

# **GEOPHYSICAL REPORT**

on an

## **INDUCED POLARIZATION SURVEY**

### **PIL NORTH PROJECT**

Latitude 57° 19' N, Longitude 126° 55' W

Omineca Mining Division, N.T.S. 94E/7W

British Columbia, Canada

### **Finlay Minerals Ltd.**

Vancouver, B.C.

Canada

Survey by

***SJ GEOPHYSICS LTD.***

Report by

***S.J.V. CONSULTANTS LTD.***

E. Trent Pezzot, Geophysicist.

November 28, 2003

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## **1. SUMMARY**

During August and September, 2003, SJ Geophysics Ltd. completed approximately 16.6 km of Induced Polarization surveying on Finlay Mineral Ltd.'s PIL property in the Toodoggone area of B.C. The survey extended reconnaissance exploration on copper-gold porphyry targets mapped in previous years.

The Central Zone, previously tracked to line 3600N has been extended some 2 km along strike to line 5600N. While the zone appears to be narrowing it is still considered open to the northwest. A small satellite chargeability anomaly has been identified to the southeast of the central zone.

A sharp resistivity contact is mapped immediately northeast and below the Central Zone chargeability anomaly. This suggests the anomaly may be related to a regional geological contact.

The Northeast Zone was previously thought to terminate near line 3200N. This recent surveying shows anomaly continues to the northwest to line 4000N and it is considered open along strike.

A previously undetected trend, delineated by a unit exhibiting elevated resistivity and elevated chargeability, is mapped along the eastern halves of lines 3600N to 5200N. This response is present at depth (>150 metres) and likely maps a discrete geological unit.

2D inversion processing completed on these data provided significant refinements to the interpretation over what was available from pseudosection plots alone. Consequently, IP data gathered in previous years was reprocessed and inverted. Plan maps provided with this report combine all of the IP results available.

As a result of this inversion processing, previous drill recommendations should be re-evaluated.

## **2. INTRODUCTION**

Finlay Minerals Ltd. commissioned SJ Geophysics to conduct a magnetometer and IP survey across a portion of their PIL Claims in the Toodoggone mining camp. Similar surveys conducted in 2001 and 2002 outlined four high chargeability zones that fit the copper-gold porphyry target being sought. These latest surveys were intended to extend geophysical coverage for these targets along strike to the northwest.

Previous surveys were conducted on ~400 metre spaced lines from 800N to 3600N. The current survey extended coverage (at 400 metre spacings) from line 4000N to line 5600N. In addition, Line 3600N was extended to the northeast in order to better delineate the Northeast zone and Line 2500N was surveyed to fill-in a gap in the previous coverage.

The IP data was processed through the UBC Inversion program (DCIP2D) to produce interpreted depth sections. Encouraging results prompted an extension of the inversion processing to include data from the previous surveys.

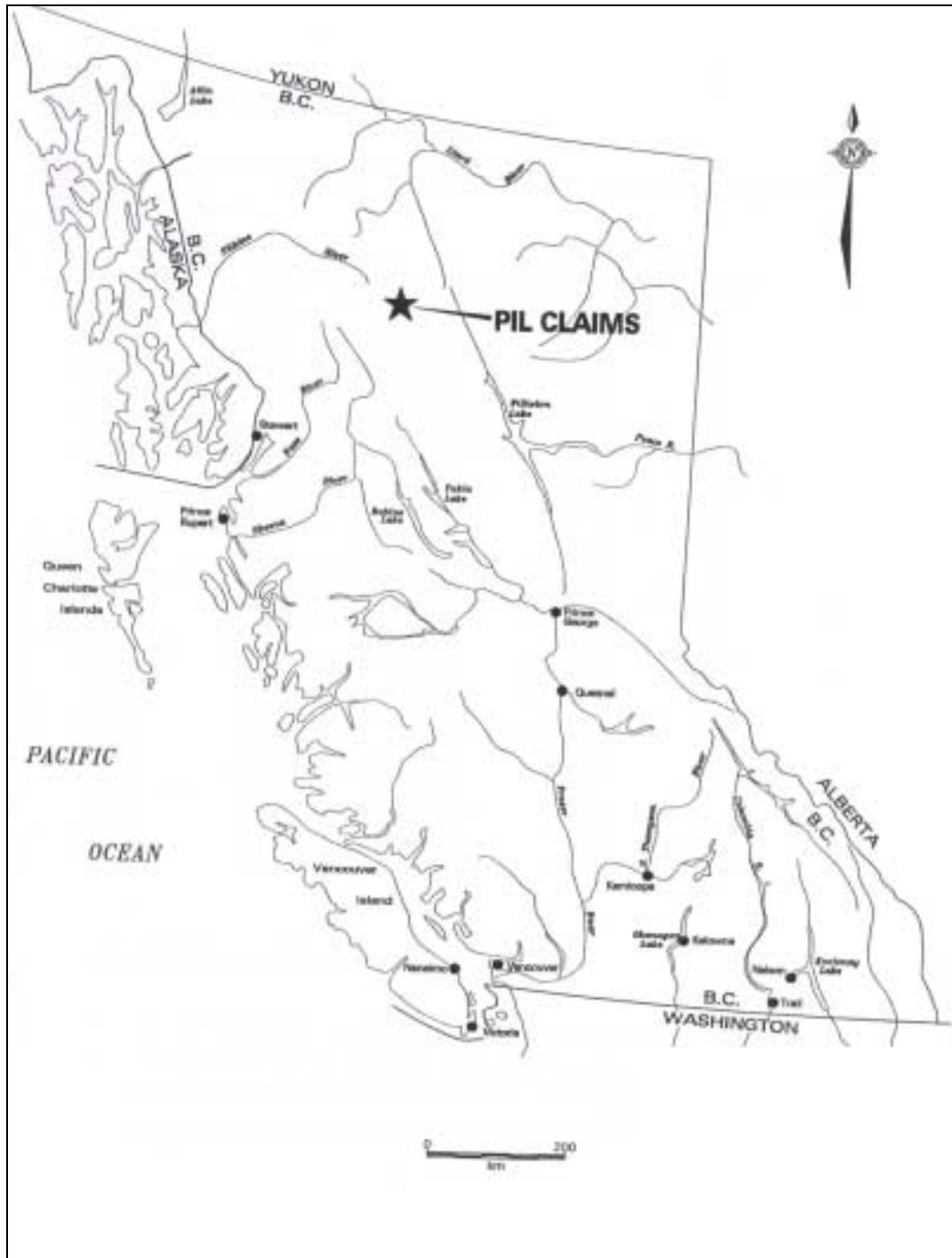
Due to time and logistical constraints, the proposed magnetometer surveying was cut from the program.

## **3. PROPERTY LOCATION AND ACCESS**

The Pil claims lie within a broad region of prospects and mines known as the Toodoggone mining camp. They straddle the west flowing Jock Creek, north of the Finlay River. The claims are located in the Omineca Mining Division on NTS map Sheet 94E/7W. Approximate geographical coordinates are latitude 57° 19' N and longitude 123° 55' W. Access to the property for this survey was by helicopter from the Finlay camp, located north of the Kemess mine.

The camp facilities are accessible by road from the city of Prince George. Under normal driving conditions the trip takes approximately 12 hours. From Prince George, travel north towards the town of Mackenzie (130 km) along Hwy 97. Just before MacKenzie, turn left onto the Finlay Forest service road that leads northwest to the Kemess mine. The turnoff towards Kemess mine is at a junction called Windy Point. Approximately 320 kilometres down the road, you must take a left and follow a gravel road that leads to the Sturdee Strip landing pad. At Sturdee Strip, you take a right and

drive for about 5 km and then turn right onto another gravel road. Follow this road until you come to an old shaft mine, unmarked, right beside a river. Go through the old mining site and follow the small gravel road alongside the river for about 10 km to reach the Finlay camp.



**Figure 1: Location Map**

#### 4. CLAIM HOLDINGS

The following claim information was provided by Finlay Minerals Ltd.

| Tenure Number | Claim Name | Issue Date | Good Standing To | Units |
|---------------|------------|------------|------------------|-------|
| 308127        | PIL 1      | 14-MAR-92  | 31-Jan-04        | 8     |
| 308128        | PIL2       | 14-MAR-92  | 31-Jan-04        | 20    |
| 316950        | PIL 4      | 29-MAR-93  | 31-Jan-04        | 20    |
| 316951        | PIL 5      | 29-MAR-93  | 31-Jan-04        | 15    |
| 316952        | PIL 6      | 29-MAR-93  | 31-Jan-04        | 12    |
| 316953        | PIL 7      | 29-MAR-93  | 31-Jan-04        | 20    |
| 316955        | PIL 9      | 29-MAR-93  | 31-Jan-04        | 16    |
| 316956        | PIL 10     | 29-MAR-93  | 31-Jan-04        | 18    |
| 316957        | PIL 11     | 29-MAR-93  | 31-Jan-04        | 20    |
| 319649        | PIL 12     | 21-JUL-93  | 31-Jan-06        | 20    |
| 319650        | PIL 13     | 21-JUL-93  | 31-Jan-04        | 20    |
| 340215        | PIL 20     | 16-SEP-95  | 31-Jan-04        | 9     |
| 340216        | PIL 21     | 16-SEP-95  | 31-Jan-04        | 16    |
| 340217        | PIL 22     | 16-SEP-95  | 31-Jan-04        | 16    |
| 340218        | PIL 23     | 17-SEP-95  | 31-Jan-04        | 18    |
| 340219        | PIL 24     | 16-SEP-95  | 31-Jan-04        | 1     |
| 340220        | PIL 25     | 16-SEP-95  | 31-Jan-04        | 1     |
| 340221        | PIL 26     | 16-SEP-95  | 31-Jan-04        | 1     |
| 340222        | PIL 27     | 16-SEP-95  | 31-Jan-04        | 1     |
| 340223        | PIL 28     | 16-SEP-95  | 31-Jan-04        | 1     |
| 340224        | PIL 29     | 16-SEP-95  | 31-Jan-04        | 1     |
| 340225        | PIL 30     | 16-SEP-95  | 31-Jan-04        | 1     |
| 340226        | PIL 31     | 16-SEP-95  | 31-Jan-04        | 1     |
| 340227        | PIL 32     | 16-SEP-95  | 31-Jan-04        | 1     |
| 340228        | PIL 33     | 16-SEP-95  | 31-Jan-04        | 1     |
| 370563        | LIP 1      | 28-JUL-99  | 31-Jan-04        | 1     |
| 370564        | LIP 2      | 28-JUL-99  | 31-Jan-04        | 1     |
| 370565        | LIP 3      | 28-JUL-99  | 31-Jan-04        | 1     |
| 370566        | LIP 4      | 28-JUL-99  | 31-Jan-04        | 1     |
|               | 29 CLAIMS  |            | TOTAL UNITS      | 262   |

*Table 1 – Claim Information*

As this report is being prepared as an addendum to a larger report, readers are referred to this parent document for a map showing the claim outlines.

## **5. GENERAL GEOLOGY**

The Toodoggone area lies within the Intermontane Belt, between the east end of the Stikine Arch in the north and the Skeena Arch in the south. Geology along the east-northeast margin of the Stikine Terrane is dominated by successive volcano-plutonic arcs which were constructed from Permian time, but most importantly, during the late Triassic and early Jurassic. The Toodoggone area lies within a north-northwest trending corridor of Mesozoic island-arc magmatism.

Two supracrustal units are important hosts of mineralization in the Toodoggone mining camp. Volcanics of the Takla Group host the Kemess Mine (Kemess South deposit), and numerous porphyry prospects. The Toodoggone Formation of the Hazelton Group is the most important stratigraphic unit in terms of epithermal precious metal deposits.

## **6. FIELD WORK AND INSTRUMENTATION**

SJ Geophysics' crew consisted of 5 employees. Ron Sheldrake, John Wilkinson and Brian Chen mobilized from Eskay Creek to Finlay camp from August 22 to August 23. While the remaining two, Travis Montgomery and Chad Barzen mobilized from Vancouver and were picked up at the Prince George airport on August 23 and joined the crew for the 12 hour drive into the Finlay camp.

Induced Polarization production started on August 24 and was completed on September 9 with only one day of standby, a result of the helicopter being grounded due to fog and a low cloud base. The IP data were collected using IRIS instrumentation: a VIP 3000 kW transmitter (Tx) with 2 seconds on, 2 seconds off duty cycle, and an ELREC-10 receiver (Rx). In addition to the IP measurements, clinometer readings and gps control points were taken along each line. An expanding pole-dipole array was used and for each line a new remote current (infinity) electrode was laid out. Steel rods were used for the electrodes. Discussions of the geophysical methods used on these surveys are discussed in Section 7 "Geophysical Techniques".

Each evening in the field camp, the day's IP measurements, gps readings and clinometer data were downloaded to the field computer system where it was then processed and the quality of the data was reviewed. At the completion of each line, a



preliminary inversion of the IP measurements for that line was run and shown to the Finlay Minerals' field representative.

IP surveying was completed on 8 line segments totalling 16.625 kilometres as shown below in Table 2: Line Parameters. The sequence of surveying was Line 44N, 40N, 52N, 48N, 56N, 25N, an extension to Line 40N and Line 36N. Line spacing was 400m for lines 36N through 56N and Line 25N further south was off on its own. Highly resistive ground was encountered at the western end of Line 40N. As a result of the inability to record adequate data for this region, stations 1500E to 2000E were dropped from the program. Also with Line 40N, the initial 2D inversion results and pseudosections showed a high chargeability feature at the east end of line. After showing this to the clients' field representative, the client decided to extend Line 40N 500m to the east and also to extend Line 36N (from a previous survey) by 650m.

| <i>Line</i> | <i>Start Station</i> | <i>End Station</i> | <i>Length (m)</i> | <i>Azimuth</i> |
|-------------|----------------------|--------------------|-------------------|----------------|
| 25N         | 1200E                | 4500E              | 3300              |                |
| 36N         | 4700E                | 5350E              | 650               |                |
| 40N         | 2025E                | 4400E              | 2375              |                |
| 40N Ext.    | 4400E                | 4900E              | 500               |                |
| 44N         | 1000E                | 3500E              | 2500              |                |
| 48N         | 1000E                | 3600E              | 2600              |                |
| 52N         | 1300E                | 3500E              | 2200              |                |
| 56N         | 1500E                | 4000E              | 2500              |                |

*Table 2: Line Parameters*

As a result of timing constraints caused by other commitments for the crew and snow on the prospect the magnetic survey has been postponed. Syd Visser, SJ Geophysics' president, talked to Robert Brown (Finlay Minerals Representative) and agreed to put off the survey until next year.

## **7. GEOPHYSICAL TECHNIQUES**

### **7.1. IP Method**

The time domain IP technique energizes the ground surface with an alternating square wave pulse via a pair of current electrodes. On most surveys, such as this one, the IP/Resistivity measurements are made on a regular grid of stations along survey lines.

After the transmitter (Tx) pulse has been transmitted into the ground via the current electrodes, the IP effect is measured as a time diminishing voltage at the receiver electrodes. The IP effect is a measure of the amount of IP polarizable materials in the subsurface rock. Under ideal circumstances, IP chargeability responses are a measure of the amount of disseminated metallic sulfides in the subsurface rocks.

Unfortunately, there are other rock materials that give rise to IP effects, including some graphitic rocks, clays and some metamorphic rocks (serpentinite for example) so, that from a geological point of view, IP responses are almost never uniquely interpretable. Because of the non-uniqueness of geophysical measurements it is always prudent to incorporate other data sets to assist in interpretation.

Also, from the IP measurements the apparent (bulk) resistivity of the ground is calculated from the input current and the measured primary voltage.

With regard to precision, IP/Resistivity measurements are generally considered to be repeatable within about five percent. However, they will exceed that if field conditions change due to variable water content or variable electrode contact.

IP/Resistivity measurements are influenced, to a large degree, by the rock materials nearest the surface (or, more precisely, nearest the measuring electrodes), and the interpretation of the traditional pseudosection presentation of IP data in the past have often been uncertain. This is because stronger responses that are located near surface could mask a weaker one that is located at depth.

### **7.2. Inversion Programs**

“Inversion” programs have recently become available that allow a more definitive interpretation, although the process remains subjective.

The purpose of the inversion process is to convert surface IP/Resistivity measurements into a realistic “Interpreted Depth Section.” The use of the inversion routine is a subjective one because the input into the inversion routine calls for a number of user selectable variables whose adjustment can greatly influence the output. The output from the inversion routines do assist in providing a more reliable interpretation of IP/Resistivity data, however, they are relatively new to the exploration industry and are, to some degree, still in the experimental stage.

The inversion programs are generally applied iteratively to, 1) evaluate the output with regard to what is geologically known, 2) to estimate the depth of detection, and 3) to determine the viability of specific measurements.

The Inversion Program (DCINV2D) used by the SJ Geophysical Group was developed by a consortium of major mining companies under the auspices of the UBC-Geophysical Inversion Facility. It solves two inverse problems. The DC potentials are first inverted to recover the spatial distribution of electrical resistivities, and, secondly, the chargeability data (IP) are inverted to recover the spatial distribution of IP polarizable particles in the rocks.

The Interpreted Depth Section maps represent the cross sectional distribution of polarizable materials, in the case of IP effect, and the cross sectional distribution of the apparent resistivities, in the case of the resistivity parameter.

## **8. DATA PRESENTATION**

The geophysical data from this survey are displayed in three formats, as indicated below.

### **8.1. Pseudosections**

Apparent resistivity and chargeability data is presented in a false colour contour pseudosection format. Calculated values are posted in a cross-sectional display along the survey line, with the location and depth determined by the location of the current and potential electrodes at the surface. This is a standard type of display for IP data. It has the advantage of including an estimate of the depth to anomalous bodies in addition to the surface location. One major disadvantage of this display is that it is often misinterpreted by untrained viewers as a geological cross-section. This is

misleading because the display includes patterns that are produced as a direct result of the electrode array geometry.

Pseudosections are presented in page size format in Appendix 3 and at a 1:5000 scale in map pockets at the back of this report.

## **8.2. Inverted Depthsections**

As described above (section 7.2) the IP data is processed through an inversion program that outputs one possible subsurface distribution of resistivity and polarizable materials that would produce the observed data. These results are presented in a false-colour depthsection similar to the pseudosection display. The main difference however, is that the geometric effects of the electrode arrays have been removed and these displays can be directly interpreted as geological cross sections.

Depthsections are presented in two formats: page sized along with the pseudosections in Appendix 3 and as 1:5000 scale plots in map folders at the back of this report. In the page size formats, final results are plotted with 2 colour scales. One is a customized scale based on a linear distribution about the range of values observed on that line. This display usually provides the best resolution of localized features on the line. The second is a standardized scale of 150 – 15,000 ohm-m for the resistivity and 0 – 36 msec for the chargeability. These images allow for a better comparison of the relative strengths and amplitudes of the anomalies and trends across the grid.

## **8.3. Plan Maps**

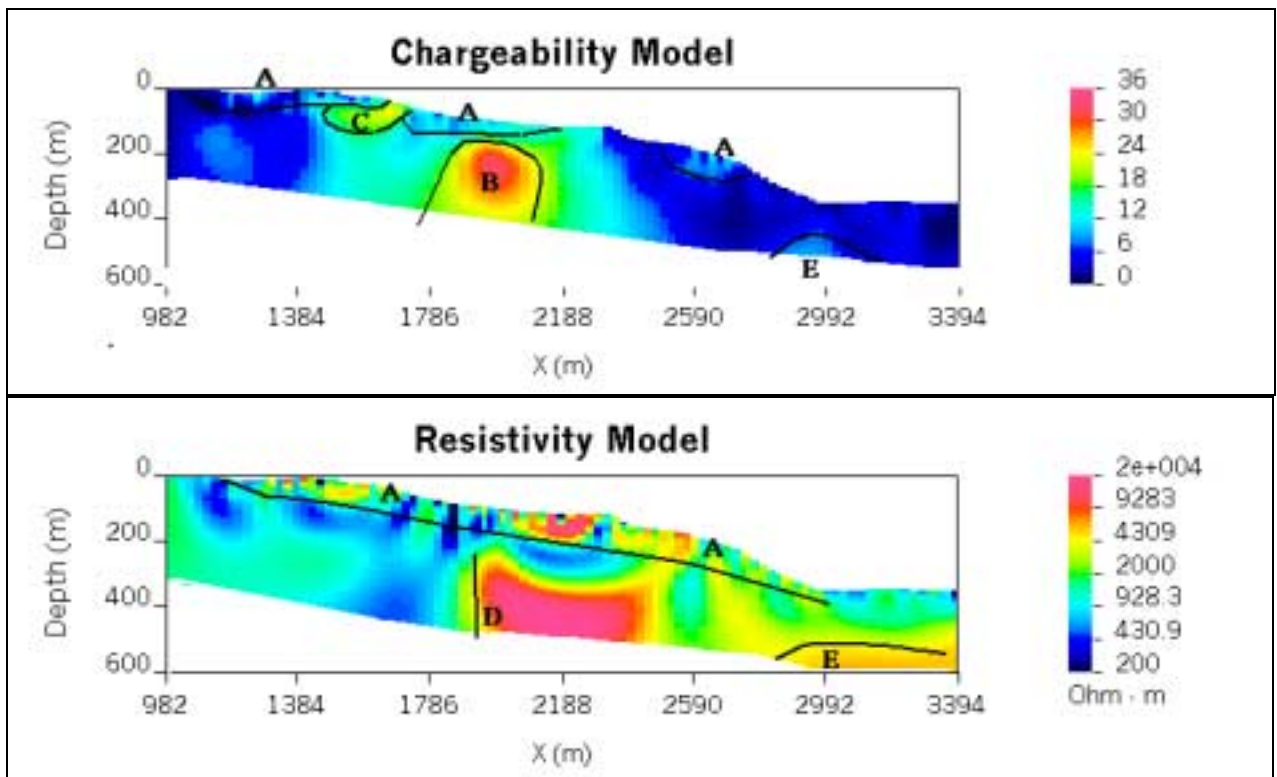
False colour contour maps of the inverted resistivity and chargeability results can be produced for selected depths. Data is positioned using UTM coordinates and plotted over a topographic base map provided by Finlay Minerals Ltd. This display illustrates the areal distribution of the geophysical trends, outlining strike orientations and possible fault offsets.

Plan maps are plotted for both resistivity and chargeability at depths of 20m, 100m and 150m below surface at a 1:10000 scale and included in map folders at the back of this report.

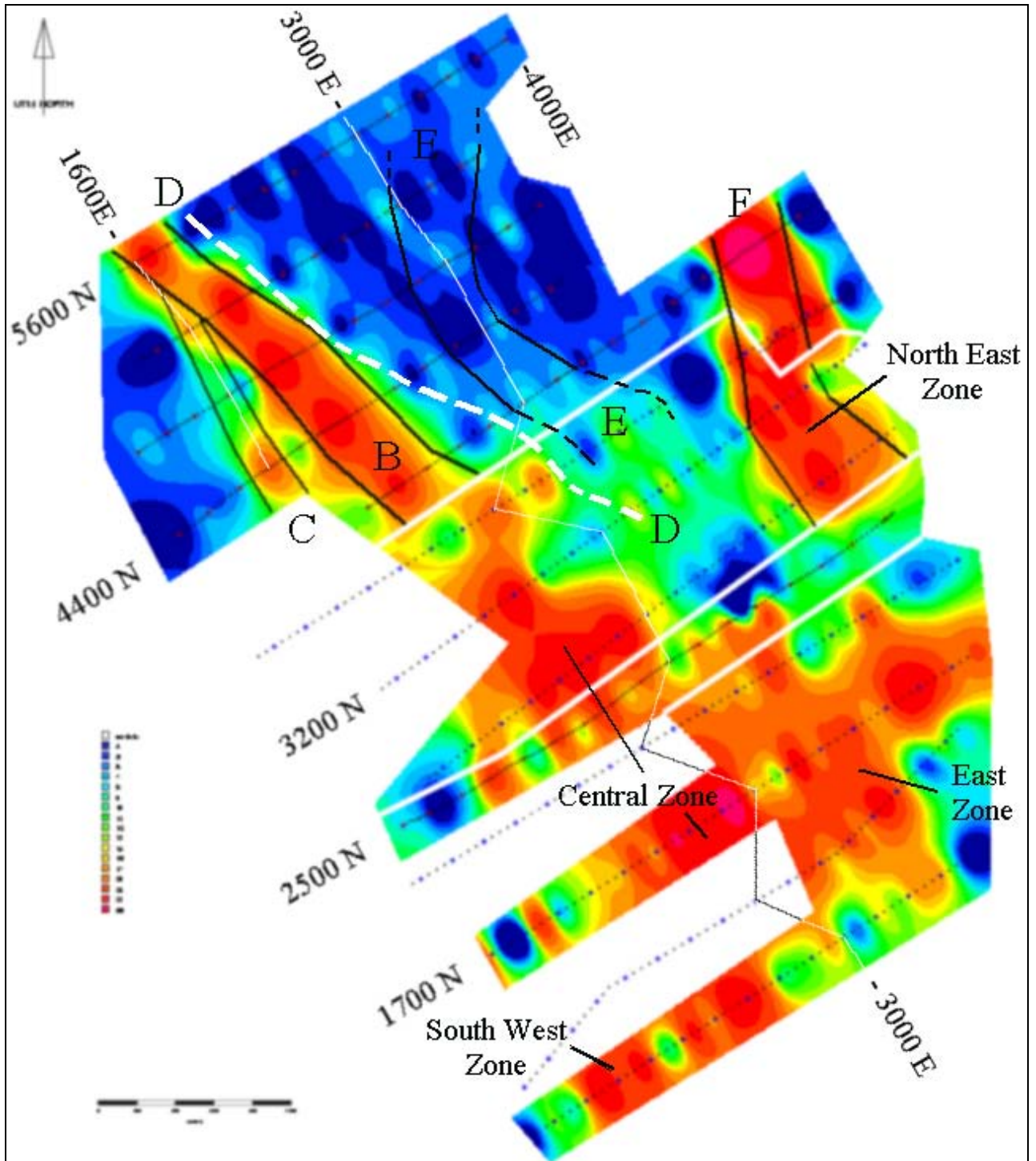
## 9. INTERPRETATION

### 9.1. 2003 Survey

The resistivity data, and to a lesser extent the IP data, typically exhibits highly variable data in the near surface (0 – 50 metres). These responses are difficult to correlate across the widely spaced lines and are interpreted as reflections of highly weathered rocks. There are however several strong chargeability and resistivity anomalies evident in the data that extend across several lines and are interpreted as mapping discrete geological trends. Figure 2 shows the inversion depthsections for line 4400N and are annotated to highlight the anomalies discussed below. Figure 3 shows the interpreted chargeability distributions at 100 metres depth for both the 2003 and previous years' data. The anomalous trends are clearly evident on this display.



**Figure: 2 – Interpreted Chargeability and Resistivity Depth Sections – Line 4400N**  
(Annotated with Anomaly and Trend Identifiers)



**Figure: 3 – Interpreted Chargeability at 100 metres depth**  
(Annotated with Anomaly and Trend Identifiers)

The continuation of the Central Zone is clearly outlined as a high chargeability trend (B) evident on lines 4400N to 5600N and is associated with a low resistivity background. The zone appears to narrow from about 400 metres width on line 4400N to 100 metres width on line 5600N but is considered open to the northwest. The zone likely outcrops on lines 4000N, 5200N and 5600N but is probably covered by a 25 to 50 metre thick layer of weathered material (A) exhibiting highly variable resistivity on lines 4400N and 4800N.

A small, near surface chargeability feature (C) that flanks the Central Zone to the southeast on line 4400N seems to plunge to depth to the northwest and merge with Central Zone near line 5200N.

A sharp resistivity contact (D) separating low resistivities to the SW from high resistivities to the NE appears to run roughly parallel to the Central Zone. It is generally immediately east of and below the chargeability target (B). There is no obvious correlation between the higher resistivity zone to the northeast and any chargeability signature. This feature does not reach the surface and generally appears between 100 and 150 metres depth. It is likely mapping a buried geological unit.

Another zone (E) is mapped on the eastern halves of lines 5200N to 4000N. It is different from the other target zones in that it is characterized by elevated resistivities as well as elevated chargeability and is buried at about 150 metres depth. The inversion results only identify the top of this unit. This response likely reflects a discrete geological unit.

The Northeast Zone (not shown on the depthsection for line 4400N) exhibits similar characteristics to the Central Zone. Based on previous work, this zone was thought to terminate in the vicinity of line 3200N. A strong chargeability anomaly (F) was mapped on line 4000N along the projected strike of this zone. This prompted the crew to survey a northeast extension of line 3600N which confirmed the anomaly in this location. The Northeast Zone is now mapped from line 2700N to 4000N and is still considered open along strike to the northwest.

## **9.2. Previous IP Survey**

Most of the data for the previous IP surveys were provided as processed digital files in geosoft format. Data for the western portions of lines 1200N and 2200N were

unavailable. The data was reformatted for input the DCIP2D inversion program. Detailed topographic information was not recorded for these lines. While the topographic base map provides some information, it is not detailed enough to warrant inclusion in the process and these lines were inverted as if the ground surface was flat. Consequently, while the presence of anomalous chargeability and resistivity units will be detected, there could be some lateral displacement in their locations.

As a general comment, the inversion results have refined the previous interpretation. Anomalous chargeability trends located from the pseudosection displays have generally been shifted slightly, usually towards the potential electrodes. This is a fairly common result stemming from the removal of the geometric effects of the electrode array present in the pseudosection presentation.

These inversion results have been included in the plan maps provided and are also presented in depthsection format in Appendix 3.

## **10. CONCLUSIONS & RECOMMENDATIONS**

The 2003 IP survey results complement the previous work. The Central Zone IP anomaly has been traced for a further 2 kilometres along strike to the northwest. The zone appears to narrow and plunge to depth in this direction but is still considered open to the northwest. Resistivity data suggests the eastern edge of this zone is controlled by a geological contact.

The Northeast Zone IP anomaly has been traced for a further 600 metres along strike to the northwest and is considered open in that direction.

A small chargeability anomaly has been located that runs sub-parallel to the Central Zone from line 4400N to 5200N. This anomaly appears to be at or near the surface on line 4400N and plunge to depth to the northwest, in the same manner as observed on the Central zone. At line 5200N these two trends are close enough that the IP responses appear to merge.

An unidentified geological unit, characterized by elevated resistivities and elevated chargeabilities is mapped at about 150 metres depth on the eastern halves of lines 4400N to 5000N.



2-D inversions of the IP data gathered over the previous years has refined the interpretation on the southern portion of the property.

The recommendation by Lloyd (2002) to test the strongest chargeability zones by drilling remains valid, however the locations he cited in his report should be re-evaluated and based on the recent inversion results.

Using the same criteria, the extensions to the Northeast and Central Zones mapped by this latest survey warrant drilling. The inversion depthsections should be correlated with existing geological and geochemical data in order to prioritise targets. Drilling should be planned to intersect the highest chargeability sections of the target zones at orientations dependent on the apparent dip of the source bodies as reflected by the depthsections.

Because of the widely spaced electrodes used to measure the IP effect, the interpretation of IP data produces generalized and smoothed targets. Consequently, it is a common practice to test IP anomalies with a fence of drill holes that cross the target area. Negative results from a single drill hole are not considered conclusive evidence to warrant dismissal of an interpreted target.

Respectfully submitted,

Per S.J.V. Consultants Ltd.

E. Trent Pezzot, B.Sc., P.Geo,

Geophysics, Geology

**11. APPENDIX 1 – STATEMENT OF QUALIFICATIONS – E. TRENT  
PEZZOT**

I, E. Trent Pezzot, of the city of Surrey, Province of British Columbia, hereby certify that:

- 1) I graduated from the University of British Columbia in 1974 with a B.Sc. degree in the combined Honours Geology and Geophysics program.
- 2) I have practised my profession continuously from that date.
- 3) I am a registered member of the Association of Professional Engineers and Geoscientists of British Columbia.
- 4) I have no interest in Finlay Minerals Ltd. or any of their subsidiaries or related companies, nor do I expect to receive any.

Signed by: \_\_\_\_\_

E. Trent Pezzot, B.Sc., P.Geo.

Geophysicist/Geologist

## **12. APPENDIX 2 – INSTRUMENT SPECIFICATIONS**

### **12.1. IRIS VIP-3000IP Transmitter**

#### **Output Ratings**

|                               |   |
|-------------------------------|---|
| Output power:                 | 3000 VA maximum.  |
| Output voltage:               | 3000V maximum, auto voltage range selection.  |
| Output current:               | 20 ma to 5A, current regulated to better than 1 %.  |
| Dipoles:                      | 9, push button selected.  |
| Output connectors:            | Uniclip connectors accept bare wire or plug of up to 4 mm diameter.   |
| Waveforms:                    | see figure 4.1.   |
| Fall times:                   | better than 1 msce in resistive load.   |
| Time domain:                  | preprogrammed on and off times from 0.25 to 8 seconds, by factor of 2.<br>Other cycles programmable by user.<br>Automatic circuit opening in off time.                    |
| Frequency domain:             | Preprogrammed frequencies from 0.0625 Hz to 4Hz, by factor of 2.<br>Alternate or simultaneous transmission of two frequencies.<br>Other frequencies programmable by user. |
| Time and frequency stability: | 0.01 %<br>1 PPB optional  |

#### **Other**

|                        |   |
|------------------------|---|
| Display:               | Alphanumeric liquid crystal display.  |
| Power source:          | 175 to 270 VAC, 45-450 Hz, single phase.  |
| Operating temperature: | -40 to +50o C.  |
| Protection:            | short circuit at 20 $\Omega$ ,<br>open loop at 60 000 $\Omega$ ,<br>thermal,<br>input overvoltage and undervoltage. |
| Remote control:        | full duplex RS232C, 300-19 200 bps.   |
| Dimensions (h w d):    | 410 x 320 x 240 m   |
| Weight:                | 16kg.   |

## 12.2. IRIS ELREC 10 IP Receiver

### Technical:

|  |   |
|--|---|
| Input impedance:   | 10 Mohm   |
| Input overvoltage protection up to 1000V   |   |
| Automatic SP bucking with linear drift correction  |   |
| Internal calibration generator for a true calibration on request of the operator                     |   |
| Internal memory:   | 3200 dipoles reading  |
| Automatic synchronization and re-synchronization process on primary voltages signals whenever needed |   |
| Proprietary intelligent stacking process rejecting strong non-linear SP drifts                       |   |
| Common mode rejection:   | More than 100 dB (for $R_s = 0$ )   |
| Self potential (Sp)  | : range: -15V - + 15V<br>: resolution: 0.1 mV                                 |
| Ground resistance measurement range:   | 0.1-100 kohms   |
| Primary voltage  | : range: 10 $\mu$ V - 15V<br>: resolution: 1 $\mu$ V<br>: accuracy: typ. 1.3% |
| Chargeability  | : resolution: 10 $\mu$ V/V<br>: accuracy: typ. 0.6%                           |

### General:

|  |             |
|--|-------------|
| Dimensions:  | 31x21x25 cm |
| Weight (with the internal battery):                              | 9 kg        |
| Operating temperature range:                                     | -30°C -70°C |
| Case in fiber-glass for resisting to field shocks and vibrations |             |

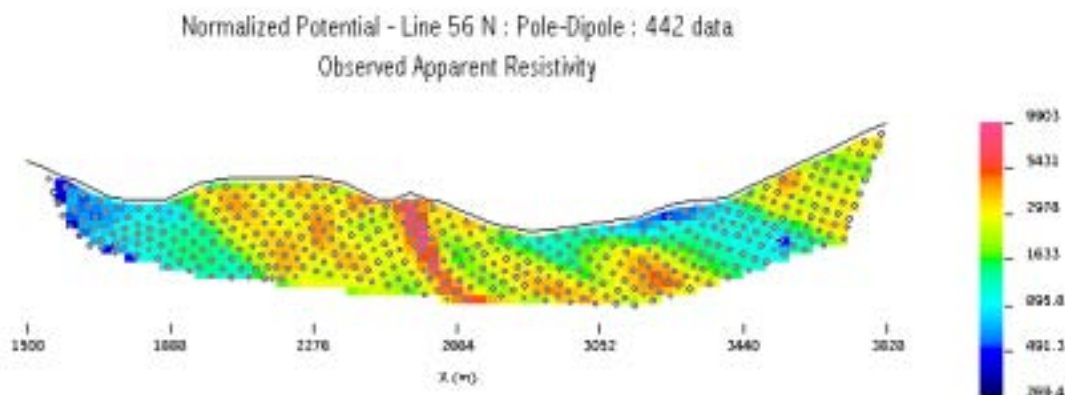
**13. Appendix 3 – IP Pseudosections and Depthsections**

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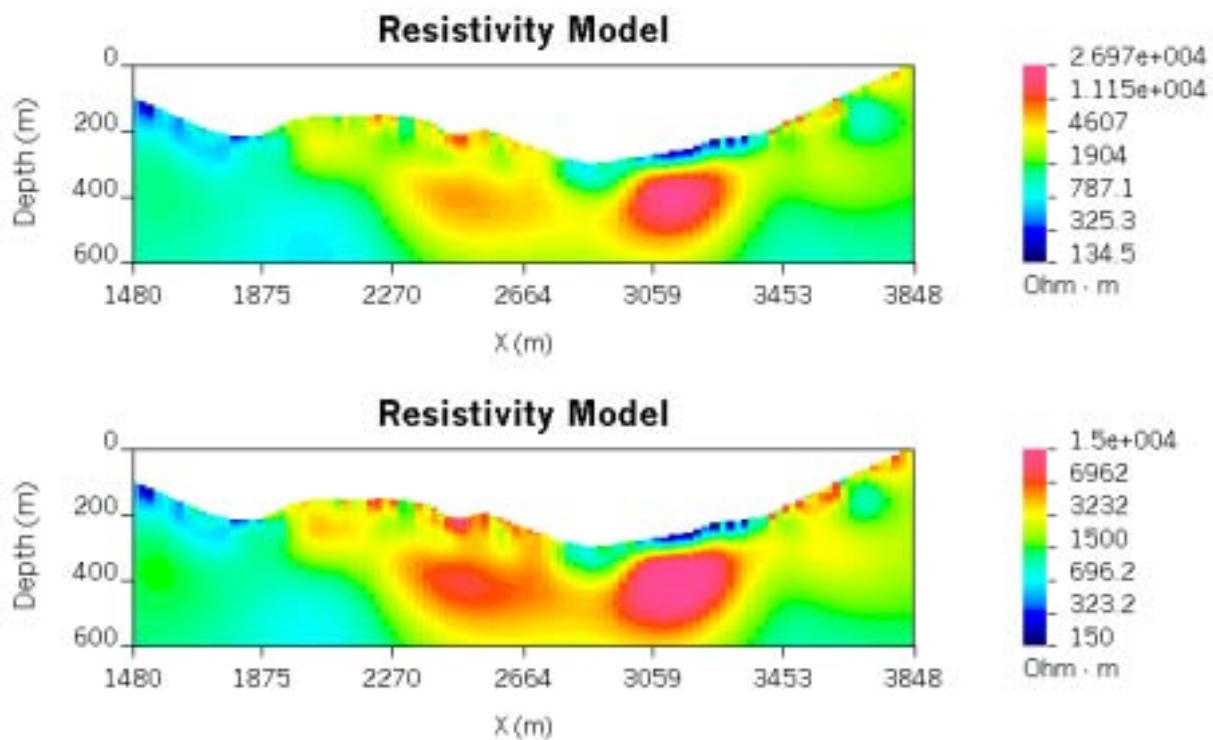
## 2003 Data

### Line 5600N Resistivity

Pseudosection

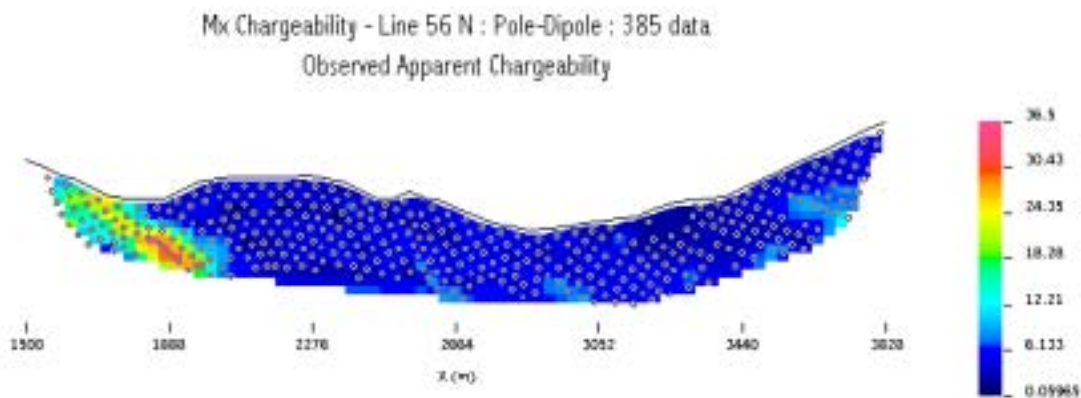


Inversion Depth Section

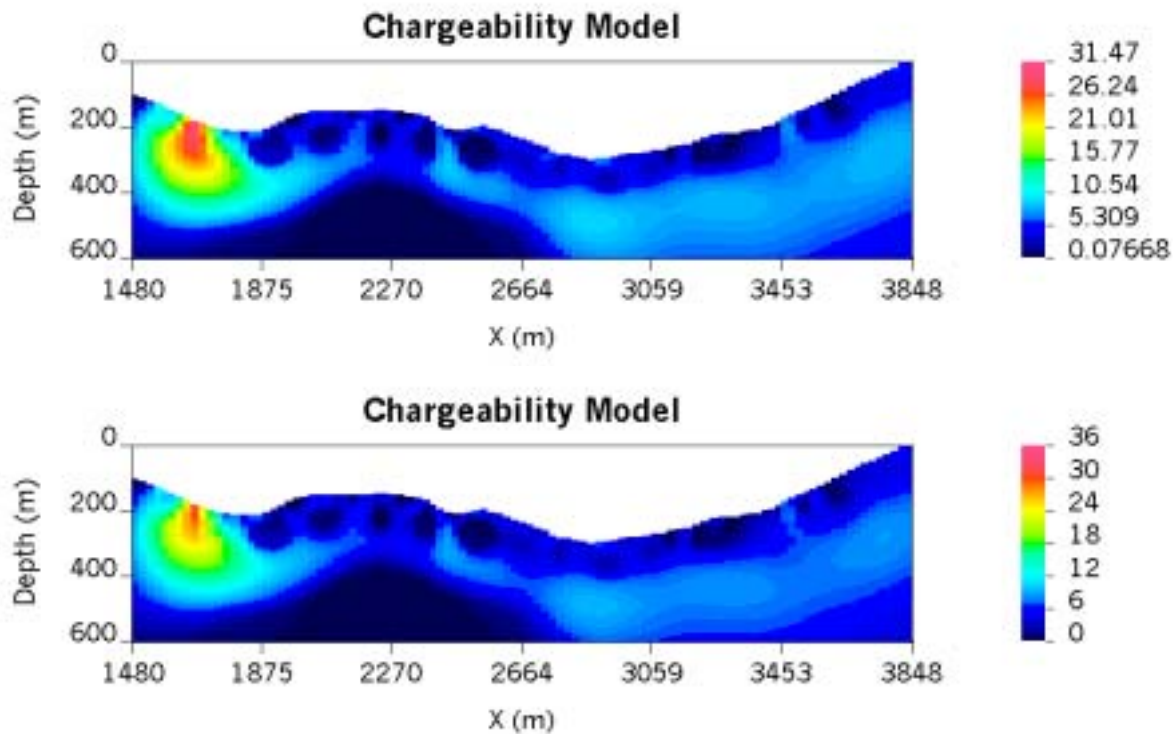


## Line 5600N Chargeability

Pseudosection

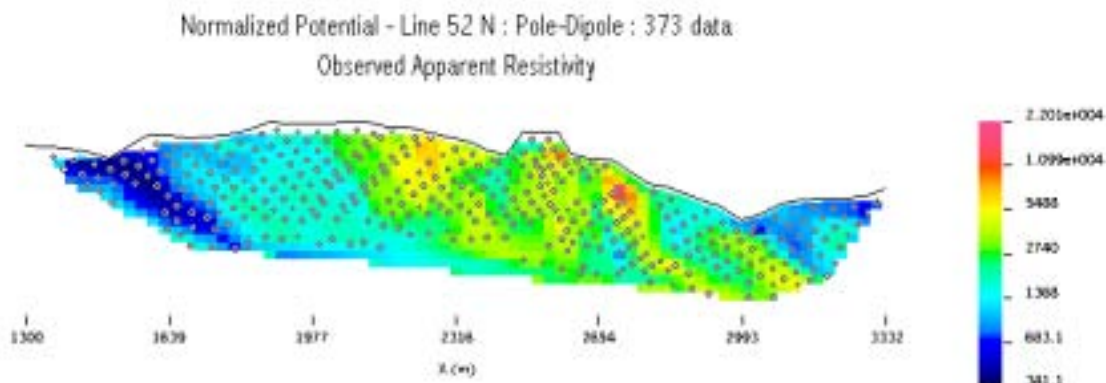


Inversion Depth Section

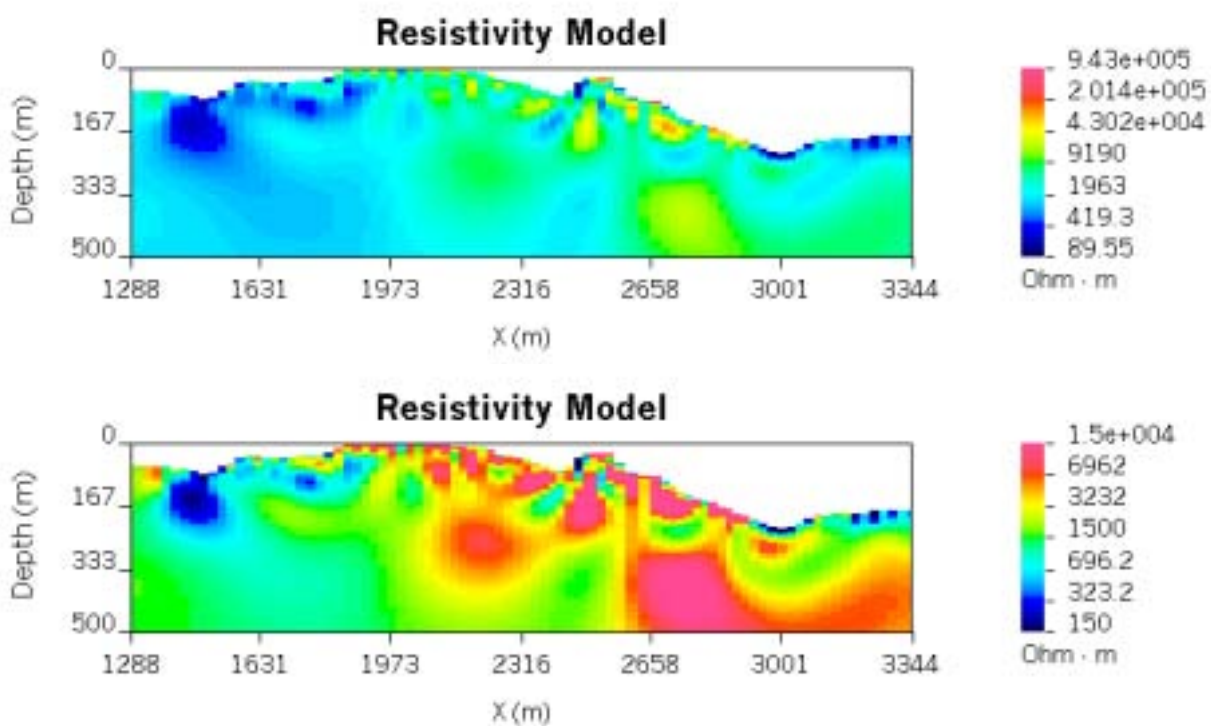


## Line 5200N Resistivity

Pseudosection



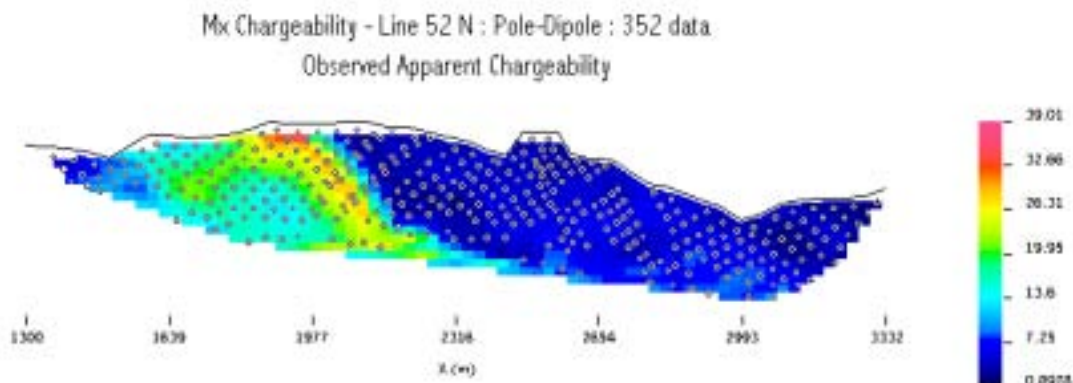
Inversion Depth Section



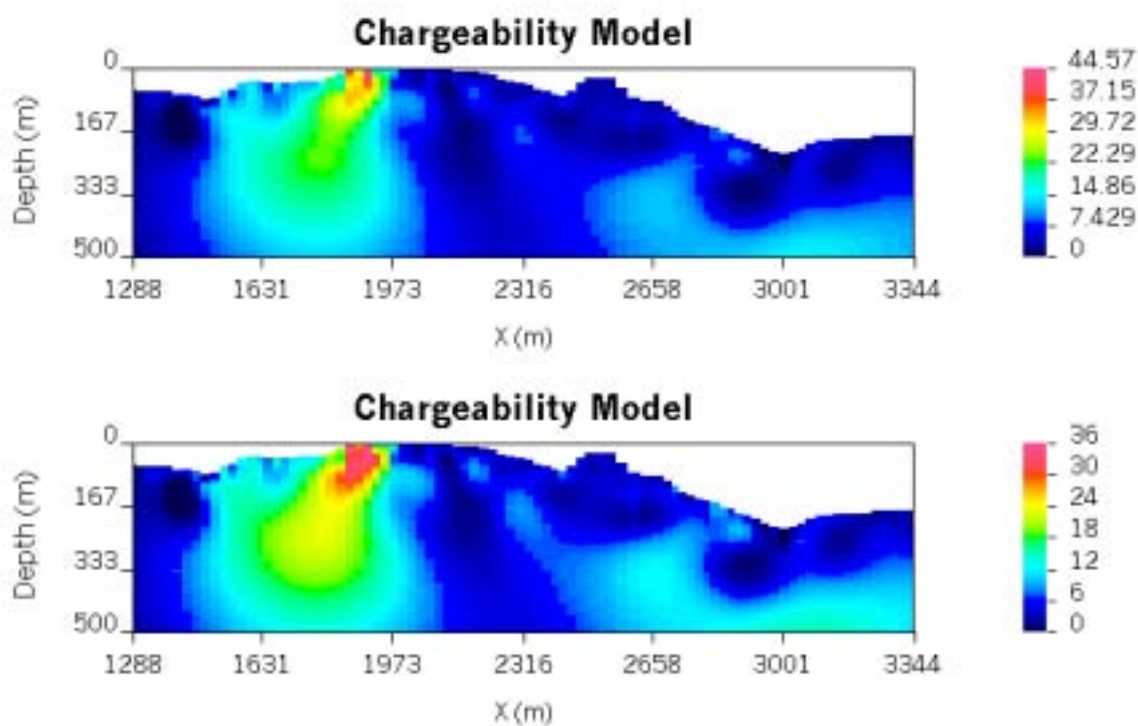


## Line 5200N Chargeability

Pseudosection

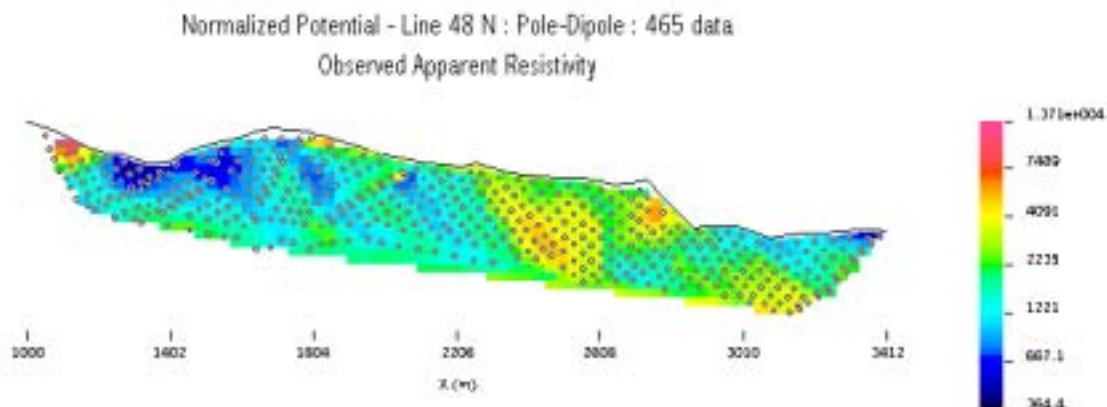


Inversion Depth Section

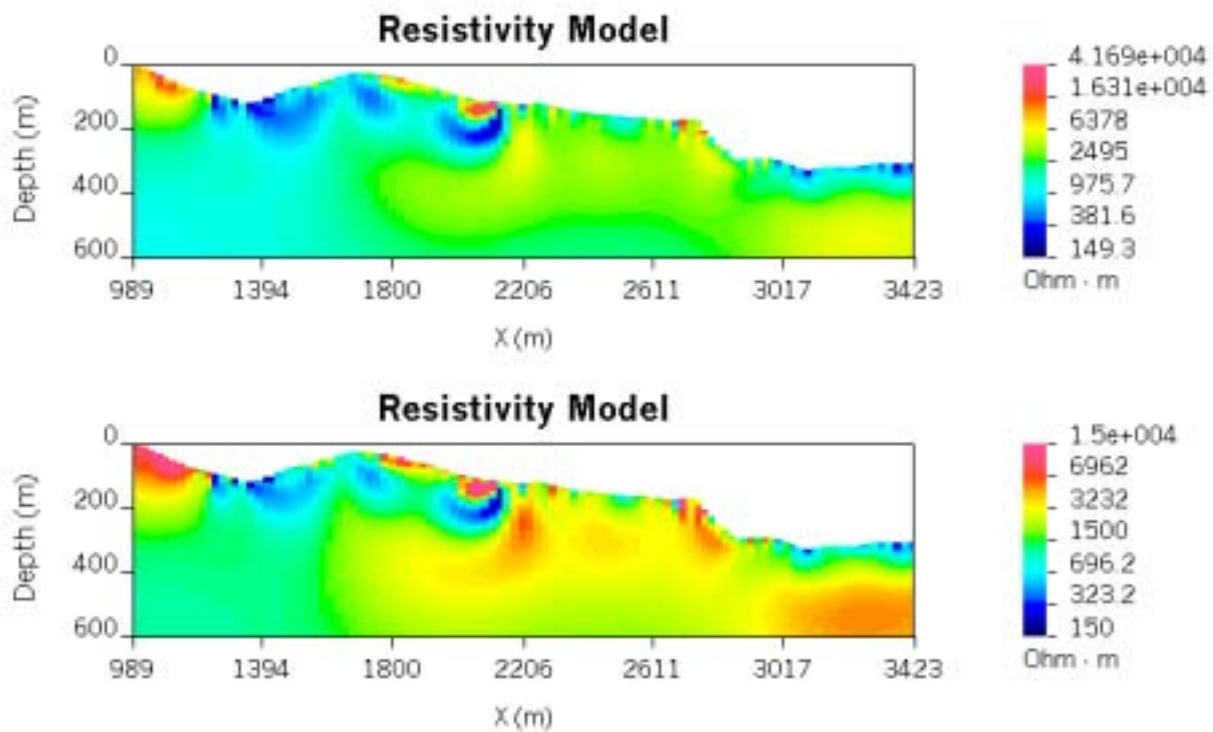


## Line 4800N Resistivity

Pseudosection

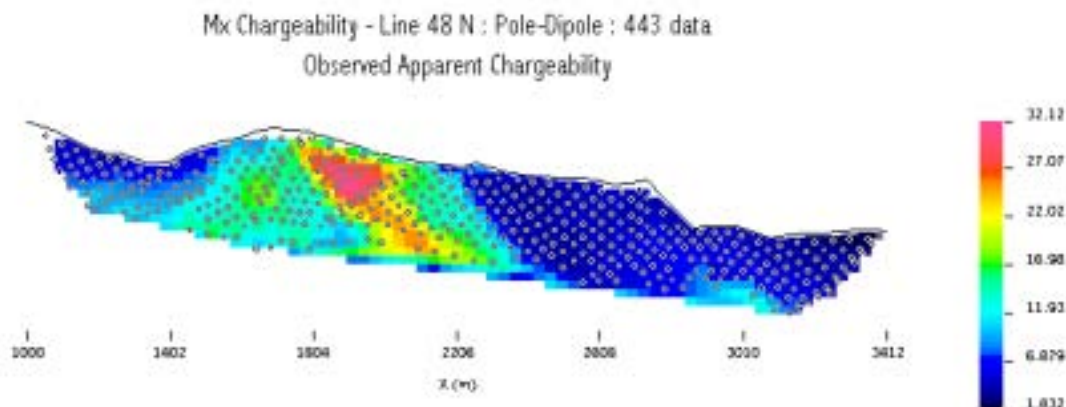


Inversion Depth Section

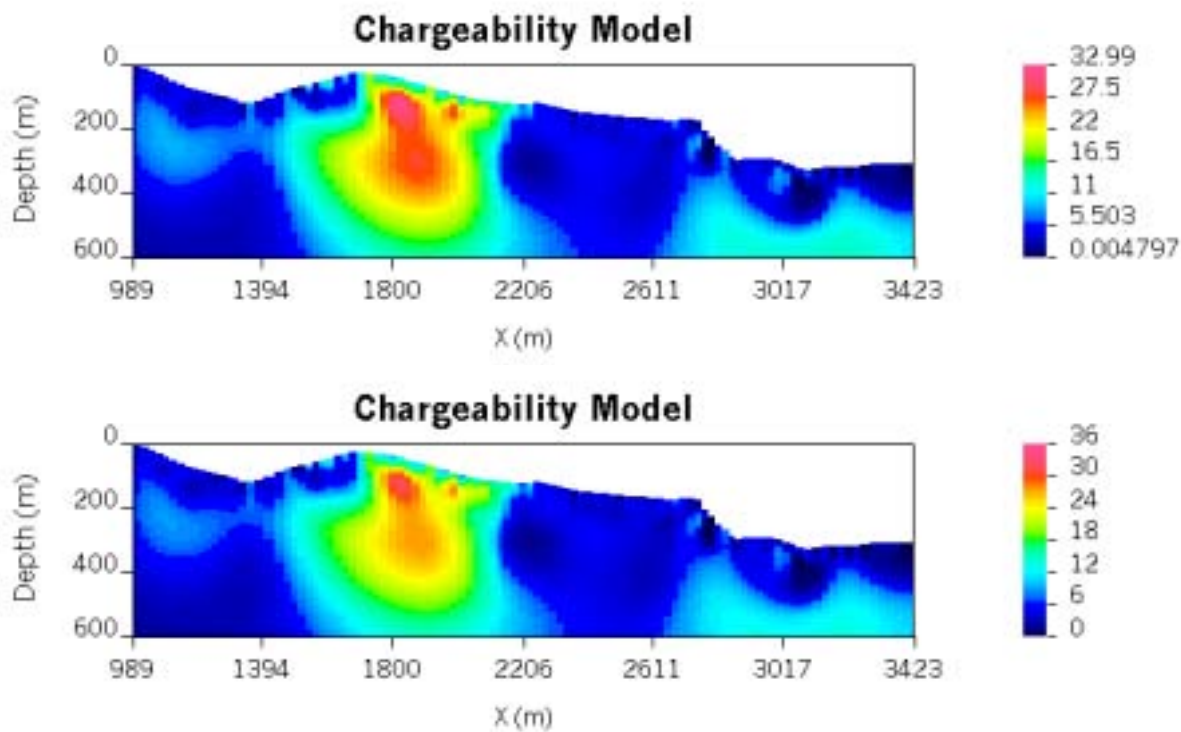


## Line 4800N Chargeability

Pseudosection

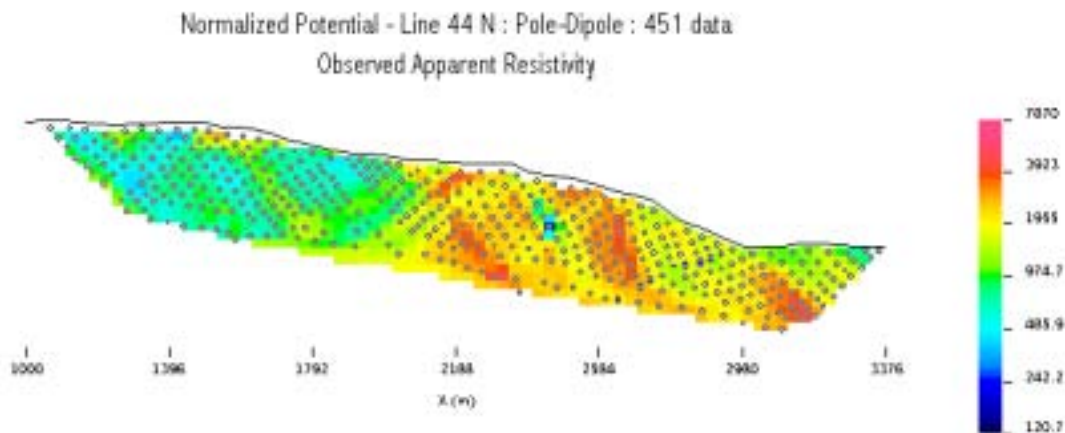


Inversion Depth Section

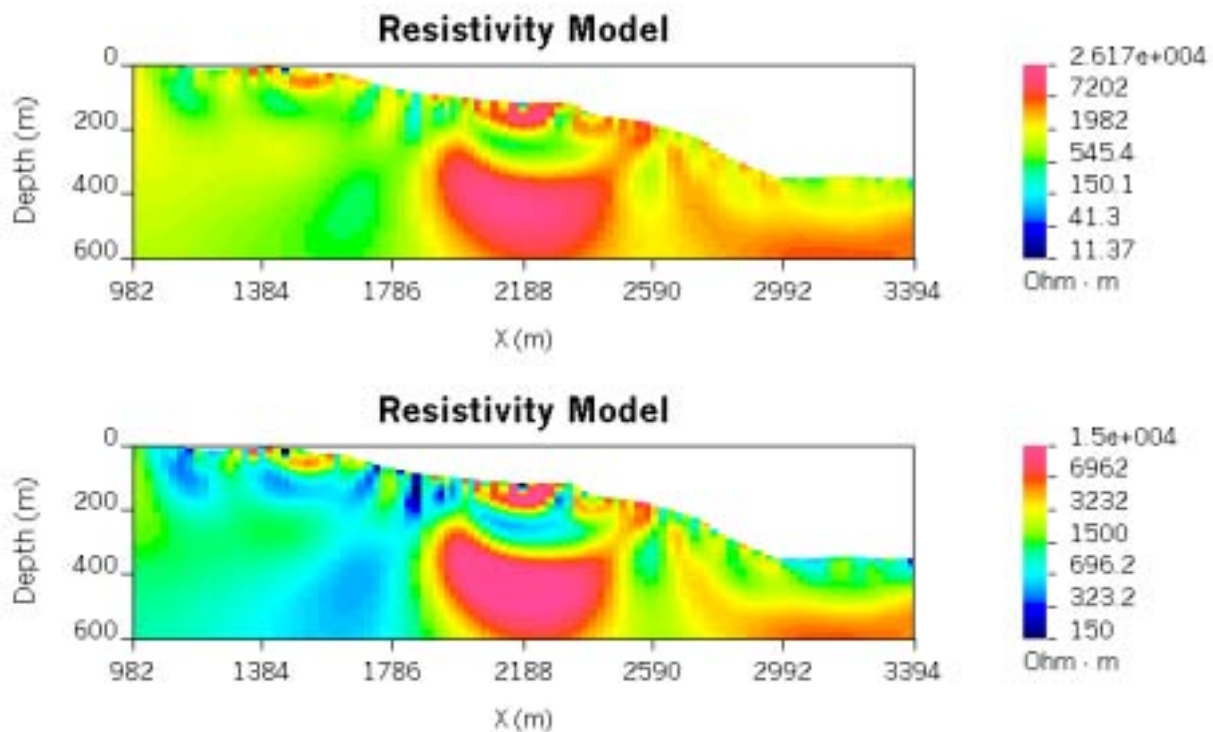


## Line 4400N Resistivity

Pseudosection

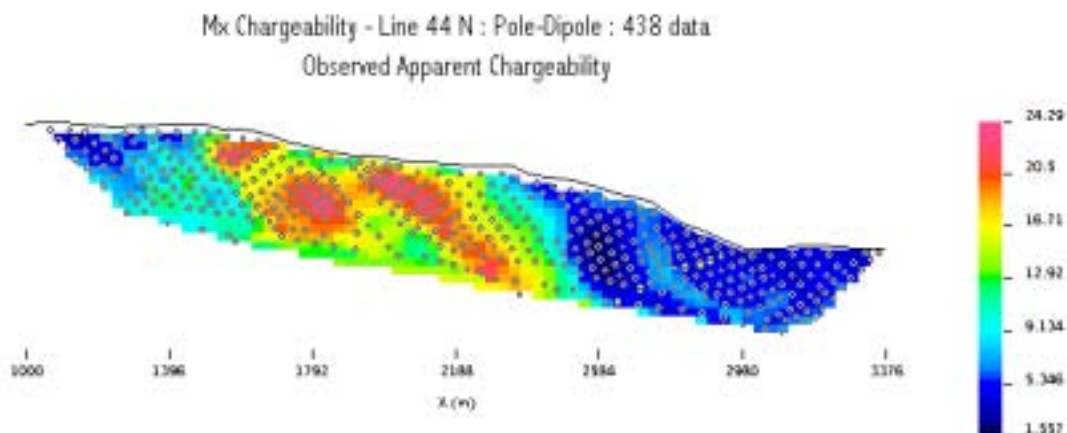


Inversion Depth Section

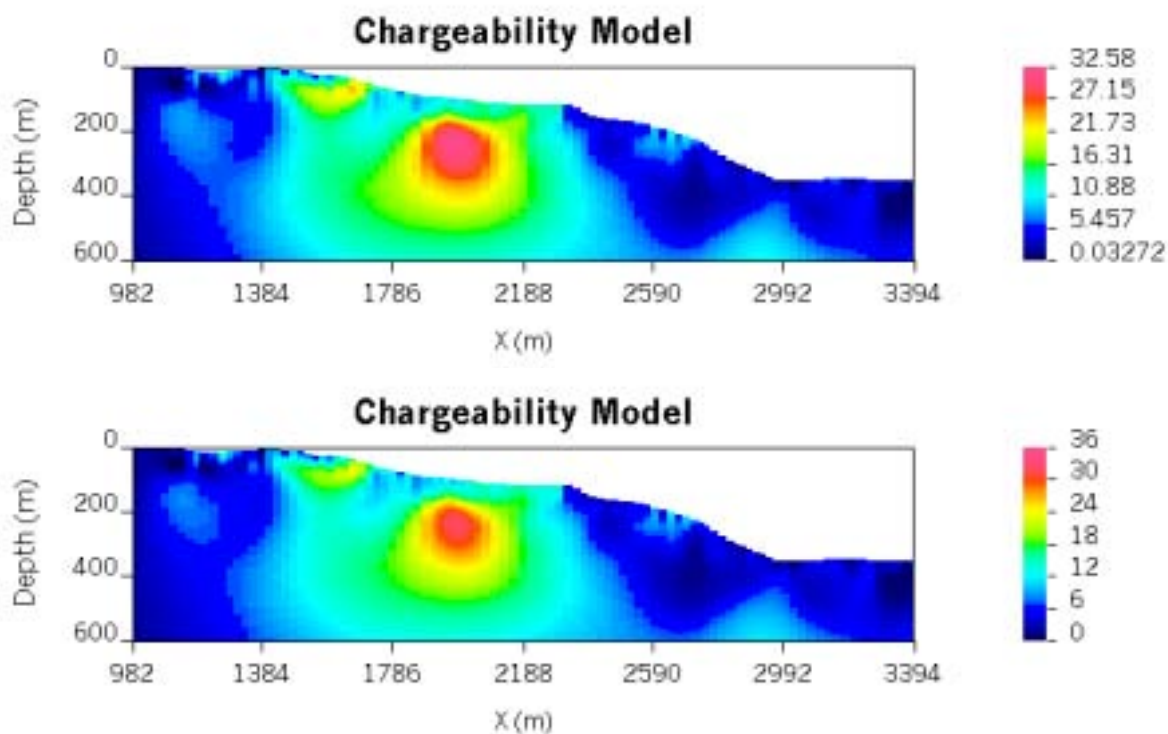


## Line 4400N Chargeability

Pseudosection



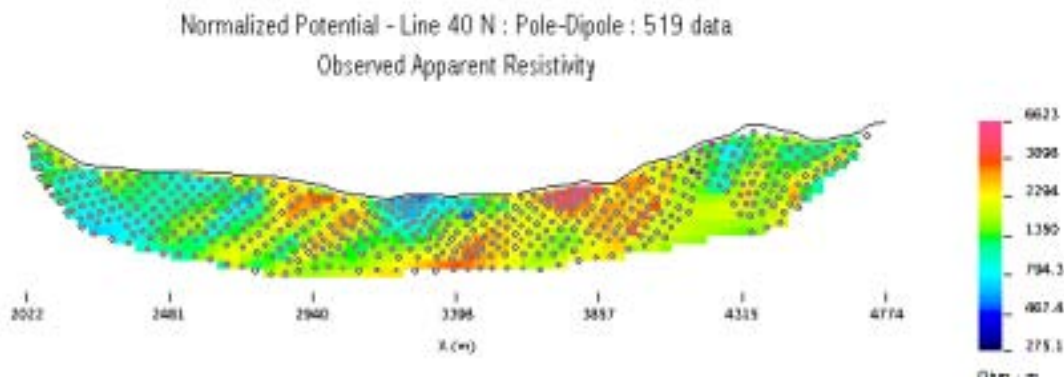
Inversion Depth Section



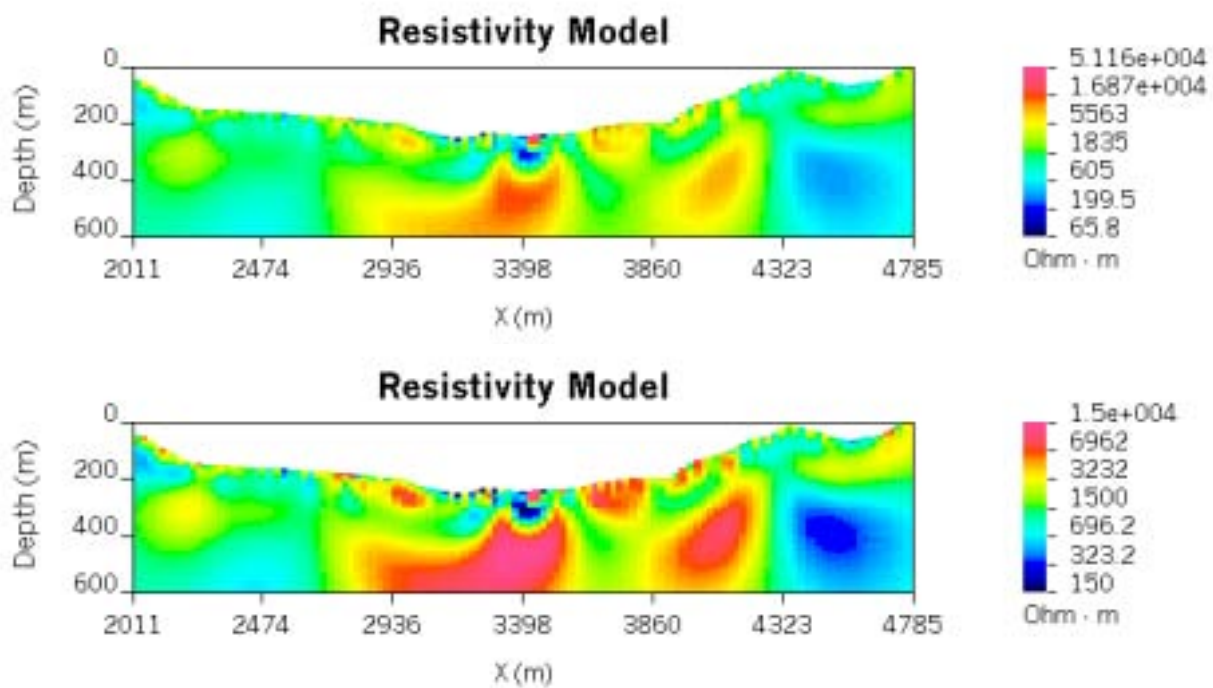


## Line 4000N Resistivity

Pseudosection

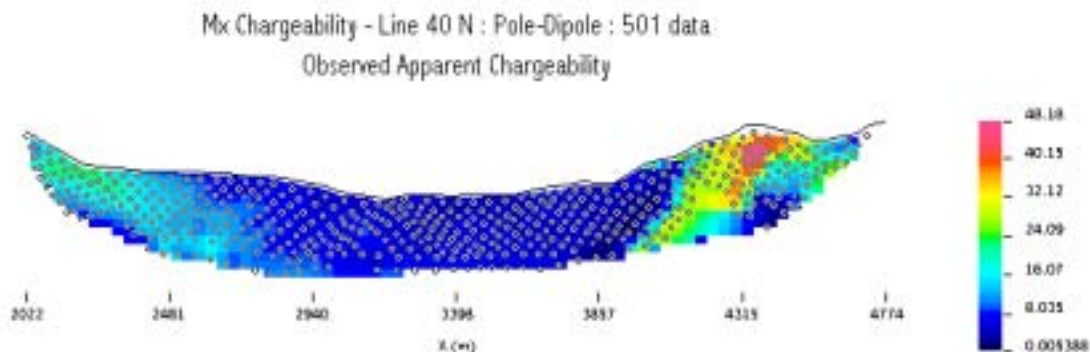


Inversion Depth Section

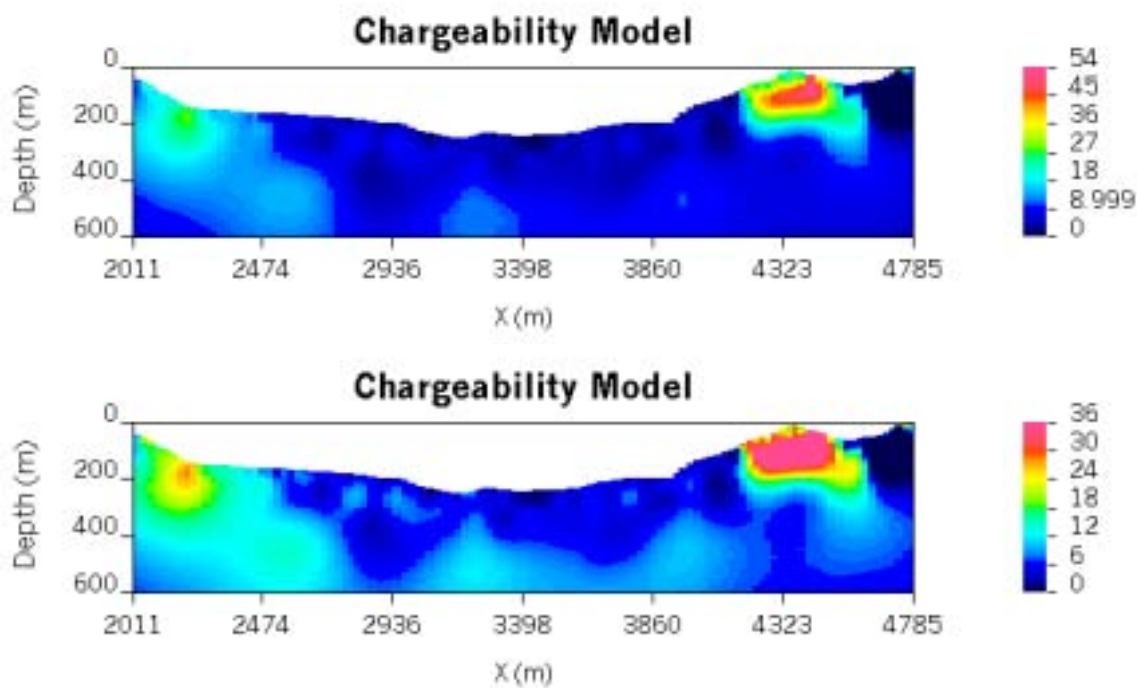


## Line 4000N Chargeability

Pseudosection

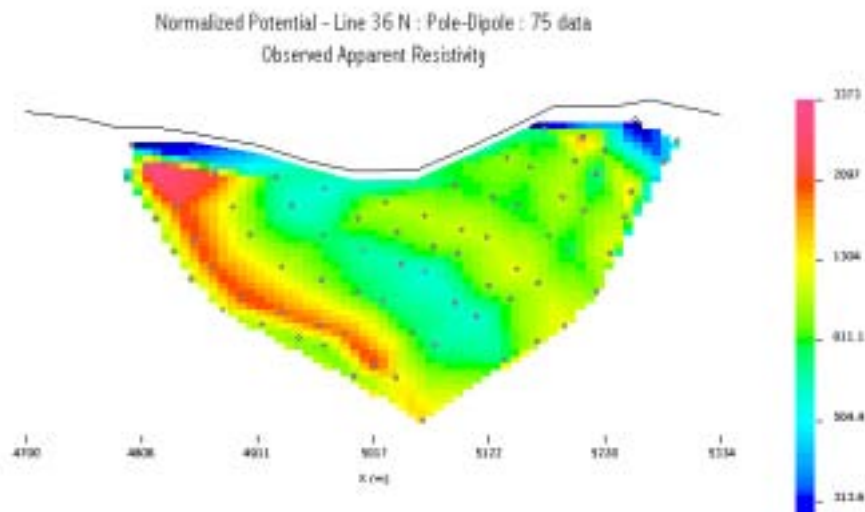


Inversion Depth Section

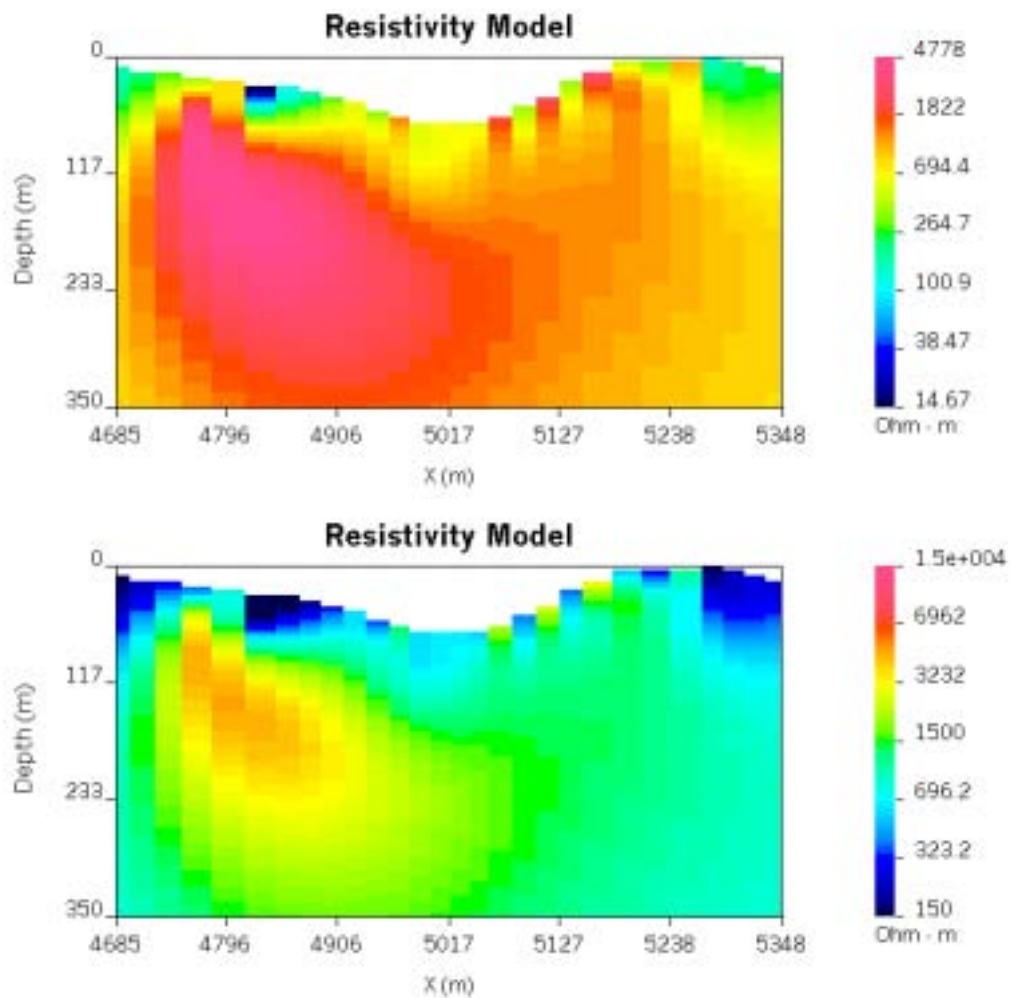


## Line 3600N Resistivity

Pseudosection



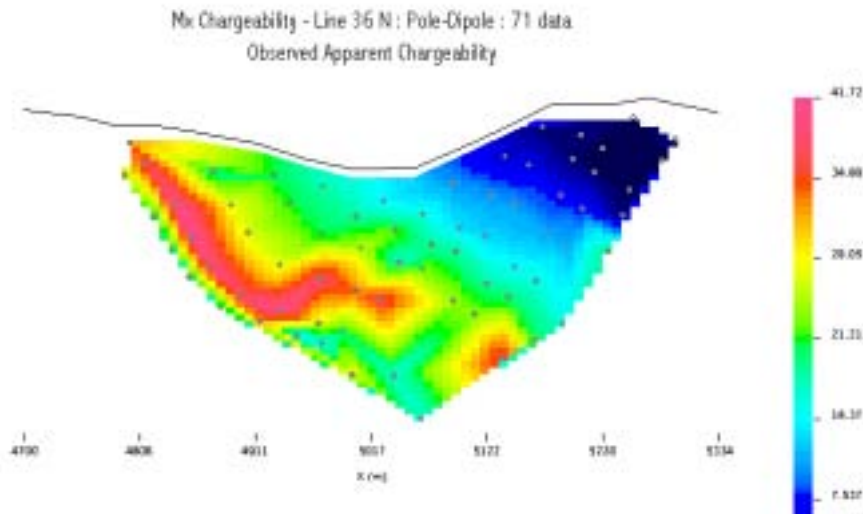
Inversion Depth Section



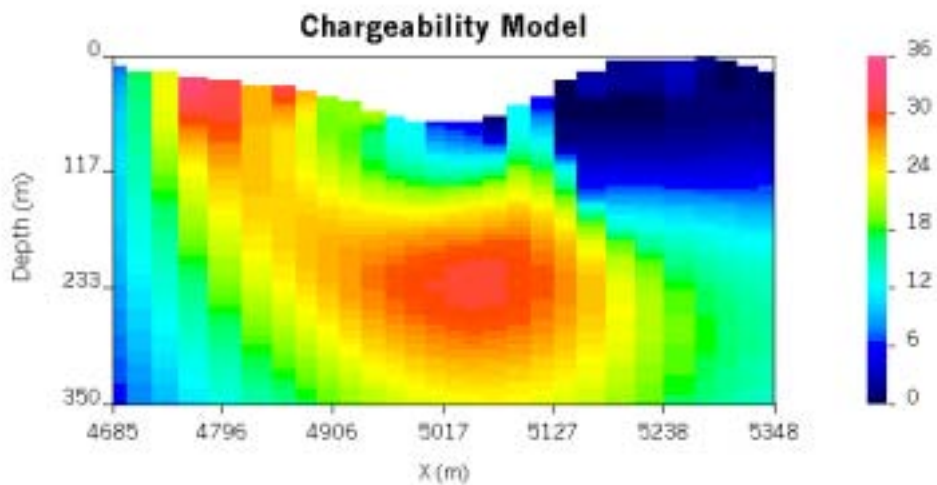
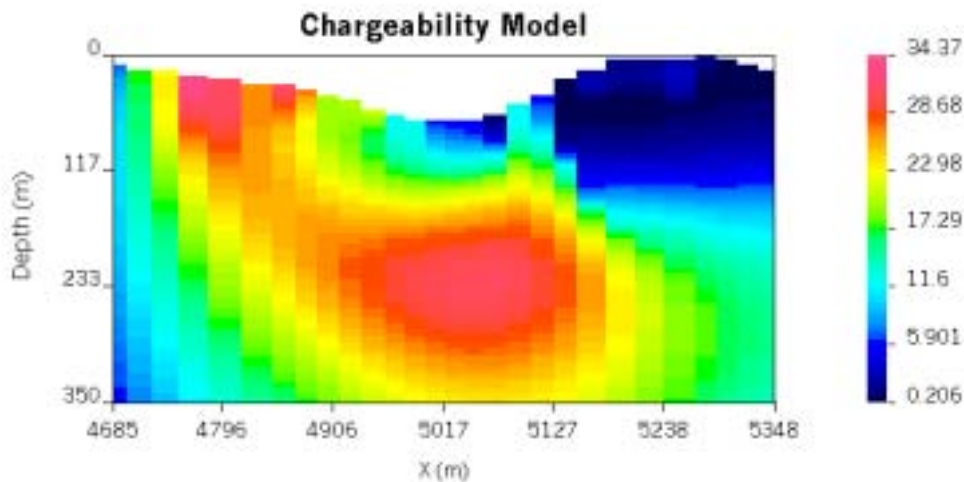


## Line 3600N Chargeability

Pseudosection

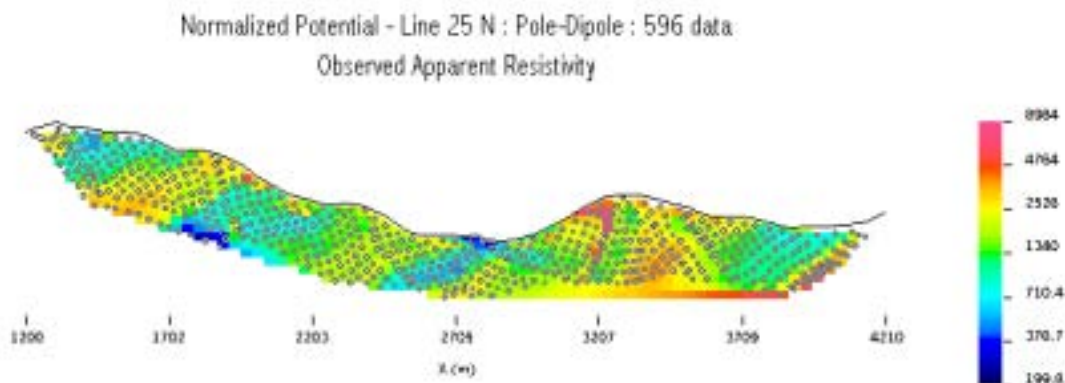


Inversion Depth Section

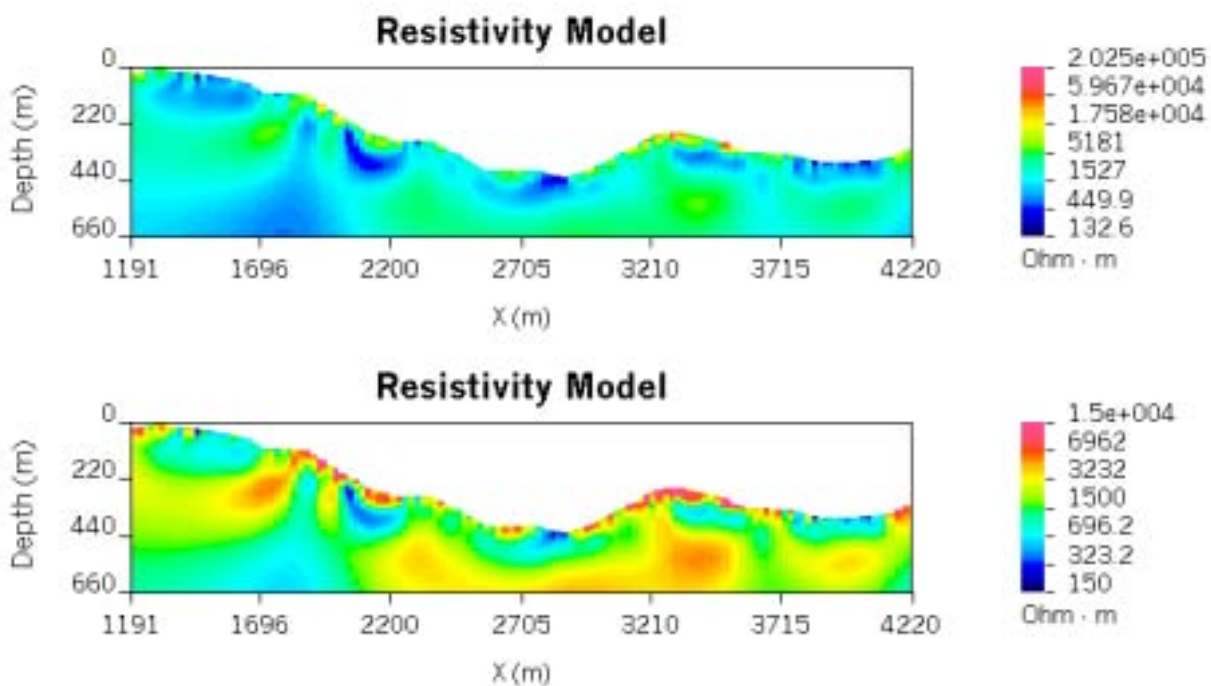


## Line 2500N Resistivity

Pseudosection

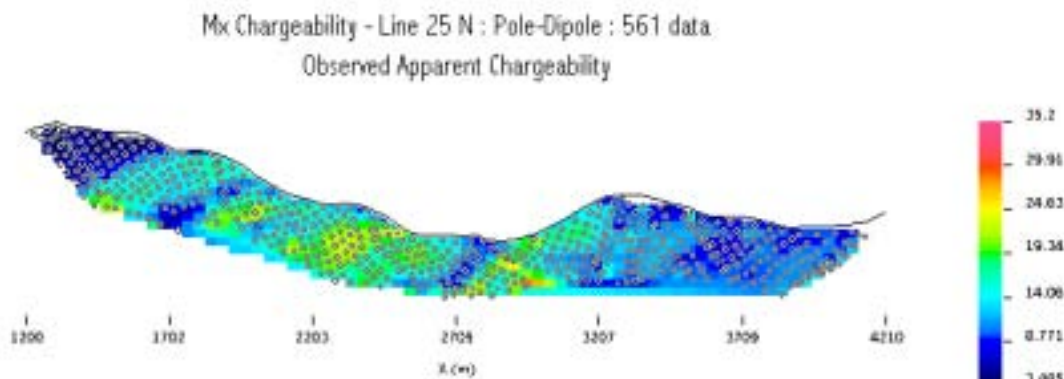


Inversion Depth Section

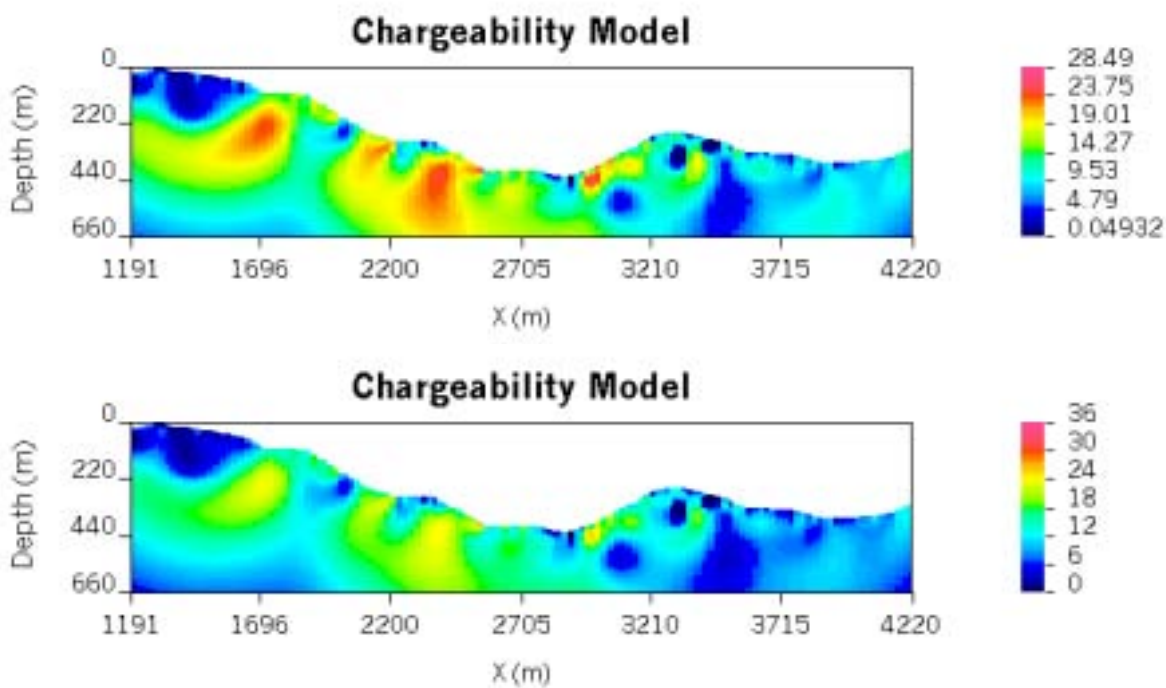


## Line 2500N Chargeability

Pseudosection



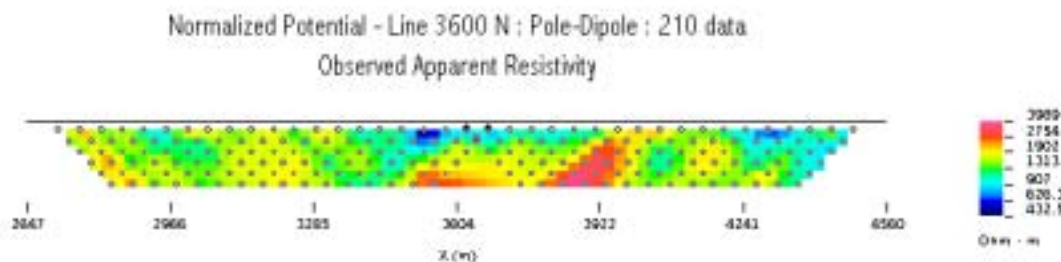
Inversion Depth Section



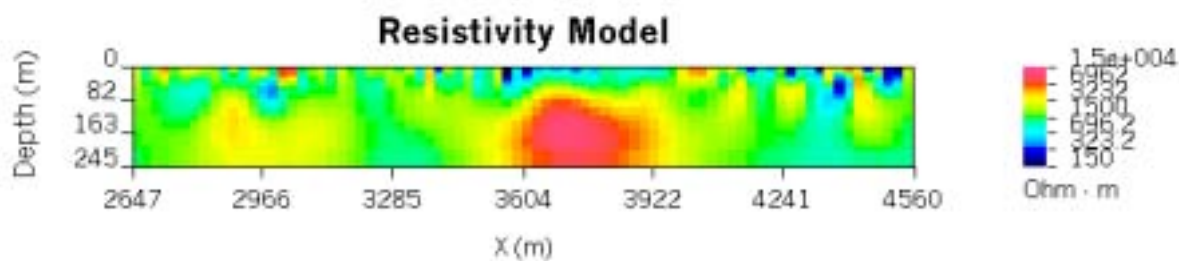
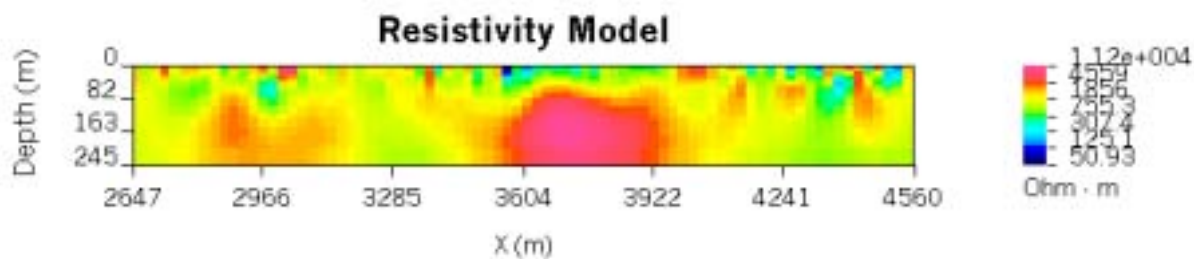
## 2001 – 2002 Data

### Line 3600N Resistivity

Pseudosection

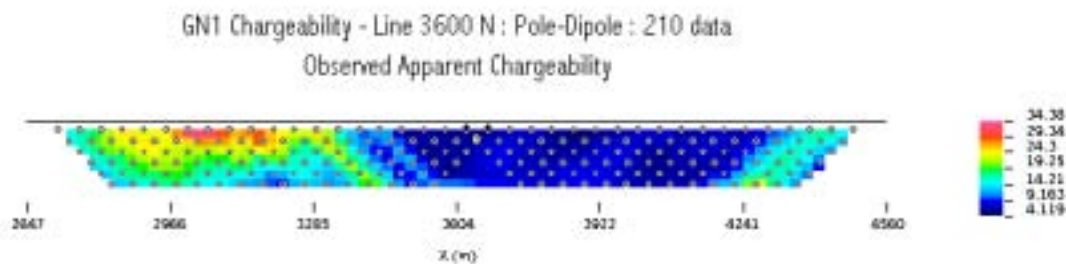


Inversion Depth Section

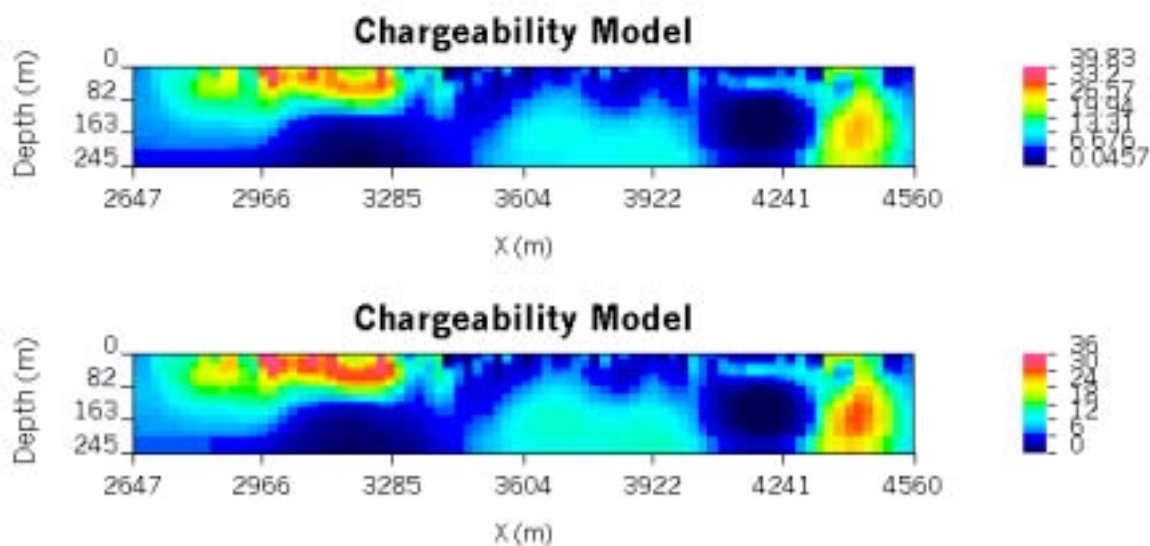


## Line 3600N Chargeability

Pseudosection



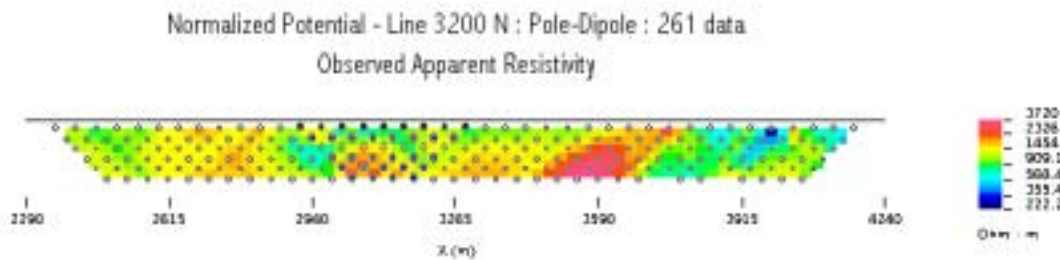
Inversion Depth Section



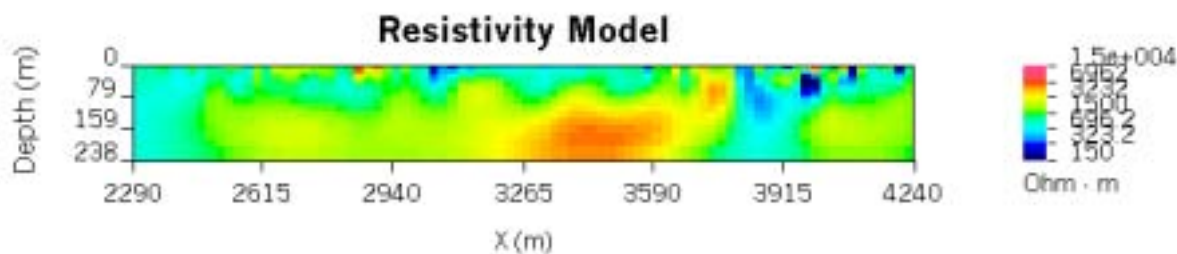
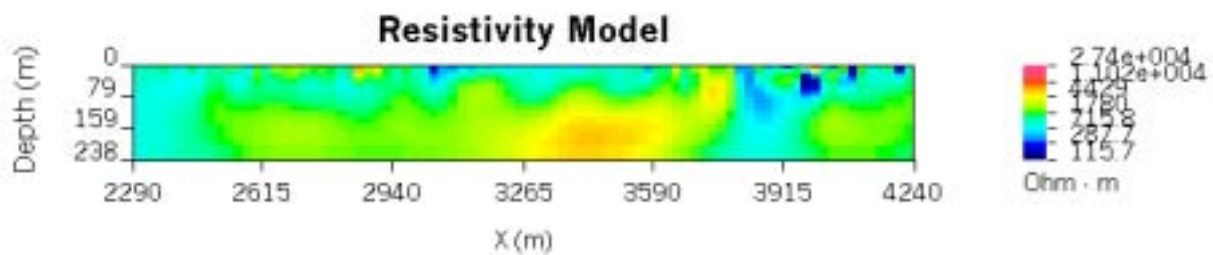


## Line 3200N Resistivity

Pseudosection

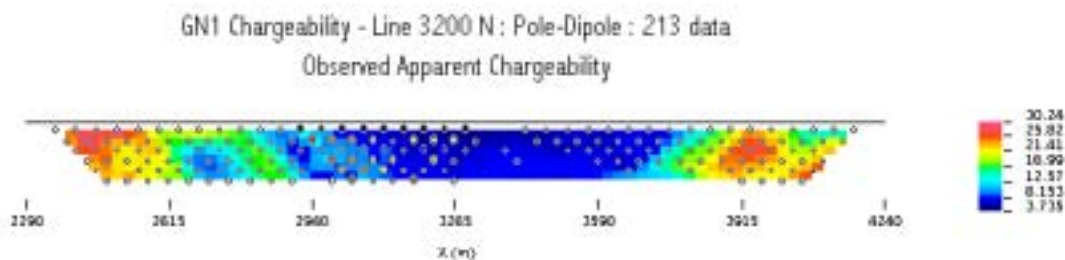


Inversion Depth Section

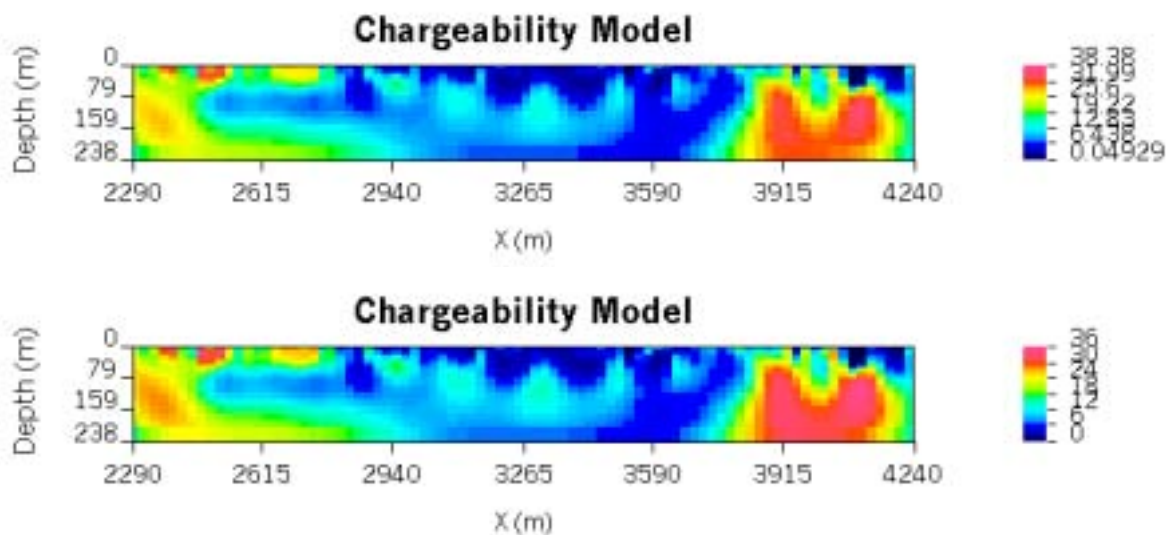


## Line 3200N Chargeability

Pseudosection

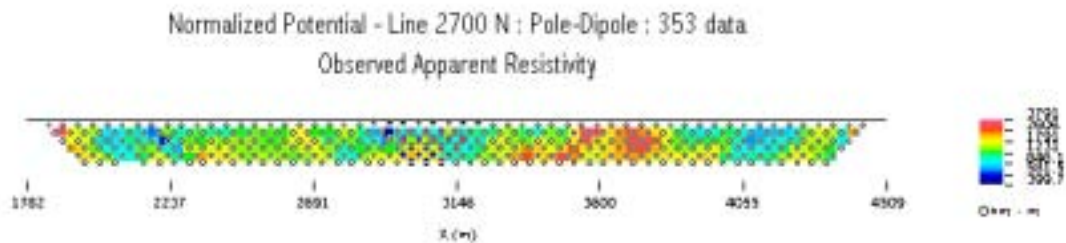


Inversion Depth Section

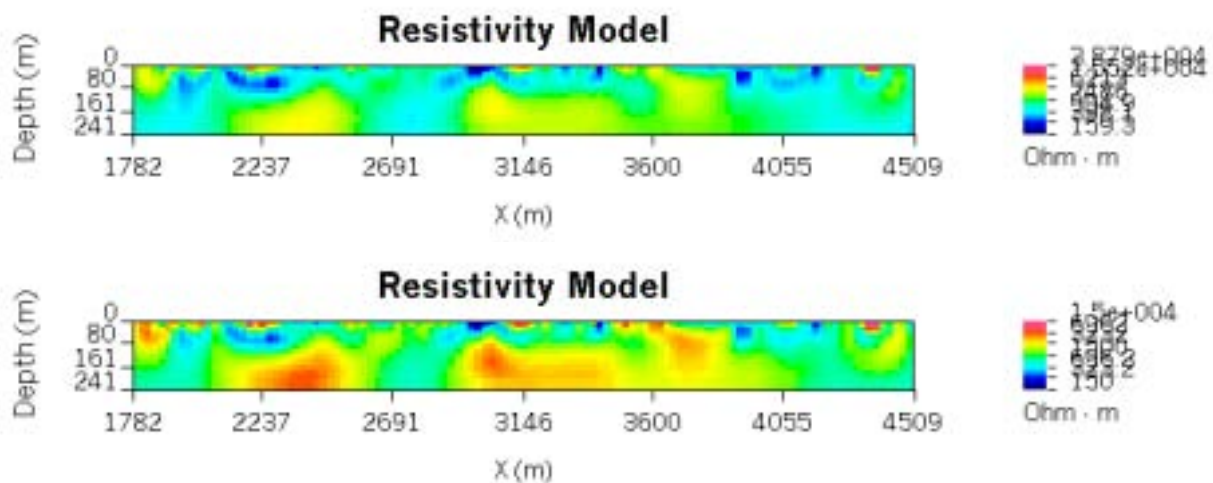


## Line 2700N Resistivity

Pseudosection



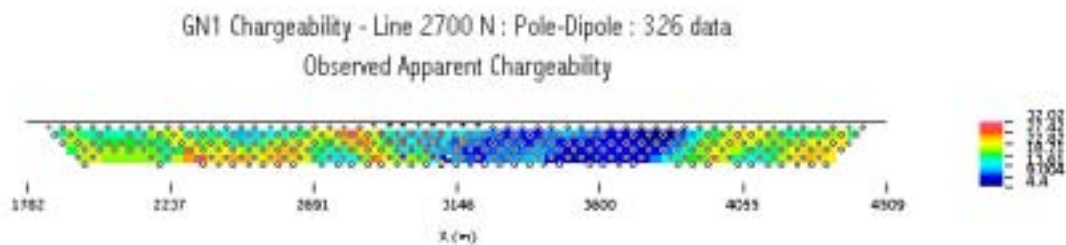
Inversion Depth Section



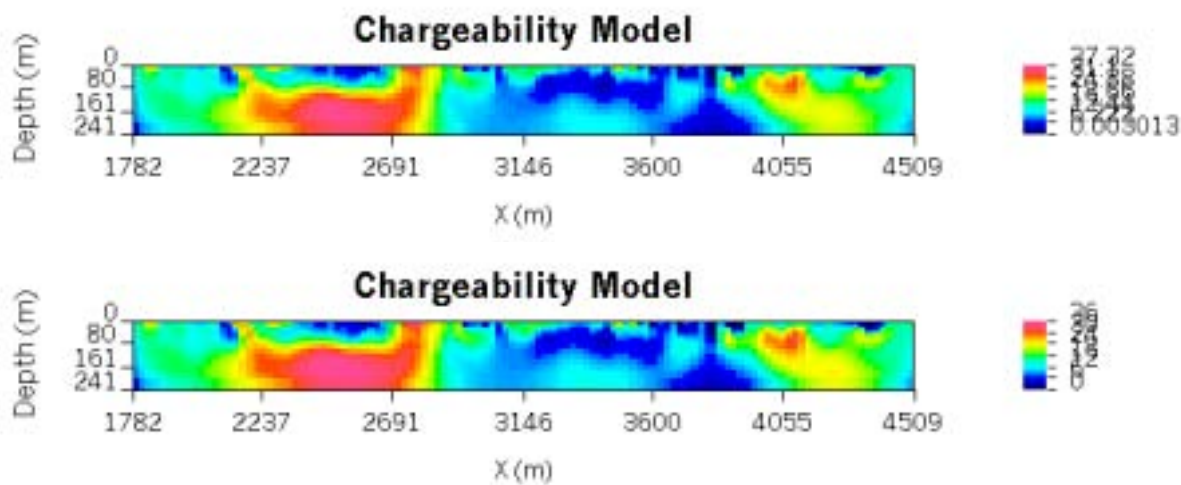


## Line 2700N Chargeability

Pseudosection

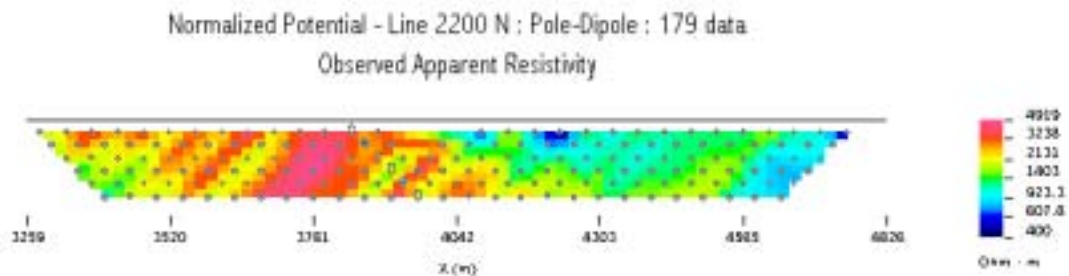


Inversion Depth Section

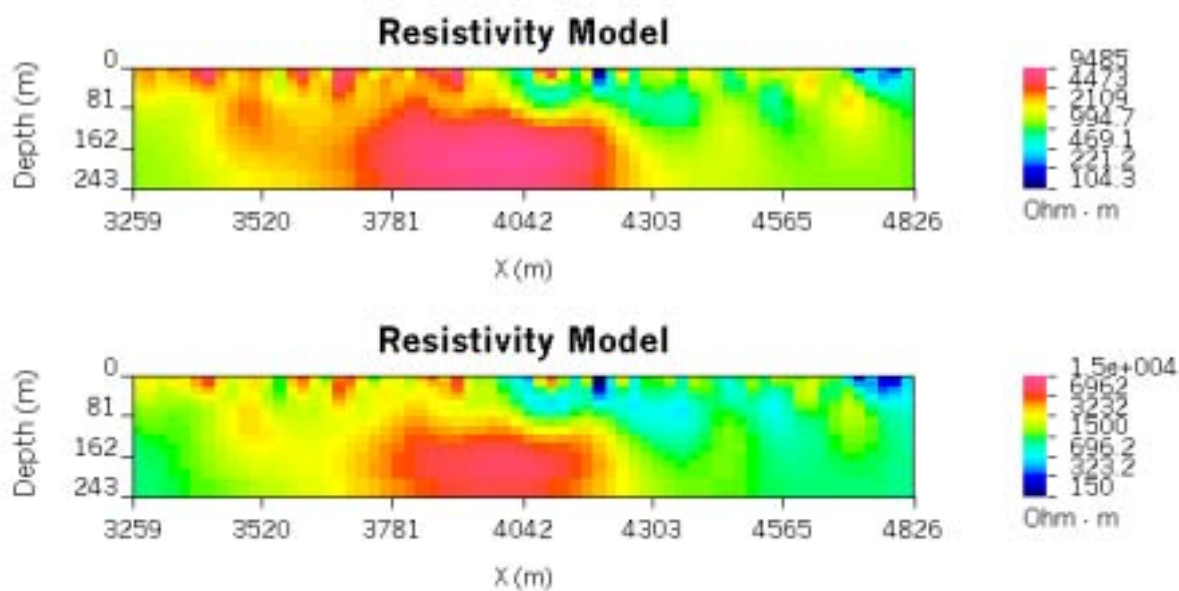


## Line 2200N Resistivity

Pseudosection

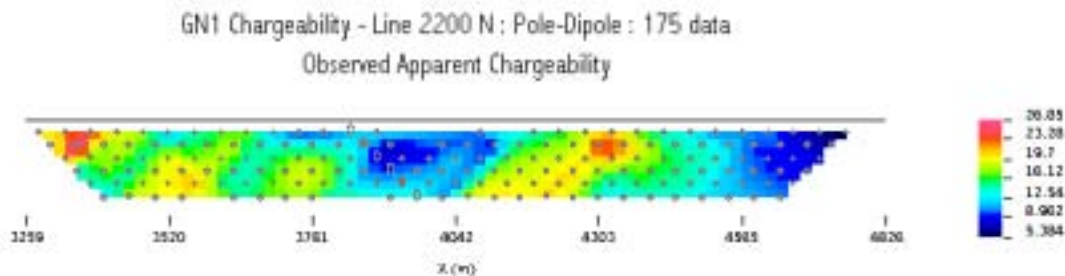


Inversion Depth Section

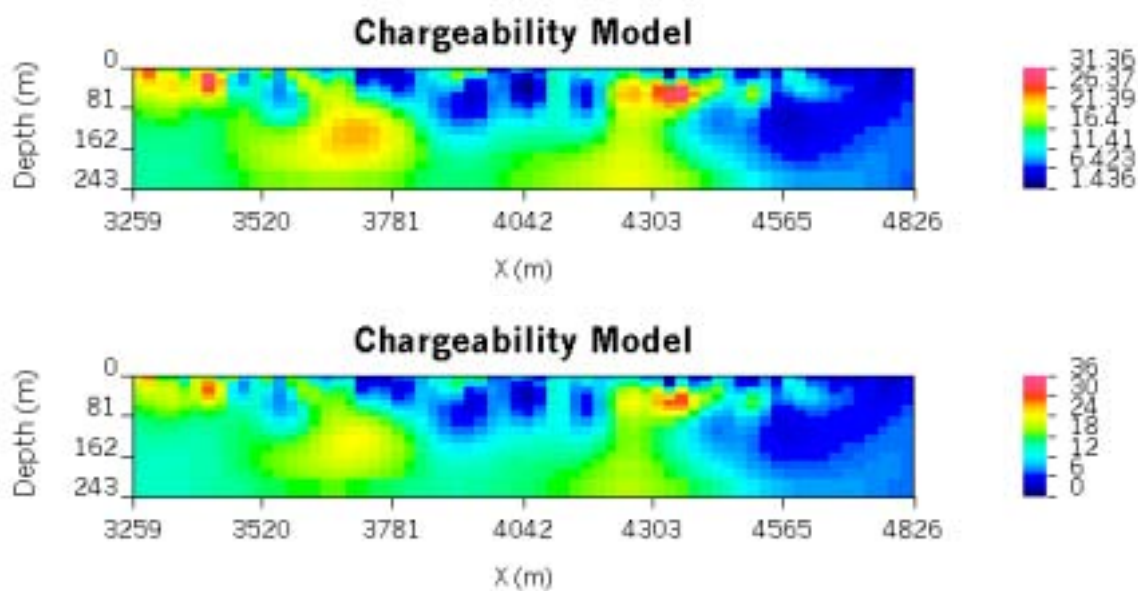


## Line 2200N Chargeability

Pseudosection

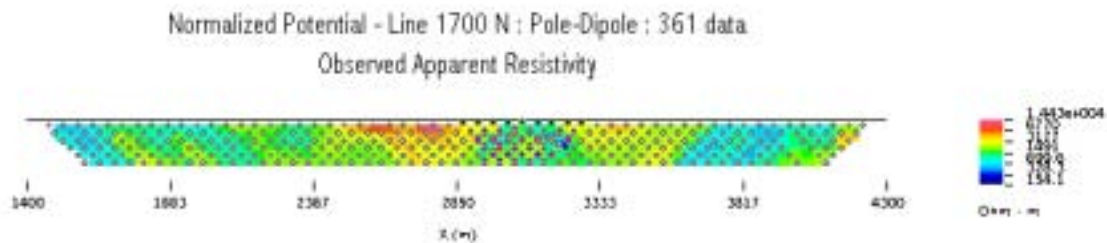


Inversion Depth Section

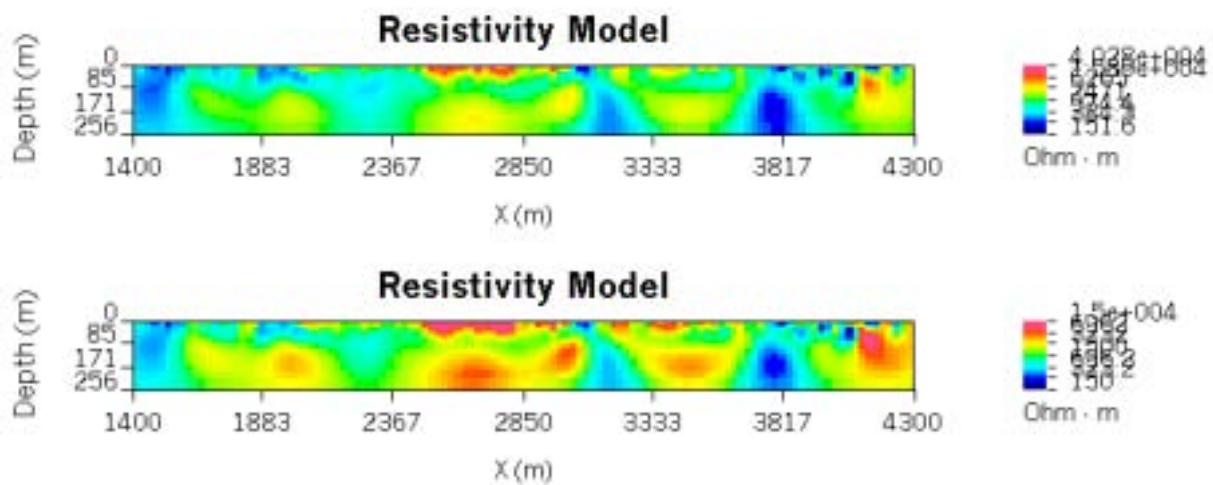


### Line 1700N Resistivity

## Pseudosection

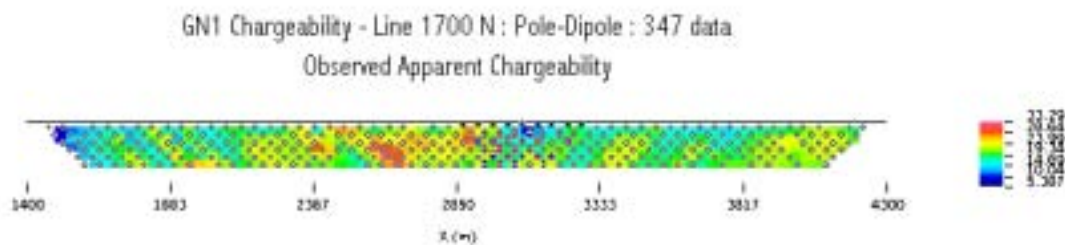


## Inversion Depth Section

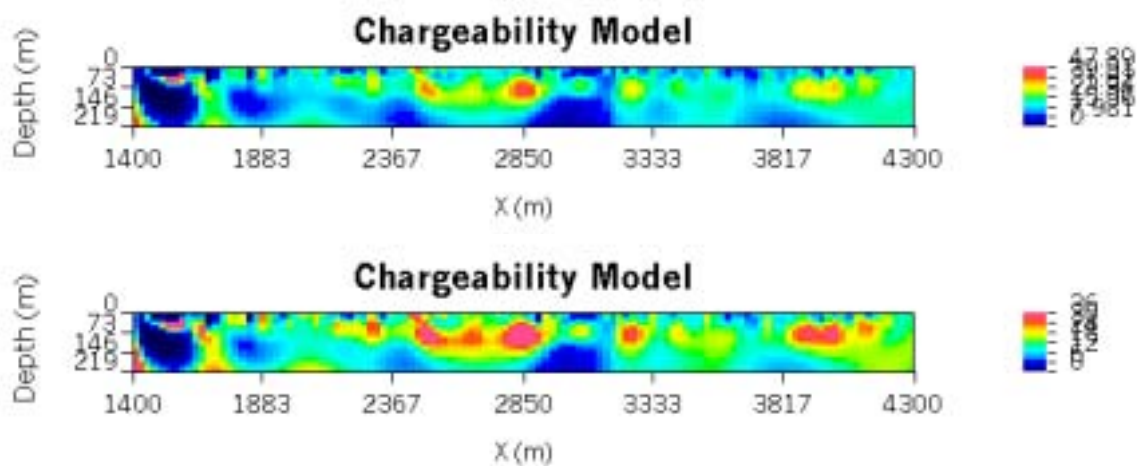


## Line 1700N Chargeability

Pseudosection

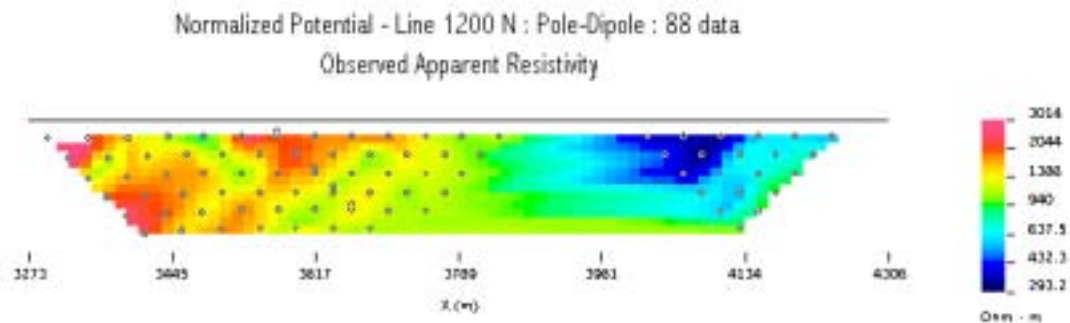


Inversion Depth Section

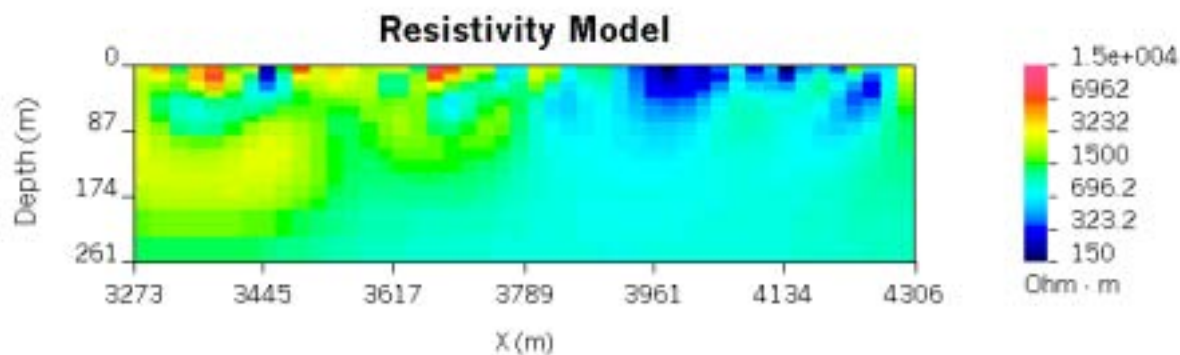
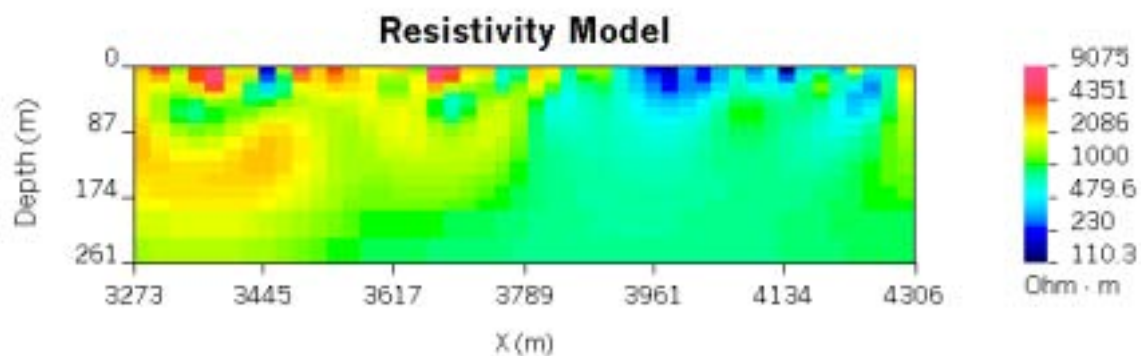


## Line 1200N Resistivity

Pseudosection



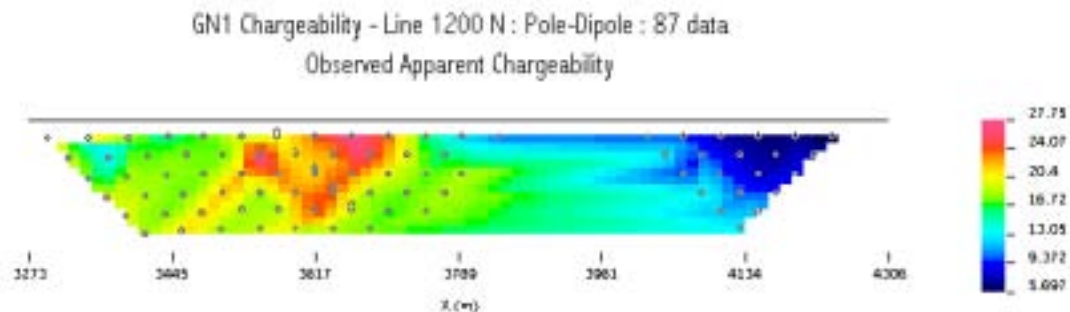
Inversion Depth Section



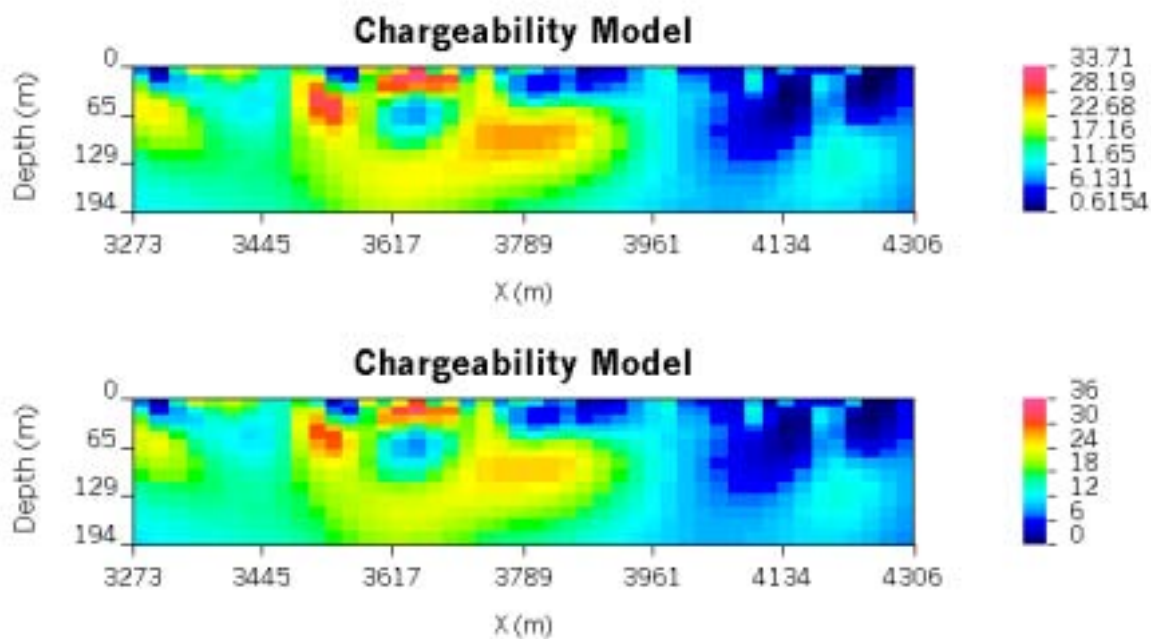


## Line 1200N Chargeability

Pseudosection

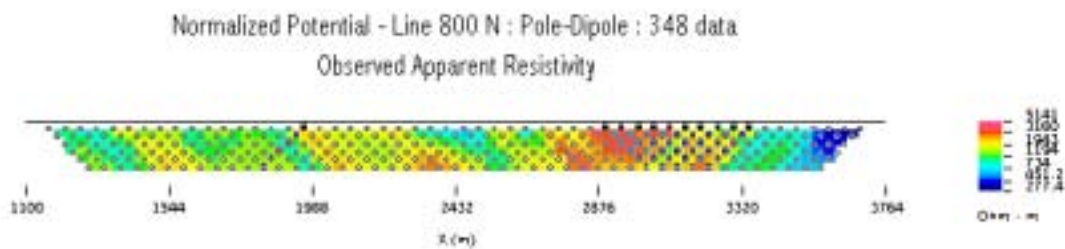


Inversion Depth Section

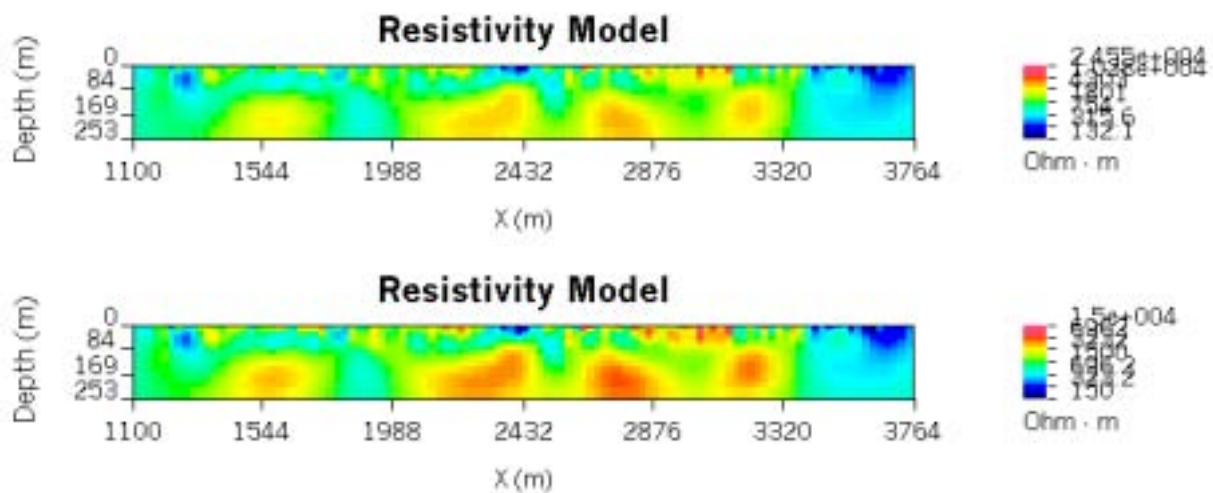


## Line 800N Resistivity

Pseudosection



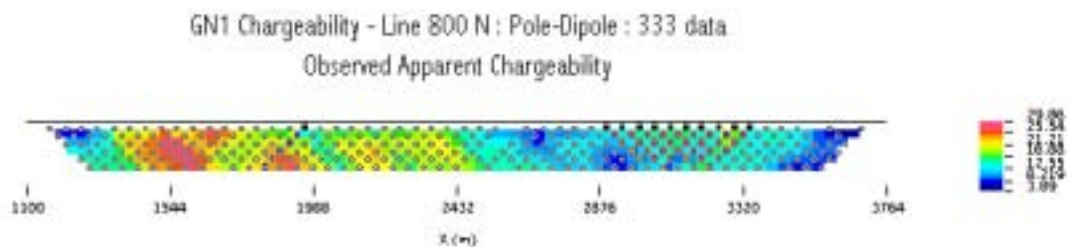
Inversion Depth Section



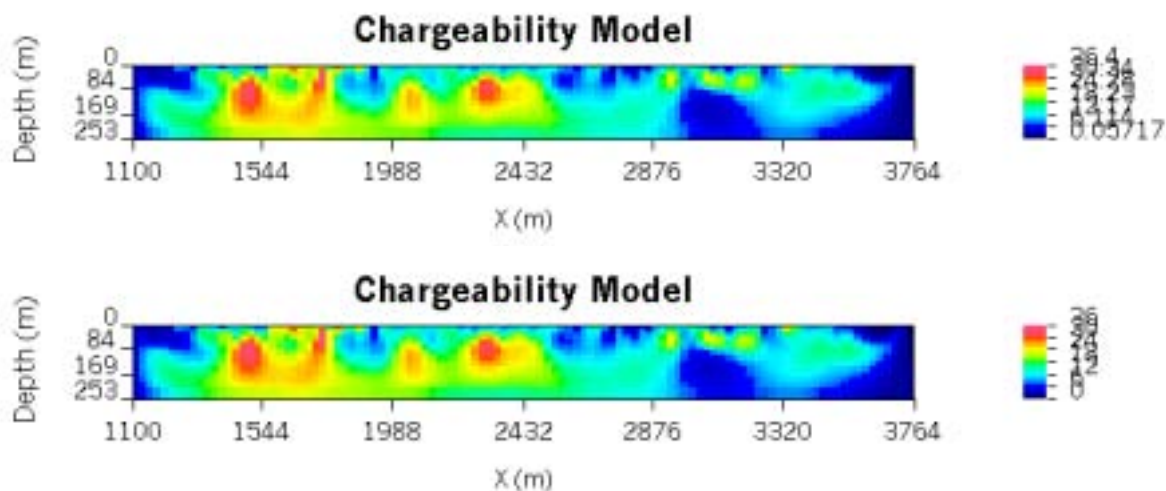


## Line 800N Chargeability

Pseudosection

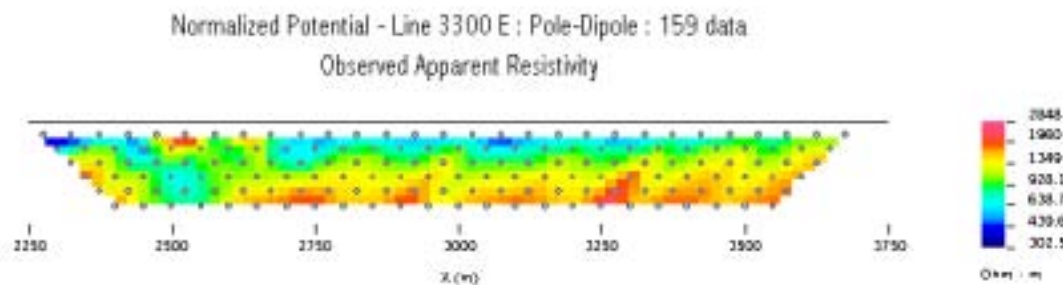


Inversion Depth Section

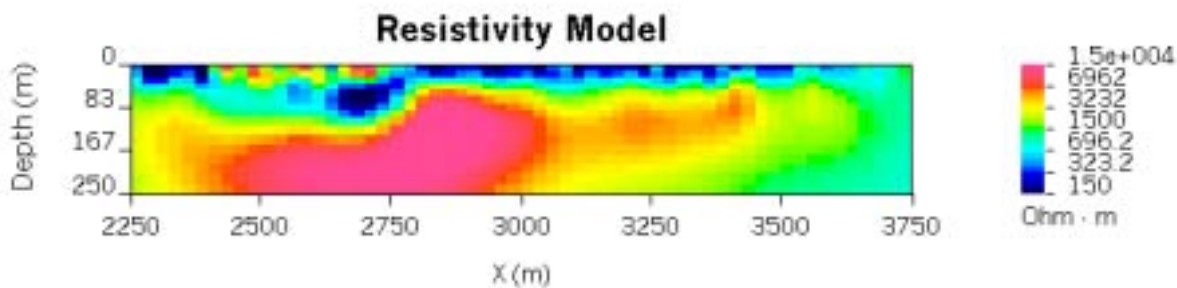
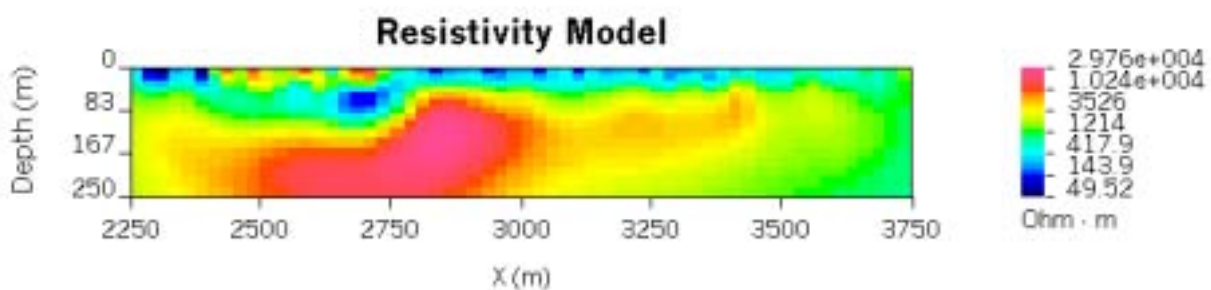


## Line 3300E Resistivity

Pseudosection

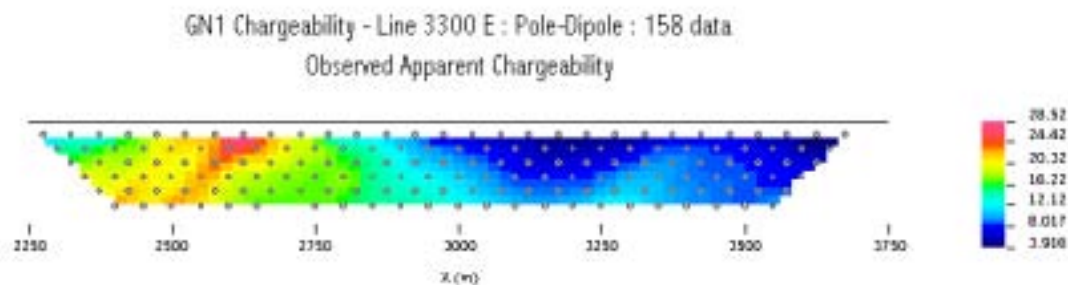


Inversion Depth Section



## Line 3300E Chargeability

Pseudosection



Inversion Depth Section

