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REVIEW OF GEOPHYSICAL DATA MAMMOTH PROJECT, NTS.82F/6 NELSON MINING DIVISION SOUTH CENTRAL BRITISH COLUMBIA FOR ROSSMIN EXPLORATIONS LTD.

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GEOLOGICAL SURVEY BRANCH ASSESSMENT REPORT

REVIEW OF GEOPHYSICAL DATA

MAMMOTH PROJECT, NTS.82F/6

NELSON MINING DIVISION

SOUTH CENTRAL BRITISH COLUMBIA

FOR

ROSSMIN EXPLORATIONS LTD.

BY

DELTA GEOSCIENCE LTD.

MAY 2, 1997.

GRANT A. HENDRICKSON, P.GEO.

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INTRODUCTION

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At the request of Rossmin Explorations Ltd., Delta Geoscience has conducted a new data presentation and review of the geophysical data presently available on the Mammoth property. This valuable data had been collected by previous operators of the claims, however the data presentation was guite poor - a fact which was detrimental to its understanding.

This well known property is located in the Nelson Mining Division of south central British Columbia, approx. 15 kilometers south of the city of Nelson, NTS.82F/6. The property has a long history with previous exploration work dating back to the turn of the century.

The property lies within the Jurassic age Rossland Group of volcanics and sedimentary rocks that have been intruded by late Jurassic dioritic porphyry bodies, which are possibly co-magmatic with the volcanics. Later intrusive activity (granodiorites), probably related to the Bonnington Pluton which is located immediately adjacent to the west side of the present grid, is known to crosscut and/or assimilate the Rossland volcanics (Elise Formation) and overlying sediments (Hall Formation).

The geology of this complex property is described in much better detail in two geologic reports prepared for Rossmin by independent consultants. These reports are referenced at the back of this report.

The property scale geology for this interesting old property is poorly understood due to extensive areas of relatively thin overburden cover. Geophysical data does allow us to infer the probable geology in overburden covered areas.

Mineralization on the property tends to be hosted in altered intrusive and/or the Elise volcanics, although the coarser portion of the Hall Formation could also be mineralized where it is fractured and/or sheared.

Three exploration models are currently in vogue for this property:

- 1) Copper gold porphyry.
- 2) Shear zone hosting quartz sulphide veins.
- 3) Copper gold skarn.

Recent papers by the government geologists suggest the structural control to the shear zone hosted precious metal vein deposits in the Mammoth survey area are related to compressional tectonics, followed by a period of more open folding and extensional faults. The latter two would have controlled the orientation and distribution of the veins. The vein mineralization would have been emplaced in these pre-existing structures during later Jurassic intrusive activity.

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FIGURE1B CLAIM MAP FOR PROPERTY AREA

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TABLE 1: MAMMOTH PROJECT, CLAIM INFORMATION

Claim Name	Units	Tenure No.	Current Expiry Date	
TNT L 14695	1	232603	July 7. 2003	
TNT FR 14880	1	232604	u "	
MAMMOTH FRACTION LI5034	1	232605	u 1 9	
MAMMOTH NO. 4 LI5035	1	232606	u "	
MAMMOTH NO. 3 LI5036	1	232607	16 19 -	
MAMMOTH NO.2 L14694	1	232630	March 13, 1998	
MAMMOTH L.14692	1	232601	July 7, 2005	
MAMMOTH NO.1 L14693	1	232602	4 ^{, , , _} ,	
COPPER 1	1	341590	Oct. 30, 1997	
COPPER 2	1	341591	64 yı	
COPPER 3	1	341592	4 9	
COPPER 4	1	341593	4 P.	
MAMMOTH 5	18	350530	Sept. 13, 1997	
МАММОТН 6	18	350531	4 19	
MAMMOTH 7	1	350540	4 11	
MAMMOTH 8	1	350541	4 7	
МАММОТН 9	1	350542	ч т	
MAMMOTH 10	1	350543	66. 19	
MAMMOTH 11	1	350544	4 11	
MAMMOTH 12	1	350545	4 9	
MAMMOTH 13	1	350546	4 7	
MAMMOTH 14	1	350547	et 11	
MAMMOTH 15	1	350548	6L 17	
MAMMOTH 16	1	350549	н т	
MAMMOTH 17	1	350550	44 39	
MAMMOTH 18	1	350551	K W	
MAMMOTH 19	1	350532	Oct. 29, 1997	
MAMMOTH 20	1	350533	Sept. 12, 1997	
MAMMOTH 21	1	350534	44. 79	
MAMMOTH 22	1	350535	4 19	
MAMMOTH 23	1	350536	"	
MAMMOTH 24	1	350537	14 Y	
MAMMOTH 25	1	350538	4. п	
MAMMOTH 26	1	350539	44 m	
MAMMOTH 27	1	350552	44	
MAMMOTH 28	1	350553	Sept. 19, 1997	

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EQUIPMENT AND DATA QUALITY

Modern state of the art geophysical instruments were used to collect the data over this property. Data quality is generally good. The report by C.M.E. referenced at the back of this report, outlines the equipment used in more detail.

DATA PRESENTATION

For this report, the original data has been redigitized and plotted in various formats to:

- a) present all the data to a common scale (1:4000).
- b) display the data in solid colour plans and sections.
- c) improve the processing and display of the data to assist interpretation.

The new data formats are as follows:

- a) contoured plans.
- b) contoured sections.
- c) composite plan of the Resistivity, Magnetics and gridded Induced Polarization data.
- d) posted data.

The contour plans and sections are also presented (in colour) at a reduced scale (approx. 1:12700) for convenience in viewing the data in a single page format.

Contour plans and sections give a good spatial view of the data's intensity and continuity in three dimensions. The gradient array sections in particular help to define the depth and shape of the body causing the anomalous response.

The multiple gradient array approach has an important advantage over the dipole-dipole array in the horizontal and vertical resolution of anomalies, while maintaining an excellent depth of investigation.

The colour plans and sections that accompany this report have all been prepared so that lowest numerical values are the blue colour, whereas the higher numerical data values are red to purple. The colour spread shown on the Induced Polarization and Resistivity sections apply to that section line only. Using one colour range standard for all the sections results in a loss of definition for certain sections, thus was not used. The colour contour plans do however accurately show the relative magnitude of the response from line to line.

SURVEY PROCEDURE

The report by C.M.E. referenced at the back of this report documents the field procedures used in the collection of the data.

Some comments on the survey methodology are in order.

Dipole-Dipole Survey:

The 25m dipole size, in conjunction with readings at six N spacings, has resulted in a shallow looking (80 meter), but detailed (horizontal resolution) survey. The double peaking effect that occurs with the dipole-dipole array over shallow depth extent limited bodies, is quite apparent in the data. Complications in interpretation of the data arose when adjacent zones interfered with each other thru the double peaking process.

This type of survey is quite expensive to run in rugged topography. The dipole-dipole data has been draped over the topography to improve the presentation, thus the depth scale is the elevation above sea level.

Gradient Array Survey:

The gradient array electrode configuration was used for the bulk of the I.P/Resistivity work. This was a good choice since this array is more cost effective in rugged terrain. The basic grid coverage was completed with a current electrode separation (AB) of 1500 meters. The potential electrode separation (MN) was fixed at 50 meters. The focus depth of this size array is approx. 260 meters. Targets in the 100 to 300 meter depth range will be well revealed by this survey, however shallow targets of moderate depth extent will be out of focus and poorly detected.

Note that it is essential to keep the "MN" distance as small as signal levels will permit, in order to achieve good horizontal resolution of narrow vein systems. A smaller "MN" size would have been more appropriate for this survey.

Overlap on each reading should be 50%. Unfortunately for this survey, the data was taken at 50 meter intervals with 50 meter dipoles, thus no overlap. The larger measuring dipole, in conjunction with large moves, has resulted in reduced resolution of narrow, near surface veins of limited depth extent. Survey coverage of this size grid would have required several gradient array blocks. A shift in the data value often occurs when switching current electrode locations (blocks). This shift can be evaluated by overlapping the block coverage. The continuity and amplitude of the data between blocks does appear good for this survey, thus it's probable the contractor just averaged the overlapping lines of data to achieve a base level.

The gradient array chargeability and resistivity sections were constructed by surveying each line with four different size gradient arrays (AB's), focused at different depths. Current electrode separation varied from approx. 200 meters to 1500 meters. The data from each different current electrode separation (AB's) was assigned a focus depth of investigation based primarily on the current electrode separation.

The focus depth of investigation the contractor chose for this survey is 0.2 times the current electrode separation. At least 60% of the current transmitted into the ground would have been flowing above this depth, thus it was a good choice for the display depth.

The depth of investigation of a particular array size can be accurately determined from the simultaneous inversion of chargeability and resistivity depth sounding data, in conjunction with geologic knowledge of the stratigraphy from borehole and surface outcrops. The information obtained from one survey area can generally be applied to similar geologic and geophysical settings elsewhere.

The depth scale shown on the Gradient array I.P/Resistivity sections is the depth below the ground surface and assumes flat topography. When detail topographical information (elevations) are available, the data should be draped with the topography for a more accurate presentation. In this case, the depth scale becomes the elevation above sea level. This step is essential when dealing with shallow dipping or flat lying mineralized zones in steep terrain, however it is not as important when searching for near vertical mineralized zones.

Magnetic Field Strength Surveying:

A common feature of the interesting Porphyry occurrences in the Rossland Group is fracture controlled chalcopyrite-pyrite and magnetite mineralization within the porphyritic intrusives and spatially related volcanics. The above is obviously the basis for the extensive Induced Polarization, Resistivity and Magnetic surveys that have been conducted within the Rossland Group. Strong magnetic anomalies appear to be indicative of the Diorite Porphyry.

In addition, intense hydrothermal alteration processes along shear zones can destroy existing magnetite mineralization, therefore very localized magnetic field strength lows, peripheral to or along suspected shear zones, could be indicative hydrothermal activity that might have created quartz sulphide veins.

Pyrrhotite mineralization in veins may give quartz sulphide veins a magnetic signature, although not nearly as strong as magnetite mineralization. Pyrrhotite in veins is extremely variable and erratic in its magnetic susceptibility, thus the presence or absence of a magnetic anomaly is not definitive.

The magnetic data was acquired at 25 meter intervals along lines 100 meters apart. This reading density is barely adequate for the exploration models and in the future, readings should be acquired at 10 or 12.5 meter intervals, with the line interval changed to 50 meters in areas of interest.

DISCUSSION OF THE DATA

A perusal of the Resistivity Plans, Figs #3 & 10, helps give a basic understanding of the grid geology. The very high resistivity mass centered at 3800N, 2500E, is very likely related to the Bonnington Pluton, which obviously has a very low sulphide content, as indicated by the very low chargeability response.

irregular N-S trending contact zone between the An Bonnington and the Elise Formation volcanics exists at approx. however it appears the Bonnington has partially 2700E, assimilated the Elise Formation. The lower part of the Elise Formation appears to have a very low sulphide concentration with respect to the upper part of the Elise. A strong continuous linear, but narrow I.P. response probably marks the contact of the Upper Elise Formation with the Hall sediments. This I.P. anomaly (contact zone) is likely due to the presence of a major shear structure (at approx. 3000E), that should be extensively prospected along its course.

The partial assimilation of the Elise Formation by the Bonnington Pluton postulated above, appears to have partly destroyed a major shear zone at approx. 2800E. This shear is revealed by the relatively low N-S oriented resistivities at this location. At its southern end, this shear is called the Marcus To the north, this postulated shear zone passes thru the Shear. Mammoth skarn zone, however appears to be obliterated by later intrusive activity immediately to the north of the Mammoth. A northeast trending splay may have developed just south of the This splay joins the main shear zone postulated Mammoth zone. above at approx. 3050E, 4200N. An ancillary, but much smaller shear (the Stephie) is also evident at approx. 2900E in the resistivity data.

A third, very significant silicified shear zone appears to cut thru the Hall Formation sediments at approx. 3700E. The eastern flank of this proposed broad shear zone has a significant Induced Polarization response, whereas the western flank has a weaker, but still significant I.P. response. A possible explanation for the increased sulphide content on the east side of this proposed shear may be due to the existence of a coarser, perhaps more permeable, sedimentary unit of the Hall Formation. It is not unlikely to expect the shear to occur along a lateral facies change within the Hall Formation. To the north, this postulated shear zone appears to be faulted to the east by a northwest trending fault. The broad area of increased I.P. response centered within a 300 meter radius of 3400E, 4050N, is probably related to a well mineralized sill-like intrusive body at depth, contained within folded Hall Formation rocks. The northwest trending fault mentioned above is probably related to the intrusion of this body. This hypothesis will be discussed further in the detail section of each section line. It is important to note that all of the geophysical data supports the concept of a major flexure in the stratigraphy centered around this anomaly.

The magnetic survey data indicates the diorite porphyry is generally strongly magnetic and is often intruded along faults and shear zones to form narrow lenticular anomalies.

The large generally northwest oriented magnetic low centered around 4600N, 3200E, may be indicative of a large felsic intrusive body, possibly of very late Jurassic age. This area has received no I.P/Resistivity coverage.

The very strong localized magnetic low centered at 4050N that occurs along the postulated shear at 3000E, is an interesting feature that may be indicative of hydrothermal alteration along the shear for a considerable distance (400+ meters). The proximity of this magnetic low to the Mammoth showing should be noted. The linear N-S oriented narrow magnetic low centered at 4000N, 3700E, may also be a feature indicative of hydrothermal activity at the faulted north end of the relatively broad shear zone postulated to exist at 3700E.

The magnetic data can indicate where a potential auriferous shear zone has been partially destroyed by the injection of large dykes of porphyritic mafic rock (diabase and/or diorite). The attenuation of the I.P. response over the postulated 3000E shear zone at 3500E, may be indicative of the above concern.

The isolated magnetic low at 3100E and centered at 2950N is coincident with a modest, but distinct I.P. response that could be indicative of a narrow vein system possibly a splay off the proposed 3000E shear zone.

A discussion of each line where a data section was recorded follows:

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Dipole-Dipole Section Line 4300N:

- a very weak near surface I.P. anomaly occurs at 2940E. Since this anomaly has very little depth extent, it was not detected by the deep looking gradient I.P. survey. This response may mark the northern end of the Mammoth zone mineralization.
- a strong near surface I.P. anomaly occurs at 3040E that has appreciable depth extent (75 meters). This response was also detected by the deep gradient I.P. survey as part of the postulated major N-S shear defined at approx. 3000E.
- from approx. 3250E to 3550E, there are a series of near surface narrow depth limited I.P. anomalies that have the appearance of narrow sulphide veins above a deeply buried mineralized intrusive body.
- some moderate fault dislocation appears to occur at 3400E.

Dipole-Dipole Section Line 4200N:

- again, a weak near surface I.P. anomaly of poor depth extent occurs at 2950E. This response is likely arising from the Mammoth zone mineralization.
- a strong near surface I.P. anomaly of appreciable depth extent (100 meters). occurs at 3000E. This response was also detected by the deep gradient I.P. survey as part of the postulated major N-S shear.
- a series of narrow depth extent limited near surface I.P. anomalies exist between 3200E and 3450E. This series of anomalies again has the appearance of narrow veins over a deeply buried mineralized intrusive body.
- a strong deeply buried (50 meters) anomaly occurs at 3600E at the eastern edge of the data. Moderate fault dislocation at approx. 3550E may be responsible for isolating this anomaly from the abovementioned cluster.

Dipole-Dipole Section Line 4100N:

- a weak near surface I.P. response of poor depth extent occurs at 2930E. This anomaly is also related to the Mammoth zone mineralization. The poor depth extent of this anomaly is clearly shown in the corresponding gradient array section. a much stronger near surface I.P. response of moderate (50m) depth extent occurs at 2980E. This anomaly is part of the postulated major N-s shear zone defined by the deep gradient survey.

from approx. 3160E to 3560E a series of narrow near surface I.P. responses of moderate to good depth extent occur. As mentioned before, these responses could originate from a deeply buried mineralized intrusive body. This possibility will be discussed further in the discussion of the deep gradient array section for L.4000N.

- the low near surface resistivities encountered around 3500E suggest the Hall Formation at this point is probably carbonaceous argillite with little silicification.

Gradient Array Section, Line 4100N:

- unfortunately the gradient section is not as laterally extensive as the dipole-dipole section. It is important to extend the deep looking gradient sections further to the east.
- the poor depth extent of the anomaly at 2930E (Mammoth zone mineralization), is clearly illustrated in the section.
- the moderate depth extent of the postulated shear zone anomaly at 2980E is also well illustrated, however there is some indication this anomaly exists to a deeper level, although much attenuated. The resistivity data clearly shows this anomaly is hosted in very high resistivity rocks, likely the Elise Formation.
- another narrow near surface I.P. response of moderate depth extent was detected at 3060E. This response was not clear in the dipole-dipole data, which is a reflection of the higher resolution of the gradient array. The anomaly lies on the contact between the Elise Formation and the Hall sediments and is likely part of the postulated N-S shear zone.
- at approx. 3160E, a narrow strong near surface I.P. anomaly of good depth extent exists. Near the surface, this anomaly shows a steep west dip, however becomes vertical at depth.

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- two I.P. anomalies very similar to the above also occur at 3260E and at 3350E, however the latter response is only partially defined by the present survey.
- the source of these narrow vein-like responses must lie at depth, most probably a mineralized intrusive.

Gradient Array Section Line 4000N:

- a strong broad I.P. response of limited (50m) depth extent exists from approx. 2800E to 2980E. This anomaly is likely due to Mammoth mineralization and a mineralized shear zone at 2925E. The resistivity suggests this shear dips steeply to the east in the near surface, then becomes vertical.
- a second strong, but narrow I.P. response of moderate depth extent (80m) exists at 3020E. This anomaly probably marks the contact of the Elise Formation with the Hall Formation, but also is part of the N-S major shear zone postulated at 3000E.
- a modest broad I.P. response of poor depth extent occurs at 3230E.
- the approximate horizontal contouring seen in the resistivity data between 3200E and 3700E, is an interesting A near surface (70m) synclinal shaped, very low feature. resistivity zone exists between 3300E and 3700E. Beneath low resistivity zone (at 150m depth) the is a thick, relatively higher resistivity zone (700 ohm-m). This zone likely caused by a large sill-like well mineralized is – porphyry body dipping gently to the west. A flat lying low resistivity zone occurring over a thick strongly polarizable high resistivity can cause the strong negative I.P. responses seen at shallow depth in this section.

Further to the west, this anomaly appears extensively faulted, although the apparent fault zones are strongly silicified as evident from the high resistivities. To the east, this anomaly appears to come very close to the surface between 3625E and 3750E. The question also arises here of a possible roof zone, with the Hall Formation acting as a trap to contain sulphide rich solutions within an underlying porphyry.

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The major N-S shear zone postulated for the approx. 3700E area offsets the above anomaly and may, in part, be responsible for the stronger near surface I.P. responses in this area.

- a strong narrow near surface steeply east dipping I.P. response of excellent depth extent occurs at 3830E. This anomaly is hosted in very low resistivity Hall Formation rocks. The I.P. response is closely coincident with a narrow zone of high resistivities to the west, which suggests silicification along a structure.
- a broad strong I.P. response was also detected at the eastern end of the line, in conjunction with higher resistivities. The top of this anomaly lies at some depth (30 meters), and it has only been partly defined by the survey.

Several strong narrow magnetic responses occur between 3000E and 3800E. The poor line to line correlation of these anomalies and interferences from multiple, closely spaced responses, suggests this important area has been poorly sampled by the magnetic survey. A more detailed magnetic survey would help in the evaluation of this area.

The stronger magnetic responses are probably related to near surface magnetite mineralization, possibly mafic dikes. The correlation of many of the weaker magnetic responses with I.P. responses suggests pyrrhotite mineralization in very narrow (cm scale), near surface fracture zones within the Hall Formation. This geophysical feature supports the concept of a buried, well mineralized porphyry body at depth.

The significant sulphide showing, the Linny (4025N, 3525E), lies within this apparent mixed sequence of Hall Formation and porphyry intrusives. The apparent high background sulphide content of the immediate area (3000E to 3700E), as indicated by the I.P. data, is an important consideration. More detail I.P. section work is required on adjacent lines to fully develop this exploration target.

Gradient Array Section Line 3800N:

a narrow, but strong, very steeply east dipping I.P. anomaly of good depth extent occurs at 3710E. This anomaly lies within a weakly conductive host rock, probably the Hall Formation argillite. On the immediate west flank of this anomaly, there is a thin high resistivity anomaly, possibly due to silicification along the major N-S shear postulated for the area. The proximity of this significant I.P. response to the shear zone is an important consideration. The resistivity data also suggests a moderate vertical displacement of the rock units occurs along the proposed shear with the eastern side uplifted.

A deep, thus much weaker I.P. anomaly occurs at 3890E. This anomaly pinches out toward the surface.

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Gradient Array Section Line 3300N:

- a narrow strong I.P. response of good depth extent occurs at This anomaly does not appear to subcrop. 3000E. The anomaly lies within low resistivity rocks, but immediately on the flank of very high resistivity rocks to the west. The anomaly at 3000E is part of the major N-S shear postulated for the area. This section should have been extended to the west (to 2600E), to cover the southern projection of the Marcus mineralization. The fact that no I.P. response appears to occur with the Marcus mineralization is due to the depth focus of the gradient array in conjunction with limited depth extent of the mineralization. A shallower looking array would probably have detected an I.P. response.
 - a second narrow, but strong I.P. response of good depth extent occurs at 3100E. This anomaly also does not appear to subcrop, although immediately to the northeast is the Sarah mineral showing. The I.P. anomaly lies with low resistivity rocks, but appears intimately associated with a large zone of near surface high resistivity rocks that do not have good depth extent. This zone of high resistivity centered at 3150E, terminates at approximately the 150 meter depth, where it is underlain by very low resistivity (Hall formation?) rocks. It is important to note that the southern extension of this significant I.P. anomaly correlates directly with an arcuate N-S trending magnetic low.

a strong near surface I.P. response of good vertical depth extent also occurs at 3740E. This response lies within low resistivity rocks, but immediately on the east flank of a strong vertical zone of high resistivity. This zone of high resistivity is likely part of the broad major N-S shear proposed for the area.

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CONCLUSIONS AND RECOMMENDATIONS

The geophysical survey work has outlined four apparent shear zone structures that deserve further exploration work. The location of these structures is shown on the page size plan maps of the Chargeability and Resistivity (Figs. #2 & 3). As these structures come very close to surface, detailed prospecting and trenching along their course may prove helpful. If too much overburden is encountered, a shallow drilling program is warranted.

The Marcus shear mineralization appears to have strike length potential, both to the north and south.

The Mammoth skarn mineralization does not appear to have good depth extent, although the apparent shear zone immediately to the east (in the magnetic low area), is an interesting exploration target.

The Sarah showing lies to the immediate northeast of a significant narrow lenticular I.P. response that correlates directly with a distinct magnetic low extending further to the south.

The Linny showing exhibits some very high grade mineralization and lies near the centre of the area postulated to contain a buried well mineralized sill-like porphyry body approx. 350 meters in diameter. More detail I.P/Resistivity section work on adjacent lines would help develop this proposed drill target.

The present I.P/Resistivity survey coverage should be expanded to the same size as the Magnetic survey. In addition, small detail grids in areas of high interest would help delineate trenching and/or drill targets better. The data sampling intervals should be reduced to 10 meters, both for the I.P/Resistivity and Magnetics. The focus depth of the I.P. survey must be more seriously considered in the exploration of shear zone quartz sulphide vein systems.

In this geologic setting (shear zone hosted quartz sulphide veins), hydrothermal activity in a well developed porous and permeable portion of a structure, would be the important ore formation process. Perhaps the best geophysical evidence that this process has occurred, would be long linear closely coincident zones of increased I.P. and Resistivity resulting from silicification and sulphide deposition.

HENDRICKSON Grant A. Hendifckson, P.Geo.

STATEMENT OF QUALIFICATIONS

Grant A. Hendrickson.

- B.Science, University of British Columbia, Canada, 1971. Geophysics option.
- For the past 26 years, I have been actively involved in mineral exploration projects throughout Canada, the United States, Europe and Central and South America.
- Registered as a Professional Geoscientist with the Association of Professional Engineers and Geoscientists of the Province of British Columbia, Canada.
- Registered as a Professional Geophysicist with the Association of Professional Engineers, Geologists and Geophysicists of Alberta, Canada.
- Active member of the Society of Exploration Geophysicists, European Association of Geoscientists and Engineers, and the British Columbia Geophysical Society.

Dated at Delta, British Columbia, Canada, this $\underline{2}$ day of \underline{MAY} , 1997.

FNDRICKSON CAGEN Grant A. Hendr P .Geo.

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Fig. #23-4.





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		Posted data, (mV/V) Gradient array data, AB=1500m, MN=50m Iris & Huntec instruments data provided by client DELTA GEOSCIENCE LTD, fig #9B



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2300	2400	2500	2600	2700	2800	2900	3000	3100	3200	3300	3400
0-+	+	+	+	+	+	+	+	+	+	+	+
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+	+	+	+	+	+	+	+	÷	+	+	+
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2300	2400	2500	2600	2700	2800	2900	3000	3100	3200	3300	340
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2300	2400	2500	2600	2700	2800	2900	3000	3100	3200	3300	3400
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3- +	9003 - +	15806 - 19822 -	8675 - +	8752 · + 11154 ·	7899 +	18787 + 8312	1979 +	128 + 1120	1265 + ·	⁵⁵ + ⁵⁵	₽ • +
2200	2400	2500	2600	2700	2800	2900	3000	3100	3200	3300	3400

2600	2700	2800	2900	3000	3100	3200	3300	3400	3500	3600	3700
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			28.8 12.7	38.9 - 35.8 - 28.8 -	20.5 36.5 34.9	6.4 3.2- 12.7-				22	10.2 ·
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+ - 500	+	+	19.4 +	29.8 95.7 34.2 +	20.5 57.7 17 15	18.5 +	+	+	+	28 - 38.7∔ 4 1 -	2,3,3 2,3,1 19,1
+	5.9 + ^{8.2}	^{8.5} +	24 .2- +	34.7- 34.5	37.2 11.9- +	11.9 + 14.2-	9.5- +	15.4 +	8. + ^{9.3}	21.2- 21.2-	19.9- +
2600	2700	2800	2900	3000	3100	3200	3300	3400	3500	3600	3700

2600	2700	2800	2900	3000	3100	3200	3300	3400	3500	3600	3700
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+ -100	+	+	1040 - 15 2068 - <mark>9</mark> 11	101 15 963 15	1388 108 748 48 53 + 48	2179 2505 · 17	+	+	+	ې + م	⁴ ⁴ ⁴ ⁴
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2600	2700	2800	2900	3000	3100	3200	3300	3400	3500	3600	3700

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GEOLOGICAL SURVEY BRANCH ASSESSMENT REPORT 28, Scale 1:4000 100 150 50 200 250 (meters) ROSSMIN EXPLORATIONS LTD MAMMOTH PROPERTY CHARGEABILITY SECTION, LINE 4200N NELSON / SALMO AREA, BRITISH COLUMBIA Contour interval 2 mV/V Dipole-Dipole data, a=25 m, N 1-6 Iris instruments data provided by client DELTA GEOSCIENCE LTD. fig # 23

