ASSESSMENT REPORT

on a

3D Induced Polarization and Resistivity Survey

for WYN DEVELOPMENTS INC. RECEIVE on the AUG 1 6 2004 **RAND PROJECT** Gold Commissioner's OBJEUSTRY MOUNTAIN AREA VANCOUVER, B.C. Lillooet BC, Canada Latitude 50° 38'N LONG LINK EV BRANCH REPORT GEOLUSESS Mining Zone: Kamloops NTS sheet: 92I/12 Surre SJ Geophy t d 🏅 Report by

E. Trent Pezzot, B.Sc., P.Geo.

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	SJ Geophysics Ltd. / S.J.V. Consultants Ltd. 11762 - 94th Ave., Delta, B.C. Canada i tel: (604) 582-1100 fax: (604) 589-7466 e-mail: sydv@sjgeophysics.com

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1 <u>SUMMARY</u>

SJ Geophysics Ltd. was commissioned by Wyn Developments Inc. to conduct an induced polarization survey across a portion of their Blustry Mountain Project, in south-central BC. This area is being explored for its potential to host a gold and silver enriched epithermal alteration system similar to the Blackdome deposit located some 100 km to the northwest. It was the intention of the survey to help in this exploration by mapping geology and (hopefully) directly detecting exploration targets.

The survey was configured as a 3-D array with current and potential electrodes located on adjacent survey lines, spaced at 100 metre intervals. This configuration allows for the application of 3-D interpretation techniques, including 3-D inversion algorithms.

Combinations of resistivity and chargeability characteristics have outlined 3 distinct geological regimes across the survey area. A large portion of the northeastern corner of the grid (Lines 1600N - 2400N) is covered by a thin (50m thick) cap of highly resistive material. This overlies a 100m thick layer of highly variable material that include several pods of extremely conductive and chargeable material. Basement rocks in this area appear to relatively uniform, exhibiting low resistivity and elevated chargeability. The second regime is mapped from 1500N to 900N. It is also characterized with a resistive cap which often occurs as two or more thin layers. The underlying rocks exhibit low resistivity and low chargeability and contain a few isolated anomalies. The third regime covers the southwest corner of the grid. It is characterized by scattered zones of variable chargeability and resistivity in the top 75 metres. At depth the geophysical responses become more uniform and reveal two structural trends: N15⁰W and N45⁰E.

There are several lineations and trends that are mapped as abrupt discontinuities of a particular geophysical parameter. These are likely representing sharp geological contacts or fault zones. There are several pods of extremely high resistivity that can be interpreted as areas of silica flooding. Several pods of anomalously high chargeability have been identified that could represent disseminated sulphide mineralization.



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Figure 1 Figure 1: Location map

2 INTRODUCTION

This report documents the results from a 3-D IP survey completed across the Blustry Mountain Project area, east of Lillooet, B.C. The area is being explored for a gold and silver enriched epithermal alteration system. The intention of the survey was twofold: first to assist in the geological mapping and second, to identify specific anomalies consistent with the exploration model.

3 LOCATION AND ACCESS

The property lies 18 km east of Lillooet, south-central British Columbia, in the Kamloops Mining Division and NTS 92I/12. It is centred on latitude $50^{0}29'31"N$ and longitude $122^{0}19'30"W$. (Figure 1)

There was no direct road access to the area of the IP survey grid. A tent camp was set up on the property and the crew and equipment were ferried in from the Lillooet airport using a Long Ranger provided by Valley Helicopters. Snowmobiles were used for transportation across the grid area.

4 <u>Property</u>

The Blustry project lies within the Rand Property, which is comprised of 44 staked, fourpost mineral claims, totalling 668 claim units (16,700 hectares). The claims are contiguous and details of the mineral titles are listed in Table 1 and illustrated on Figure 2.

Table 1. Status of Blustry Mineral Titles					
Claim Name	# of Units	NTs Map Sheet	Record #	Expiry date	Mining Division
1 Solomon 1	20	92I/07/1,07/2	402638	04/05/13	Lillooet
2 Solomon 2	12	921/07/2	402639	04/05/13	Lillooet
3 Solomon 3	12	92I/07/2	402640	04/05/13	Lillooet
4 Solomon 4	20	921/07/1,07/2	402641	04/05/13	Lillooet
5 Comstock 1	20	921/07/2	402634	04/05/13	Kamloops
6 Comstock 2	20	921/07/2	402635	04/05/13	Kamloops
7 Comstock 3	20	921/07/2	402636	04/05/13	Kamloops
8 Comstock 4	20	921/07/2	402637	04/05/13	Kamloops
9 Homestake 1	20	921/07/2	402642	04/05/13	Kamloops
10 Homestake 2	20	92I/07/2	402643	04/05/13	Kamloops
11 Homestake 3	20	921/07/2,06/2	402644	04/05/13	Kamloops

12 Homestake 4	20	921/07/2,06/2	402645	04/05/13	Kamloops
13 Anaconda 1	18	92I/06/2	402646	04/05/13	Kamloops
14 Anaconda 2	4	92I//06/2	402647	04/05/13	Kamloops
15 Anaconda 3	3	92I/06/2	402648	04/05/13	Kamloops
16 Anaconda 4	18	92I/06/2	402649	04/05/13	Kamloops
17 Eureka I	18	92I/06/2	402650	04/05/13	Kamloops
18 Eureka 2	18	92I/06/2	402651	04/05/13	Kamloops
19 Motherlode 1	18	921/06/2	402659	04/05/13	Kamloops
20 Motherlode 2	18	92I/06/2	402660	04/05/13	Kamloops
21 Motherlode 3	18	921/06/2,05/2	402661	04/05/13	Kamloops
22 Motherlode 4	18	921/06/2,05/2	402662	04/05/13	Kamloops
23 Blustry Mountain 1	15	92I/06/2	395042	04/07/18	Kamloops
24 Blustry Mountain 2	20	921/06/2	395043	04/07/18	Kamloops
25 Blustry Mountain 3	15	92I/06/2	395044	04/07/18	Kamloops
26 Blustry Mountain 4	20	92I/06/2	395045	04/07/18	Kamloops
27 Kalgoorlie 1	4	92I/05/2	402663	04/05/13	Kamloops
28 Kalgoorlie 2	3	921/05/2	402664	04/05/13	Kamloops
29 Kalgoorlie 3	18	92I/05/2	402665	04/05/13	Kamloops
30 Kalgoorlie 4	16	92I/05/2	402666	04/05/13	Kamloops
31 Eureka 3	12	92I/05/2	402652	04/05/13	Kamloops
32 Eureka 4	12	92I/05/2	402653	04/05/13	Kamloops
33 Bonanza 1	20	921/05/2	402667	04/05/13	Kamloops
34 Bonanza 2	20	921/05/2	402668	04/05/13	Kamloops
35 Bonanza 3	20	921/05/2	402669	04/05/13	Kamloops
36 Bonanza 4	20	921/05/2	402670	04/05/13	Kamloops
37 SL 1	16	92I/05/2	402680	04/05/13	Kamloops
38 SL 2	18	921/05/2,05/3	402681	04/05/13	Kamloops
39 SL 3	18	921/05/2,05/3	402682	04/05/13	Kamloops
40 SL 4	16	921/05/2	402683	04/05/13	Kamloops
41 Cripple Creek 1	20	921/05/3	402684	04/05/13	Kamloops
42 Cripple Creek 2	20	921/05/3	402685	04/05/13	Kamloops
43 Cripple Creek 3	20	92I/05/3,04/3	402686	04/05/13	Kamloops
44 Cripple Creek 4	20	921/05/3,04/3	402687	04/05/13	Kamloops
Total	668				

No examination of the claim posts was made by any of the geophysical crew. The following comments are copied from a previous assessment report on this property, authored by Paul Metcalfe and dated September, 2003. According to personnel from Wyn Developments Inc. these statements are still valid and accurately describe the property status.

"An examination of the relevant claim forms by G. Nicholson found these to be in order. All claims are therefore in good standing. The Blustry Mountain 1-4 claims are 100% owned by J. Shearer and are in their second anniversary year. The remaining claims are presently held by Mr.

R. Krause for the Rand Syndicate. Mr. Krause is a director of Wyn Developments. These claims are in their first anniversary year.

Two separate agreements exist, one with J.T. Shearer (dated July 1, 2003) to option 100% of the Blustry Mountain 1-4 claims and the second (dated July 17, 2003) with the Rand Syndicate to acquire by purchase 100% of all claims held by the Syndicate and surrounding the Blustry Mountain Claims."

5 <u>Geology</u>

The geological descriptions provided below have been copied from a Summary Geology Report on the Rand Project, written by Paul Metcalfe and dated September, 2003.

5.1 Regional Geology

A geological map of the Blustry and surrounding areas is shown in Figure 3. Despite the apparently comprehensive nature of the map, it is based upon mapping carried out by Duffell and McTaggart (1952) and Trettin (1961); smaller studies by Mortimer (1987) and Read (1988a, 1988b, 1990) have augmented the broader regional mapping. The area was compiled as part of the Geological Survey of Canada's Terrane Assemblage Map by Monger and Journeay (1994).

The Blustry property lies on the east side of the Fraser Fault, which experienced Eocene strike-slip movement of approximately 80 km and which forms a geological boundary to the west. The basement to the area comprises rocks of the Permo-Triassic Cache Creek Complex, which are bounded to the southwest by granodioritic intrusive rocks associated with the Mount Martley and Tiffin Creek stocks.

The Blustry property itself is shown on Figure 3 to be underlain by calc-alkaline volcanic rocks of the Lower Cretaceous Spences Bridge Group. Outliers of Eocene volcanic rocks assigned to the Kamloops Group occur to the east. This mapping is not entirely correct; Richards (1984a) in a report on previously held ground noted that mineralization was hosted by rocks which he assigned to the "Tertiary Kingsvale Group".

Certainly the Spences Bridge Group is not prospective for epithermal deposits, nor are siliceous volcanic rocks common in that stratigraphic unit. However, the Kingsvale volcanic rocks are cited as being Upper Cretaceous in age (Preto 1979) and their definition and extent are not well constrained. This author correlates the volcanic rocks described by Richards with outliers of Eocene volcanic rocks, 45-50 Ma in age which are exposed to the east of the claim

group. These are identified as Eocene Kamloops Group on the regional geology map, but later work (BC Geofile 2000-3) assigns these volcanic rocks as "unnamed".

To the north and west, the volcanic rocks hosting the Blackdome low-sulphidation epithermal deposit are identified as Eocene to Oligocene, uncorrelated with either Kamloops or Ootsa Lake Groups. For now the latter, uncorrelated terminology will be used for the target units, pending mapping and more precise correlation.

Regional structural geology in the area is as little documented as stratigraphy. Brittle faults cross the property, with two prominent strike directions, parallel (northwesterly) and crudely perpendicular (northeasterly) to the structural grain of the Canadian Cordillera. Normal movement is apparent on several of the faults by the lateral juxtaposition of the Eocene volcanic rocks against older rocks.

5.2 Property Geology

The authors reiterate here that they visited the Blustry property area only briefly. The following is, therefore, abridged from or a summary of previous reports (Richards 1984a, Gonzalez and Leshow 1987).

No formal geological map exists on a property scale for the area now covered by the Blustry property. As noted above, regional mapping by the Geological Survey of Canada (Duffel and McTaggart 1952) is over 50 years old and subsequent mapping by the British Columbia Geological Survey Branch (Mortimer 1987, Read 1988a, 1988b, 1990) did not cover the entire area. The following is a summary based on three assessment reports (Richards 1984a,b, Gonzalez and Lechow 1987) describing geochemical and geophysical surveys of areas now covered by the central part of the property.

Previous authors have noted that the Blustry mineral claims are underlain by volcanic rocks of the lower Cretaceous Spences Bridge Group. This Group is composed mainly of an accumulation of lavas and pyroclastics rocks. Most of the lavas are porphyritic and are fine to coarse-grained rocks of various colours. The colours are red, green, mauve, purple, brown grey, white and black. This unit is not considered prospective for economic mineralization.

In the vicinity of Blustry Mountain, dacites and minor rhyolites apparently intrude or overlie the Spences Bridge Group and are intruded by a north-easterly trending dyke swarm of creamy pink, weakly feldspar hornblende phyric andesite. Gabbroic rocks intrude the volcanic sequence to the southwest of Blustry Mountain (Richards 1984a,b) and a small plug of syenite, possibly a

coarser-grained equivalent of the pink feldspar-phyric dykes has been observed south of Cairn Peak (on the original 80-unit claim group).

Basaltic volcanic rocks of the Kamloops Group are found to the east of the property, near Hat Creek. In Hat Creek valley, a thick section of sedimentary rocks is preserved in a graben that is floored by Eocene volcanic rocks (Richards, 1984a,b). The volcanic rocks on the Blustry property have been variously assigned to the Kamloops Group (Monger and Journeay 1994), Late Cretaceous to Early Tertiary Kingsvale Group (Richards 1984a,b, Gonzalez and Leshow 1987) or "uncorrelated Tertiary". It is more convenient, in the absence of information, to regard the siliceous volcanic rocks as uncorrelated volcanic rocks of the Early Tertiary (probably Eocene); a belt of Eocene rocks of composition similar to those reported at Blustry extends southerly from the Blackdome Mine.

5.3 Styles of alteration

Zones of alteration are strongly controlled by structure. The most prominent structural trend is northeasterly while north-northwesterly trends also appear to have influenced the localization of alteration. These structural trends are thought to reflect Lower Tertiary translation and extensional tectonics that are well developed within this area.

The northeasterly trending dyke swarm is associated with a clay-sulphide zone that is developed over an area 4500 metres long and as wide as 1500 metres. Within the clay-sulphide zone area areas of silicification (silica flooding) which host precious metal and minor base metal mineralization.

5.4 Styles of mineralization

Several types of mineralization were identified and described by Richards (1984a) and were also described by Gonzalez and Lechow (1987). Quartz breccias with quartz crystal-lined vugs and intense silicification of included wallrock have been noted in float. Sulphide content is generally less than 1 or 2 percent but tetrahedrite, galena and other silver-coloured sulphides have been recognized with fine-grained pyrite.

A second type of silica flood occurs as dark grey quartz veins in parallel bands, commonly 2 mm wide but in places attaining a width of several centimeters. These compose as much as 70%, but on average 10%, of rock volume. This mineralization is developed in an area 50 to 100 m wide and 200 to 300 m long.

A third type of silicification occurs in rhyolite breccia with moderate clay alteration and less than 3 percent void space. The rhyolite breccia contains local zones with silicified fragments and with grey quartz partly filling the vugs. Silica flooding also occurs within the rhyolite and is accompanied by intense clay alteration.



Figure 3: Regional Geology Map

6 PREVIOUS WORK

The descriptions of previous work provided below have been copied from a Summary Geology Report on the Rand Project, written by Paul Metcalfe and dated September, 2003.

In 1984 a geochemical survey was initiated by Ryan Exploration, a division of U.S. Borax, and designed to provide geochemical data over the area considered to be the best target (Richards 1984a). A total of 1,076 samples were collected of which 3 were stream sediments, 85 were rock chips, and 988 were soil samples. Results indicated several areas of highly anomalous values in antimony, arsenic, copper, lead, mercury, molybdenum and zinc, coincident with anomalous gold and silver values.

In 1987 Aerodat Ltd. of Mississauga, Ontario was commissioned by Kangeld Resources to conduct an airborne geophysical survey over the property. This survey consisted of a low-level, helicopter-supported programme which included a frequency VLF-electromagnetic system, a high sensitivity caesium vapour magnetometer. Results of this survey were used to control the grid placement for a 1987 soil sampling programme conducted by Mark Management Ltd. (Gonzalez and Lechow 1987).

In 1987 Mark Management Ltd. under the direction of Archean Engineering conducted a soil geochemical survey over a grid area of 900 m x 1000 m in size. A total of 349 soil samples were collected and analyzed by Chemex Labs Ltd. using an ICP geochemical analytical technique. In general, anomalous values for Au, Ag, As, Cu, Hg, Mo, Sb, Pb, and Zn outlined an open-ended zone 650 m long by 220 m wide (Gonzales and Lechow 1987).

7 <u>GEOPHYSICAL TECHNIQUES</u>

7.1 IP Method

The time domain IP technique energizes the ground surface with an alternating square wave pulse via a pair of current electrodes. On most surveys, such as this one, the IP/Resistivity measurements are made on a regular grid of stations along survey lines.

After the transmitter (Tx) pulse has been transmitted into the ground via the current electrodes, the IP effect is measured as a time diminishing voltage at the receiver electrodes. The IP effect is a measure of the amount of IP polarizable materials in the subsurface rock. Under ideal circumstances, IP chargeability responses are a measure of the amount of disseminated metallic sulfides in the subsurface rocks.

Unfortunately, there are other rock materials that give rise to IP effects, including some graphitic rocks, clays and some metamorphic rocks (serpentinite for example) so, that from a geological point of view, IP responses are almost never uniquely interpretable. Because of the non-uniqueness of geophysical measurements it is always prudent to incorporate other data sets to assist in interpretation.

Also, from the IP measurements the apparent (bulk) resistivity of the ground is calculated from the input current and the measured primary voltage.

With regard to precision, IP/Resistivity measurements are generally considered to be repeatable within about five percent. However, they will exceed that if field conditions change due to variable water content or variable electrode contact.

IP/Resistivity measurements are influenced, to a large degree, by the rock materials nearest the surface (or, more precisely, nearest the measuring electrodes), and the interpretation of the traditional pseudosection presentation of IP data in the past have often been uncertain. This is because stronger responses that are located near surface could mask a weaker one that is located at depth.

7.2 <u>3-D IP Technique</u>

Three dimensional IP surveys are designed to take advantage of the interpretational functionality offered by 3-D inversion techniques. Unlike conventional IP, the electrode arrays are no longer restricted to in-line geometry. Typically, current electrodes and receiver electrodes are located on adjacent lines. Under these conditions, multiple current locations can be applied to a single receiver electrode array and data acquisition rates can be significantly improved over conventional surveys.

The following figure (Figure 2) shows the typical array setup used for this survey near the centre of a line. Current electrodes are shown in blue dots and receiver locations in red triangles.



Figure 4: IP Array Setup

The array was adapted at the ends of the lines and for very short lines but typically used one of the standard receiver electrode configurations described below. Once the easternmost current electrodes were used, the entire array would shift 200m to east). The locations of the current electrodes for the starts and ends of the lines normally extended past the receiver electrodes but depended on the extent of the adjacent survey lines.

7.3 Inversion Programs

"Inversion" programs have recently become available that allow a more definitive interpretation, although the process remains subjective.

The purpose of the inversion process is to convert surface IP/Resistivity measurements into a realistic "Interpreted Depth Section." The use of the inversion routine is a subjective one because the input into the inversion routine calls for a number of user selectable variables whose adjustment can greatly influence the output. The output from the inversion routines do assist in providing a more reliable interpretation of IP/Resistivity data, however, they are relatively new to the exploration industry and are, to some degree, still in the experimental stage.

The inversion programs are generally applied iteratively to, 1) evaluate the output with regard to what is geologically known, 2) to estimate the depth of detection, and 3) to determine the viability of specific measurements.

The Inversion Program (DCINV3D) used by the SJ Geophysical Group was developed by a consortium of major mining companies under the auspices of the UBC-Geophysical Inversion Facility. It solves two inverse problems. The DC potentials are first inverted to recover the spatial distribution of electrical resistivities, and, secondly, the chargeability data (IP) are inverted to recover the spatial distribution of IP polarizable particles in the rocks.

The 3-D Inversion technique involves defining a 3-dimensional space directly beneath the survey grid and dividing it into discrete, rectangular cells (mesh). Additional "padding" cells are defined to the sides of the block to move edge effects away from the area of interest. A model is constructed by assigning physical attributes (conductivity and chargeability in the case of the IP survey) to each cell in the mesh. A theoretical response to this model is calculated for the electrode array configuration and the results compared to the field measurements. The model is adjusted until the errors between the theoretical and field measurements are minimized to user defined limits.

Viewers are available that allow one to examine the solutions as 3-dimensional objects from various perspectives. The block model can also be cut and viewed as either vertical or horizontal cross-sections and correlated to drilling results.

8 <u>FIELD SURVEYS AND INSTRUMENTATION</u>

The SJ Geophysics Ltd. IP crew consisted of six men: geophysicist Lee Gulbransen, and technicians Jon Jacobson, John Wilkinson, RJ Ewen, Colin Russell and Cameron Wallace. The survey ran from April 11 to May 1, 2004 and included 2 days mob/de-mob and 19 days of survey production, less approximately ¹/₂day down time due to equipment problems.

The survey grid consisted of 25 lines (00N to 2400N), oriented NW-SE and nominally spaced at 100 metre intervals. The grid was established by compass and tight chain by a crew of two line cutters who worked in advance of the geophysical crew. The survey lines were of variable length, ranging from 450 to 1200 metres and totalled approximately 19.4 km in length. Stations were flagged at 50 metres along these lines.

During the course of the survey, the geophysical crew took GPS readings at critical points along these lines and measured topographic slopes between stations in order to accurately define the electrodes locations.

The survey used an Iris VIP 3000 transmitter, a GDD TXII transmitter and the Iris Elrec 10 receiver. The technical specifications are listed in Appendix 3.

The IP survey was configured in a modified pole-dipole, 3-D array as described in section 7.2 above.

9 DATA PRESENTATION

The geophysical data from this are displayed in three formats, as indicated below. Plan maps are registered to the NAD83, Zone 10N UTM coordinate system and plotted at a scale of 1:5,000. Cross sectional profiles are registered to the local grid coordinate system (lines and stations). Elevations are in absolute values of metres above sea-level. This report has been prepared in two formats: hardcopy and electronic. In the bound hardcopy formats the plates and figures are located in map pockets at the back of the report. In the electronic format the plates and figures are provided as separate pdf (Adobe Portable Document Format) files.

9.1 <u>Plan Maps</u>

A plan map showing the location of the survey grid with respect to UTM, topography and claims is provided as Figure 2.

False colour contour maps of the interpreted resistivity and chargeability data were produced for depths of 25, 50, 75, 100, 150, 200 and 250 metres below the surface.

9.2 Cross Sections

Cross sections showing both the interpreted resistivity and chargeability were extracted from the 3-D model along each survey line (00N to 2400N) and are presented as false coloured plates at a 1:5000 scale.

9.3 <u>3-D Block Model</u>

Various 3-D perspective plots are included in the text of this report to illustrate anomalous responses interpreted from the data.

10 DISCUSSION OF RESULTS

The exploration target is a gold and silver enriched epithermal system, possibly associated with volcanism. These deposits are recognized by the presence of silicification (quartz veins), sulphides and hydrothermal alteration minerals, resulting from extensive fluid flow through rocks along structural features. At Blustry Mountain, intense clay alteration and silica flooding are both reported in an area of anomalous geochemistry that is typical of this target. The IP

technique was chosen as an efficient means of mapping both the extreme resistivity variations and sulphide accumulations expected from this environment.

The resistivity and chargeability components of the IP survey have outlined at least 3 distinctive geological regimes the area. Differing lithologies are typically delineated by a unique combination of amplitude and character of the two parameters. In addition, faults, contacts and structural features are typically mapped as discontinuities or patterns within or along larger, regional trends.

There are two surface, high resistivity zones located in the NE corner of the survey grid, both approximately 50 metres thick. One extends from line 2400N to 1600N and appears as a cap along the crest of the mountain. The other is mapped from line 1500N to 900N and appears to cover a NE facing slope. The 3-D perspective plot and two cross-sections shown below highlight these units.



Figure 5: 3-D Perspective Plot of Resistivity – view from SW – shows high resistive caps (shown as red areas) along crest and east slope of Blustry Mountain.



Figure 6(a) Line 2000N Cross-Sections through resistivity model showing layered nature of the two surface high resistivity zones in NE section of survey grid.

The area north of line 2000N, which is partially covered by one of these surface layers of very high resistivity material, exhibits a distinct geophysical signature. Cross-sections (running NE-SW through the inversion block) show the basement units (>150m depth) have a relatively uniform clevated chargeability values and low resistivities. The rocks between this basement and the resistive surface layer are generally quite variable and include several pods with high chargeability and very low resistivity. The southern and southeastern edges of this lithology appear to be abrupt contacts, possibly associated with faulting. There is also a distinct vertically oriented pipe-like body which occurs as a high resistivity and low chargeability unit intruding through this lithology. This is evident on the castern side of resistivity cross section for line 2000N shown above.



Figure 7a Chargeability

Figure 7b Resistivity



The second resistivity cap (1500N to 900N) often appears as two distinct layers (as shown on the cross section for line 1000N above). It generally overlies a low resistivity and low chargeability rock unit. There a few chargeability anomalies scattered across the area, typically directly below the resistive cap, but they are generally very small, isolated features and don't exhibit any clearly recognizable pattern. There is one notable exception of a moderate chargeability high and coincident strong resistivity low mapped across three lines (1200N to 1000N). It most clearly outlined on line 1100N where it appears be about 100metres wide (4300E to 4400E) and 200 metres high (100m to 300m depth). This anomaly appears to coincide with a NE trending fault zone that cuts across this area.

The southwestern portion of the survey grid (south of line 1000N) is characterized by small, scattered zones of variable chargeability and resistivity in the top 75 metres. Based on their size and amplitudes, several of these zones are considered anomalous and warrant further examination. There are no clear patterns evident in these near surface responses but at greater depths, two trend orientations become apparent: $\sim N15^0W$ and $\sim N45^0E$. It is not clear whether these trends are mapping structural breaks or lithological strikes but they appear to be significant features and that can be traced for hundreds of metres. There are a couple of resistivity trends within this general area that suggest the rock units dip moderately towards grid east.

11 <u>Recommendations</u>

There are a number of areas outlined by this IP survey that exhibit characteristics consistent with the exploration target model of an epithermal alteration system. While some of these will undoubtedly warrant further examination based on their own merit, it is recommended that these results first be correlated with existing geochemical, geological and geophysical data to determine whether any of them should be given a higher priority based on collaborating evidence.

There are several large scale geological structures and trends evident in the IP data. These should be evaluated in conjunction with other mapping tools (geology, geochemistry, airborne magnetics) to help construct more accurate geological maps of the area.

Respectfully submitted,

Per S.J.V. Consultants Ltd.

E. Trent Pezzot, B.Sc., P.Geo,

Geophysics, Geology

12 <u>APPENDIX 1 – STATEMENT OF QUALIFICATIONS – E. TRENT PEZZOT</u>

- I, E. Trent Pezzot, of the city of Surrey, Province of British Columbia, hereby certify that:
 - 1) I graduated from the University of British Columbia in 1974 with a B.Sc. degree in the combined Honours Geology and Geophysics program.
 - 2) I have practised my profession continuously from that date.
 - 3) I am a registered member of the Association of Professional Engineers and Geoscientists of British Columbia.
 - 4) I have no interest in Wyn Developments Inc. or any of their subsidiaries or related companies, nor do I expect to receive any.

Signed by: _____

E. Trent Pezzot, B.Sc., P.Geo. Geophysicist/Geologist

13 <u>APPENDIX 2 - COST BREAKDOWN</u>

The following cost breakdown is broken into two sections. The first section describes the costs invoiced to Wyn Developments Inc. by SJ Geophysics Ltd. that are directly applicable to the Blustry Mountain IP Project. The second section includes other costs incurred by Wyn Developments Inc. These costs were provided by Robert Krause of Wyn Developments Inc. for inclusion as related expenses for this project. The author has no direct knowledge of these expenses.

Item	Qty	Description	Discount	Unit Price	Total
		Field Survey			
1	1	Mob/deMob - April 11, May 1		3000	3,000.00
	2 13.5	IP Equipment with 2 operators - Apr 14, 16-26, 1/2 of 27, 28		1525	20,587.50
	13.5	Helper		185	2,497.50
4	13.5	Helper		185	2,497.50
4	5 13.5	Helper		185	2,497.50
		Standby weather for item 2- Apr 15, 1/2 of 27, 30	0.25	1525	2,859.38
7	2.5	Standby weather for item 3- Apr 15, 1/2 of 27, 30	0.25	185	346.88
	2.5	Standby weather for item 4 - Apr 15, 1/2 of 27, 30	0.25	185	346.88
9		Standby weather for item 5- Apr 15, 1/2 of 27, 30	0.25	185	346.88
10		Sat phone		20	340.00
11		4 x 4 truck	0.5	125	1,062.50
	1	April 29 - no charge day			
13	10	grid preparation - 5 men, 2 days each - Apr 12,13		200	2,000.00
		Field SubTotal			38,382.52
		Office			
14	105	Technicians - Data processing, Computer Processing		40	6,250.00
15	32	Geophysicist - Interpretation, Report		85	2,720.00
16	1	Report Printing, Map plotting - 6 sets		635	635.00
		Office SubTotal			9,605.00
		Field + Office Subtotal			47,987.52
		GST			3,359.13
		Total			51,346.65

Section 1: SJ Geophysics Expenses

Section 2: Wyn Developments Inc. Direct Expenses (as reported by Robert Krause, July 21, 2004)

Expense Type	Date:	Amount	
Assay and Recording	5/10/2004	7,459.00	
Camp Supplies	3/15/2004 to 6/15/2004	15,540.71	
Engineering and Geological	3/15/2004 to 6/15/2004	47,422.84	
Equipment rentals	4/21/2004	3,600.00	
Helicopter and Equipment	3/15/2004 to 5/3/2004	42,707.70	
Total		116,730.25	

14 APPENDIX 3 - INSTRUMENT SPECIFICATIONS

14.1 IRIS VIP-3000 IP Transmitter

Output Ratings

Technical:

Input impedance: 10 Mohm Input overvoltage protection up to 1000V Automatic SP bucking with linear drift correction				
Internal calibration generator for a true	calibration on request of the operator			
Internal memory:	3200 dipoles reading			
	chronization process on primary voltages signals whenever needed			
Proprietary intelligent stacking process				
Common mode rejection:	More than 100 dB (for Rs =0)			
Self potential (Sp)	: range: -15V - +15V			
	: resolution: 0.1 mV			
Ground resistance measurement				
range:	0.1-100 kohms			
Primary voltage	: range: 10µV - 15V			
	: resolution: 1µV			
	: accuracy: typ. 1.3%			
Chargeability	: resolution: 10µV/V			
	: accuracy: typ. 0.6%			

General:

Dimensions:	31x21x25 cm
Weight (with the internal battery):	9 kg
Operating temperature range:	-30°C -70°C
Case in fiber-glass for resisting to field	shocks and vibrations

14.2 IRIS Elrec 10 IP Receiver

Technical:

Input impedance:	10 Mohm		
Input overvoltage protection up to 1000)V		
Automatic SP bucking with linear drift	correction		
Internal calibration generator for a true	calibration on request of the operator		
Internal memory:	3200 dipoles reading		
Automatic synchronization and re-synchronization process on primary voltages signals whenever needed			
Proprietary intelligent stacking process rejecting strong non-linear SP drifts			
Common mode rejection:	More than 100 dB (for Rs =0)		
Self potential (Sp)	: range: -15V - + 15V		
-	: resolution: 0.1 mV		

Ground resistance measurement	
range:	0.1-100 kohms
Primary voltage	: range: 10µV - 15V
	: resolution: 1µV
	: accuracy: typ. 1.3%
Chargeability	: resolution: 10µV/V
	: accuracy: typ. 0.6%

General:

Dimensions:31x21x25 cmWeight (with the internal battery):9 kgOperating temperature range:-30°C -70°CCase in fiber-glass for resisting to field shocks and vibrations

14.3 GDD Tx II IP Transmitter

Input voltage:	120V / 60 Hz or 240V / 50Hz (optional)
Output power:	1.4 kW maximum.
Output voltage:	150 to 2000 Volts
Output current:	5 ma to 10Amperes
Time domain:	Transmission cycle is 2 seconds ON, 2 seconds OFF
Operating temp. range	-40 [°] to +65 [°] C
Display	Digital LCD read to 0.001A
Dimensions (h w d):	34 x 21 x 39 cm
Weight:	20kg.















.5

2.5

44.5




























Interpreted Resistivity (Ohm-m)











Interpreted Chargeability (mv/v)



Interpreted Chargeability (mv/v)



3100 3200 3300 3400































Survey by: SJ Geophysics Ltd. IP Inversion by: S.J.V. Consultants Ltd. Processing Date: May, 2004 Projection: UTM Datum: NAD83 Zone: 10 Mapping Date: Jun, 2004 Wyn Developments Ltd. Blustry Mountain Grid Lilloet, BC Canada **3D IP Survey 3D IP Inversion model** Interpreted Resistivity (Ohm-m) 150m Below Surface

SJ Geophysics Ltd.









Survey by: SJ Geophysics Ltd. IP Inversion by: S.J.V. Consultants Ltd. Processing Date: May, 2004 Projection: UTM Datum: NAD83 Zone: 10 Mapping Date: Jun, 2004



SJ Geophysics Ltd.



50m Below Surface









