2005 GEOLOGICAL and GEOCHEMICAL REPORT ON THE KIZMET, LJ, SUTLAHINE, EMU, LAW, BS-J, TUNJONY, AND PLUM PROPERTIES

VOLUME 1 – TEXT & APPENDICES

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Gold Commissioner's Office
VANCOUVER, B.C.

Located in the Sutlahine River Area,
Atlin Mining Division
British Columbia, Canada

NTS: 104K-07, 104K-08W
104K-10, 104K-11,
and 104K-14

058° 35' north Latitude
132° 85' west Longitude

Owned By:

Barrick Gold Incorporated
&
Rimfire Minerals Corporation

Work Performed By:
Barrick Gold Inc
Suite 700, 1055 West Georgia Street
P.O. Box 11120
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Submitted by

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March 24th, 2006

EXECUTIVE SUMMARY

The Kizmet project is a joint venture between Barrick Gold Inc. and Rimfire Minerals Corporation covering a continuous area approximately 70 km X 20 km in size located approximately 120 km southeast of Atlin BC. The project covers a linear belt of Late Cretaceous volcano-plutonic centres thought to be prospective for high sulphidation epithermal mineralization.

The objective of the project in 2005 was to secure a large land package of highly prospective terrain followed by rapid, priority based reconnaissance/evaluation style exploration to determine the potential of the belt. Individual targets deemed to have sufficient potential for >3 MOz Au deposits were to be moved to a more detailed phase of exploration. Areas found not to be prospective were to be dropped following the field season.

Fieldwork in 2005 did not identify any areas of advanced argillic alteration or high sulphidation epithermal mineralization. 848 rock samples were taken of which 52 contained significant levels of precious or base metals; all appear to be representative of other unimportant styles of mineralization including base metal rich massive sulphide veins, quartz +/- arsenopyrite-sphalerite veins, and quartz-chalcopyrite veining. Some of these may be related to weak porphyry style mineralization in the southwestern areas. A regional scale stream sediment survey was conducted to confirm and improve upon BC government RGS survey results from 1987. The results from the 2005 survey appear to confirm the presence of anomalous gold in the belt as well as outline more restricted basins containing epithermal suite elements. The strongest multielement anomaly sources the Thorn property; an intermediate to high sulphidation epithermal prospect owned by third parties.

Although results from the field season were generally poor, several drainages were highlighted during the 2005 regional stream sediment survey. An area in the far northwest part of the Kizmet claim block contains 8 contiguous drainages

that have Au, Sn or As, Sb +/- Ag-Pb-Zn anomalies. This area is in need of further investigation. All of the current mineral claim cells should be cut down to include only the most prospective areas and maximize applicable 2005 assessment expenditures. The focus of any future work should be shifted away from Andean-style high sulphidation epithermal bulk mineable deposits towards "Thorn style" structurally controlled intermediate-high sulphidation epithermal mineralization.

The 2005 Kizmet project assessment expenditures for 2005 are $281,345.57.

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INTRODUCTION

This report describes exploration work carried out on the Kizmet, LJ, Sutlahine, Emu, Law, BS-J, Tunjony, and Plum properties in 2005. The project is located approximately 120 kilometres southeast of the town of Atlin, in Northwest British Columbia (Figure 1). In late 2004, Barrick identified this area as having high potential to host high sulphidation epithermal deposits. This impetus for this was partly based on the historic discoveries of this style of mineralization at the Thorn property, roughly located in the center of the 2005 project area. Subsequently, a land package was staked though BC Mineral Titles Online staking and the Kizmet Joint Venture (JV) agreement was completed with Rimfire Minerals Corporation. The claims are jointly owned by Barrick Gold Inc and Rimfire Minerals Corporation while Barrick was the operator in 2005. The 8 adjoining properties are collectively known as the Kizmet project (Figure 2) and will be referred to in this manner throughout this report.

Within the Kizmet project claim groups, Barrick Gold Inc. completed reconnaissance style prospecting and sampling over the period from June 9 to Aug 28, 2005. Fieldwork focused on high priority areas identified using RGS multi-element stream sediment anomalies, historic Minfile mineral occurrences and geologic stratigraphy favourable for epithermal style gold mineralization.

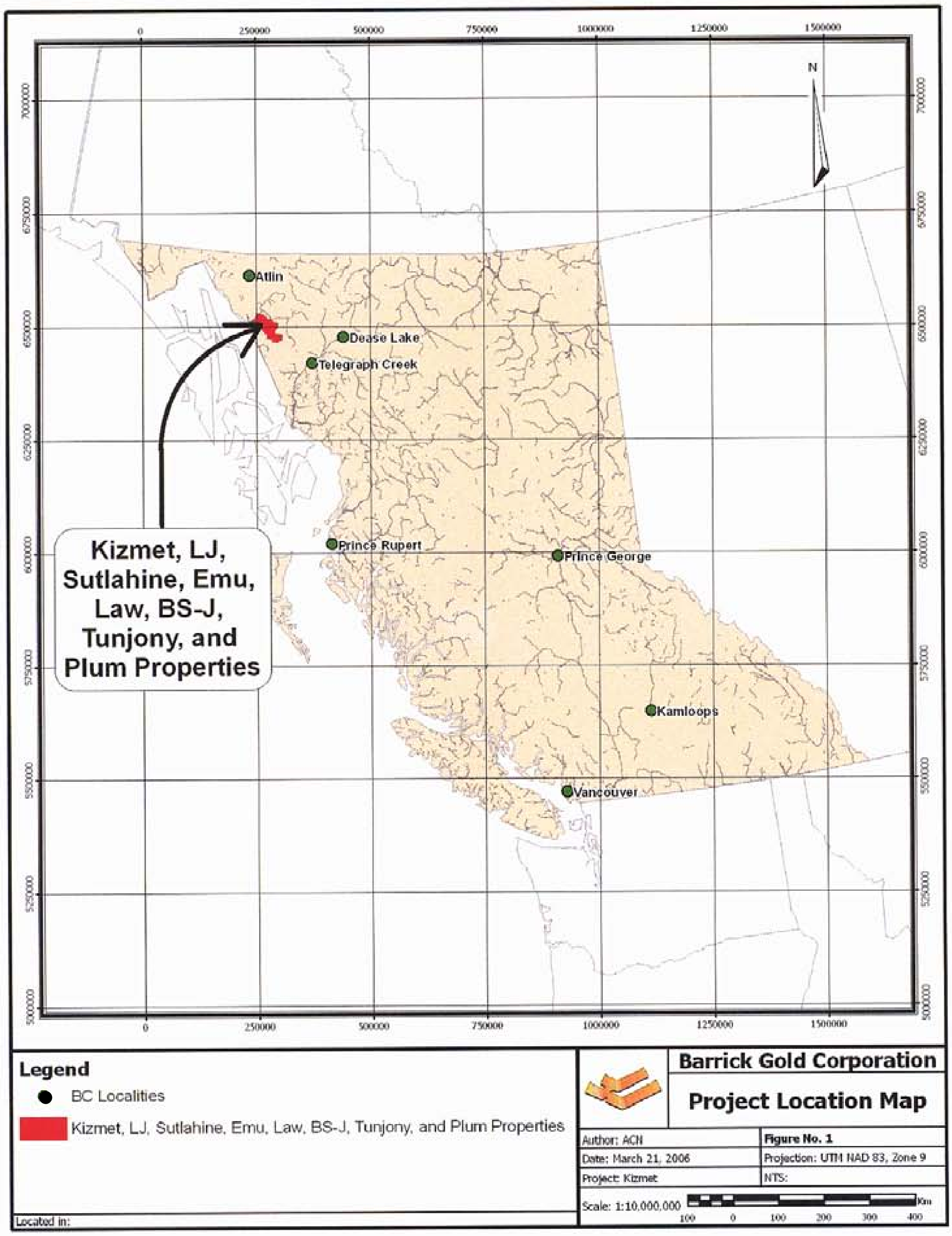

Figure 1: Project Location Map.

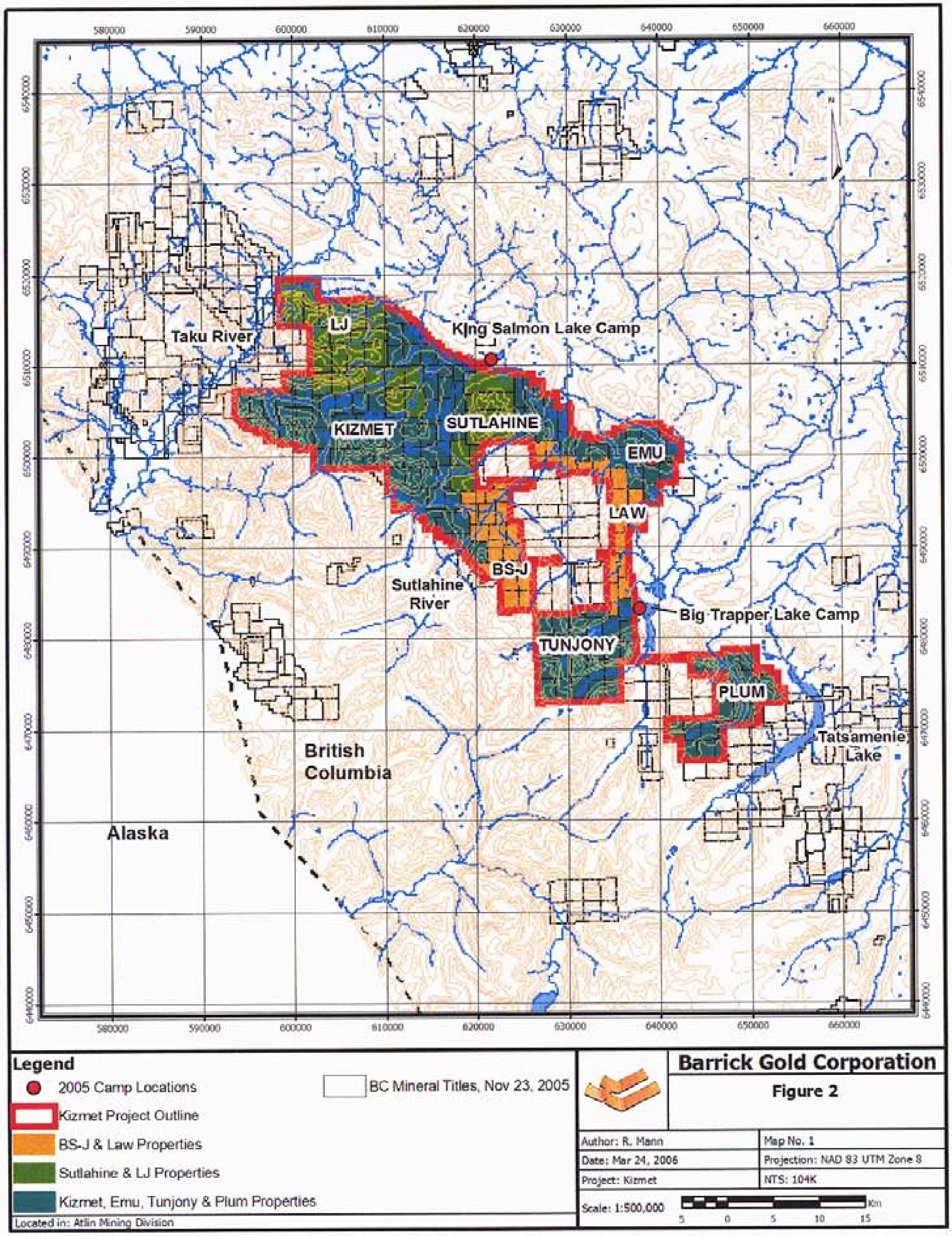

Figure 2: Kizmet project – Property location map.

2.0 LOCATION, ACCESS AND PHYSIOGRAPHY

The Kizmet main claim body lies within the Coast Mountain Range of Northwestern British Columbia, approximately 120 km southeast of Atlin, 135 km northwest of Telegraph Creek and 165 km west of Dease Lake (Figure 1). The property is located in British Columbia's Atlin Mining Division, and is centred at 58°35' north latitude and 132°50' west longitude.

The Kizmet property is accessible by helicopter and by float plane. Helicopter bases exist at Atlin or Dease Lake. King Salmon Lake, located along the north-eastern margin of the Kizmet project claim body is suitable for float equipped aircraft as is Big Trapper Lake in the southeastern end of the project area. An airstrip located at the dormant Tulsequah Chief Mine (10 km northwest of the Kizmet project claim block) provides access by plane on wheels. The Golden Bear Mine road, 85 km to the southeast, provides the closest road access.

The Kizmet project claim block is bounded to the northwest by the Taku River, to the east by Oneway Lake and to the southeast by Tatsamenie Lake. Major water bodies are tributaries of the Sutlahine, Inklin and Taku Rivers. Elevations range from 60 m on the Taku River flood-plain to over 2380 m. Exploration in 2005 focused between 900 and 2200 meters elevation.

Below tree line, at approximately 925 m elevation, the Kizmet area is dominated by mature hemlock, spruce and locally fir. Open patches of devil's club, tag alder and lightening triggered burn are common. Summer and winter temperatures are best described as moderate. Annual rainfall may exceed 200 cm and several metres of snow commonly fall at higher elevations. The ideal exploration season for the Kizmet project claims extends from early June until late October.

3.0 CLAIM STATUS

The Kizmet project claim block consists of 170 contiguous claims (listed in Table 1). The Kizmet project claim block is located in the Atlin Mining Division, British Columbia and is approximately 70 Km long (NW-SE axis) by 20 Km wide. The claim block covers 956 square kilometres (95600 Ha) (Figure 2; Map 1).

In January 2005, Barrick Gold Inc. entered into a joint venture agreement with Rimfire Minerals Corporation that included all mineral titles included in Table 1. Work was carried out by Barrick Gold Inc. As of January, 2005 Current government records indicate that these claims are in good standing until the respective expiry dates listed in Table 1.

TABLE 1: Kizmet Project – List of Claims and Indicated Fieldwork

CLAIM NAME	TENURE NO.	DATE RECORDED	EXPIRY DATE	CLAIM STANDING	WORK DONE ON CLAIM
martin1	501026	2005/JAN/12	2006/JAN/12	GOOD	NO
Richard1	501109	2005/JAN/12	2006/JAN/12	GOOD	YES
Martin2	501119	2005/JAN/12	2006/JAN/12	GOOD	YES
Richard2	501128	2005/JAN/12	2006/JAN/12	GOOD	NO
Richard3	501144	2005/JAN/12	2006/JAN/12	GOOD	NO
Richard4	501163	2005/JAN/12	2006/JAN/12	GOOD	YES
martin2	501187	2005/JAN/12	2006/JAN/12	GOOD	YES
RDB1	501215	2005/JAN/12	2006/JAN/12	GOOD	NO
martin3	501218	2005/JAN/12	2006/JAN/12	GOOD	NO
RDB2	501241	2005/JAN/12	2006/JAN/12	GOOD	NO
martin4	501244	2005/JAN/12	2006/JAN/12	GOOD	NO
DM	501257	2005/JAN/12	2006/JAN/12	GOOD	NO
	501268	2005/JAN/12	2006/JAN/12	GOOD	NO
Richard5	501284	2005/JAN/12	2006/JAN/12	GOOD	YES
RDB3	501292	2005/JAN/12	2006/JAN/12	GOOD	NO
emily1	501296	2005/JAN/12	2006/JAN/12	GOOD	YES
RDB4	501302	2005/JAN/12	2006/JAN/12	GOOD	NO

martin5	501304	2005/JAN/12	2006/JAN/12	GOOD	NO
Richard6	501325	2005/JAN/12	2006/JAN/12	GOOD	NO
emily2	501340	2005/JAN/12	2006/JAN/12	GOOD	YES
RDB5	501342	2005/JAN/12	2006/JAN/12	GOOD	NO
garage	501348	2005/JAN/12	2006/JAN/12	GOOD	NO
Richard7	501356	2005/JAN/12	2006/JAN/12	GOOD	NO
Coldones	501365	2005/JAN/12	2006/JAN/12	GOOD	YES
martin6	501375	2005/MAY/17	2006/JAN/12	GOOD	NO
RDB6	501387	2005/JAN/12	2006/JAN/12	GOOD	NO
Richard8	501398	2005/JAN/12	2006/JAN/12	GOOD	YES
emily3	501402	2005/JAN/12	2006/JAN/12	GOOD	NO
Waldo	501407	2005/JAN/12	2006/JAN/12	GOOD	NO
RDB7	501417	2005/JAN/12	2006/JAN/12	GOOD	YES
DM03	501418	2005/JAN/12	2006/JAN/12	GOOD	YES
Richard9	501431	2005/JAN/12	2006/JAN/12	GOOD	YES
RDB8	501452	2005/JAN/12	2006/JAN/12	GOOD	NO
emily4	501464	2005/JAN/12	2006/JAN/12	GOOD	YES
Richard10	501468	2005/JAN/12	2006/JAN/12	GOOD	YES
RDB9	501487	2005/JAN/12	2006/JAN/12	GOOD	NO
Perez	501489	2005/JAN/12	2006/JAN/12	GOOD	NO
Makena	501501	2005/JAN/12	2006/JAN/12	GOOD	YES
RDB10	501517	2005/JAN/12	2006/JAN/12	GOOD	NO
Olowalu	501532	2005/JAN/12	2006/JAN/12	GOOD	NO
LBlock	501541	2005/JAN/12	2006/JAN/12	GOOD	NO
Giselle	501552	2005/JAN/12	2006/JAN/12	GOOD	NO
emily5	501557	2005/JAN/12	2006/JAN/12	GOOD	NO
Kanaha	501558	2005/JAN/12	2006/JAN/12	GOOD	NO
RDB11	501570	2005/JAN/12	2006/JAN/12	GOOD	NO
Hookipa	501578	2005/JAN/12	2006/JAN/12	GOOD	NO
piclkeweasel	501579	2005/JAN/12	2006/JAN/12	GOOD	NO
carmen maria1	501588	2005/JAN/12	2006/JAN/12	GOOD	NO
Kapalua	501602	2005/JAN/12	2006/JAN/12	GOOD	YES
RDB12	501605	2005/JAN/12	2006/JAN/12	GOOD	NO
DM04	501607	2005/JAN/12	2006/JAN/12	GOOD	NO
Tatiana	501616	2005/JAN/12	2006/JAN/12	GOOD	NO
carmen maria2	501627	2005/JAN/12	2006/JAN/12	GOOD	NO
RDB14	501640	2005/JAN/12	2006/JAN/12	GOOD	NO
Bohemian	501643	2005/JAN/12	2006/JAN/12	GOOD	YES
Helen Munro	501655	2005/JAN/12	2006/JAN/12	GOOD	NO
carmen maria3	501663	2005/JAN/12	2006/JAN/12	GOOD	NO
Farkle	501664	2005/JAN/12	2006/JAN/12	GOOD	NO
RDB15	501667	2005/JAN/12	2006/JAN/12	GOOD	NO
Newfoundland1	501671	2005/JAN/12	2006/JAN/12	GOOD	YES
DM05	501680	2005/JAN/12	2006/JAN/12	GOOD	NO

West Georgia	501684	2005/JAN/12	2006/JAN/12	GOOD	NO
Kamaole	501687	2005/JAN/12	2006/JAN/12	GOOD	YES
Spiff	501688	2005/JAN/12	2006/JAN/12	GOOD	NO
Wailea	501704	2005/JAN/12	2006/JAN/12	GOOD	NO
Carmen maria5	501709	2005/JAN/12	2006/JAN/12	GOOD	YES
Stickman	501713	2005/JAN/12	2006/JAN/12	GOOD	YES
DM06	501720	2005/JAN/12	2006/JAN/12	GOOD	NO
Doughboy	501725	2005/JAN/12	2006/JAN/12	GOOD	NO
shanti1	501727	2005/JAN/12	2006/JAN/12	GOOD	YES
Honolua	501729	2005/JAN/12	2006/JAN/12	GOOD	NO
Spencer	501741	2005/JAN/12	2006/JAN/12	GOOD	NO
Napili	501742	2005/JAN/12	2006/JAN/12	GOOD	YES
shanti2	501743	2005/JAN/12	2006/JAN/12	GOOD	NO
Kaanapali	501754	2005/JAN/12	2006/JAN/12	GOOD	NO
Shanti3	501756	2005/JAN/12	2006/JAN/12	GOOD	NO
Jan	501766	2005/JAN/12	2006/JAN/12	GOOD	NO
shanti4	501781	2005/JAN/12	2006/JAN/12	GOOD	YES
Fleming	501794	2005/JAN/12	2006/JAN/12	GOOD	NO
Salmon skin	501796	2005/JAN/12	2006/JAN/12	GOOD	NO
DM08	501802	2005/JAN/12	2006/JAN/12	GOOD	NO
BigBeach	501811	2005/JAN/12	2006/JAN/12	GOOD	YES
metta1	501821	2005/JAN/12	2006/JAN/12	GOOD	NO
Cabernet	501823	2005/JAN/12	2006/JAN/12	GOOD	YES
Slaughterhouse	501836	2005/JAN/12	2006/JAN/12	GOOD	YES
Malbec	501837	2005/JAN/12	2006/JAN/12	GOOD	YES
Sangiovese	501856	2005/JAN/12	2006/JAN/12	GOOD	YES
metta2	501861	2005/JAN/12	2006/JAN/12	GOOD	YES
BabyBeach	501864	2005/JAN/12	2006/JAN/12	GOOD	YES
Shiraz	501877	2005/JAN/12	2006/JAN/12	GOOD	YES
metta3	501879	2005/JAN/12	2006/JAN/12	GOOD	NO
Prindle	501884	2005/JAN/12	2006/JAN/12	GOOD	YES
metta4	501892	2005/JAN/12	2006/JAN/12	GOOD	NO
Dr. Zhivago	501893	2005/JAN/12	2006/JAN/12	GOOD	NO
Warpigs	501894	2005/JAN/12	2006/JAN/12	GOOD	YES
Hana	501904	2005/JAN/12	2006/JAN/12	GOOD	YES
Dr. Strangelove	501909	2005/JAN/12	2006/JAN/12	GOOD	YES
Polywog	501920	2005/JAN/12	2006/JAN/12	GOOD	YES
ms1	501926	2005/JAN/12	2006/JAN/12	GOOD	YES
SweetLeaf	501934	2005/JAN/12	2006/JAN/12	GOOD	YES
Barbapapa	501943	2005/JAN/12	2006/JAN/12	GOOD	YES
The Lorax	501947	2005/JAN/12	2006/JAN/12	GOOD	NO
Chandra2	501949	2005/JAN/12	2006/JAN/12	GOOD	NO
Barbamama	501952	2005/JAN/12	2006/JAN/12	GOOD	NO
Olivia	501962	2005/JAN/12	2006/JAN/12	GOOD	YES

Snagglepuss	501969	2005/JAN/12	2006/JAN/12	GOOD	NO
Elizabee	501975	2005/JAN/12	2006/JAN/12	GOOD	YES
Crux	501978	2005/JAN/12	2006/JAN/12	GOOD	NO
DM12	501980	2005/JAN/12	2006/JAN/12	GOOD	NO
Chandra3	501983	2005/JAN/12	2006/JAN/12	GOOD	NO
Quandry	501986	2005/JAN/12	2006/JAN/12	GOOD	NO
Chloe	501989	2005/JAN/12	2006/JAN/12	GOOD	YES
Kerfuffle	501994	2005/JAN/12	2006/JAN/12	GOOD	NO
Chandra4	502003	2005/JAN/12	2006/JAN/12	GOOD	NO
IronMan	502006	2005/JAN/12	2006/JAN/12	GOOD	YES
DM14	502012	2005/JAN/12	2006/JAN/12	GOOD	YES
Namaste1	502021	2005/JAN/12	2006/JAN/12	GOOD	NO
Bananaboat	502025	2005/JAN/12	2006/JAN/12	GOOD	NO
Gambo	502032	2005/JAN/12	2006/JAN/12	GOOD	NO
DM15	502035	2005/JAN/12	2006/JAN/12	GOOD	NO
namaste2	502039	2005/JAN/12	2006/JAN/12	GOOD	NO
Endurance	502047	2005/JAN/12	2006/JAN/12	GOOD	NO
DM16	502051	2005/JAN/12	2006/JAN/12	GOOD	NO
Namaste3	502061	2005/JAN/12	2006/JAN/12	GOOD	NO
DM17	502062	2005/JAN/12	2006/JAN/12	GOOD	NO
DM18	502075	2005/JAN/12	2006/JAN/12	GOOD	NO
Dan Macniel	502077	2005/JAN/12	2006/JAN/12	GOOD	NO
Namaste4	502079	2005/JAN/12	2006/JAN/12	GOOD	NO
Ruggedly Handso	502085	2005/JAN/12	2006/JAN/12	GOOD	NO
Namaste5	502087	2005/JAN/12	2006/JAN/12	GOOD	NO
DM19	502088	2005/JAN/12	2006/JAN/12	GOOD	YES
Guacamole	502100	2005/JAN/12	2006/JAN/12	GOOD	YES
Prana1	502111	2005/JAN/12	2006/JAN/12	GOOD	NO
Paranoid	502113	2005/JAN/12	2006/JAN/12	GOOD	YES
Beachfront	502114	2005/JAN/12	2006/JAN/12	GOOD	NO
Prana2	502119	2005/JAN/12	2006/JAN/12	GOOD	NO
Orchid	502120	2005/JAN/12	2006/JAN/12	GOOD	NO
Prana3	502130	2005/JAN/12	2006/JAN/12	GOOD	NO
DM20	502138	2005/JAN/12	2006/JAN/12	GOOD	NO
Snowblind	502141	2005/JAN/12	2006/JAN/12	GOOD	YES
Prana4	502143	2005/JAN/12	2006/JAN/12	GOOD	NO
Tumbler	502144	2005/JAN/12	2006/JAN/12	GOOD	NO
DM21	502150	2005/JAN/12	2006/JAN/12	GOOD	NO
Ahimsa1	502156	2005/JAN/12	2006/JAN/12	GOOD	NO
Embryo	502160	2005/JAN/12	2006/JAN/12	GOOD	NO
Hobbs	502161	2005/JAN/12	2006/JAN/12	GOOD	YES
Ahimsa2	502165	2005/JAN/12	2006/JAN/12	GOOD	NO
Squeak	502175	2005/JAN/12	2006/JAN/12	GOOD	NO
Solitude	502178	2005/JAN/12	2006/JAN/12	GOOD	YES

Squawk	502181	2005/JAN/12	2006/JAN/12	GOOD	NO
DM22	502185	2005/JAN/12	2006/JAN/12	GOOD	NO
ElectricFuneral	502187	2005/JAN/12	2006/JAN/12	GOOD	YES
Ahimsa3	502192	2005/JAN/12	2006/JAN/12	GOOD	NO
Squick	502195	2005/JAN/12	2006/JAN/12	GOOD	NO
Qwerty	502206	2005/JAN/12	2006/JAN/12	GOOD	NO
DM23	502208	2005/JAN/12	2006/JAN/12	GOOD	NO
Supernaut	502210	2005/JAN/12	2006/JAN/12	GOOD	YES
DM24	502224	2005/JAN/12	2006/JAN/12	GOOD	NO
MS-3	502228	2005/JAN/12	2006/JAN/12	GOOD	NO
DM25	502234	2005/JAN/12	2006/JAN/12	GOOD	NO
MS-4	502235	2005/JAN/12	2006/JAN/12	GOOD	NO
Cornucopia	502240	2005/JAN/12	2006/JAN/12	GOOD	YES
Jimmy Buffet	502245	2005/JAN/12	2006/JAN/12	GOOD	NO
SabraCadrabra	502251	2005/JAN/12	2006/JAN/12	GOOD	YES
MARGARITAVILLE	502255	2005/JAN/12	2006/JAN/12	GOOD	NO
Screetchville	502259	2005/JAN/12	2006/JAN/12	GOOD	NO
SmellsLike	502267	2005/JAN/12	2006/JAN/12	GOOD	YES
LASTONE	502275	2005/JAN/12	2006/JAN/12	GOOD	NO
TeenSpirit	502276	2005/JAN/12	2006/JAN/12	GOOD	NO
K15-4ee	502362	2005/JAN/13	2008/JAN/12	GOOD	NO
	502779	2005/JAN/13	2010/DEC/31	GOOD	YES
	502792	2005/JAN/13	2008/DEC/31	GOOD	NO
	502801	2005/JAN/13	2010/DEC/31	GOOD	YES
	502803	2005/JAN/13	2010/DEC/31	GOOD	YES
	502812	2005/JAN/13	2008/DEC/31	GOOD	YES
	502815	2005/JAN/13	2010/DEC/31	GOOD	YES
	502834	2005/JAN/13	2008/DEC/31	GOOD	YES
	502840	2005/JAN/13	2008/DEC/31	GOOD	NO
	502841	2005/JAN/13	2008/DEC/31	GOOD	YES
	502848	2005/JAN/13	2008/DEC/31	GOOD	YES
	502926	2005/JAN/13	2008/DEC/31	GOOD	NO
	502964	2005/JAN/13	2008/DEC/31	GOOD	NO
	502966	2005/JAN/13	2008/DEC/31	GOOD	YES
	502967	2005/JAN/13	2008/DEC/31	GOOD	YES
	502968	2005/JAN/13	2008/DEC/31	GOOD	YES
	502971	2005/JAN/13	2008/DEC/31	GOOD	YES
	502972	2005/JAN/13	2008/DEC/31	GOOD	NO
	502973	2005/JAN/13	2008/DEC/31	GOOD	NO
	502975	2005/JAN/13	2008/DEC/31	GOOD	NO
	502984	2005/JAN/13	2008/DEC/31	GOOD	NO
	502988	2005/JAN/13	2008/DEC/31	GOOD	NO
	502989	2005/JAN/13	2008/DEC/31	GOOD	NO

	502991	2005/JAN/13	2008/DEC/31	GOOD	NO
	502992	2005/JAN/13	2008/DEC/31	GOOD	NO
	502997	2005/JAN/13	2008/DEC/31	GOOD	YES
	504171	2005/JAN/13	2008/DEC/31	GOOD	YES
LINK	512800	2005/JAN/12	2006/MAY/17	GOOD	NO

- Expiry date of claims is subject to government approval of assessment work covered by this report.

4.0 EXPLORATION HISTORY

The claim blocks comprising the Kizmet project area have undergone discontinuous exploration between 1969 and 2004. Historic exploration has focused primarily upon gold, silver and base metal veins in volcanic rocks and associated intrusive stocks and plugs where metal mineralization was known to occur locally. This work has been summarized in Table 2.

20 historic ARIS reports exist within the Kizmet project claim group. Previous work includes government and industry sponsored regional stream sediment geochemical campaigns, geological mapping, prospecting, grid establishment, soil sampling and ground based geophysical surveys.

In 1969 Geophoto Services conducted a geochemical sampling program in addition to an EM survey on the historic LC claims. Sample ER-238 consists of a massive sulphide vein within a hornblende monzonite intrusive body. This sample returned 9.6% Cu, 10.67% Pb, and 11.2% Zn. Silt samples collected in the region did not return values above background. A ground-based EM survey was not helpful in tracing this vein under cover (Estabrooks, 1969).

A geologic mapping program was carried on the BS, Mad Nut and J historic claim groups between August 1st and August 29th, 1970 by the Taku Syndicate. Copper and molybdenite mineralization was found associated with northeast trending faults and joints systems in a quartz-monzonite host. Although the mineralization was widespread based upon mapping and ground magnetic response, geochemical sampling failed to indicate an economic concentration (White, 1970).

Anglo Canadian Mining collected rock samples on the historic Jak and Joly claims after finding base metal veins in float during a regional exploration program in the area. Anglo Canadian geologists found Ag and Au mineralization in Stuhini felsites, King Salmon Formation meta-sediments and Sloko Group volcanics. The best assay was found at the Zohini Vein and consists of 0.057 Oz/ton Au, 189.5 Oz/ton Ag 9.5% Zn, and 4.8% Pb over 5 feet (Payne, 1980).

Comaplex Resources explored the GO claims in 1981. Comaplex drilled 9 holes totalling 972.63 meters and dug 6 trenches in order to investigate a series of mineralized fracture veins within Stuhini Group felsite and porphyry intrusives. The best drill hole consisted of 2.13 meters @ 0.027 Oz/ton Au, 2.030 Oz/ton Ag. Drilling results outlined a strike length of 205 meters of mineralization and a down dip extent of 150 meters. The grade and thickness was noted as sporadic but no ore grade intersections were established (Lintott,1981).

In 1980 Mefeking Minerals performed and implemented a geochemical and geophysical survey that was subsequently followed up with 3 diamond drill holes. Three holes were drilled to determine the depth extent of vein style surface mineralization found at the contact with a 'basic' dyke and monzonite country rock. After drilling 3 holes Mefeking Minerals concluded that mineralization does not widen along strike nor does it extend to depth (Estabrooks and Dalidowicz 1989).

In 1982 Chevron Standard spent 4 man-days sampling the EMU claims. Anomalous gold was found in quartz-chalcedony, pyrite, galena, sphalerite veins found within fault zones. Chevron abandoned exploration on the EMU claims because no anomalous gold values were obtained in the Takwahoni sediment or intrusive rocks on the claims but noted that a sample extracted from narrow quartz vein contained 17 ppm Ag and 675 ppb Au.

Chevron Standard also explored their ROD claims in 1982. Walton, 1983 indicated that geochemical results were encouraging and recommended follow up work including detailed geologic mapping and detailed soil or talus fine sampling (Shannon and Thicke, 1982).

Noranda carried out a one-day reconnaissance of the same ground in 1986 (Reid, 1987). They reported bleached and silicified zones flanking felsic dykes with maximum values of 70 ppb Au, 13.2 ppm Ag, 1.3% Pb and 6200 ppm As (Ried, 1987).

Georgia Resources Inc conducted soil geochemistry on the Lis-2 claims which in order to investigate upper mudstones and siltstones of the upper Triassic King Salmon formation which are cut by diorite-monzonite intrusives. The soil geochemistry proved successful and outlined highly anomalous values of gold, silver, copper, lead, zinc, arsenic and antimony. Detailed infill soil geochemistry was recommended (Lambert, 1988).

In 1988 Tahltan Holdings explored the Ant Bing and Sam claims for gold and silver. Tahltan Holdings was exploring Gold and silver mineralization within quartz +/- chalcedony +/- calcite veining and stockwork zones, quartz carbonate haloes and gossans. No more detail could be found on this work (Freeze et al, 1989).

In 1989 the Golden Met property, owned by Interex Development Corp. and United Cambridge Mine Ltd. was prospected in order to follow up soil geochemical anomalies and rock grab samples collected while the property was staked in addition to an anomaly defined by Chevron in 1982. After a brief prospecting campaign no cause for the Chevron anomaly could be found (Estabrooks and Dalidowicz 1989).

In 1989 Tech Exploration Ltd. filed assessment for work done on the historic Bush and Whatnot claims totalling 36 units. Mineralization in this area was restricted to quartz-carbonate veins and consisted of disseminated pyrite, malachite staining and minor chalcopyrite. Geochemical sampling in these areas concentrated on quartz veins in float and produced anomalies which Tech recommended follow up work to trace them to outcrop (Schellenberg, 1989).

In 1989 Cominco Ltd. explored the historic Bryar claims where work was concentrated on contour soil sampling in the vicinity of a gossanous outcrop suspected to be the source of mineralized float. Soil sampling delineated an Au and As anomaly. Rock sampling on the north portion of the claims yielded anomalous gold values from arsenopyrite bearing veins in quartz-biotite-feldspar porphyry. Cominco Ltd. recommended additional prospecting (Smith, 1989)

In 1989 Equity Silver Ltd. explored the Plum claims that covered some of the same ground explored by Chevron in 1983. Unfortunately very few of the soil and rock samples collected in this area proved to be anomalous. The few anomalies produced from this program were described as being weak anomalies. Dynes, (1989) concluded the Chevron soil anomaly outlined in 1983 does not extend to the east. Dynes, (1989) recommended that more work should be done to follow up the western anomaly (Dynes, 1989).

In 1989 Tahltan Holdings Ltd. explored the Law claims where geologic mapping and geochemical surveys (bulk heavy mineral) were carried out to determine previously undetected mineralization on this property. Dynes, (1989) concluded that geological and geochemical results indicate the presence of an epithermal system anomalous in gold and recommended follow up exploration (Dynes, 1989).

Cominco Ltd. explored the LJ claims in 1990. Explored previously reported gold values taken from quartz rich veins in shears consisting of 22.4 ppm Au over 1 meter, 40 ppm Au over 0.20 meters and 10.4 ppm Au over 0.25 meters. Cominco outlined two zones of mineralization 750 meters apart. Zone 1 was characterized by fault associated pyrite-arsenopyrite veins with proximal quartz carbonate vein breccias and Zone 2 consisted of quartz sericite veins hosted in shears. Average grades from Zone 1 consisted of 139.31 ppb Au, 0.78 ppm Ag and 6296.25 ppm As. Average grades from Zone 2 are reported as 901.2 ppb Au, 1.46 ppm Ag and 3130 ppm As. Cominco recommended no further follow up work on this property due to low Au an Ag values (Aspinall and Strain, 1991).

Solomon staked a the King claims in 1990 corresponding to the western half of the current Sutl 5-12 property and extending westward. They concentrated on two gossans [NOTE: Gossan B is on Sutl 1-4 claims on the current area covered by the Thorn property to the southeast]: Zone C in the same cirque on Sutl 5 and 7 where Taku and Noranda did their work; and Zone A on Sutl 9 and 11. In Zone A, Solomon recognized a highly silicified and pyritic hornblende granodiorite, with traces of molybdenite, intruding sediments and andesitic pyroclastics rocks, themselves silicified near the contacts. Several styles of mineralization were recognized in Zone A, including a 1m fault zone (max. 100 ppb Au, 416 ppm Ag, 7.1% Cu) and skarn (max. 210 ppb Au, 103 ppm Ag, 0.87% Cu, 1.52% Pb and 4.36% Zn). A silt sample with 1220 ppb Au was taken from the projected strike of a lineament hosting a 5m chalcedonic quartz-carbonate breccia. One-sixth of the soil samples from Zone A exceeded 50 ppb Au, with maximum values of 490 ppb Au, 130.6 ppm Ag, 20632 ppm Cu, 3556 ppm Pb and 5691 ppm Zn. In Zone C, Solomon mapped hornfelsed pelitic rocks intruded by locally sericitized hornblende granodiorite, quartz-feldspar porphyry, quartz porphyry and monzonite. The more leucocratic varieties contain disseminated molybdenite. The most spectacular gossans are in creek gullies, exposed for up to 100m; exposures are fractured, sheared, silicified and locally argillized. Solomon described three 15cm quartz-carbonate veins from Zone C, the best of which returned 4121 ppm Cu, 1918 ppm Pb, 7627 ppm Zn, 167 ppm As, 9 ppb Au and 38 ppm Ag (Aspinall, 1991).

In 1991 Teck collected 191 soil samples on the Bush and Whatnot claims. Results included several moderately anomalous zones which were thought to be indicative of an auriferous source to the north which would likely be hosted in veinlets or shear zones within. It is not clear from the assessment report which lithology the rock samples were collected from (Betmanis, 1991).

In 1991, Omega carried out limited mapping and geochemical sampling on their claim group immediately east of Solomon's King property. Results were generally low, with maximum soil values of 70 ppb Au, although a base metal anomaly (max. 1,215 ppm Zn, 3,253 ppm Pb, 535 ppm Cu and >2,000 ppm As) was reported within a gossan straddling the contact zone between granodiorite, quartz-feldspar porphyry and volcanic rocks (Chapman, 1991).

Golden Rule Resources staked the Thorn 1-5 claims in May of 1991 based upon results of a government regional geochemical survey indicating the presence of Au, Ag, Pb, and Zn. Gossanous material collected from the Thorn claims was not considered economically mineralized. Evans, 1991 suggested that special attention should be paid to the Thorn 2, 3 and 4 claims where they are underlain by Stuhini Volcanics and proximal to two RGS stream sediment anomalies grading 326 ppb Au and 730 ppb Au respectively. Evans, 1991 recommended follow up sampling for all of the Thorn Claims (Evans, 1991).

In 1992 Consolidated Parkland Resources and Slocan Development Corp. Ltd. filed assessment on the Sutlahine property (Park 10 – 13 Claims). Work consisted of follow up work to low grade copper mineralization discovered by the Taku Syndicate in 1970 (White, 1970). Work consisted of regional geological mapping, prospecting, rock sampling, soil and slit sampling. Copper mineralization was discovered in several locations on the Park 10-13 claims. Mineralization has been described by Crowe and Haynes, 1992 as veinlets and fracture filling in weakly altered quartz monzonite. Silver values as well as elevated As and Pb were often associated with Cu mineralization. Soil anomalies were delineated, some of which were associated with hornfelsed sediments. Follow up work was recommended for the 1992 soil anomalies in the form of detailed soil sampling, prospecting and mapping (Crowe and Haynes, 1992).

In 1992 International Suneva Resources Ltd. explored the Borg property (Green, Orange, Black and Red claims) for base and precious metal mineralization. The Borg property is underlain by two age-distinct monzonite intrusions, mafic and felsic dykes. Several highly anomalous rock samples (both base and precious metals) were found on the property and were associated with four types of mineralization. The four types of mineralization include silicified and locally pyritized Cretaceous-Tertiary monzonite; northwest-north–north-east trending, narrow quartz veins developed predominantly within older diorite – quartz monzonite proximal to the younger monzonite; quartz veins developed along-side mafic dykes; and quartz veins, quartz stockwork and quartz breccias developed marginal to felsic dykes. Values up to 0.680 Oz/t Au, 191.8 ppm Ag and 5,040 ppm Cu, 17,340 ppm Pb and 26,215 ppm Zn have been returned from mineralized grab samples (Crowe,1992).

In 1992 Georgia Resource Inc. returned to do follow up work on the Lis 1-4 claims. This property is located within the upper Triassic Stuhini sediments and volcanic rocks in the Tulsequah area. The Lis1-4 claims are bisected by the King Salmon thrust fault and associated sulphide mineralization consisting of arsenopyrite, chalcopyrite, pyrite, galena, sphalerite, and pyrrhotite. Rock, silt and soil geochemical samples were collected. Anomalous values in gold, silver, copper and zinc showed that sulphide mineralization was widespread. Grab samples (base metal rich fracture veins) on the Liz1-4 claims returned 6.98 gpt

Au, 351.3 gpt Ag, 2.5% Cu, 69.95% Pb and 1.2% Zn. Follow up work was recommended to bring this property to the drill stage (Terry, 1992).

In July 1998 10 days of work were done on the Check-Mate 2 mineral claim, formerly Chevron's Inlaw claims. The goal of this exploration was to evaluate the low grade bulk tonnage gold potential of the claim, confirm Chevron's 1984 soil geochemical anomaly and to propose a model for gold mineralization. The highest gold from a rock sample returned 704 ppb Au. Reconnaissance mapping indicated that Chevron's Fe-carbonate alteration zone was hosted by Late Triassic Stuhini breccias, agglomerates and extended 3800 m in a NW trend. The model proposed for the Check-mate 2 claims was an early Tertiary epithermal structurally controlled system near an ancient paleosurface marked by two unconformities associated with Fe-carbonate alteration (Aspinall, 1998).

In 1998 XPlorer Gold Corp. conducted magnetic and EM surveys ion the Kap 9-10 and King 2-5 claims in the Tulsequah Area, northern British Columbia. XPlorer Gold Corp. was exploring for Besshi-type VMS deposits. Although two HLEM anomalies were reported they were not deemed significant and no follow up work was recommended (Lee Carmen, 1998).

TABLE 2: Kizmet Project Exploration History

Owner/Historic Claim	CLAIMS	Geochemistry	Geophysical	Drilling Trenching	Author	REPORT No.	Expenditures
***Geophoto Services, 1969**	LC	34 Stream/Soil Samples	EM Survey	30 by 10 foot area trenched	Estabrooks EM	2060	Unknown
Taku Synd, 1970	MAD NUT		40 Line Miles Mag		White LG	2537	$8,501.50
***Taku Synd, 1970**	BS-J	43 Rock chip samples			White LG	2648	$10,245.50
***Taku Synd, 1970**	BS-J		Magnetometer Survey		White LG	2649	$4,254.00
Anglo Canadian Mining Corporation, 1980	JAK JOLY	~ 19 Assays			Payne John G	9048	$4,254.00

Comaplex Resources International, 1981	GO			6 Trenchs 9 DDH	Lintott K	9495	$208,583.73
***Mafeking Minerals 1980 Partnership, 1980**	JONY		Ground EM	3 Diamond Drill holes (J-1 to J-3) totalling 139 feet	Estabrooks EM Dalidowicz F	9578	$140,559.00
***Chevron Standard, 1982**	EMU	12 Rock samples			Shannon K Thicke M	11108	$2,148.80
***Chevron Standard, 1982**	ROD	115 Soil samples, 28 Rock samples			Walton Godfrey	11819	$5,641.45
Noranda, 1987	Sut	14 silts, 12 talus fines, 22 rocks, 4 panned concentrates			Reid (1987)	15477	
Georgia Resources Inc. 1988	Lis 2	61 Soil Samples			Lambert Ellen	17517	$5,627.62
Tahltan Holdings, 1988	Vardis 1-4				Freeze JC Dynes WJ Wetherill JW	17908	
Interex Development	Golden Met Golden Met 2	5 Rock samples 25 Soil samples			Thompson W	18926	$8,789.31
Teck Ex, 1989	Bush Whatnot	24 Rock Samples 42 Soil Samples			Schellenberg G	19259	$7,987.35
Cominco Ltd, 1989	Bryar	11 Rock Samples 50 Soil Samples 10 Silt Samples			Smith Scott William	19326	$4,522.50
Equity Silver Ltd., 1989	Plum 1-4	6 Talus Heavy Mineral Samples 25 Rock Samples 6 Silt Samples			Dynes WJ	19375	$7,926.46
Tahltan Holdings,1989	Law 1-2	20 Rock Samples 3 Bulk Heavy Mineral Concentrate Samples 2 Silt Samples			Dynes WJ Wetherill JW	19377	$6,648.43
Cominco Ltd, 1990	LJ	167 Soil Samples 16 Rock Samples			Aspinall C Strain DM	20433	$29,895.97
Solomon Resources Ltd, 1991	Wahb 3-4		VLF 7 Survey Lines		Terry M	21435	$4,466.20

Solomon Resources Ltd, 1991	King 10-14 King 2-4	69 Rock Samples 313 Soil Samples 13 Silt Samples			Aspinall C	21530	$87,625.18
Teck Ex, 1991	Bush Whatnot	5 Rock Samples 194 Soil Samples			Betmanis AI	21718	$12,239.65
Omega Gold Corp, 1991	D 1-8	Unknown - Prospecting			Chapman J	21907	$38,283.95
Omega Gold Corp, 1991	A 1-8	79 Rock Samples 47 Soil Samples 7 Silt Samples			Chapman J	21908	$80,000.00
Golden Rule Resources Ltd, 1991	Thorn 1-5	21 Rock Samples			Evans BT	21968	$22,643.00
Consolidated Parklane Res, 1992	Park 10-13	109 Soil Samples 12 Silt Samples			Haynes LR Crowe GG	22208	$21,281.00
Int Suneva Res, 1992	Green Black Red	47 Rock Samples 85 Soil Samples			Crowe GG	22268	$12,888.42
Aspinall Clive, 1998	Check-Mate 2	51 Rock samples			Aspinall NClive	25669	$12,876.14
Georgia Resources Inc, 1992	Lis 1-4		Mag/VLF 45 Line KM.		Terry M	22384	$51,975.00
XPlorer Gold Corp, 1998	King 3, Red Cap 3		5.05 Km of Ground Mag and HLEM		Lee Carmen C	25458 and 25460	$30,021.52
Total Cost							$833,443.18

5.0 2005 EXPLORATION PROGRAM

The objectives of the program in 2005 were:

1. To discover economically viable high sulphidation (HS) epithermal Au deposits greater than 3 MOz in size within the Kizmet area or downgrade the mineral potential of the belt.
2. To systematically evaluate the potential of predefined targets through a priority based reconnaissance style exploration program.
3. To reduce the large land position to include only areas of significant mineral potential.

Compilation of all relevant geological data, interpretation of compiled results, and target generation resulted in identification of ~45 initial targets within the Kizmet area by June 1, 2005. Field examination was conducted from June 9 to August 28, 2005 and was designed to be carried out in two phases. The first was designed to quickly and efficiently evaluate the potential of the target to host a large high sulphidation deposit. The second was designed to conduct detailed work on any targets advancing past phase 1, with the goal of advancing the target to target delineation or possibly drill testing stage. Approximately 26 of the original targets received significant fieldwork in 2005. No targets showed sufficient potential to advance to Phase 2 detailed fieldwork in 2005.

Fieldwork emphasised rapid target evaluation through prospecting, rock +/- soil and silt geochemical sampling, reconnaissance mapping, and identification of specific features associated with high sulphidation epithermal mineralization. Mapping focussed on verification of the nature and extent of known or newly discovered geologic units and determining their hosting potential for mineralization. This mapping resulted in identification of several areas underlain by previously unmapped Late Cretaceous to Paleogene aged volcanics and their intrusive equivalents. Short Wave Infrared spectrometry (SWIR) was used to determine alteration mineralogy and identify any advanced argillic and argillic alteration zones. All rock samples were analysed using a Portable Infrared Mass Spectrometer (PIMA).

A belt-scale stream sediment survey was also carried out with the intent of providing an improved regional perspective of mineral potential at a 15 km^2 drainage scale.

During the 2005 program, crews camped on the perimeter of Trapper Lake and King Salmon Lake (Figure 2). Fieldwork was completed through helicopter supported daily traverses by a 4 person field crew based out of small camps located at strategic locations in the belt. Transportation to the property was carried out by a Hughes 500 helicopter, which was operated by Northern Air Support.

All samples collected were shipped to ALS Chemex Laboratories in Vancouver. 848 rock samples were collected. Analyses requested consisted of ALS Lab Code Au-ICP21 (gold fire Assay by ICP-AES) and ALS Code ME-ICP41m (34 elements by aqua regia acid digestion and ICP-AES plus Hg (0.01-100 ppm) by cold vapour).

46 stream sediment samples were collected in the 2005 program. Analyses requested consisted of ALS Lab Code ICP-21 (gold fire Assay by ICP-AES) and ME-MS41 (50 elements by aqua regia digestion and combination ICP-AES and ICP-MS analysis)

31 soil samples were collected and analysed according to ALS Lab Code ICP-21 (gold fire Assay by ICP-AES) and ME-MS41 (50 elements by aqua regia digestion and combination ICP-AES and ICP-MS analysis)
Geochemical analyses are listed in Appendix IV and V.

The Kizmet claim group has been divided into 7 geographic areas broadly representing the original claim groups submitted for exploration permits as well as the main target areas for the project. To facilitate map making maps have been printed at scales representing these areas (see Map 2).

Software programs used in support of the exploration of the property and the preparation of the report are: ArcGIS V.9.0, Microsoft Word and Excel 2003, Adobe Acrobat 7.0, and AcQuire V3.8.7.

6.0 REGIONAL GEOLOGY

By far the most up to date and eloquently summarized description of the regional geology of the Kizmet project area was described by Adam Simmons (Simmons, 2005) and is reproduced below. Figure 3 depicts the regional geology and project boundaries.

Cretaceous and Tertiary continental arc related magmatic rocks in western Canada and Alaska were emplaced into Paleozoic and Mesozoic volcanic arc rocks and related sedimentary rocks accreted onto the western North American continental margin no later than early Middle Jurassic in the northern cordillera (Mihalynuk, 1999; Bacon, 1990; Hart, 1995). This region is collectively underlain by the Stikine, Cache Creek and Quesnel tectonostratigraphic terranes; the Stikine Terrain being the host terrane to the Kizmet project area. These terranes are believed to represent a continuous ~1400km long island arc developed outboard of western ancestral North American margin along a northwest trending subduction zone during and prior to the Late Carboniferous (Mihalynuk, 1994). This subduction zone experienced a Late Permian to Late Triassic reconfiguration following the initial stage of collapse of the Slide Mountain basin during east verging contraction (Nelson, 1993). The subduction zone was re-established outboard of Quesnellia and Stikinia by the Late Triassic, and the magmatic products are represented by the Stuhini and Nicola-Takla Groups. During this time these terranes began to experience oroclinal bending in response to the collision of the Cache Creek plateau with Quesnellia and Stikinia.

Extensive arc volcanism and plutonism dominated Stikinia and Quesnellia during the Upper Triassic. These Triassic volcanic rock sequences are generally considered to be very similar, grading from subalkaline volcanic rocks (Stuhini and Nicola-Takla Groups respectively) in the south to volcaniclastic sedimentary rocks (Lewis River Group and Nazcha Formation respectively) in the north

(Gabrielse, 1969). The Upper Triassic arc, near the proposed oroclinal hinge is

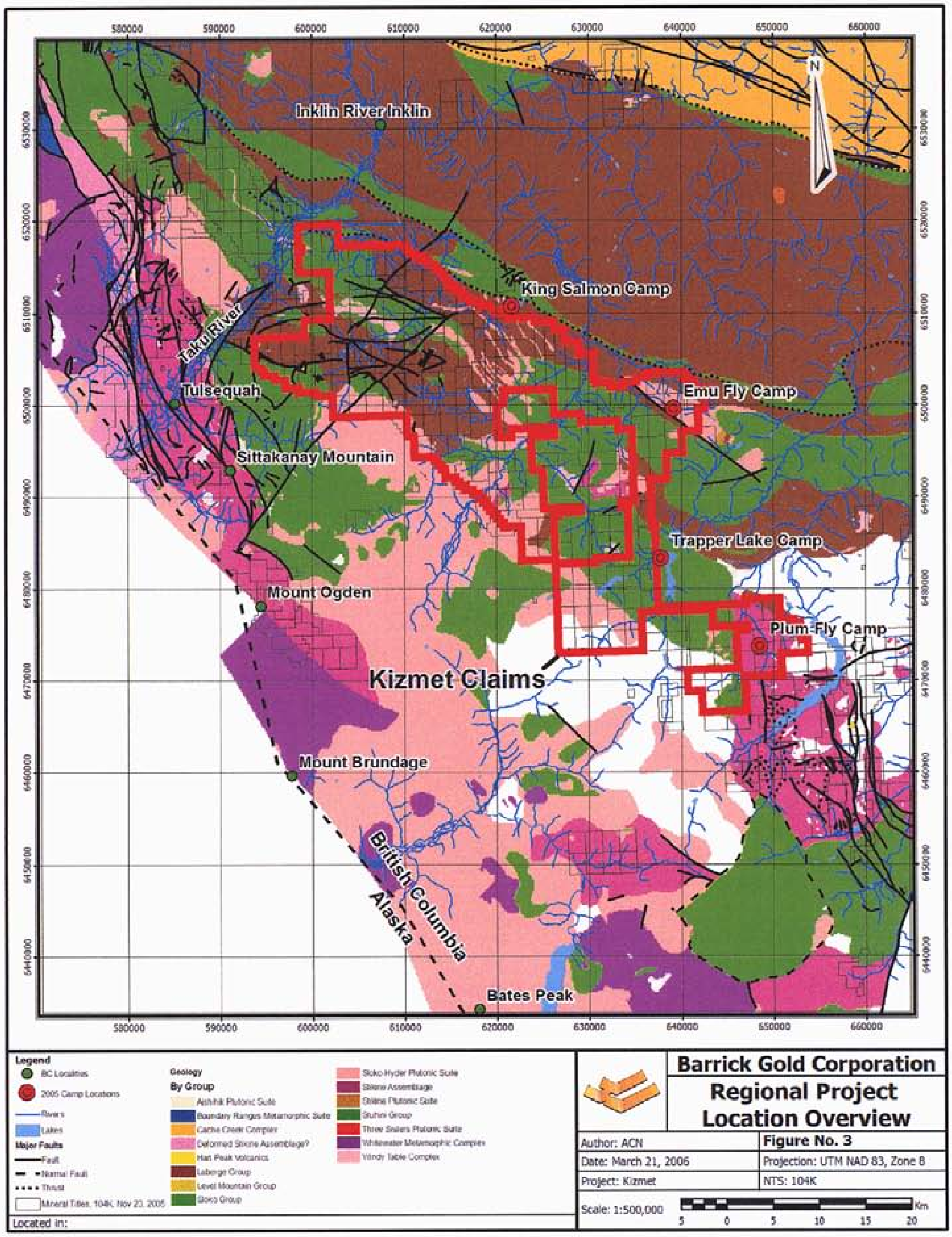

Figure 3: Regional geology and extent of Kizmet project claim blocks.

dominated by arc derived sedimentary rocks whereas volcanic rocks dominate elsewhere along the arc (Gabrielse, 1969). This time period represents the tectonic closure of the Cache Creek Ocean brought on by oroclinal bending of the Stikinia-Quesnellia arc. Monger et al. (1991) and Jackson (1992) provide evidence for the linkage of Stikinia and Cache Creek Terranes at this time while the linkage between Cache Creek and Quesnellia (albeit in southern British Columbia) is provided by Monger (1984).

Lower to Middle Jurassic Laberge Group clastic sedimentary rocks were deposited on Upper Triassic Stuhini Group rocks throughout the region. Mihalynuk (1999) proposed that these sedimentary rocks recorded dissection of the Stuhini arc at the hinge zone of the oroclinal bend. South of the hinge zone, continued subduction under Stikinia and Quesnellia resulted in voluminous calcalkaline Hazelton Group volcanism. Nixon et al. (1993) established that Quesnellia was emplaced along the western margin of North America by 186 Ma. Locally, clastic sedimentary rocks are composed of detritus from older Laberge Group strata, indicating rapid uplift in limited areas. This exhumation and subsequent erosional event may mark the initiation and collapse of the Whitehorse Trough (Mihalynuk, 1999). Southwest verging thrust faults are critical to the development of the arc during this time. The southwest verging movement of the King Salmon Thrust (Thorstad and Gabrielse, 1986) facilitated the emplacement of the Laberge Group and the clastic foredeep of the proto-Bowser Basin, which began in the latest Toarcian to Aalenian time (Ricketts, et al., 1992; Mihalynuk, 1999).

The Stikine, Quesnel and Cache Creek Terranes were discrete tectonic elements separated by subduction or collision zones prior to the Middle Jurassic. Thrusting of Quesnellia over western North America and thrusting of Cache Creek over Stikinia was the result of oroclinal collapse. Shortening in the vicinity is believed to be in excess of 50%. (Mihalynuk, 1999). The Cache Creek and Stikine Terranes are stitched together by 172 Ma (undeformed Fourth of July batholith; Mihalynuk et al., 1992). Metamorphic cooling of the terranes occurred ca. 172 Ma. (40Ar/39Ar plateau age Boundary Range metamorphic suite) (Smith and Mihalynuk, 1992). Laberge Basin sedimentation ceased by the latest Middle Jurassic marking the beginning of a period of tectonic quiescence and a magmatic lull lasted approximately 50 million years.

Magmatism and strike slip deformation dominated the tectonic regime of the Northern Cordilleran by the Cretaceous Period. Magmatism resumed by the latest Early Cretaceous with the onset of the Whitehorse Magmatic Epoch, which lasting from approximately 111 Ma (Hart, 1995) until 100 Ma (Mihalynuk, 1999).

The following 15 million years were relatively quiescent regionally; however Thorn Suite Magmatic rocks were deposited during approximately 15 million years within a regionally quiescent period. Windy Table Magmatic Suites were emplaced immediately after the emplacement of Thorn Suite Magmatism. The onset of Windy Table volcanism is variable along the northeast trending belt, beginning as early as 85.5 Ma at the Thorn property (Simmons et al., 2005). Volcanic rocks of this age are generally flat-lying, but may be tilted adjacent to steeply dipping normal faults (Simmons et al., 2005).

Following yet another magmatic hiatus, magmatism resumed from 58.5 - 53Ma (Mihalynuk, 1999) as part of the Sloko Magmatic Epoch consisting of intermediate to felsic magmatic rocks represented by voluminous eruptions of volcanic rocks with coeval semi-circular granitic plutons, the roots of Sloko–age volcanoes in the Coast Belt (Mihalynuk, 1999).

7.0 Property Geology

7.1 Lithology

The Kizmet project claims cover a region containing three distinct periods of magmatism (Thorn Suite, Windy Table Suite, Sloko Group) formed during Upper Cretaceous and Paleogene periods. These magmatic events were superimposed upon deformed Upper Triassic volcanics and associated sedimentary rocks of the Stuhini Group sediments of the Lower to Middle Jurassic Laberge Group. Jurassic and older rocks are believed to have predated the targeted style of mineralization in the belt and will be referred to as basement rocks below.

The oldest rocks mapped in the study area lie immediately to the northwest of Tatsamenie Lake and have been variably mapped as Upper Triassic or older fine grained sediments and associated mafic volcanic rocks (Souther, 1971), or Mississippian aged undivided mafic to intermediate volcanic rocks of the Stikine Assemblage (BCGS 1;250,000 compilation, 2004), or as a mixture of both Stikine Assemblage sediments and Upper to Middle Triassic mafic volcanic rocks assigned to the Stuhini Group (Oliver, 1993). Observations from 2005 fieldwork showed that commonly carbonate altered rocks in this area displayed similar features to Stuhini Group rocks elsewhere. Such features included carbonate (ankerite) alteration, a weakly developed foliation and minor pillow basalts. Stikine Assemblage rocks observed elsewhere typically contain stronger and multiphase foliations. If these rocks are of Paleozoic age, they represent a window into this stratigraphic package, uncommon elsewhere on the property.

These rocks are unconformably overlain by Stuhini Group basaltic to andesitic flows, pillow flows, flow breccia and minor volcanic sandstone, siltstone and greywacke. The Stuhini Group rocks are most commonly exposed in northwest trending anticlines and for the most part appear to contain very little hydrothermal alteration. Historic exploration has likely thoroughly tested this package for VHMS mineralization.

Rocks of the Stuhini Group are unconformably overlain by limestone and lesser sandstone and argillite of the Sinwa Formation. Exposures of this unit are rare on the property, mainly occurring immediately east of the Thorn property. Most commonly it occurs along the base of the King Salmon Thrust Fault with lies to the northeast of the project boundary.

These basement rocks are unconformably overlain by rocks of the Laberge Group, which in the study area was observed to be composed of mudstone to siltstone, well bedded greywacke, and extensive pebble to cobble conglomerates. The top of this package represents and erosional unconformity on which Cretaceous and younger volcanic rocks were deposited

These basement rocks were deformed during the amalgamation of Stikinia, Cache Creek and Quesnellia and their subsequent accretion onto the western flank of North America. Deformation during this period is characterized by regional sub-greenschist metamorphism, northwest trending upright, open to close folds and northwest trending thrust faults (Simmons, 2004). The basement rocks are intruded and unconformably overlain by a series of late Cretaceous to Eocene volcano-plutonic centres.

The earliest of these Cretaceous and younger magmatic events is represented by 93-88 Ma porphyritic granodiorite dykes and stocks, most noticeably the Thorn Stock located on the Thorn property. This stock is host to extensive intermediate to high sulphidation epithermal mineralization. Rocks of this age more commonly exist as dykes and sills on the Kizmet project claims.

The next magmatic event is represented by several volcano-plutonic complexes identified by Simmons et al., (2004) and have since been classified as belonging to the 86-80 Ma Windy Table Suite. Exposures of this unit are somewhat limited, consisting of erosional volcanic remnants, commonly 1-3 km in diameter and near coeval sills, dykes, and stocks. Compositions range from basaltic to rhyolitic. Features such as welded tuffs and the presence of petrified wood indicate the volcanics are subaerial in nature. Work by Simmons et al., (2004) and Simmons, (2005) has linked intermediate to high sulphidation mineralization at the Thorn property to this suite through reasonably extensive age dating of alteration minerals associated with the mineralization. 2005 fieldwork identified several other areas likely underlain by subaerial volcanic tuffs of this suite and are shown on several maps included with this report.

Simmons, (2005) noted that recent work by Mihalynuk et al., (1999, 2003) and Simmons et al., (2004, 2005a, b) have begun to define a distinctive NNW trend to this Late Cretaceous volcano-plutonic arc, extending from at least the Golden Bear Mine area to the B.C.-Yukon border. Significantly, this belt is coincident with a belt of anomalous Au, Ag, Cu, Pb & Zn in stream sediments sampled by the federal-provincial Regional Geochemical Stream Sediment Survey. The Late Cretaceous portion of the Coast Batholith has limited E-W spatial distribution with an estimated maximum width of 20 km at these latitudes (Simmons, 2005).
The Windy Table Suite volcanic units were targeted in the 2005 exploration campaign as they are believed to be the most prospective host for large scale epithermal mineralization.

Paleogene magmatism belonging to the Sloko Group has been recognized near Lisadale Lake by Simmons, (2004) and is estimated to represent the last magmatic event in the area, dated at ~55-58 Ma. Several targets near Niagara Mtn in the Kizmet property (KZ-2, KZ-3, KZ-4) are likely underlain by rocks of this age but as no age dating was performed in 2005, this age association cannot be verified. Rocks in these areas tend to form mesa like buttress with >300m cliff faces that commonly display relatively flat lying volcanic strata intruded by both sills and dykes of unknown ages.

A lithological classification developed by Simmons (2004) represents the most modern stratigraphic interpretation for the area and provided a good guide to geological mapping in 2005. This interpretation in the form of a legend is summarized in Table 3.

Table 3: Kizmet Project Lithologic Units

TERTIARY
INTRUSIVE ROCKS

Basalt/andesite dykes: fine-grained, dark green to brown, weakly magnetic, aphyric or feldspar-phyric, calcite amygdules common

Sloko Suite Intrusive Rocks

Coarse-grained quartz-feldspar-biotite porphyry: 15–40% anhedral 1–5mm feldspar (plagioclase) , 5-10% euhedral equant 2-4mm glassy quartz and 5–15% euhedral equant 3–6mm biotite phenocrysts, in a fine grained matrix

LATE CRETACEOUS
Windy Table Suite Intrusive Rocks

1) Monzonite and diorite: feldspar porphyritic, biotite and hornblende are both present, quartz forms irregular shaped crystals between feldspar grains
2) Biotite-hornblende granodiorite: fine- to coarse-grained, equigranular, local miarolitic cavities

Windy Table Suite Subaerial Volcanic and Related Sedimentary Rocks

1) Dacitic/andesitic tuff, lapilli tuff and block tuff: Maroon to grey-brown, matrix-supported
2) Unwelded rhyolitic tuff and agglomerate: lithic clasts comprise 10-35% of the rock ranging from lapilli to block size and dominated by granitic intrusive rocks and felsic volcanic rocks
3) Welded lapilli tuff or crystal tuff: dominantly rhyodacitic with 5-15% chlorite replaced pumice fragments and variable amounts of lithic fragments (up to 40%), crystal tuff dominated by 1-4mm euhedral feldspar phenocrysts and 1-3mm rounded quartz eyes
4) Dacite to andesite flow
5) Rhyolite flow
6) Volcaniclastic rocks: finely bedded, rounded silt to sand sized particles and commonly associated with accretionary lapilli horizons

DIORITE TO QUARTZ DIORITE PORPHYRY

1) Coarse-grained feldspar-quartz-biotite porphyry: 15–40% anhedral 1–5mm feldspar, 15–30% euhedral equant 3-6mm glassy quartz and 5–15% euhedral equant 3–6mm biotite phenocrysts
2) Fine-grained feldspar-quartz-biotite porphyry: 30% anhedral 0.5–2mm feldspar, 0–5% subhedral 2–4mm quartz and 5% euhedral equant 4mm biotite phenocrysts

LOWER TO MIDDLE JURASSIC

Laberge Group Clastic Sedimentary Rock

1) Cobble conglomerate: clasts range in size from pebble to boulder, but is generally cobble sized, commonly matrix supported, and clasts types are dominated by either mafic volcanic rocks or felsic granitic rocks, which typically don't occur together
2) Siltstone, shale and argillite: finely bedded and often preserve primary sedimentary features and contain abundant fossils (mainly ammonites and bivalves)
3) Sandstone: typically feldspathic arenite, but may contain variable amounts of quartz and lithic fragments, most often silica cemented with lesser carbonate cement
4) Limestone: typically skarned, dolomitized and recrystallized, clastic sedimentary input is evident by the "dirty" nature of the strata, rare fossils occur away from recrystallized and skarned areas

UPPER TRIASSIC

Sinwa Formation Limestone and Lesser Clastic Rocks

1) Limestone
2) Argillite
3) Boulder conglomerate containing volcanic and intrusive rocks

Stuhini Group Mafic Volcanic Rocks

1) Pillow basalt
2) Andesitic lapilli tuff
3) Massive andesite: dark green, aphyric, aphanitic to fine-grained
4) Feldspar-augite porphyry: dark green, fine- to medium-grained, sparse <1mm feldspar and augite phenocrysts

MARINE SEDIMENTARY ROCK

1) Interbedded siltstone, feldspathic arenite and wacke: well-bedded
2) Argillite

7.2 Structure

Mapping during the 2005 field season confirmed observations made by Simmons (2004) that indicated Triassic and Jurassic strata of the Stuhini and Laberge Groups are variably deformed within the project area. These rocks are open to closed folded and with fold hinges trending NNW. This folding event is related to regional compression during the accretion of the Stikine Terrane onto the western margin of North America (Simmons, 2004). This compressional event ceased by ca. 175Ma, as evidenced by undeformed plutons which have intruded the deformed strata. Late Cretaceous volcanic rocks appear to have a spatial association with the area where anticlines plunge towards each other creating basins (Simmons, 2004). This relationship appears to apply near Lisadale Lake based on anticlines mapped by Souther (1971), but not at the Thorn property where Late Cretaceous volcanic deposition and preservation may be related to the location a synclinal structure also mapped by Souther (1971).

Several later steeply dipping normal faults offset Late Cretaceous rocks. Two prominent sets of steeply dipping normal faults trending 240° and 270° create major physiographic lineaments in the project area that may extend for several 10's of kilometers (Simmons, 2004). One such feature may extend westward from King Salmon Lake through the Joly and Jak areas.

7.3 Mineralization and Alteration

Within the Kizmet project area, a Late Cretaceous magmatic arc has been recently associated with both significant epithermal mineralization and hydrothermal alteration at the Thorn property. Mineralization there consists of high sulphidation epithermal veins and breccia hosted Ag-Pb-Zn-(Au-Cu). Outside of the Thorn property, recent fieldwork by Rimfire Minerals (2004) has identified Au-Ag bearing silicified vein breccia and silicified host rock, base metal rich carbonate veins, sedimentary hosted Au bearing disseminated pyrite, Cu skarn, and Fe-Zn skarn. (Simmons, 2004) in areas that are currently included in the Kizmet project. Other significant mineralization along the belt is known at the Golden Bear mine where Carlin-like mineralization (Poulsen, 2000) appears to be associated with Late Cretaceous magmatic rocks, as well as at the Tulsequah Chief past producing mine, where VHMS style mineralization is associated with Upper Triassic volcanics. These areas are located approximately 20 km southeast and 20 km northwest of the Kizmet project area respectively.

Fieldwork conducted by Barrick in 2005 identified various types of base metal and Ag-rich massive sulphide veins and Au-bearing quartz veins, but failed to identify any significant epithermal systems.

For the purposes of this report, the Kizmet project area has been divided into 7 areas of interest that collectively contain 26 individual targets worked in 2005 (Map 2). Each area has been discussed individually with respect to mineralization, alteration and mineral potential. Significant results have been provided as tables for each region. In this report significant results are defined by >50 ppb Au, >2.5 ppm Ag, and >1000 ppm Pb, Zn, or Cu.

7.3.1 LJ Area

The LJ area was staked in 2004 by Rimfire Minerals as part of their regional exploration program. Work in 2005 by Barrick focussed on testing areas proximal to Windy Table Formation stratigraphy, regional stream sediment anomalies, large scale structures and anomalous geochemistry from the 2004 program. Several types of mineralization were located, mostly at historic showings such as the Joly and in areas marked by historic documentation (Map 3). Mineralization included arsenopyrite +/- sphalerite and quartz vein sets near the Joly showing which ran as high as 17.45 ppm Au (KZ05R3001). This area has been adequately drill tested by at least 5 historic drill holes. In the vicinity of Mount Lester Jones, large accumulations of snow, poor weather and steep topography made exploration difficult but lack of large alteration zones, and poor rock sample results limit the potential for a large high sulphidation epithermal deposit in this area. Alteration throughout this area noticeably lacked in high temperature/low pH minerals such as kaolinite, dickite, pyrophyllite, and alunite. Illite and smectite were the only abundant clay minerals and did not represent an anomaly within the belt. In the event significant mineralization is ever located in this area, topography immediately surrounding Mount Lester Jones may be rugged enough to prohibit mine development.

One target **(LJ-1)** equivalent to the Lisadale Lake Target of Simmons, 2004, showed potential during compilation due to the presence of the largest preserved Windy Table Suite Volcanic succession within the Kizmet project area, and anomalous regional stream sediment (RGS) samples. Phase-1 evaluation identified two distinct areas of volcanism separated by a major fault (Map 3). The eastern area consists of predominantly unaltered subaerial volcanics while the western area consists of gossanous, weakly to moderately silicified, and locally weakly clay altered subaerial volcanics (Photo 1). The gossanous appearance is due to oxidation of the 1-3% contained disseminated pyrite. These volcanics were mapped in part as sediments during 2004 Rimfire Minerals fieldwork and generally ignored. Volcanics overlie gently dipping gossanous and locally siliceous sediments similar in overall appearance to the gossanous volcanics. PIMA analysis has identified illite, smectite, and silica alteration in the gossanous volcanic areas. Significant brecciation was also observed on the property including one large (2m X 25m) crosscutting breccia in a cliff face within the volcanics (Photo 2). Fifty rock samples were taken at the target. Of these only 1 returned significant values. Sample KZ05R0031 consisted of a secondary (possibly fault related) breccia containing 2% disseminated arsenopyrite in the matrix. Fragments appeared to be angular siltstone and possibly banded volcanic material. An area of similar coloured rock was seen above in bluff but appears to be limited in size.

Table 4: LJ – Significant Rock Sample Results

Sample Number	Width (cm)	Au (ppm)	Ag (ppm)	Pb (ppm)	Zn (ppm)	Cu (ppm)	As (ppm)	Hg (ppm)
KZ05R0002	GRAB	**0.064**	**2.7**	222	45	30	790	0.12
KZ05R0031	FLOAT	**1.64**	**3.4**	**2010**	104	66	>10000	0.005
KZ05R3001	GRAB	**17.45**	**25.6**	955	**2080**	142	>10000	0.09
KZ05R3003	FLOAT	**0.307**	0.6	32	68	6	6130	0.01
KZ05R3036	FLOAT	0.003	**2.8**	20	77	166	25	0.03
KZ05R3037	FLOAT	0.001	**5.2**	306	334	41	19	0.6
KZ05R3040	FLOAT	**0.092**	**11**	**1390**	**5860**	120	>10000	0.05

Photo 1: LJ-1 - Gossanous volcanic cap, looking northwest. Mt Lester Jones in background.

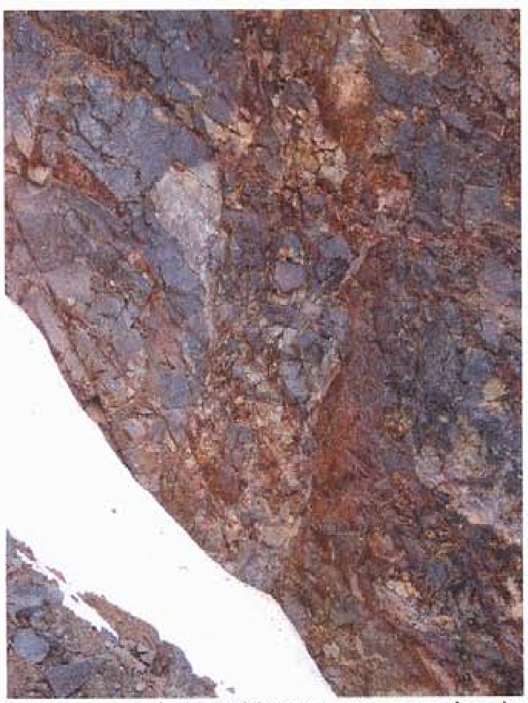

Photo 2: LJ-1 - 2 X 25m crosscutting breccia in a cliff face 40m from a major fault.

7.3.2 Sut Area

The Sut target area encompasses a large massif bounded by King Salmon Lake to the north and the Sutlahine River drainage to the south and east (Map 4). The King Salmon Lake target from Simmons, 2004 is included within this area. The property-scale geology is composed predominantly of Jurassic Laberge Group clastic sediments overlain by minor erosional remnants of Windy Table Suite volcanics; these are heavily intruded by both Windy Table Suite and Thorn Stock-aged sills and dykes (Map 4). Portions of this target are highlighted by anomalous RGS samples.

Fieldwork on this large target (~50km2) consisted of 14.5 person days in 2005 and identified isolated occurrences of small scale structurally controlled and dyke related mineralization. Several significant results of 101-273 ppb Au indicate some Au endowment but occurrences were always limited to zones less than a meter. The 1.3 ppm Au sample of "unaltered sediment" (276527) reported from the Rimfire 2004 campaign, was located and found to correlate to a 25 cm thick semi-massive pyrite vein. The sample does contain Au but mineralization is limited to 25 cm in width. Two brecciated rock samples containing dickite were identified as well as one sample with a pyrophyllite, illite, quartz assemblage (not compatible). These samples and several more containing kaolinite alteration outline a 2.0 km X 0.5 km area of increased hydrothermal activity (Map 4). The pyrophyllite sample is the only example of this mineral in the Kizmet project area. Silicification occurs in at least one significant area as several 10 m+ thick strongly silica replaced sedimentary horizons. These initially appeared prospective for mineralization but fieldwork indicated that there is no early stage leaching and no anomalous metal content.

Large gossans tend to occur at lower elevations all around the massif and are believed to generally represent silica and disseminated pyrite alteration of the basement rocks (Photo 3). Silica alteration is pervasive but not similar to that in high sulphidation systems. Large clay altered zones are absent as is Au mineralization. Table 5 and Map 4 summarize the significant results from the Sut area.

There appears to be little potential for a large high sulphidation deposit but smaller scale epithermal mineralization could be present in the area. The target is large and several gossanous areas remain poorly examined.

Photo 3: Sut-1 - West slope of massif, facing northwest and overlooking the Mad minfile showing. Minor clay alteration shows up as brown or light grey areas along slopes. Windy Table Suite subaerial volcanics were identified as erosional remnants on ridge tops.

Table 5: SUT Significant Rock Sample Results

Sample Number	Width (cm)	Au (ppm)	Ag (ppm)	Pb (ppm)	Zn (ppm)	Cu (ppm)	As (ppm)	Hg (ppm)
KZ05R0018	FLOAT	0.0005	**5.6**	**1100**	530	34	6	0.26
KZ05R0019	FLOAT	0.0005	**5.5**	**7770**	**7680**	8	23	0.04
KZ05R0020	FLOAT	0.003	**8**	876	167	21	4	0.05
KZ05R1228	GRAB	**0.273**	1.3	66	44	12	2080	0.2
KZ05R2014	CHIP(0.3m)	**0.225**	**4**	325	**1970**	109	3330	0.28
KZ05R2027	FLOAT	**0.061**	**2.9**	15	63	104	824	0.12
KZ05R2031	GRAB	0.005	**33.7**	**20900**	76	13	6	0.61
KZ05R2235	GRAB	**0.926**	2.2	25	79	37	10000	0.04
KZ05R2241	FLOAT	0.001	1	223	**10000**	82	28	0.83
KZ05R3009	FLOAT	0.007	**9.7**	**1345**	157	83	>27	0.89
KZ05R3277	GRAB	**0.101**	0.7	1	43	72	92	0.09
KZ05R3281	GRAB	**0.188**	0.4	7	10	8	167	0.34
KZ05R3282	FLOAT	**0.124**	**5.3**	**1475**	83	40	76	0.17

7.3.3 KZ Area

The KZ target area is spatially defined by the mountainous ranges southwest of the LJ claim block that form a strip of Barrick-staked claims running from the Sutlahine River towards the Taku River drainage. Geologically the target is defined by the occurrence of several erosional remnants of undeformed "Sloko Group" volcanics. Mapping in 2004 indicates that there is potential for these volcanics to be of Windy Table Suite affinity, increasing their mineral potential. The northwestern most peaks are too rugged and glaciated to be explored so work focussed on the basement rocks and glacial till surrounding their bases.

Fieldwork in 2005 identified multiple buttress shaped mountains with caps of subaerial volcanics preserved at the summits (Photo 4). The extent of volcanics is generally less than a few square kilometres, with stratigraphic thicknesses commonly less than 200m. Extensive dyking was observed along with minor magmatic and possibly local hydrothermal brecciation. Two samples containing minor base metal veining were discovered.

One target (KZ-1) allowed easier access to the Late Cretaceous – Tertiary volcanics (Map 5). These proved to be composed of multiple horizons of ash and lapilli tuff as well as pyroclastic flow deposits. The rocks were pristine and showed no substantial hydrothermal alteration. Three samples containing sphalerite, galena, and chalcopyrite in veins were discovered that contained 0.109 - 1.145ppm Au and up to 3260 ppm Ag (Map 5). These veins were located near the contact between the subaerial volcanics and the basement Laberge Group sediments and appear to cross through both units. Unfortunately, the extent of this veining is limited to a few metres. No argillic or advanced argillic alteration assemblages were identified on the KZ target areas, indicating very little potential for a high sulphidation deposit. Table 6 summarizes significant rock sample results from the KZ area.

Photo 4: Buttress shaped peaks of the northwest KZ target area. Several hundred metres of preserved subaerial volcanics cap basement rocks. Strong dyking and sill intrusion is evident.

Table 6: KZ Significant Rock Sample Results

Sample Number	Width (cm)	Au (ppm)	Ag (ppm)	Pb (ppm)	Zn (ppm)	Cu (ppm)	As (ppm)	Hg (ppm)
KZ05R0081	FLOAT	0.003	**3.5**	**3940**	208	93	40	0.01
KZ05R1033	GRAB	0.0005	0.5	331	**1195**	21	10	0.08
KZ05R3029	FLOAT	**0.482**	**569**	**20000**	**130000**	**5560**	5260	14.45
KZ05R3030	FLOAT	**1.145**	**3260**	**22700**	749	**3630**	10000	3.12
KZ05R3031	FLOAT	**0.109**	**285**	**4000**	899	**1655**	5650	0.59
KZ05R3032	GRAB	0.014	**31.4**	**1240**	419	69	539	0.21
KZ05R3033	GRAB	0.007	**4.2**	77	116	28	176	0.14
KZ05R3055	GRAB	0.004	**61.2**	**25200**	**2090**	1755	38	0.06

7.3.4 BS-J Area

The BS-J target area claims are spatially defined by the northwest trending high mountainous terrain along the west side of the Thorn property (Photo 5). Geologically the target is defined by known gossans and is proximal to an Au RGS anomaly (Map 6). Four target areas were visited in 2005 but no significant results were returned from the 39 rock samples. Two of the areas (**BS-J-1** and **BS-J-4**) did contain kaolinite alteration in granites, which dominated the area. The Au RGS anomalies were not sourced but a new area of either Sloko Group or Windy Table Group volcanism was identified at **BS-J-3** (Map 6). The area is not likely to host a >3 MOz Au high sulphidation epithermal deposit.

Photo 5: BS-J-3 Looking west over granitoid dominated stratigraphy.

7.3.5 Emu Area

The Emu area represents a large target (~75 km2) defined by the presence of both gossans and possibly clay alteration on trend with mineralization at the Thorn property. Fieldwork in 2005 identified large areas of feldspar +/- biotite porphyritic intrusive rocks overlain by erosional remnants of probable Windy Table Suite volcanic tuffs and ash likely representing distal facies to volcanic deposits at the Thorn (Photo 6). Basement Laberge Group clastic sediments were identified locally at relatively low elevations. Alteration in the target area was typified by significant kaolinite replacement of feldspar phenocrysts over a very large area (>4 km2) (Photo 7) as well as stronger replacement of the intrusive bodies near faults. In August three days of follow-up fieldwork focussed on areas not visited earlier in the year, and identified many significant exposures of subaerial volcanics (Photo 8) as well as continued extensive clay alteration (Map 7). Unfortunately, no advanced argillic alteration was located and only two rocks returned anomalous results. Sample KZ05R1207 sampled granitoid float containing galena, sphalerite, in calcite veining and is likely representative of mineralization at the minfile showings in the area (Griz, Griz 3). Table 7 summarizes the significant results from the Emu area.

The lack of anomalous stream sediment samples, advanced argillic alteration and mineralized rocks suggests there is very little potential for a >3 MOz Au high sulphidation epithermal deposit in this area.


Photo 6: Variably clay altered feldspar +/- biotite porphyritic intrusive rocks exposed in deep ravine on the Emu target.

Table 7: EMU Significant Rock Sample Results

Sample Number	**Width (cm)**	**Au (ppm)**	**Ag (ppm)**	**Pb (ppm)**	**Zn (ppm)**	**Cu (ppm)**	**As (ppm)**	**Hg (ppm)**
KZ05R1061	FLOAT	**0.105**	0.4	11	9	28	77	0.1
KZ05R1207	FLOAT	**0.144**	**51**	**10400**	**14100**	178	1110	4.74

Photo 7: Weakly gossanous talus slope in the Emu-3 area.

Photo 8: Erosional remnants of shallowly dipping lapilli and ash tuffs in the Emu 3 area.

7.3.6 Law Area

The Law area is spatially defined by a strip of mountainous land immediately north of both Tunjony and Big Trapper Lakes and south of the Thorn and Inlaw properties (Map 7 & 8). Geologically the target is defined by an abundance of Late Cretaceous – Paleogene undeformed "Sloko Group" subaerial volcanics, and anomalous RGS drainage basins.

Fieldwork in 2005 identified one area northwest of Big Trapper Lake containing multiple
2 m X 3 m to 30 m X 75 m areas of kaolinite, silica +/- dickite alteration (Photo 9). This argillic alteration assemblage indicates significant hydrothermal activity in the area. Alteration appears to be related to fault structures, lithological contacts, fold hinges, and dykes based on field observations. This area is underlain by a variably eroded 5-200m thick sequence of lapilli to block and ash flow tuffs believed to form the south facing slope of a presently eroded volcanic center, theoretically located ~1-5km to the north. These rocks are unconformably underlain by weakly folded Stuhini Group Mafic volcanics. The basement volcanics contain both the clay alteration mentioned above as well as ankerite alteration locally. Unfortunately the highest gold value returned from this area was 71 ppb.

An area along the ridge northwest of Tunjony Lake and adjacent to a large glacier (Map 8) was found to contain north-south trending vuggy quartz veins from 1-100 cm thick and commonly spaced 5-25 m apart over a distance of ~200 m. These veins locally contain Au values up to 1.025 ppm Au and may explain the anomalous (Au) drainage basin immediately north of these samples. Minor dickite alteration is associated with these veins. Au is restricted to the veins and doesn't extend into the country rock. Table 8 summarizes the significant assays from the Law area.

Although the Law area contains visible argillic alteration, gossanous areas, and is proximal to favourable stratigraphy, it is unlikely that there is a large high sulphidation epithermal deposit present. This is due to the lack of significant Au contained in rocks, and insufficient preservation of the volcanic package.

Table 8: LAW Significant Rock Sample Results

Sample Number	Width (cm)	Au (ppm)	Ag (ppm)	Pb (ppm)	Zn (ppm)	Cu (ppm)	As (ppm)	Hg (ppm)
KZ05R0189	GRAB	0.002	0.3	3	129	**1585**	3	0.01
KZ05R0338	FLOAT	**0.1**	0.5	1	30	80	483	1.29
KZ05R1080	GRAB	**0.071**	1.3	13	35	57	34	0.05
KZ05R1255	FLOAT	**0.068**	0.7	4	4	9	29	2.07
KZ05R2178	GRAB	0.039	1.3	5	60	**6730**	8	0.18
KZ05R2244	GRAB	0.003	0.3	5	**1065**	98	50	0.43
KZ05R2303	FLOAT	**0.057**	0.4	4	8	5	40	0.01
KZ05R3077	GRAB	**0.676**	0.9	12	5	46	12	0.01
KZ05R3078	FLOAT	**1.025**	**4**	181	46	15	141	0.18
KZ05R3079	FLOAT	**0.245**	2.4	309	95	15	26	0.57
KZ05R3098	GRAB	0.004	1.9	3	21	**1860**	6	43.2

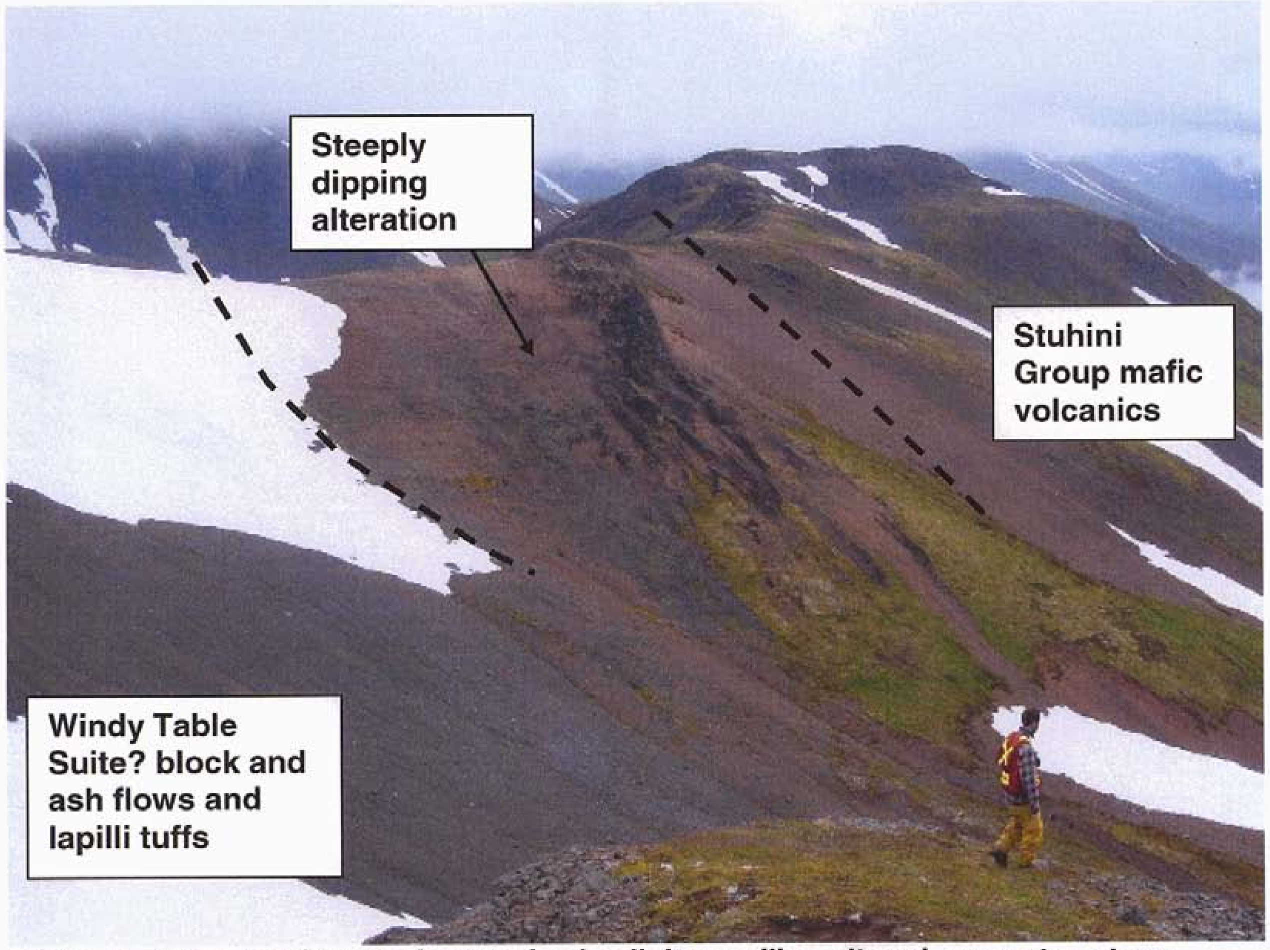

Photo 9: View looking northeast of a kaolinite + silica altered zone at a steep structure marking the contact between Late Cretaceous to Paleogene subaerial volcanics and Stuhini Group mafic volcanic basement. The core of the alteration zone is silicified.

7.3.7 Tun Area

The Tun Area is spatially defined by the claim group south of Tunjony Lake and northeast of the Metla property (Map 8). Geologically the target is defined by gossans, and an area containing anomalous property scale stream sediment samples identified through compilation.

Fieldwork in 2005 identified several rocks with anomalous Au +/- Cu in the Tun 4 area. These are associated with weak Cu-Au porphyry mineralization and weak silica alteration in one of several identified intrusive units. Table 9 summarizes significant results from rock sampling.

At the **Tun-5** target, undeformed subaerial volcanic remnants were identified at elevations higher that 2130 m. At lower elevations within intrusive rocks, gossans contained silica +/- illite, smectite, muscovite alteration but unfortunately did not contain any significant mineralization.

No evidence of epithermal mineralization was identified in the Tun area, greatly reducing the potential for a >3 MOz Au high sulphidation epithermal deposit.

Table 9: TUN Significant Rock Sample Results

Sample Number	Width (cm)	Au (ppm)	Ag (ppm)	Pb (ppm)	Zn (ppm)	Cu (ppm)	As (ppm)	Hg (ppm)
KZ05R2134	FLOAT	**0.715**	**6.5**	5	46	**9390**	15	0.01
KZ05R2135	GRAB	**0.071**	0.8	4	36	292	1	0.005
KZ05R2141	FLOAT	0.033	1.2	64	112	**2030**	14	0.03

7.3.8 Plum Area

The Plum target is defined by several highly anomalous (up to 494 ppb Au) drainages (RGS 1987). This area lies between the Metla and Tatsa properties and is underlain by somewhat different stratigraphy that the rest of the Kizmet project region. Stratigraphy is dominated by thick mafic volcanic packages containing minor pillow textures and intercalated sediments. These rocks have been mapped both as Triassic Stuhini Group mafic volcanics and Paleozoic Stikine Assemblage by other workers (Map 9)

The dominant features in this target area are multiple orange-brown gossans 100 m^2 to 1 km^2 size (Photo 10). They mark pervasively ankerite and locally silica altered mafic volcanics and local banded calcite vein zones. No significant Au mineralization was found in this style of alteration. Minor chalcopyrite and pyrite bearing calcite veins were located in, and adjacent to, these alteration zones at high elevations on the south side of the valley. Based on a historic work compilation of assessment reports, these veins typically contain anomalous gold and possibly account for the RGS stream sediment anomaly. These zones are similar to those immediately south on the Tatsa property, which although not of high sulphidation style, are interesting due to their size (up to 4 km^2) and intensity of hydrothermal alteration. Six kilometres of contour soil sampling (B-horizon) was conducted to identify any mineralization not located through prospecting. Results of this survey indicate a general trend of increasing gold to the southwest although no strong anomalies were identified (Map 9). The Barrick regional stream sediment survey results from 2005 indicate that the Plum drainages are not anomalous, conflicting with the RGS results. Five grab samples of silt (~50 grams) were taken from one of the Plum drainages in an effort to verify RGS results. The 5 samples were not anomalous in Au but did show a south-westerly trend of increasing Au similar to that in the soils lines. Table 10 summarized significant results from rock sampling in the Plum area.

The ankeritic style of alteration in mafic volcanics is not conducive to high sulphidation epithermal mineralization.

Photo 10: View of ankerite altered zone within mafic volcanics. This style of alteration dominates the basement rocks south of Trapper Lake.

Table 10: PLUM Significant Rock Sample Results

Sample Number	Width (cm)	Au (ppm)	Ag (ppm)	Pb (ppm)	Zn (ppm)	Cu (ppm)	As (ppm)	Hg (ppm)
KZ05R0135	FLOAT	**0.2**	**4.8**	1	26	**3570**	368	25.6
KZ05R0138	FLOAT	0.009	0.5	3	24	2790	91	0.03
KZ05R0293	FLOAT	**0.279**	**87**	2	249	**34400**	421	0.8
KZ05R2268	GRAB	0.0005	**3.8**	5	406	**9100**	4450	1.31
KZ05R3236	FLOAT	**0.195**	0.2	1	15	**3600**	62	0.17
KZ05R3239	FLOAT	0.029	0.6	1	30	**5870**	5	0.03
KZ05R3240	FLOAT	**0.158**	**18.1**	6	29	27	131	1.01
KZ05R3243	FLOAT	0.027	**3.6**	13	129	837	3100	12.2

8.0 REGIONAL STREAM SEDIMENT GEOCHEMISTRY

A regional stream sediment survey was conducted in the Kizmet project area during the 2005 field season. The objective of the survey was to improve upon the accuracy and comprehensiveness of the 104K RGS stream sediment survey conducted by the BC geological survey in 1987. It was also designed to confirm the anomalies highlighted by that survey within the area of interest.

Barrick's stream sediment survey was carried out with the aid of a Hughes 500D helicopter during the period of July 16-25th. This timing allowed us to sample approximately 1-2 weeks after peak water level in the streams had occurred. Drainages of ~ 15 km^2 were targeted to maintain a consistency in drainage size, (a problem with the RGS survey) although the range of sampled drainage sizes did vary significantly also in the 2005 survey due to natural drainage patterns. Drainage sizes ranged from 5.08 km^2 to 121.91 km^2, and averaged 24.00 km^2 although only one drainage >61.07 km^2 was sampled (La Jeune Creek, Thorn property, Figure 4).

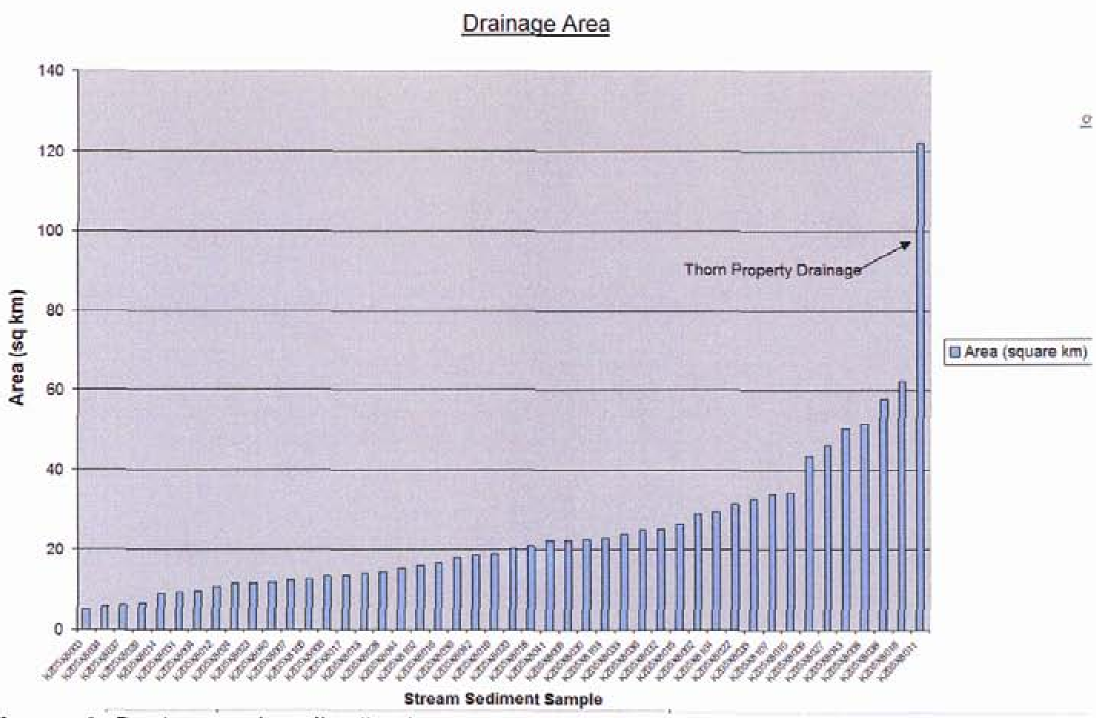

Figure 4: Drainage size distribution

Altogether, 41 samples of ~10kg in size were collected within the Kizmet project area. This represents a subset of a larger regional scale survey completed by Barrick in 2005. Sample media consisted of sediment passing 1mm mesh taken from the relatively high energy areas of gravel bars, and most commonly, the downstream toe of gravel bars. Alluvial deposits in this environment consisted of boulders, cobbles, sand and some silt trapped between large grain fractions. The muddy water acquired though sieving was retained and returned to camp to settle for 24 hours at which time any sediment separated from the water was collected and added to the sample. The samples were sent to ALS Chemex where they were dried, homogenized, screened to -80 mesh and split into 3 equal portions. One portion was then screened to -150 mesh and analysed by Au fire assay and 50 element ICP. All remaining fractions and splits were retained by Barrick for future reference. Results of the survey confirmed several of the strongest RGS anomalies and disputed several others. For the purpose of this report, anomalous Au is defined as greater than 75th percentile (130 ppb) while other elements were considered anomalous when above the 90th percentile. Percentile values were calculated based on the entire 2005 regional stream sediment survey which includes sample outside of the Kizmet project claims.

The strongest anomaly in the survey highlighted La Jeune Creek, downstream of the Thorn showings. This sample (1.34 ppm Au, 1 ppm Ag) was the strongest Au anomaly of the survey also contained the most robust multi-element epithermal signature, containing anomalous Au, Ag, Cu, Pb, Zn, As, Sb, Te, Bi (Map 10). The strength and elemental signature of this anomaly suggests that the mineralization on the Thorn property may be the most significant in size of any epithermal style occurrences in the Kizmet project area.

An interconnected cluster of 8 anomalous drainages is located in the northwest corner of the Kizmet project area, south of King Salmon Creek and covers the eastern half of the LJ claim block. The anomalous area extends further to the south and to the west of the **KZ-1** target (Map 10). These drainages vary in elemental abundances but can generally be broken into two types: 1) Au + Sn or 2) a partial epithermal suite (As, Sb +/- Ag, Zn, and Pb). The elemental zonation in the far northwest group forms a pattern of Au, Sn anomalies which likely occur distal to the theoretical source with the other elements occurring proximal to the theoretical source. These anomalies were not fully sourced during the 2005 campaign but several occurrences of mineralization were located including: possible porphyry style near the Red Cap showings, Au and Ag veins at the Joly showing, and base metal rich veins located by Simmons in 2005. The drainage south of the **LJ-1** target and southwest of Lisadale Lake is heavy forested and little work was done in the area. Several mineralized rock samples were taken from the periphery of this drainage but no mineralization was identified of sufficient size to account for this anomaly.

The drainage immediately west of Big Trapper Lake was anomalous in Au, Hg, and Ba showing a moderate epithermal signature. Extensive phase one fieldwork was conducted on the north side of this drainage and although very gossanous, no significant mineralization was found. Perhaps Au from mineralized quartz veins north of Tunjony Lake has been able to pass though the Tunjony Lake sediment trap.

The drainages to the northeast of the Thorn property, in the Emu target area are enriched in Ba. No source for this anomaly is known.

9.0 DISCUSSION and CONCLUSIONS

The 2005 program met the predefined objectives: i.e., it systematically evaluated the prioritized pre-defined targets, it downgraded the mineral potential of the belt by significantly reducing the likelihood of a discovery of a large high sulphidation epithermal deposit, and it provided enough information to reduce the size of the initial mineral claim package. The 2005 program has confirmed that the belt is anomalous with respect to precious and base metal abundances, evidenced by several strong anomalies identified by the 2005 regional stream sediment survey and multiple Au, Ag or base metal-rich vein showings. However, very little evidence was seen to support the potential for large high sulphidation systems. There are no known advanced argillic alteration zones and no mineralization was identified that contained high sulphidation epithermal mineralogy or textures.

Although the belt appeared to be a likely candidate to host high sulphidation mineralization, (supported by the known occurrences on the Thorn property) several factors may hinder the exploration potential of these systems. The main factor is suggested to be the current erosional levels. Generally, areas underlain by Windy Table Suite volcanics are erosional remnants, consisting of small (less than 4 km^2) isolated caps of volcanic rocks. Commonly these are comprised by block and ash flows as well as ash to lapilli tuffs, which likely represent medial to distal lithofacies. The exception is at the Thorn property where mineralization is located proximal to a thick (>900 m) series of volcanic units including substantial flow deposits indicating proximity to volcanic vent(s). Only the Lisadale Lake **(LJ-1)** target area studied in 2005 appears to contain significant flow facies volcanics and has preservation of a significant thickness of volcanic strata. Other targets generally contain a thin, discontinuous veneer of tuffs representing distal volcanic facies and indicate less potential for epithermal systems.

Another negative factor is the difficult terrain in the least explored areas. While much of the project area lies in well exposed mountainous topography, much of this is now believed to be well prospected, leaving only the low elevation, highly vegetated, and still very rugged areas. Although difficult to explore, these areas contain the most remaining mineral potential. The drainages in the northwestern portion of the project area (e.g. LJ targets) are a good example as they contain stream sediment anomalies sourced from steep but generally forested areas. The mineralization on the Thorn property provides a good example of mineralization found in this tough topographical environment, proving that although difficult, to explore, it can be done.

10.0 RECOMMENDATIONS

Although results from the 2005 field seasons were generally poor, several areas could be considered for future work as follows:

- The 2005 Barrick regional stream sediment survey identified an area of 8 contiguous drainages in the far northwest region of the Kizmet claims that contain either Au, Sn anomalies or Au, Sb +/-Ag, Pb, Zn anomalies (Map 11). There is an apparent zonation from Au and Sn in the larger second order drainages to Sb, As +/- Ag, Pb, Ag in the smaller upstream drainage basins immediately west. This is interpreted to represent anomalies related to epithermal mineralization in the western drainages. This area might also partially source porphyry? style mineralization near the Redcap property. This area is in need of follow-up fieldwork.

- The Kizmet main claim block should be cut down to the most prospective areas, mainly defined by anomalous drainages, and the presence of similar geological features to that at the Thorn property, such as Windy Table Suite volcanics (and intrusive equivalents). Excess assessment expenditures for 2005 should be applied to 2006 to meet requirements for 2 years.

- The focus of any future work should be on a "Thorn style", structurally related, intermediate-high sulphidation, high grade deposit, rather that an Andean style high sulphidation model. Well exposed examples of this style of mineralization would likely have been discovered so any future efforts should focus on buried or more subtle mineralization.

APPENDIX I

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APPENDIX II

STATEMENT OF EXPENDITURES

STATEMENT OF EXPENDITURES

PROFESSIONAL FEES AND WAGES

Full time and contract employees worked discontinuously on the properties from June 9 to July 27th, and then from Aug 10 to Aug 28, 2005 totalling 283 person days. Report writing was conducted discontinuously between Nov 8 and Dec 15th, 2005, for a total of 21 person days. 20 office days were allocated for project preparation and research in the spring of 2005.

Full Time Staff		
Richard Mann	82.75 days @$300/day	$24,825.00
Adrian Newton	79.5 days @$300/day	$23,850.00
Martin Stewart	0.75 days @$300/day	$225.00
Rob Brown	9 days @$400/day	$3,600.00
Dave Heberlein	3 days @$400/day	$1,200.00
Contract Staff		
Yvonne Thornton	60.5 days @ $264.60/day	$16,008.30
Marc Cianci	51.5 days @ $183.60/day	$9,455.40
Greg Ross	15 days @ $156.60/day	$2,349.00
Adam Simmons	1 day @$520/day	$520.00
		$ 82,032.70

Camp costs, and all costs related to transportation, field consumables, fuel and camp food occurred from June 10-July 27 and from Aug 10- Aug 27, 2005.

LOGISTICS		
Travel Expenses		$10,493.32
Fuel		$20,129.56
Camp Costs	281.62 days @ $25/day	$7,040.52
Field Consumables, Materials & Supplies		$6,829.12
Helicopter Charters (66.75 hrs @ ave rate $886.19/hr + $1000.26 mob)		$60,153.17
Aircraft Charters		$25,843.00
Freight and Shipping		$8,386.56
Communications		$3,538.99
Camp Food		$5,343.17
		$147,757.41

Equipment Rentals		
Radios and Repeater	(June 1 – Sept 15, 2005)	$2,207.94
2000W Generator	(~July 7 – Aug 28, 2005) 51.52 days @$10/day	$515.16
		$2,723.10

TECHNICAL STUDIES

Analytical Work	
848 rock samples at an average cost of $34.08/sample	$28,895.93
31 soil samples at an average cost of $29.44/sample	$912.54
41 10 kg silt samples at an average cost of $56.54/sample	$2,318.32
5 grab silt samples at an average cost of $45.37/sample	$226.84
Maps, TRIM Topo, Publications & Reproductions	$1,952.04
PIMA Analysis and Interpretation 848 samples @ $10/sample	$8,480.00
	$42,785.67

SERVICES

Expediting	$6,046.69
	$6,046.69

Total:	**$281,345.57**

TOTAL WORK FILED FOR ASSESSMENT CREDIT: **$281,345.57**

APPENDIX III

STATEMENTS OF QUALIFICATION

STATEMENT OF QUALIFICATIONS

I, Richard K. Mann, of 1322 Apel Drive, Port Coquitlam, British Columbia, do hereby certify that:

1. I am presently employed by Barrick Gold Corporation of 700-1055 West Georgia Street, Vancouver, British Columbia. V6E 3P3.
2. I am a graduate of the University of British Columbia, Vancouver, BC (1999, BSc in Geology)
3. I have been employed in my profession as an Exploration Geologist since graduation.
4. I am duly registered as a Professional Geoscientist in training in the Province of British Columbia (#131934).
5. I have no interest in the property described herein, nor in the securities of any company associated with the property, nor do I have any plans to acquire any such interest.

Signed at Vancouver, British Columbia this 24th day of March, 2006.

Richard K. Mann, B.Sc., GIT

STATEMENT OF QUALIFICATIONS

I, Adrian C. Newton, of 1655 22nd Street, West Vancouver, British Columbia, do hereby certify that:

1. I am presently employed by Barrick Gold Corporation of 700-1055 West Georgia Street, Vancouver, British Columbia. V6E 3P3.
2. I am a graduate of Simon Fraser University, Vancouver, BC (2004, BSc in Earth Sciences)
3. I have been employed in my profession as an Exploration Geologist since graduation.
4. I am duly registered as a Professional Geoscientist in training in the Province of British Columbia (#145726).
5. I have no interest in the property described herein, nor in the securities of any company associated with the property, nor do I have any plans to acquire any such interest.

Signed at Vancouver, British Columbia this 24th day of March, 2006.

Adrian C. Newton, B.Sc., GIT