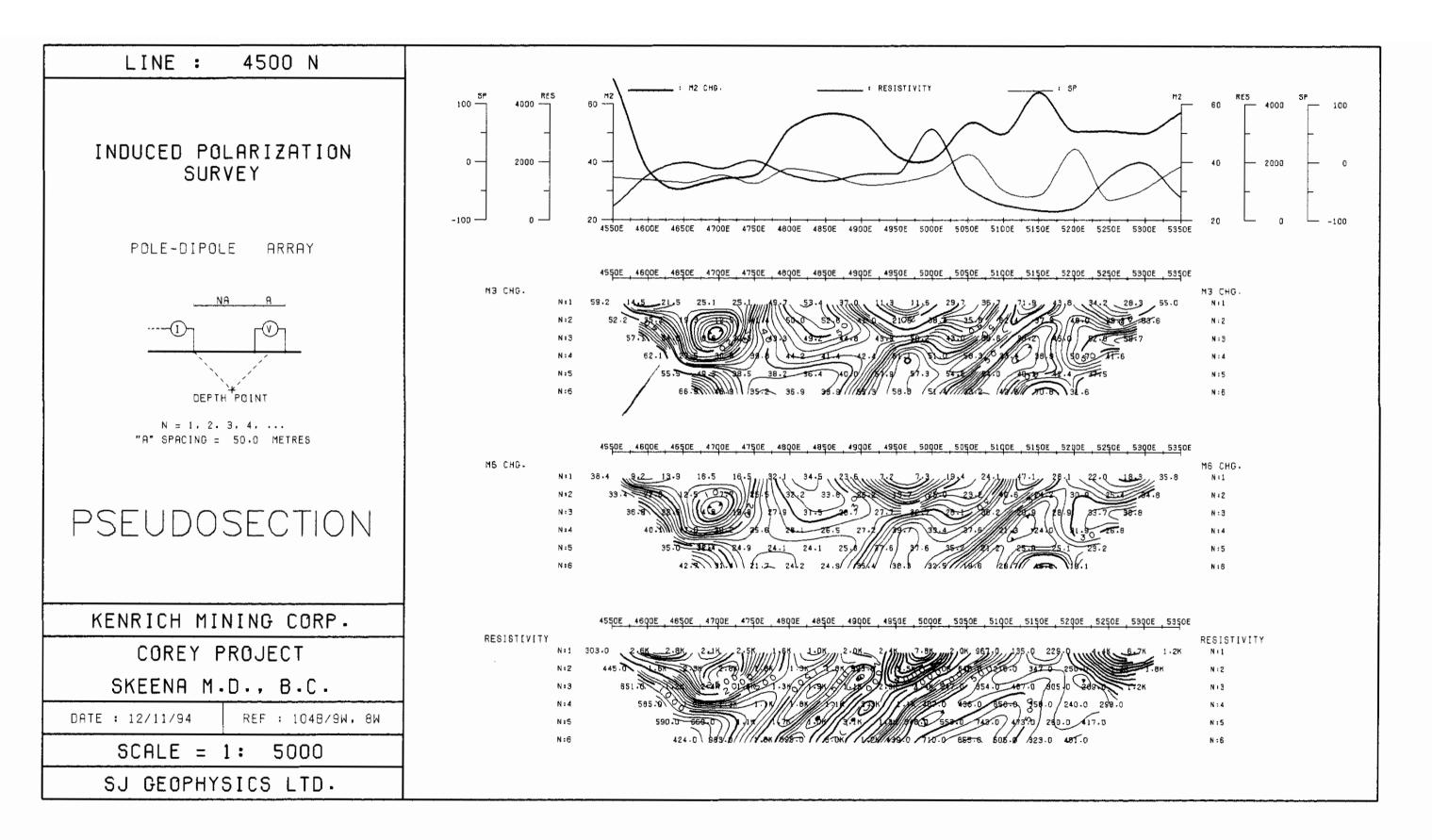
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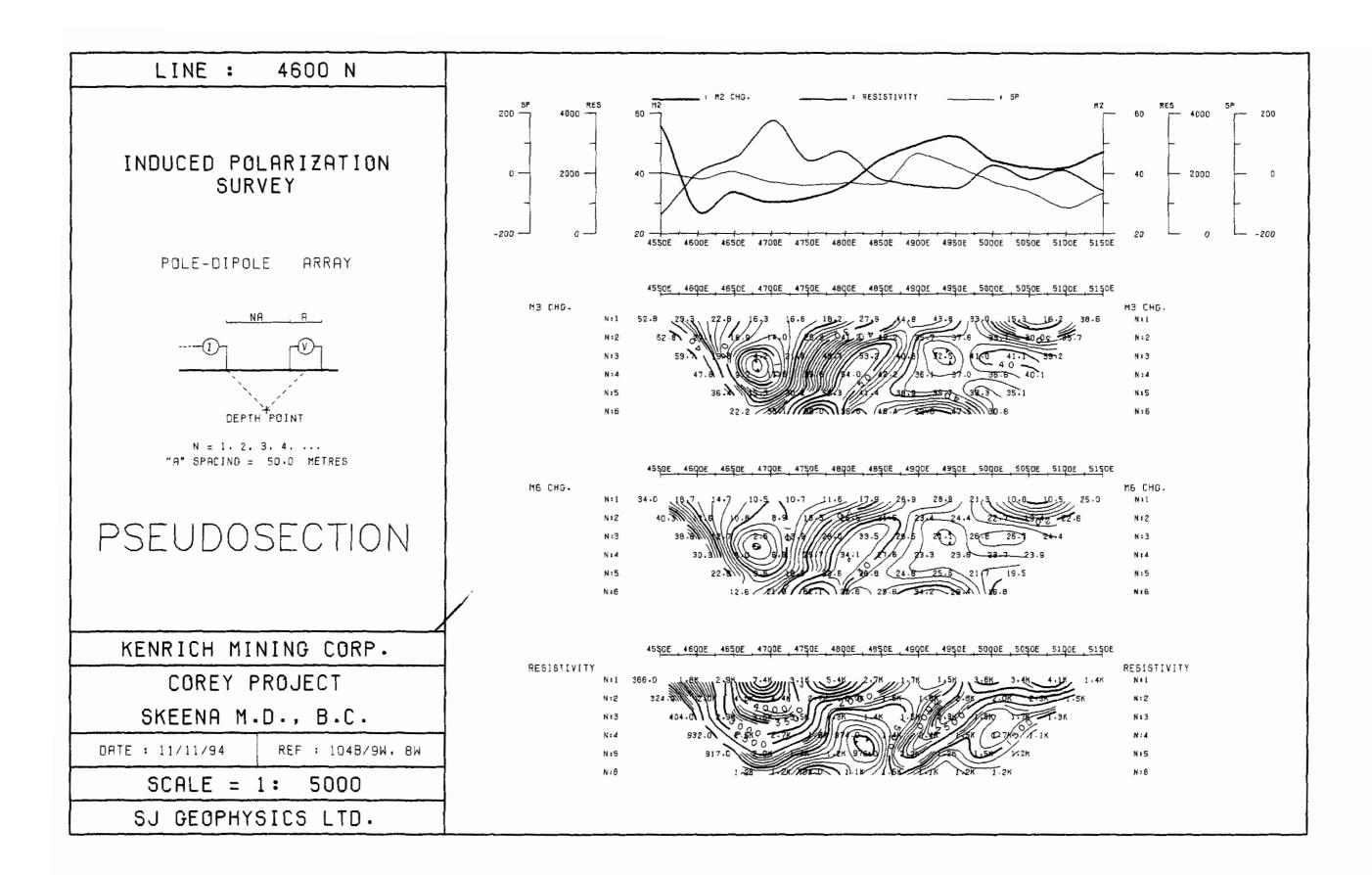
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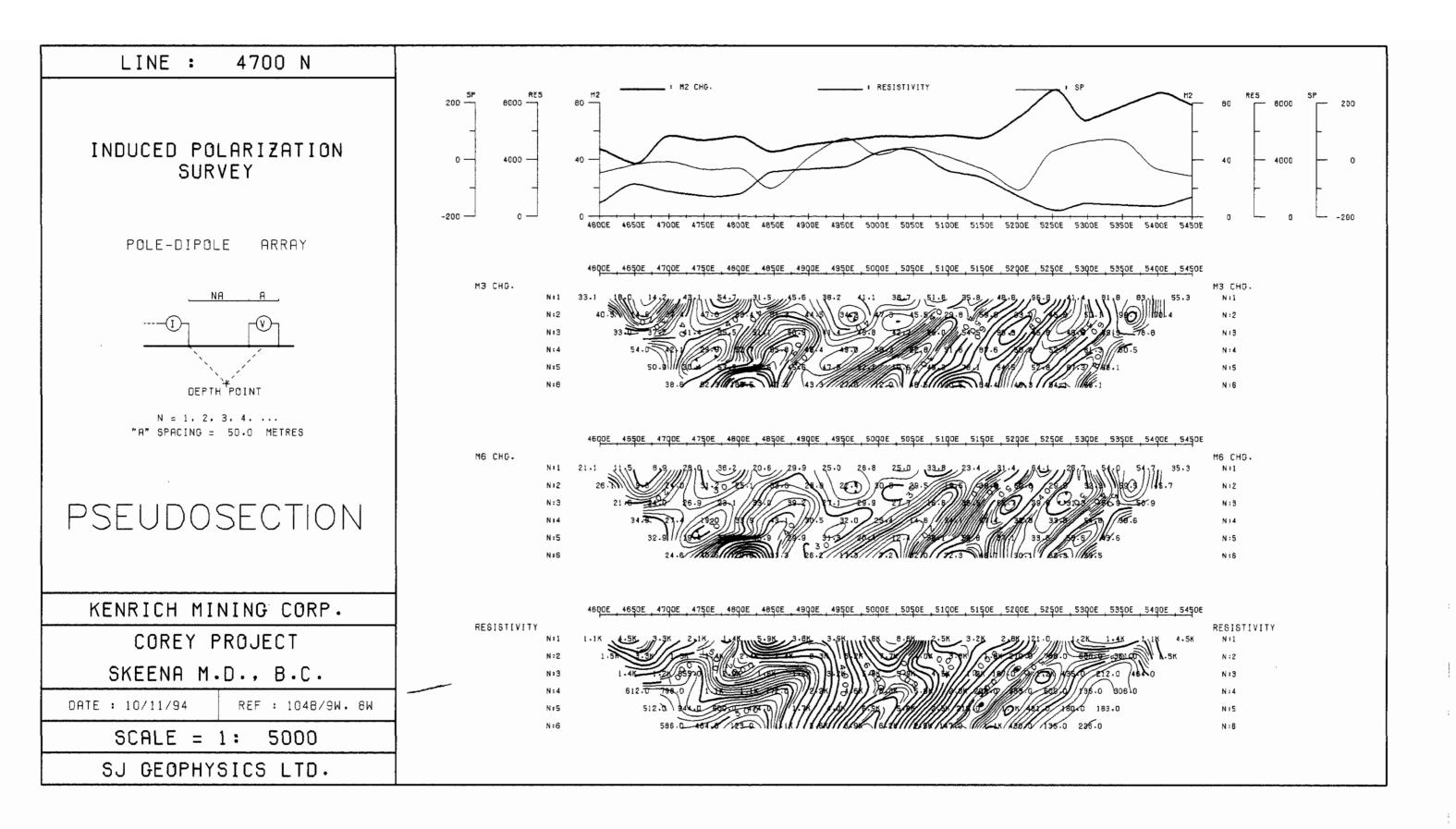
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PROPERTY - BENCH GRID
PARENT RESISTIVITY (N=3)
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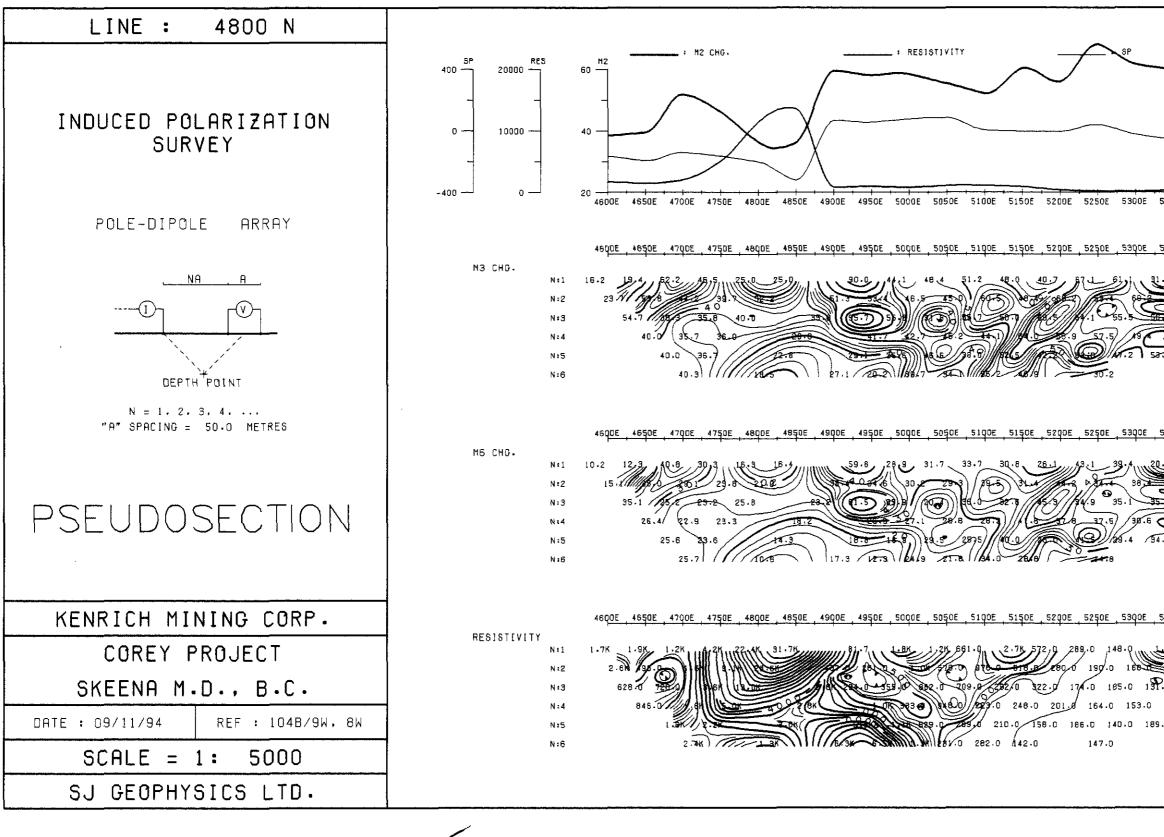




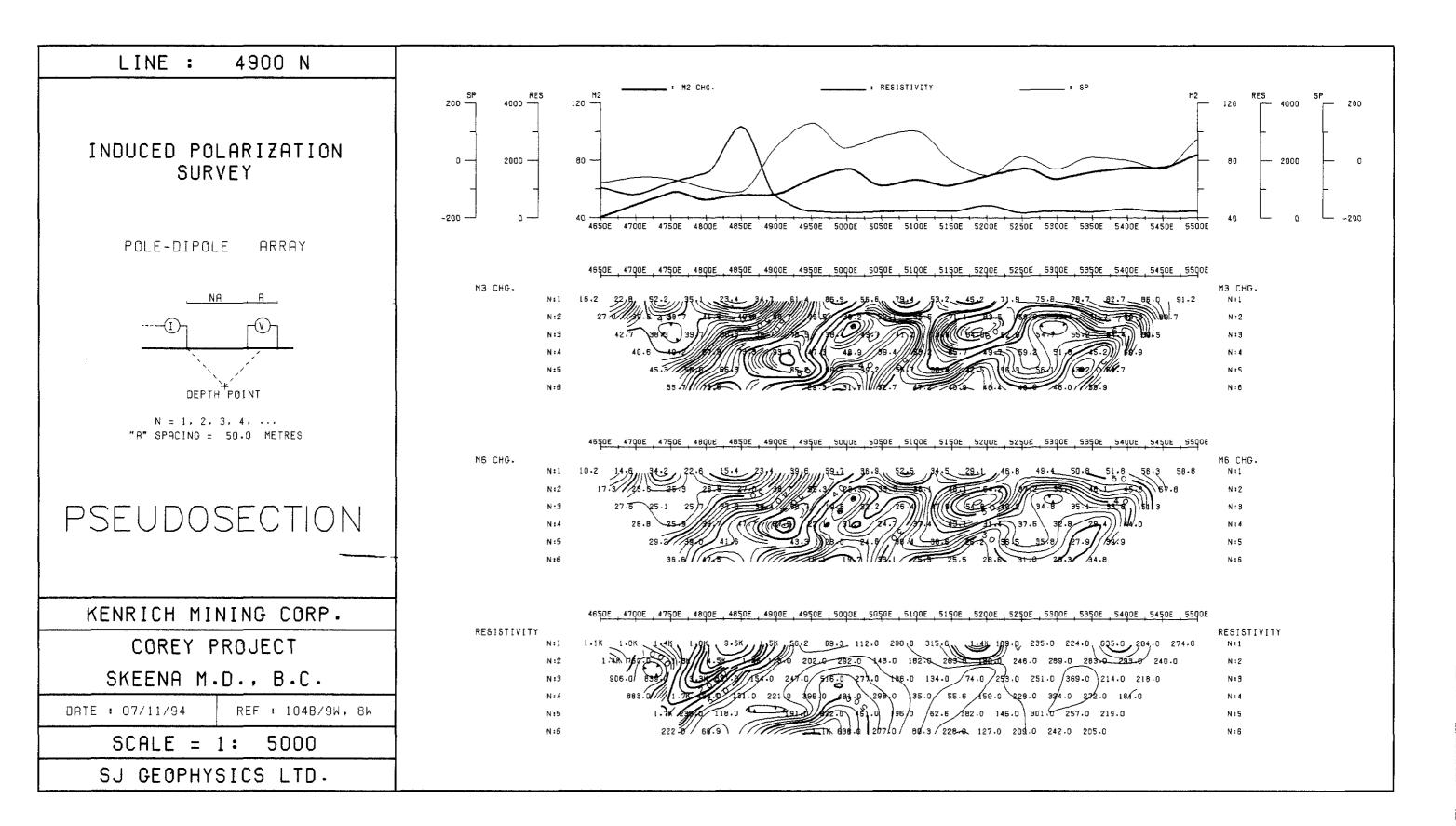
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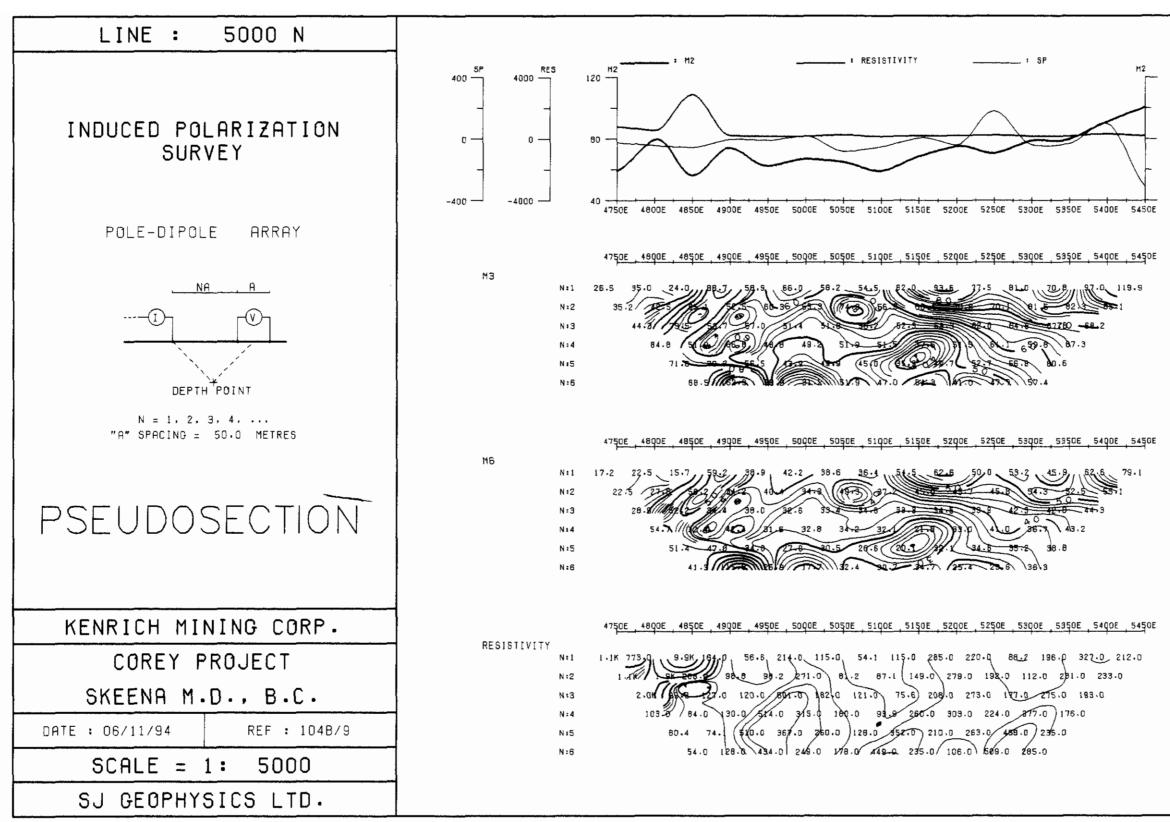


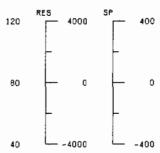




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APPENDIX B



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Chemex Labs Ltd.

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Analytical Chemists * Geochemists * Registered Assayers

212 Brooksbank Ave., North Vancouver British Columbia, Canada V7J 2C1 PHONE: 604-984-0221

To: COAST MOUNTAIN GEOLOGICAL LTD. P.O. BOX 11604 1680 - 650 W. GEORGIA ST. VANCOUVER, BC V6B 4N9

Project : COREY Comments: ATTN: JIM CHAPMAN

Page Number :1 Total Pages :1 Certificate Date: 13-JAN-95 Invoice No. :19510364 P.O. Number : Account :MHK

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Analytical Chemists * Geochemists * Registered Assayers

212 Brooksbank Ave., North Vancouver British Columbia, Canada V7J 2C1 PHONE: 604-984-0221

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To: COAST MOUNTAIN GEOLOGICAL LTD. P.O. BOX 11604 1680 - 650 W. GEORGIA ST. VANCOUVER, BC V6B 4N9 Page Number : 1-A Total Pages : 1 Certificate Date: 12-DEC-94 Invoice No. : 19432000 P.O. Number : Account : MHK

Project : COREY Comments: ATTN: DON PENNER or JIM CHAPMAN

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SAMPLE	PRI CO	-	Ag ppm Aqua R	A1 %	As ppm	Ba ppm	Be ppm	Bi	Ca %	Cđ ppm	Со ррш	Cr ppm	Cu ppm	Fe %	Ga ppm	Hg ppm	K %	La ppm	Mg %	Mn ppm	Мо ррт									
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CERTIFICATION:_



Chemex Labs Ltd.

Analytical Chemists * Geochemists * Registered Assayers 212 Brooksbank Ave., North Vancouver British Columbia, Canada V7J 2C1 PHONE: 604-984-0221

To: UUAST MUUN FAIN GEULOGICAL LID. P.O. BOX 11604 1680 - 650 W. GEORGIA ST. VANCOUVER, BC V6B 4N9

rage Number :1-B Total Pages :1 Certificate Date: 12-DEC-94 Invoice No. : [9432000 P.O. Number : Account :MHK

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Project : COREY Comments: ATTN: DON PENNER or JIM CHAPMAN

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										CERTIFICATE OF ANALYSIS				NALYSIS	A9432000
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2-04 2-05	201 229 201 229	0.03	79 30	850 1230	< 2 60	< 2 16	8 8	79 34	0.10 0.07	10 < 10	< 10 < 10	46 60	< 10 < 10	262 696	
-06 9+00N 47+50E	203 205 203 205	0.03	65 30	880 130	8 16	2 < 2	7	18	0.01	< 10 10	< 10 < 10	47 438	< 10 < 10	414 22	
9+00N 47+75E	203 205	0.01	23	510	6	< 2	14	7	0.39	10	< 10	113	< 10	186	
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9+00N 49+00E	203 205	0.01	13	400	10	< 2	4	18	0.45	< 10	< 10	185	< 10	16	
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SAMPLE	PR CO		Ag ppm	A1 %	Ås ppm	Ba ppm	Be ppm	Bi ppm	Ca %	Cđ ppm	Co ppm	Cr ppm	Cu ppm	Fe %	Ga ppm	Hg ppm	K %	La ppm	Mg %	Mn ppm	Mo ppm
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