Assessment Report for the

Mount Rice Property

Fort Steele Mining Division
N.T.S. 82 F/ 10E
Latitude 49° 42' N, Longitude 116° 32' W

for

Jasper Mining Corporation
1020, 833 - 4th Avenue S.W.
Calgary, Alberta
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of

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SUMMARY

The MOUNT RICE property is located in the western Purcell Mountains, approximately 61 kilometres west of Cranbrook, B.C. and consists of 2 Mineral Tenure Online (MTO) mineral tenures. The northwest corner of the property can be accessed by gravel Forest Service Roads (FSR) from Cranbrook / Kimberley along the St. Mary's road network.

The claims were initially acquired as they overlay an apparent magnetic anomaly evident on regional geophysics available from the BC Government MapPlace website, hosted within strata correlated to the uppermost Late Proterozoic Purcell Supergroup . Felsic intrusive lithologies correlated to the Bayonne Magmatic Suite typically have a prominent magnetic signature, either associated with the intrusion or as a halo in the immediately surrounding host rocks.

Recent work on the Mount Skelly Pluton (proposed as a possible model, located approximately 23 km to the southwest) has distinguished a three phase intrusive complex (Logan and Mann 2000) that consists of fine- to coarse-grained granites correlated to the Cretaceous Bayonne Magmatic Suite. Near contacts with sedimentary strata, the granite appears to be both finer grained and perhaps more mafic, having a darker colour. In addition, there are more xenoliths of (an) earlier phase(s) of intrusive material and rounded sedimentary inclusions. Phenocrysts of alkali feldspar are present, ranging in size from less than a centimetre to approximately 2 centimetres in diameter, within a matrix of plagioclase feldspar, quartz and biotite ± hornblende. The granite has local iron-stained veins with variable amounts of iron sulphide, predominantly as pyrite. The veins appear to occupy apparent discontinuous brittle shear zones which trend essentially north-south (± 20°). The Mount Skelly Pluton (Complex) comprises the exploration model for the properties of the Cretaceous Granite Project.

In addition, recent work on mineralization associated with intrusions has resulted in the Intrusion-Related Gold (IRG) Model. Examples include numerous examples in Alaska (i.e. Fort Knox, Pogo) and continue southeastward through the Tintina Gold Belt. Several occurrences in B.C. have been examined in a preliminary manner to evaluate Intrusion-Related Gold potential, including the Baldy Batholith and the Mt. Skelley Pluton. With reference to this model, elevated As, Bi, Sb, W are considered as "pathfinder" elements for potential IRG deposits. Furthermore, the Sanca Stock and Mount Skelly Pluton are of Cretaceous age with a prominent magnetic halo, both features characteristic of many occurrences along the Tintina Gold Belt. Several locations, including many of the documented MINFILE occurrences, may be compatible with an IRG-type model, particularly those associated with the northwestern lobe (Sanca Stock) of the exposed granitic phases.

The 2006 field program was limited to acquisition of a preliminary suite of soil samples. A total of 164 soil samples were recovered over three days on the property. Samples were submitted to Acme Analytical Laboratories for processing using the SS80 package and analysis using the Group 1EX package.

TABLE OF CONTENTS

LIST OF FIGURES

LIST OF APPENDICES

INTRODUCTION

The MOUNT RICE property is located in the western Purcell Mountains, approximately 61 kilometres west of Cranbrook, B.C. (Fig. 1 and 2) and consists of 2 Mineral Tenure Online (MTO) mineral tenures (Fig. 3). The northwest corner of the property can be accessed by gravel Forest Service Roads (FSR) from Cranbrook / Kimberley along the St. Mary's road network.

The claims were initially acquired as they overlay an apparent magnetic anomaly, evident on regional geophysics available from the BC Government MapPlace website, hosted within strata correlated to the uppermost Late Proterozoic Purcell Supergroup (Fig. 4). Felsic intrusive lithologies correlated to the Bayonne Magmatic Suite typically have a prominent magnetic signature, either associated with the intrusion or as a halo in the immediately surrounding host rocks.

Recent work on the Mount Skelly Pluton (proposed as a possible model, located approximately 23 km to the southwest) has distinguished a three phase intrusive complex (Logan and Mann 2000) that consists of fine- to coarse-grained granites correlated to the Cretaceous Bayonne Magmatic Suite. Near contacts with sedimentary strata, the granite appears to be both finer grained and perhaps more mafic, having a darker colour. In addition, there are more xenoliths of (an) earlier phase(s) of intrusive material and rounded sedimentary inclusions. Phenocrysts of alkali feldspar are present, ranging in size from less than a centimetre to approximately 2 centimetres in diameter, within a matrix of plagioclase feldspar, quartz and biotite ± hornblende. The granite has local iron-stained veins with variable amounts of iron sulphide, predominantly as pyrite. The veins appear to occupy apparent discontinuous brittle shear zones which trend essentially north-south (± 20°). The Mount Skelly Pluton (Complex) comprises the exploration model for the properties of the Cretaceous Granite Project.

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The 2006 field program was limited to acquisition of a preliminary suite of soil samples collected along three contour lines. A total of 164 soil samples were recovered over three days on the property. Samples were submitted to Acme Analytical Laboratories for processing using the SS80 package and analysis using the Group 1EX package.

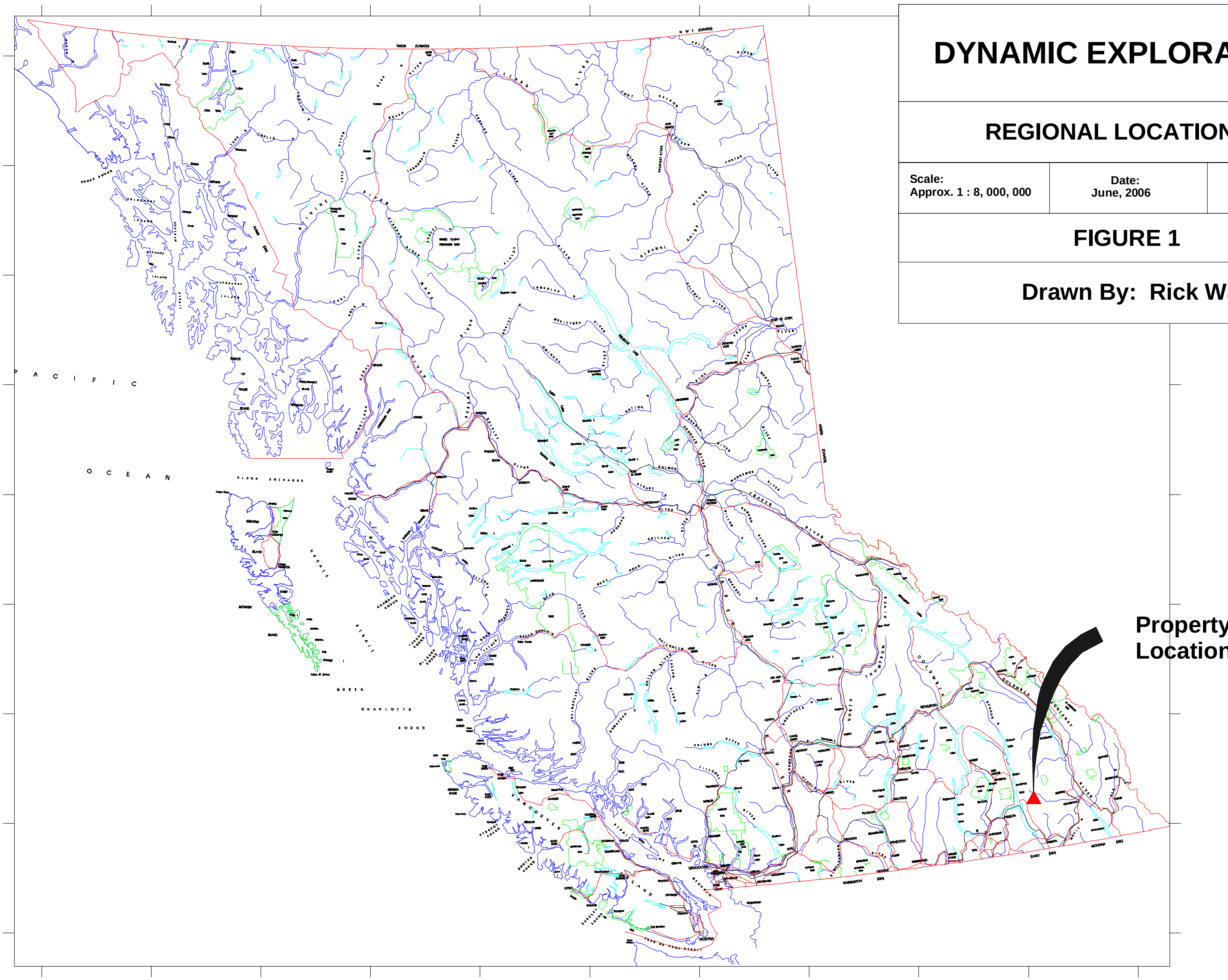
DYNAMIC EXPLORATION LTD

REGIONAL LOCATION MAP

Scale:	Date:	Mapsheet:
Approx. 1 : 8, 000, 000	June, 2006	N.T.S. 82F / 10E BCGS: 082F 068 & 078

FIGURE 1

Drawn By: Rick Walker



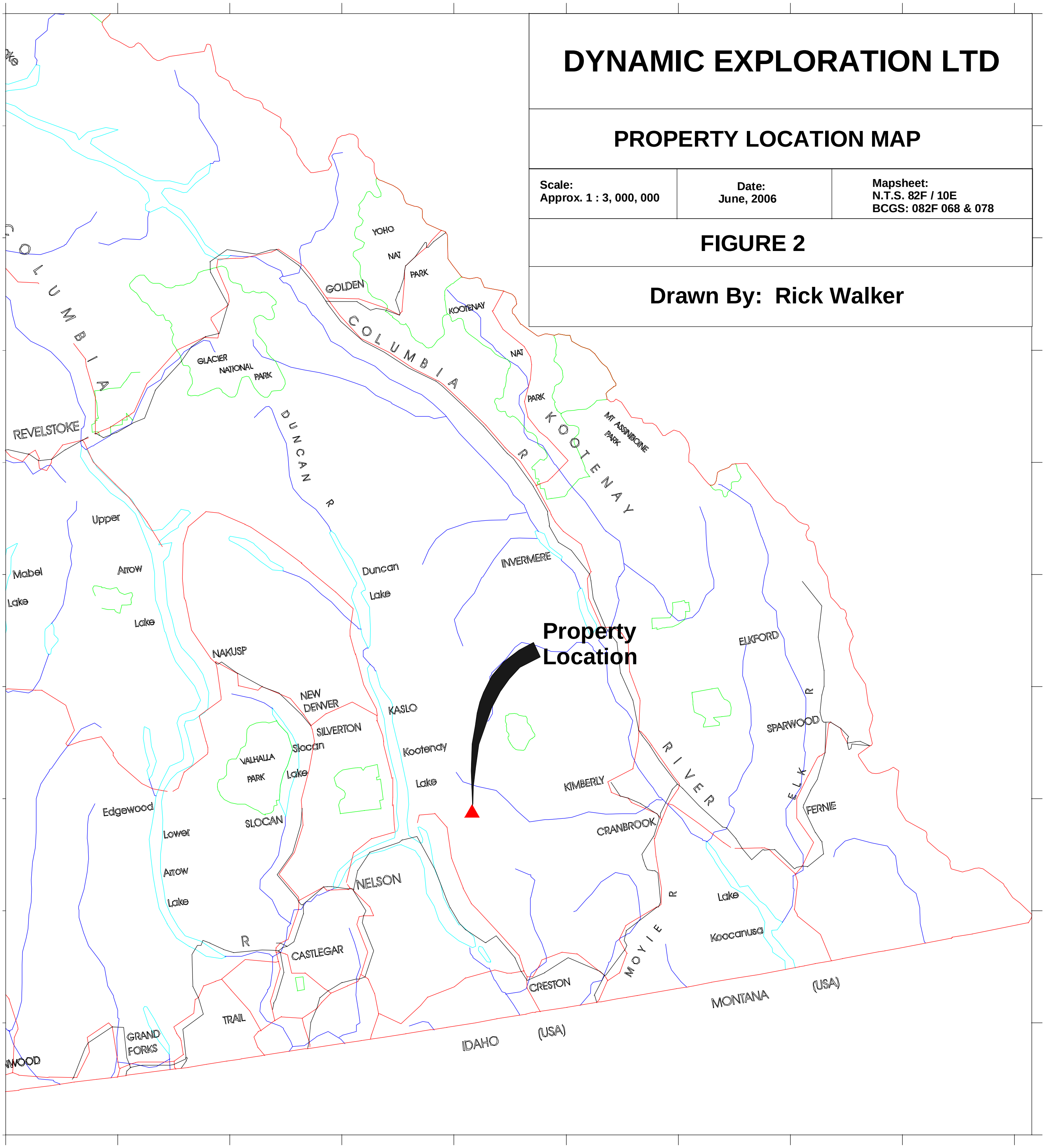
DYNAMIC EXPLORATION LTD

PROPERTY LOCATION MAP

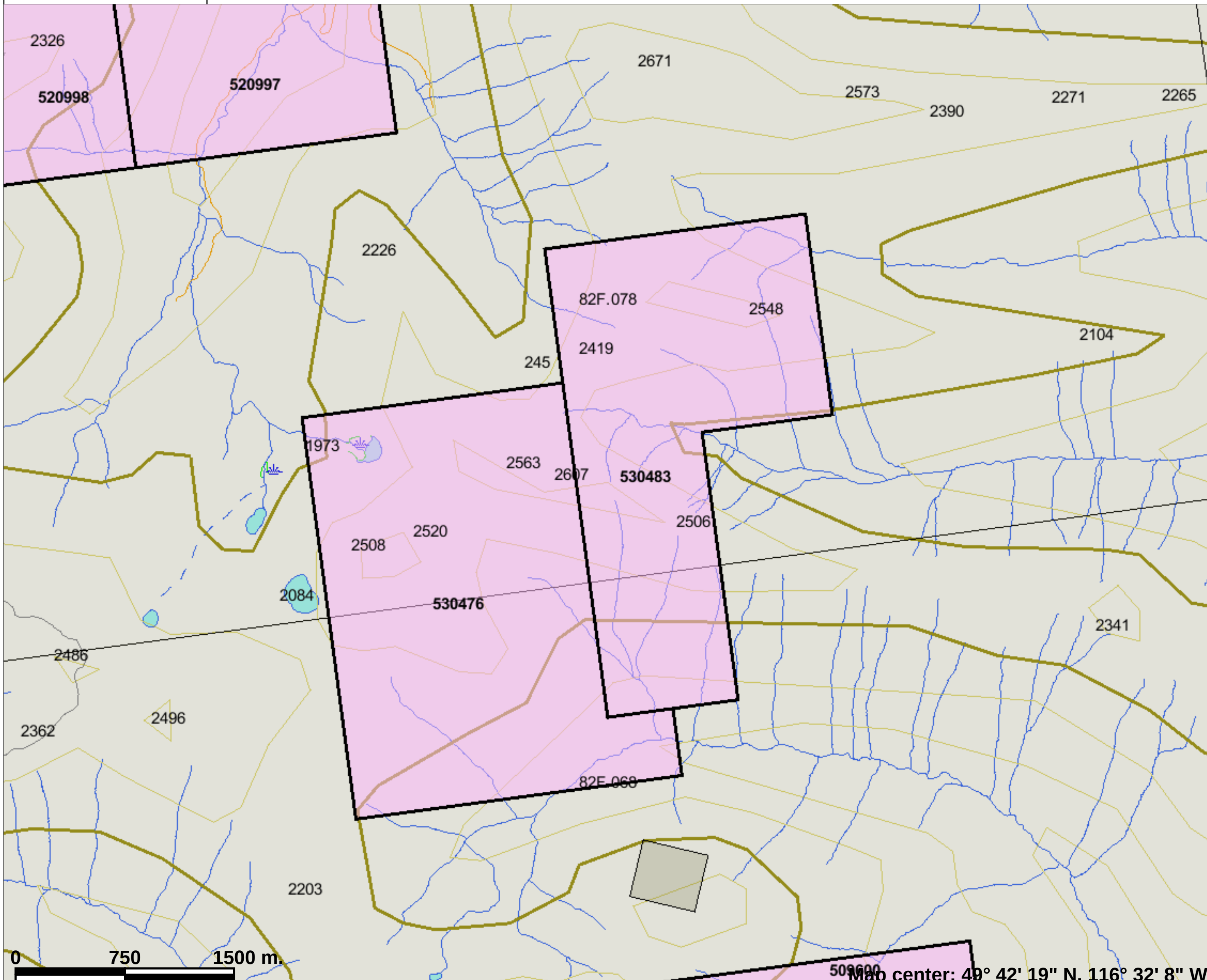
Scale: Approx. 1 : 3, 000, 000	Date: June, 2006	Mapsheet: N.T.S. 82F / 10E BCGS: 082F 068 & 078

FIGURE 2

Drawn By: Rick Walker



Internet Mapping Framework



Legend

Indian Reserves
National Parks
Parks
Mineral Tenures (Mineral - LRDW)
Mineral Claim
Mineral Lease
Reserves (Mineral - LRDW Sites)
Placer Claim Designation
Placer Lease Designation
No Staking Reserve
Conditional Reserve
Release Required Reserve
Surface Restriction
Recreation Area
Others
Mining Division (MTO)
Survey Parcels
BCGS Grid
Contours (1:250K)
Contour - Index
Contour - Intermediate
Areaof Exclusion
Areaof Indefinite Contours
Annotation (1:20K)
Transportation - Points (TRIM)
Helipad
Transportation - Lines (TRIM)
Airfield
Airport
Airstrip
Airport.Abandoned
Ferry Route
Road (Gravel Undivided) - 1 Lane

Scale: 1:41,585

This map is a user generated static output from an Internet mapping site and is for general reference only. Data layers that appear on this map may or may not be accurate, current, or otherwise reliable. THIS MAP IS NOT TO BE USED FOR NAVIGATION.

Notes: Figure 3: Mineral Tenures from MTO web-site

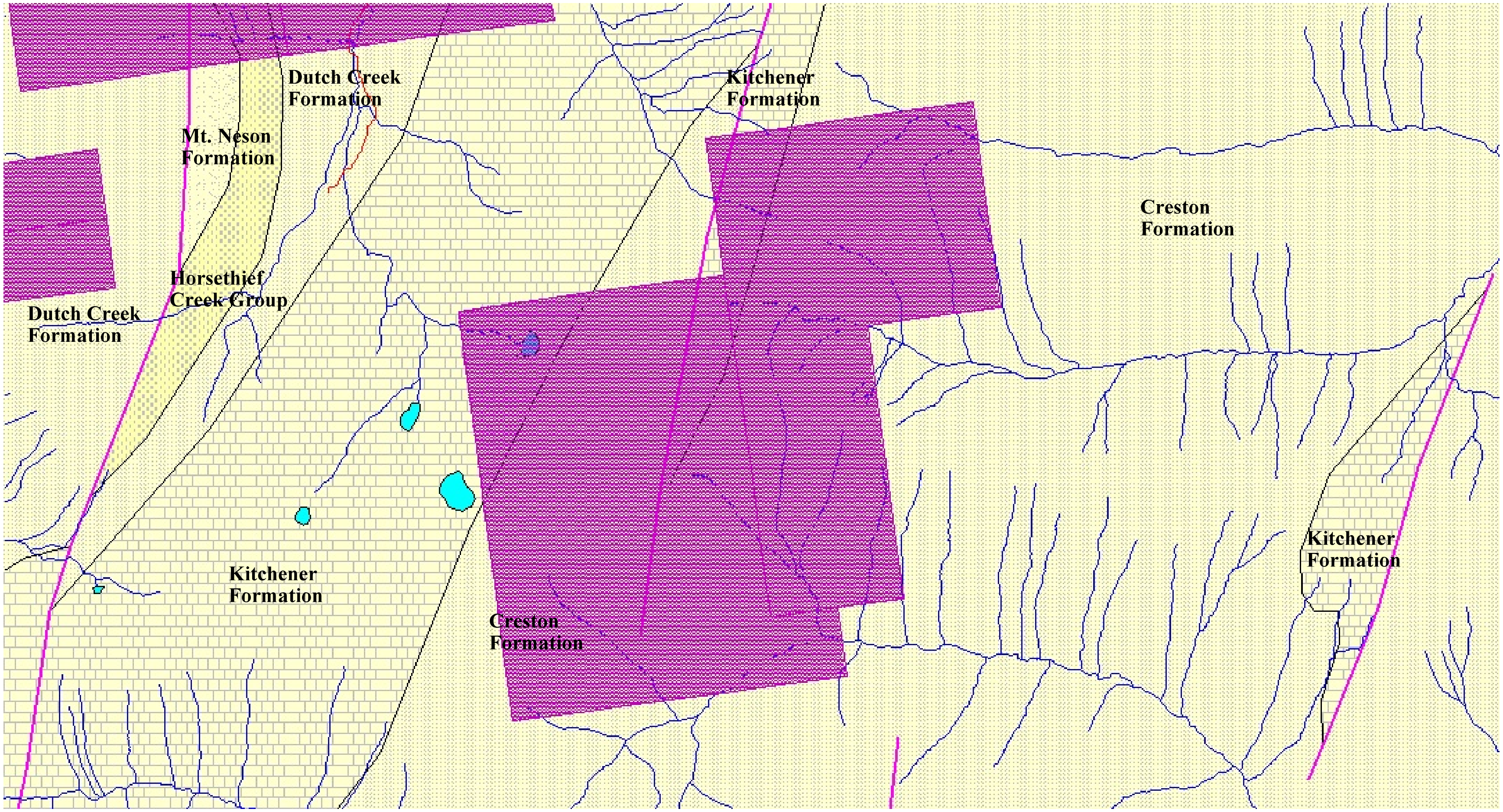


Figure 4: Local Geology map showing area covered by the Mount Rice Mineral tenures (approximate scale 1: 32,000). Fault contacts represented by dark pink lines, geological contacts by thin black lines.

LOCATION AND ACCESS

The MOUNT RICE property is located in the western Purcell Mountains (latitude 49° 42' N, longitude 116° 32' W), approximately 61 kilometres west of Cranbrook, B.C. on N.T.S. mapsheet 82 F/10E (Fig. 1 and 2). The property consists of 2 Mineral Tenure Online (MTO) mineral tenures.

The property can be accessed by gravel Forest Service Roads (FSR) from Cranbrook / Kimberley along the St. Mary's Road. The road is well maintained west of St. Mary's Lake to Km 42. At km 45, continue north along the St. Mary's FSR for approximately 12 km to a fork in the road. Take the left fork and cross over the bridge, then the left fork again, crossing over another bridge. Approximately 3 km farther along, the road crosses another bridge to the south side of the St. Mary's River. Approximately 6 km further along the road, there is a branch road on the south side of the road which provides vehicle access along the east side of Sawyer Creek. Continuing along the Sawyer Creek Road southward to the end allows access to the northern edge of the property.

Helicopter access is recommended for much of the property.

PHYSIOGRAPHY AND CLIMATE

The MOUNT RICE property is located slightly east of Rose Pass (Fig. 2). Relief in the area varies from 1,820 metres (5,570 feet) along a creek at the western edge of the property to approximately 2,520 metres (8,270 feet) on Mount Rice.

Vegetation in the area consists predominantly of coniferous trees, with deciduous trees preferentially located along the valley bottom. Undergrowth consists largely of small deciduous shrubs.

The claims are located east of Kootenay Lake in a regional topographic high, comprising the local drainage divide, and are therefore subject to heavier precipitation. As a result, the region is characterized by heavy snowfall during the winter months. The property is available for exploration from June to late October.

CLAIM STATUS

The MOUNT RICE property consists of 2 Mineral Tenure On-line (MTO) mineral tenures (Fig. 3). Significant claim data are summarized below:

Tenure Number	Claim Name	Anniversary Date		Area (ha)
530476	RICE WEST	2011 / Mar / 24		522.408
530483	RICE EAST	2011 / Mar / 24		417.855
			Total	**940.263**

*After 2006 assessment credit applied.

HISTORY

No previous work is known to the author for the property or the immediately surrounding area.

REGIONAL GEOLOGY

The only previous work undertaken pertaining to the general area of the Sawyer claims was that of Reesor (1996, 1993) for the east side of Kootenay Lake. The stratigraphy of the Purcell Supergroup strata has been well described to the east by Höy (1993) and the Purcell and Windermere Supergroup to the north by Pope (1990).

Stratigraphy

Proterozoic

Belt-Purcell Supergroup

The following has been modified from Höy (1993).

Sheppard Creek Formation (Lower Dutch Creek Formation)

The Sheppard Formation includes up to several hundred metres of stromatolitic dolomite, quartz aretrite, siltstone and argillite lying above the Nicol Creek Formation. A dramatic increase in thickness in the Skookumchuk area is accompanied by prominent facies changes in the Sheppard Formation and in the overlying Gateway and Phillips formations.

The Sheppard Formation is characterized by an assemblage of green siltite, sandy dolomite, quartz wacke, distinctive stromatolitic dolomite and oolitic dolomite layers.

West of Skookumchuk, the formation is still recognizable but is referred to as the lower Dutch Creek Formation. It comprises green siltstone and argillite with minor dolomitic siltstone and, near the top, stromatolitic dolomite. This stromatolitic sequence can be traced north of Bradford Creek and marks the contact between the lower and upper Dutch Creek. It comprises cycles of rounded and gritty quartz wackestone, overlain by oolitic, stromatolitic or massive dolomite. These cycles may contain a few thin purple argillite beds with mud cracks and locally, rip-up clasts. They are overlain by and interbedded with light green siltstone-argillite couplets, usually lenticular, laminated and graded.

Gateway Formation (Upper Dutch Creek Formation)

The Gateway Formation is defined to include siltite, argillite, arenite and dolomite between the Sheppard Formation and red and maroon siltstone and argillite of the overlying Phillips Formation. It correlates with the lower part of the upper Dutch Creek Formation northwest of Skookumchuk. The Gateway Formation comprises dominantly pale green siltstone and minor dolomitic or argillaceous siltstone.

... Salt casts and symmetrical ripples throughout the Gateway Formation suggest deposition in shallow water; dessication cracks, mud-chip breccias and oxidized facies indicate periods of subaerial exposure. ... The formation thickens rapidly to the north in the Skookumchuk area primarily as the result of an increase in the pale green siltstone component. The absence of the overlying Phillips Formation, sparse outcrop and the similarity between lithologies in the upper Gateway and lower Roosville formations make it difficult to determine the thickness and extent of the Gateway Formation to the north and west. ...

Dutch Creek Formation

The Dutch Creek Formation is defined as a group of rocks between the Purcell lavas (Nicol Creek Formation) and the Mount Nelson Formation. The lavas are not exposed in the Lardeau and Nelson east- half map areas and hence it is difficult to determine the exact thickness and extent of the Dutch Creek Formation there. It is estimated to be between 1200 and 1500 metres thick in the Windermere area and a 1300~metre section has been measured east of Kootenay Lake at Rose Pass.

In the Fernie west-half map area, the Dutch Creek Formation is only exposed northwest of Skookumchuck. The lower part of the formation is described in the section on the Sheppard Formation. The upper part includes the Gateway Formation the Roosville Formation and overlying rocks beneath the Mount Nelson Formation. The maximum thickness of the Dutch Creek Formation in the Bradford Creek area is estimated to be 4800 metres, including approximately 3300 metres of upper Dutch Creek.

The upper Dutch Creek is discontinuously exposed north of Skookumchuck. A carbonate marker bed approximately 200 metres thick occurs within the formation some 3000 metres above the Nicol Creek lavas. It is a massive, cream to tan-weathering, thick to medium-bedded dolomite and limestone unit. Crypto-algal features are present locally. The top and the base of the unit consist mainly of argillaceous silty dolomite. It is included within the Dutch Creek rather than the Mount Nelson Formation as the basal quartzite typical of the Mount Nelson is not exposed below it. Furthermore, green siltstone, black argillite and thin oolitic dolomite interbeds higher in the section probably correlate with similar facies in the Roosville Formation at Larchwood Lake.

Mount Nelson Formation

The Mount Nelson Formation comprises a thick sequence of quartzite, dolomitic argillite and siltstone that conformably overlies the Dutch Creek Formation. It was restricted to include only the lower part of the formation. The upper part, informally named the Frances Creek Formation, is separated from the Mount Nelson Formation (new) by a disconformity.

The lower Mount Nelson Formation is divisible into three members in the Mount Forster map: a basal white orthoquartzite 100 to 200 metres thick, 100 to 300 metres of buff and grey dolomites and an upper unit, to 370 metres thick, of purple and red shale with buff dolomite interbeds. The overlying Frances Creek Formation comprises thick-bedded orthoquartzite, grey dolomite and interbedded sandstone and shale.

The total thickness of the Mount Nelson Formation (new) in the Mount Forster area varies from 500 metres to 1950 metres, due partly to erosion prior to deposition of the Frances Creek Formation or Windermere Supergroup and partly to syndepositional tectonics. The Frances Creek Formation varies in thickness from

750 metres to 1020 metres. At Rose Pass east of Kootenay Lake, the entire Mount Nelson Formation is approximately 750 metres thick.

In Fernie west-half map area, the Mount Nelson Formation is only exposed at Lookout Mountain along the northern edge of the map area. It has a gradational contact with the underlying Dutch Creek Formation; phyllitic black argillite-siltstone rocks become increasingly more quartzitic and the interbeds of quartz wacke become cleaner up-section. The basal quartzite of the Mount Nelson is a clean, well-rounded and well-sorted, medium-bedded orthoquartzite containing a few thin beds of sandy dolomite. The basal quartzite is overlain by a mixture of white, green and purple quartz arenite and dolomitic sandstone, locally gritty, as well as some purplish dolomite and argillite. Locally, the diagenetic character of these maroon beds is clearly demonstrated as the colouring crosscuts bedding planes and leaves spotty remnants of light green argillite. A buff weathering sequence of dolomite overlies these quartzwacke, siltstone and argillaceous dolomite beds. This package is overlain by more green siltstone and minor purple siltstone and argillite. The total exposed thickness of the Mount Nelson Formation is approximately 400 metres.

The following has been summarized from Aitken and McMechan (1991).

Middle carbonate division

A distinctive carbonate unit comprises the middle division of the Purcell (Belt) Supergroup. To the east, in the Rocky and eastern Purcell mountains, the middle division consists of the well known Kitchener Formation In the west the middle carbonate division consists of the more basinal facies of the thick, lower subdivision of the Coppery Creek Group. The thick (1400 m) lower unit consists of dolomite interbedded with green, grey or black phyllite which grades upward to silvery and green phyllite, siltite and some carbonate.

Upper division

The strata comprising the Van Creek Sheppard, Gateway and Roosville formations of the Rocky
and eastern Purcell Mountains pass laterally into a succession of grey and green siltite, argillite and phyllite, quartzite, argillaceous dolomite and dolomite. The volcanic (Nicol Creek) and red quartzite marker (Phillips) units thin and disappear to the west, making subdivision of the upper division impractical. Therefore, the upper two units of the 'Coppery Creek' and 'La France Creek' groups are interpreted to comprise the upper division along the western Purcell Mountains.

The upper two divisions of the Coppery Creek group consists of a middle unit approximately 200 m thick comprised of thinly laminated black phyllite and grey siltite. The upper unit consists of silvery phyllite, calcareous dark grey phyllite and dolomite, with a sequence of interbedded dolomite and quartzite at the top and is approximately 300 metres thick.

The 'La France Creek group' of the western Purcell is approximately 1000 m thick, comprised of
intensely deformed and metamorphosed sediments dominated by siltite, quartzite and phyllite. The group has been subdivided into a lower unit consisting of thinly interbedded, black phyllite and grey siltite and an upper unit of grey siltite and quartzite with black phyllite and carbonate-bearing siltite and phyllite near the top. The 'La France Creek group' gradationally overlies the upper unit of the 'Coppery Creek group'. In most areas, strata of the 'La France Creek group' grade into thicker-bedded quartzite at the base of the Mount Nelson Formation.

The Mount Nelson Formation consists of a cliff-forming, basal unit of white, grey or green orthoquartzite with rare argillaceous laminae and partings, overlain by brownish red to grey-weathering impure carbonate interbedded with black, purple or red argillite and grey siltite. Stromatolites and lenses or nodules of chert occur locally within the carbonate unit. The basal orthoquartzite, up to 70 m thick, thins gradually to the south. Interbeds of green, black or red argillite are common within the upper quartzite unit and green and black argillite and siltite form the top of the preserved formation. The carbonate unit is thicker in western exposures, where it is overlain by interbedded black phyllite and grey siltite. Cream-weathering dark-coloured dolomite and brown-weathering, white dolomite, locally interbedded with black phyllite, occur at the top of the formation as preserved. Mud cracks in argillite, ripple marks in quartzite and solution-breccias in dolomite are locally common in both area.

The Mount Nelson Formation, whose maximum preserved thickness is about 1000 m is unconformably overlain by conglomerate of the Toby Formation of the Upper Proterozoic Windermere Supergroup. Evidence for small-scale, pre-Toby block faulting is found locally. Regionally, the unconformity cuts out progressively older Purcell strata southward along the western Purcell Mountains ".

The following has been modified from Pope (1990):

Van Creek Formation

The Van Creek Formation consists of coarse to medium-grained, light-grey or green to dark-green quartzites, siltstones and silty argillites. The beds have consistent thicknesses of between 20 to 50 centimetres with slightly undulose bases and truncated tops, together with internal cross and planar lamination and grading. Van

Creek quartzites grade upward into thinly bedded pale green quartzites and then into thinly interbedded 2 to 20 centimetre pale green quartzites, silts and buff weathering dolomitic silts of the Lower Gateway Formation, Hg 1 member.

Lower Gateway Formation

The Lower Gateway Formation is subdivided into two members Hgl and Hg2.

Hg 1: The contact between the Van Creek and Lower Gateway formations is gradational and in the absence of the Nicol Creek Formation can only be roughly estimated. The lowermost units of the Lower Gateway Formation are identified as where carbonate first occurs in the succession. The thin bedded quartzites in this transitional sequence are characterized by weathered pyrite, which imparts a distinctive red spotted appearance.

The Hgl member is estimated ... to be well in excess of 1000 metres thick. It consists of interbedded packages of quartzite, green siltstone and buff dolomitic siltstone and dolomite. Sedimentary structures such as cross lamination, grading, channelling and dewatering structures, are well preserved and compositional differences frequently enhance exposures. Siltstones in the dolomitic packages usually show an upwards gradation from dolomite free, finely cross-laminated silt and sand to dolomitic cross-laminated siltstone and cryptalgal to stromatolitic-laminated micritic dolomite. Bed thicknesses vary from generally 2 to 10 centimetres in the fine grained quartzite dominated lower part, to 10t o 50 centimetres in the upper dolomite dominated part of the Hg 1 member.

Hg2: The dolomite dominated upper part of the Hgl member passes into a 90-metres thick, cream to buffweathering dolomite unit. The dolomite displays cryptalgal and stromatolitic laminations, cream chert intercalations, rare halite casts and silty and sandy cross lamination. Bed thickness varies between 50 centimetres to 2 metres, and grain size varies from micrite, which is typically blue-grey, to coarse sucrose-textured, light coloured recrystallized dolomite.

Dutch Creek Formation

The boundary between the Lower Gateway Formation and the Dutch Creek Formation is characterized by a narrow zone of rusty weathering. The contact is interpreted as a parallel unconformity and the rusty weathering zone marking a hiatus.

Within the Dutch Creek Formation there is not a clearly defined stratigraphy, but four basic lithofacies (A to D) have been distinguished. Beds are usually between 2 to 20 centimetres thick and consist of fine grained quartzite and argillite in graded couplets.

Sedimentary structures include fine herringbone ripple and channel cross-laminations. The Dutch Creek Formation has a marked lack of carbonate.

Lithofacies A - Finely interlaminated green and dark grey to black graded siltstone-argillite couplets. Beds 1 - 10 cm thick.

Lithofacies B - Drab green to grey silt to fine sand quartzite and grey green to black silty argillite interbeds 5 - 20 cm thick.

Lithofacies C - Grey black argillite and siltstone with buff dolomitic siltstones.

Lithofacies D - Dark grey limestone and limey siltstone interbedded with argillite beds 10 cm to 1 m thick.

There is a great variation in thickness of the Dutch Creek Formation from an estimated 1000 metres to less than 300 metres over a lateral distance of 5 kilometres. Although the observed contact with the overlying Mount Nelson Formation is always paraconformable, the contact is very sharp and represents a major change in facies, hydrodynamic energy and sedimentary processes, and is therefore interpreted as an unconformity.

Mount Nelson Formation

The Mount Nelson Formation has been subdivided into the:

a) lower quartzite, a useful 50 to 150 metre thick marker horizon consisting of white, well-sorted, fine- to medium-grained pure quartz arenites,

b) lower main dolomite - an approximately 400 metre thick sequence which conformably overlies and is gradational with the lower quartzite, comprised of cryptalgal to stromatolitic laminated, pale grey weathering dolomites with interbedded carbonaceous argillites capped by a cream-coloured stromatolitic, crystalline cherty-dolomite unit approximately 20 metres thick overlain in sharp contact by,

c) the middle quartzite - an apple green coloured sequence consisting of massive, fine- to coarse-grained quartz arenites, impure sandstones and argillites having A-B to A-E Bouma sequences evident,

d) orange dolomite sequence - approximately 180 metres thick consisting of varicoloured buff weathering dolomitic siltstones, argillites and impure sandstones underlying bright orange-buff weathering silty and sandy crystalline dolomites with abundant cryptalgal and stromatolitic laminations and intercalated chert.

e) white markers conformably overlie the orange dolomite and are up to 70 metres thick. The white markers consist of cream, buff and silver-grey dolomites with purple, green and buff dolomitic mudstones and local interbeds of pure white magnesite up to 1 metre thick,

f) purple sequence - gradationally overlies the white markers, consisting of purple weathering dolomitic sandstones and siltstones which grade upward into purple weathering argillite. Mudchip breccias and monomict pebble conglomerates are interbedded with siltstones and argillites and the sequence is overlain by a pebble to boulder conglomerate with a purple weathering sandy argillitic matrix in sharp contact with the purple shales. The pebble to boulder conglomerate is the interpreted locus of an intraformational unconformity with a thickness between 2 and 10 metres thick,

g) upper middle dolomite - approximately 80 metres thick and similar to the lower main dolomite. It is distinguished by abundant algal allochems which are typically replaced by black chert,

h) upper quartzite - a distinctive cliff-forming unit consisting of white quartzites more than 260 metres thick (equivalent to the upper Mount Nelson Quartzite (Atkinson 1975)). The upper quartzite consists of well sorted medium- to coarse-grained, essentially pure arenites. They are distinguished from the lower quartzite on the basis of massive bedding and poorly preserved sedimentary structures.

i) upper dolomite - the uppermost unit in the Belt-Purcell exposed below the Windermere unconformity. The upper dolomite is gradational with the underlying quartzite over 10 metres consisting of interbedded purple argillite, quartzite and dolomite. The upper dolomite is comprised of pale to dark grey dolomite interbedded with quartz and dolomite pebble conglomerates with dolomitic quartz sands.

Windermere Supergroup

The Windermere Supergroup varies in thickness in the Toby Creek area, from 80 metres to over 3 kilometres and is in sharp contact with the underlying Belt-Purcell Supergroup across an unconformity with considerable topography, interpreted as a result of a local basement high, the "Windermere High" (Reesor 1973). The Windermere Supergroup was deposited above this unconformity and consists of a basal conglomeratic unit, the Toby Formation, and the overlying argillite and pebble conglomerate dominated Horsethief Creek Formation.

Toby Formation

The Toby Formation is the basal unit of the Windermere Supergroup and overlies different levels of the Belt-Purcell stratigraphy in the separate fault panels, interpreted to indicate active faulting during sedimentation (Pope 1990). Four distinct facies have been identified in the Toby Creek area but their stratigraphic position relative to one another is uncertain due to rapid lateral facies changes.

The Toby Formation consists of:

a) a basal boulder breccia lithofacies consisting of monomict clast-supported boulder breccias.

b) a diamictite lithofacies - the most commonly developed facies consisting of rounded quartzite and subangular dolomite boulders (derived from the immediately underlying Mount Nelson Formation) in a sandy argillite matrix.

c) a sparse clast diamictite lithofacies consisting of graded fine to coarse-grained, poorly sorted arenites and argillites with a minor component of rounded quartzite pebbles or cobbles.

d) a siltstone-argillite lithofacies which comprises the bulk of, and is the dominant lithology in, the upper portion of the Toby Formation, consisting of well-sorted and graded fine quartz arenites and argillites which typically exhibit complete Bouma sequences.

The Toby volcanics are the oldest igneous rocks identified in the Toby Creek area and are believed to be altered submarine basalts related to regional Hadrynian extension. The flows are holocrystalline and glomeroporphyritic basaltic andesites, having plagioclase phenocrysts in a fine-grained plagioclase groundmass.

Green metadiabase dykes have also been identified and have been interpreted as the metamorphic equivalent to the Toby volcanics. They are the most common igneous rocks and are always intruded at a high angle to bedding. They are typically altered, consisting of anhedral masses of chlorite, anhedral to euhedral carbonate and sericite and skeletal opaques. Chlorite pseudomorphs after pyroxene and amphibole have been identified. Bulk mineralogical proportions indicate these dykes were most probably originally basaltic in composition and have been subsequently hydrated.

Horsethief Creek Group

The Toby Formation is gradational into the overlying Horsethief Creek Formation, in which five lithofacies have been identified. These lithofacies define a rudimentary stratigraphy of facies within the Horsethief Creek Formation as individual lithological units are inconsistent due to rapid lateral thickness and facies variations.

The lithofacies identified in the Horsethief Creek Formation are as follows:

a) siltstone-argillite - dominant in the lower half of the Horsethief Creek Formation and separate the remaining lithofacies throughout the formation. This lithofacies consists of thick sequences of thin bedded (1 to 10 cm), graded siltstone and argillite and finely laminated (1 to 5 mm), black, green and grey argillite.

b) black carbonate - an easily traced marker used to identify and map the base of the Horsethief Creek Formation consisting of thin bedded (5 to 20 cm), dark grey to black limestone, with variable quartz sand and silt in a calcitic matrix, and thin calcareous quartz-arenite beds.

c) dolomite - buff weathering dolomite, up to 30 metres thick, dolomite pebble-conglomerate beds and dolomite supported quartzite occur throughout the Horsethief Creek Formation.

d) quartz feldspar arenites and pebble conglomerates - consist of pebble conglomerates comprised of grain-supported, moderately sorted crystalline quartz and quartz feldspar clasts with variable red jasper, green to grey argillite, quartzite and dolomite clasts in a quartz, feldspar, carbonate, sericite and chlorite matrix. Clasts are generally 1 to 2 centimetres in diameter but may exceed 10 centimetres in length. Coarse arenite beds are similar to the pebble conglomerates but have a greater proportion of matrix and are generally poorly sorted.

e) red and varicoloured argillites - are present at the top of the Horsethief Creek Formation and consist of variably coloured argillites with interbedded pink carbonate, and varicoloured impure arenites.

Mesozoic

Granitic Intrusions

Cretaceous intrusives of broadly "granitic" composition are present in a belt extending from the westernmost Rocky Mountains to Kootenay Lake, northward to the Baldy Batholith. Intrusions range from small dykes and sills to larger intrusive complexes such as the Mt. Skelly Batholith and are collectively referred to as the Bayonne Magmatic Belt (or Suite).

> "Intrusive rocks ... include a number of small post kinematic mesozonal quartz monzonite, monzonite and syenitic plutons, numerous small quartz monzonite to syenite dikes and sills probably related to these stocks, and late mafic dikes. The Kiakho and Reade Lake stocks, two of the larger of the mesozonal plutons, cut across and apparently seal two prominent east-trending faults that transect the eastern flank of the Purcell anticlinorium, and hence place constraints on the timing of latest movement on these faults.
>
> The Kiakho stock is exposed on the heavily wooded slopes of Kiakho Creek approximately 10 kilometres (west-southwest) ... of Cranbrook ... Exposures consist mainly of large, fresh angular boulders of boulder fields. Although contacts with country rock were not observed, regional mapping indicates that it intrudes clastic rocks of the Aldridge and Creston formations. The distribution of outcrops and a pronounced aeromagnetic anomaly indicate that it cuts the east-trending Cranbrook normal fault with no apparent offset. ...
>
> The Kiakho stock is similar to the Reade Lake stock with the dominant phase being a light grey, medium-grained quartz monzonite. It is generally equigranular but grades into a hypidiomorphic granular porphyritic phase with prominent plagioclase and light grey to flesh-coloured potassic feldspar phenocrysts; both are up to several centimetres in diameter in a granular groundmass of white subhedral plagioclase, light grey potassic feldspar, quartz and black hornblende" (Höy 1993).

The Bayonne Granitic Suite is a composite batholith comprised of a number of smaller Jurassic to Cretaceous age granitoid stocks and plutons which extends from near the International Boundary across Kootenay Lake. On the east side of the Kootenay Lake, the Bayonne Granitic Suite locally includes the Mount Skelly Pluton, a biotite (hornblende) monzogranite with megacrysts of potassium feldspar (Reesor 1996). Rice (1941) grouped these granitoids under the broad heading of the Bayonne Batholith, as described below.

> "The Bayonne batholith varies in composition from a granite to a calcic granodiorite; the average composition is that of a fairly alkaline granodiorite. ... Much of the rock has an equigranular texture, but a porphyritic phase occurs in many places, at some of which phenocrysts of potash feldspar 2 or 3 inches long are present. The potash feldspar may be orthoclase or microcline and in some specimens both occur. The plagioclase is oligoclase, generally well twinned and frequently in zoned crystals.

> Dark brown biotite is the only ferromagnesian mineral abundant, but grains of hornblende occur in rare instances. The usual accessories are present. Sericite and epidote are the commonest secondary minerals, but neither occur in significant amounts except where the rock has been altered.
>
> A marked feature of the Bayonne batholith is its highly variable nature. This is observable not only in the range of composition but in the appearance of the rock. Coarse-grained and fine-grained, porphyritic and non-porphyritic, pink and light or dark grey phases may occur in a single exposure, in some places in streaks and patches. Masses of pegmatite and dykes of pegmatite and aplite occur everywhere. Some of the pegmatite dykes are over 100 feet wide. A few large crystals of blue-green beryl, pink garnet, magnetite, and a little black tourmaline were seen in these pegmatites.
>
> Large inclusions of granitized sediments are locally abundant. ... These inclusions vary in size from a foot to some hundreds of feet. Alteration is severe, but the sedimentary nature of the original rock is, in most cases, still recognizable and the boundary between the granite and the inclusion is generally fairly sharp. Other inclusions or zenoliths (sic.) from a few inches to a foot long also occur, which can readily be distinguished from the first type mentioned. They parallel one another, are darker coloured, their original texture and composition has been more or less completely altered, they are fairly uniform in size, and they usually grade imperceptibly into the granite. They are more widely distributed, indeed very few exposures of any size were examined that did not contain some of these zenoliths (sic.), and in places they are extremely abundant. The zenoliths (sic.) are often most common in the porphyritic phases and scarcer in the non-porphyritic phases of the granite ...".

Cretaceous intrusions interpreted to underlie the properties comprising the Cretaceous Granitic Project are interpreted to be exemplified by the Mount Skelly Pluton, located approximately 23 km southwest of the Mount Rice property along the east shore of Kootenay Lake. Recently there has been limited mapping undertaken on the pluton as part of a regional study of the Bayonne Magmatic Belt (Logan 2002), with local sampling and mapping of the Mount Skelly Pluton and Sanca Stock (Lett et al. 2000, Logan and Mann 2000).

Mount Skelly Pluton / Sanca Stock

The dominant lithology comprising the Mount Skelly Pluton is that of a biotite granite. In areas proximal to the mapped contact between the pluton and host sediments, the grain size is slightly reduced to that of a medium- to coarse-grained granite. At low to middle elevations along the eastern portion of Sanca Creek, the granite assumes a porphyritic texture due to the presence of megacrystic alkali feldspar phenocrysts. Individual, equant crystals of white to pinkish alkali feldspar phenocrysts up to 2 cm in diameter were noted in a finer grained matrix of medium- to coarse-grained white plagioclase and biotite ± hornblende. Xenoliths are rare to absent at deeper levels within the pluton, becoming more abundant and larger both at higher elevations and along Sanca Creek to the west. Xenoliths are predominantly sedimentary, however, inclusions of finer grained, more mafic granite were noted and may have been derived from an earlier phase of the intrusion or a separate, deeper intrusion altogether.

Recent mapping and geochronology by Logan and Mann (2000) have resolved the granite exposures of the Sanca Creek area into three separate phases, specifically, the Mount Skelly Pluton and the Sanca Stock. The Mount Skelly Pluton is further sub-divided into:

1) Granite - "Fine to medium grained, equigranular biotite monzogranite. Minor aphanitic, leucocratic phases and dikes", and

2) Granodiorite - "Coarse grained biotite-hornblende granodiorite. Common euhedral megacrystic potassium feldspar and mafic (hornblende-biotite-titanite-rich) inclusions. Biotite, K-AR dates of 97.1 to 98.7 Ma

The younger Sanca Stock is described as a "Medium to coarse grained biotite granodiorite. Characteristic coarse, sub-rounded violet to grey quartz crystal aggregates. Biotite, K-Ar dates of 78.9 to 80.9 Ma". Therefore, the granites of the Sanca Creek area can be differentiated into three phases, the older Mount Skelly Pluton (at 97.1 to 98.7 Ma) and the younger Sanca Creek Stock (at 78.9 to 80.9 Ma).

Structure

Four major phases of deformation have been identified in the Toby Creek area, Helikian-Devonian extension (D1), Jurassic-Paleocene contraction (D2-D3) and Eocene extension (D4).

The first phase of deformation resulted in unconformities at the base of the Dutch Creek and Mount Nelson Formations (D1a) and the unconformity at the base of the Windermere Supergroup (D1b). Thinning of Paleozoic strata onto the Windermere High is interpreted to reflect the effects of D1c deformation together with the development of small fault-bounded sub-basins.

Contraction during the Columbian (D2) and Laramide (D3) orogenies resulted in a series of northeast vergent thrust faults and the development of a regional foliation (S1). Three major thrust sheets are evident in the Toby Creek area with one, the Mount Nelson thrust sheet, comprised of four smaller fault panels. The three major thrust sheets represent out-of-sequence faults, having propagated toward the hinterland, carried in the hanging wall of the Purcell Thrust.

Contraction during D2 and D3 produced east-vergent imbricate thrust faults and west vergent backthrusts. Many of these faults were subsequently reactivated during the fourth phase (D4) of deformation. High angle brittle faults are also a result of D4.

LOCAL GEOLOGY

Stratigraphy

The Mount Rice property is underlain by a south striking, moderately steeply west dipping panel of overturned Late Proterozoic age strata correlated to the uppermost Purcell Supergroupon the western limb of the Purcell Anticlinorium. Correlations interpret the strata as belonging to the upopermost Creston and Kitchener formations, with a fault slice coring the property (Fig. 4). In detail, displacement along the fault separating the two panels diminishes to the south, dying out into the Creston Formation.

Given the presence of the large Fry Creek Batholith approximately 8 km to the northwest, the Sawyer Creek Stock (Logan 2002; located approximately 6.5 km to the north-northwest) may be a small satellite intrusion. Alternatively, it may be a small unrelated Cretaceous intrusion, having been intruded during a regional Mesozoic intrusive event. Either way, a felsic intrusive has been documented on an adjacent property, suggesting that the magnetic anomaly evident underlying the Mount Rice property on the BC MapPlace web-site may represent a similar blind Cretaceous intrusive.

No geological mapping was undertaken on the property during the 2006 field season. As such, the author is not in a position to address possible stratigraphic correlations. The field data (soil sample) have been plotted on the digital geology for the property.

Structure

The structure of the Mount Rice property is dominated by its position on the western flank of the Purcell Anticlinorium, a north plunging fold of regional significance. The Purcell Anticlinorium is allochthonous with respect to North American cratonic basement, having been transported northeastward in the hanging wall of the Purcell Thrust. This major structure has been complicated slightly by a number of regional and local faults, discussed below with reference to the Kootenay Lake mapsheet of Reesor (1996). An early folding event has been proposed for early structures interpreted to have developed in the Late Proterozoic during the Goat River Orogeny (Höy 1993).

The prominent faults in the Mount Rice area are interpreted to be predominantly the result of the Laramide orogeny, characterized by east-verging, west-dipping thrust faults. The major fault system of the area is the St. Mary / Hall Lake fault system, interpreted to be a long lived fault initiated in the Late Proterozoic as a growth fault and periodically active at least into the Laramide orogeny. Eastward directed movement across the St. Mary / Hall Lake fault resulted in steeply dipping strata on the western limb of the Purcell Anticlinorium being juxtaposed against relatively shallowly to moderately dipping strata closer to the hinge axis.

Later thrust faults are evident in the hanging wall of the St. Mary / Hall Lake fault. The Redding Creek fault is locally significant fault. It is a west dipping, east verging thrust fault that juxtaposes middle Creston strata against the lower member of the Coppery Creek group. A number of smaller, normal faults are indicated in the hanging wall of the Redding Creek Fault, all of which appear to have minor dip (and probably strike-slip) movement. All of the faults in the hanging wall of the St. Mary / Hall Lake fault are interpreted to be older than the Cretaceous Mount Skelly Pluton (Bayonne Magmatic Belt) as all are truncated at the contact of the pluton.

2006 PROGRAM

A total of three days were spent collecting 164 soil samples on, and immediately adjacent to, the MOUNT RICE property (Rice West, Tenure 530476 and Rice East, Tenure 530483). Samples were collected along three contour lines with samples collected approximately every 50 metres (Fig. 6). The samples were taken from the “B Horizon” and placed in Kraft bags at the sample site. Sample depths ranged from 5 cm to 50 cm.

The samples were dried in Cranbrook, then shipped by Greyhound Courier to Acme Analytical Laboratories Ltd in Vancouver. Samples were analyzed using Acme’s Group SS80 package and analysis using the Group 1EX package.

The results from the soil samples recovered from the property are included in Appendix B. (Note: the Mount Rice program was completed together with the Sawyer program and, therefore, some of the results were reported separately under the Sawyer report although the sample results are mixed with those included in this report).

RESULTS

Soil Samples

To qualify the following review of the results of soil sampling, it must be remembered that the samples from the 2006 program represent a limited sub-set (only 164 soil samples). As such, only a limited review of initial analysis follows.

The property has been evaluated on a preliminary basis for copper, lead and zinc (below).

Copper

Copper analyses range between 1.4 and 81.4 ppm. Of the three lines sampled, the northernmost line along the northwest boundary of the property returned a numerous anomalous, adjacent copper results (Fig. 6). The eastern line returned a number of moderate anomalies and the southern line had a single highly anomalous result.

Lead

Lead analyses range between 5.0 and 156.5 ppm. Anomalous lead was returned from all three lines sampled (Fig. 7), with the southern and northern lines returning multiple, adjacent anomalous results.

Zinc

Zinc analyses range between 5 and 405 ppm. Zinc results are similar to those of lead (Fig. 8), however, the eastern line returned a higher proportion of moderately to highly anomalous, adjacent results.

Bismuth

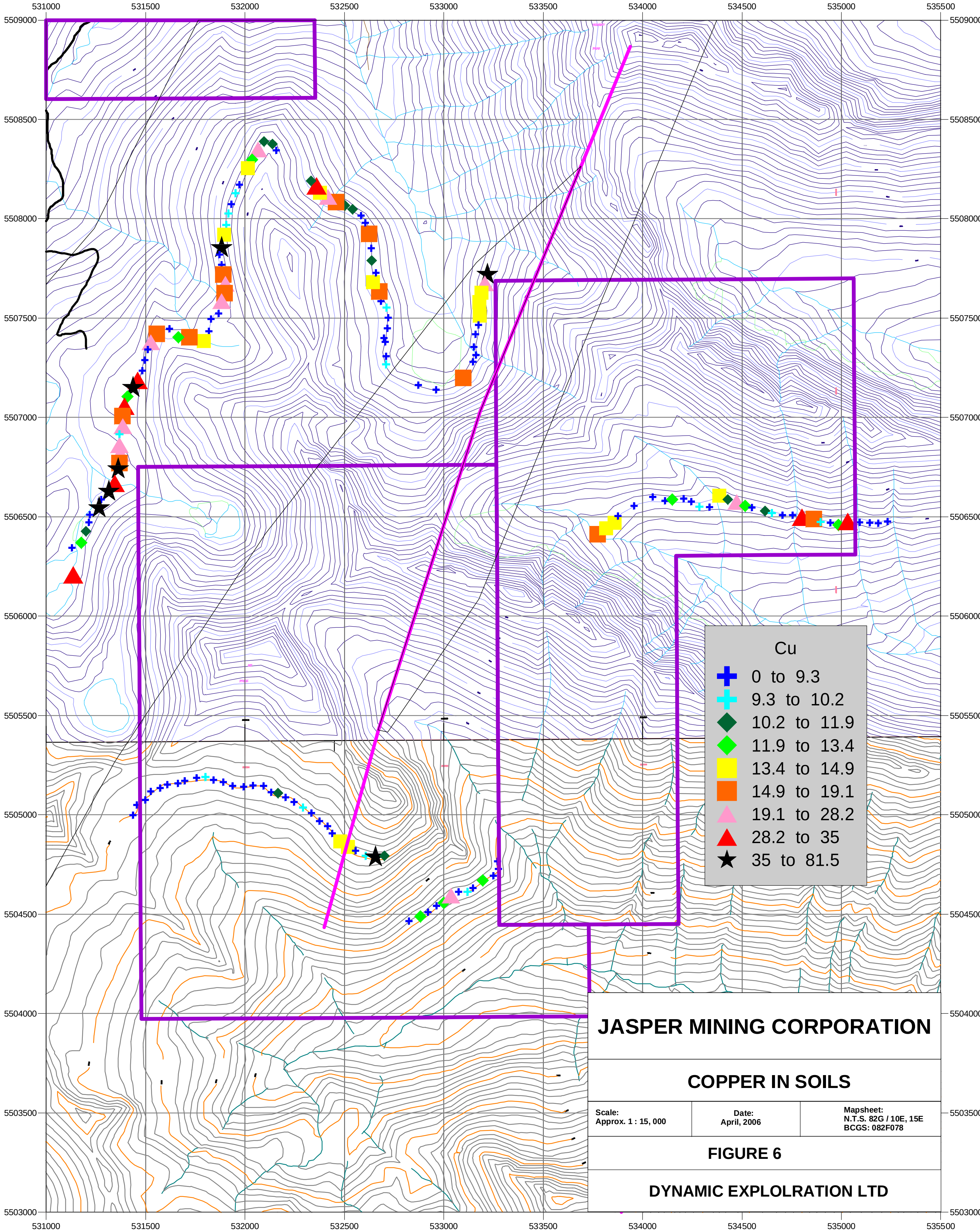
Bismuth values returned range between 0.2 and 1.4 ppm, with several multi-sample anomalies. The eastern line (Fig. 9) returned a number of moderately anomalous, adjacent results while the northern line returned moderately to highly anomalous, adjacent results.

DISCUSSION

One ongoing objective for evaluation of the properties comprising the Cretaceous Gold Project is the possibility for intrusion-related gold. As such, values for precious (Au, Ag) and base (Cu, Pb, Zn) metals are reviewed for potentially anomalous values, as well as those possibly indicative of intrusion-related mineralization (As, Bi, Sn and W). In addition, given proximity to Cretaceous intrusions (i.e. Fry Creek Batholith, Hall Lake Stock and Sawyer Stock) as well as the recently released Inferred Resource from Eagle Plains Resources' Sphinx property, immediately north of Grey Creek Pass, molybdenum is another metal of potential interest.

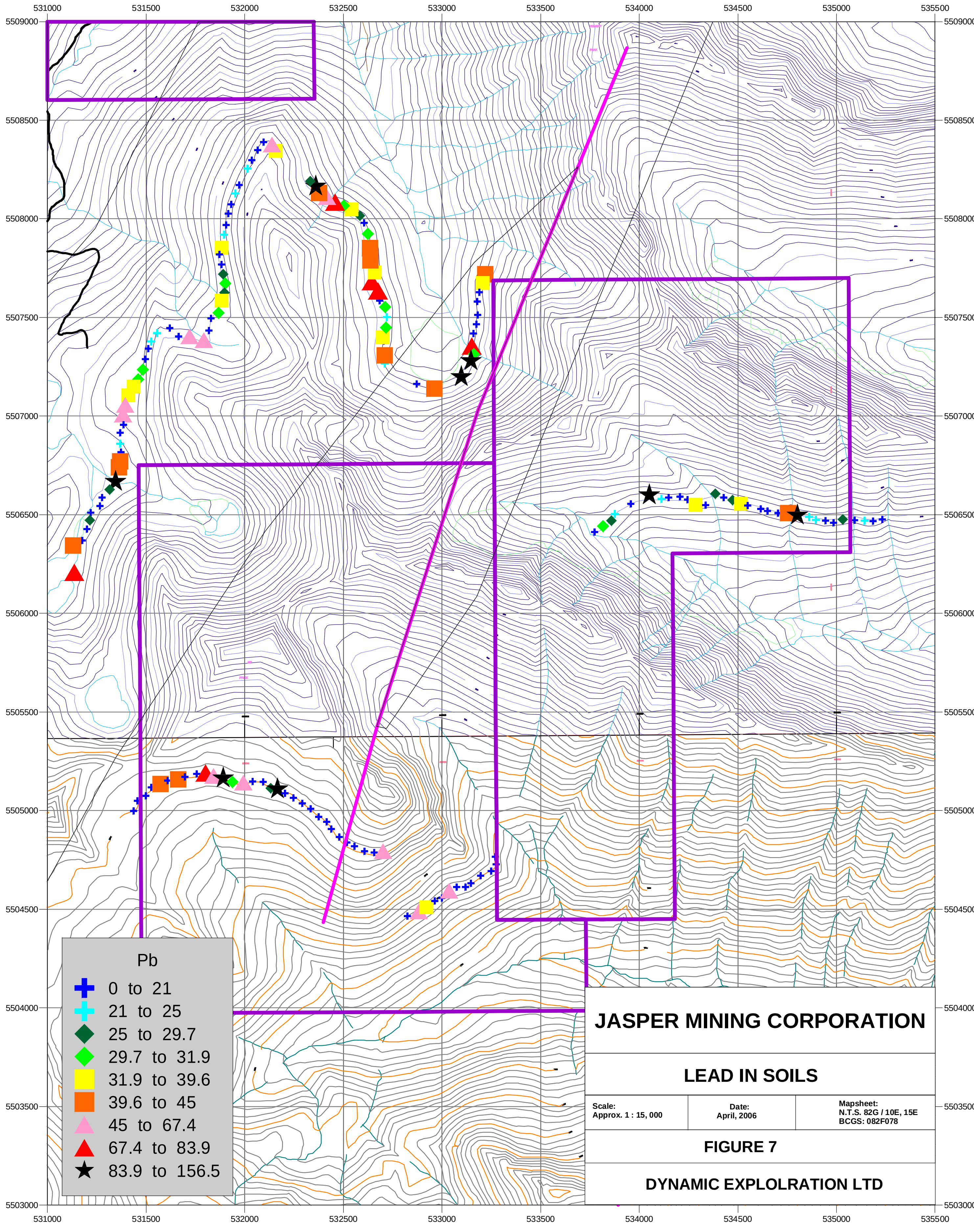
Obviously, it is difficult to reach any meaningful conclusions regarding such a large property on the basis of such a limited data set. However, there are several potentially interesting associations and trends suggested by the data, indicating that further work is warranted to assess multi-element, multi-sample anomalies along the few contour lines completed to date.

Mount Rice



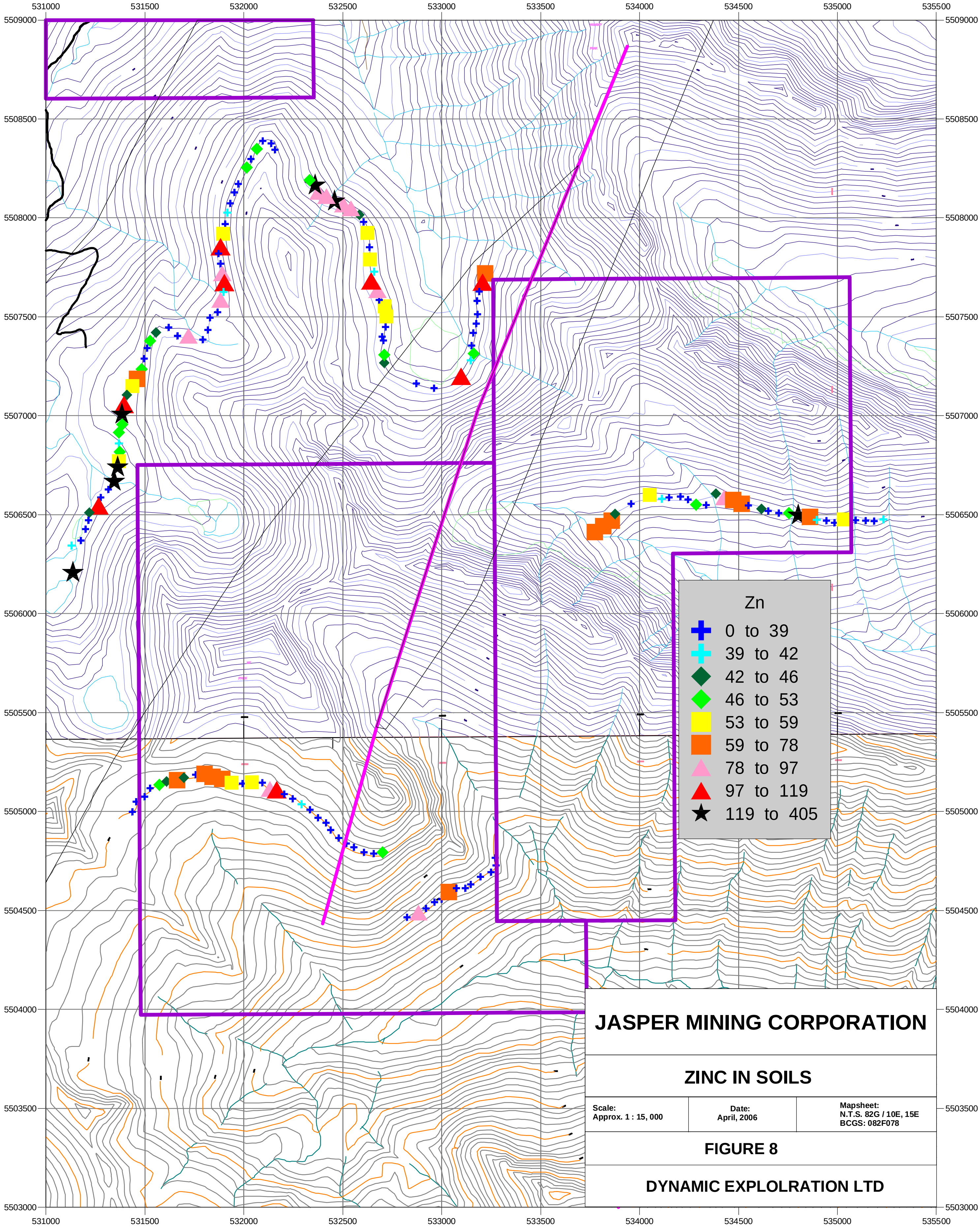
JASPER MINING CORPORATION		
COPPER IN SOILS		
Scale: Approx. 1 : 15, 000	Date: April, 2006	Mapsheet: N.T.S. 82G / 10E, 15E BCGS: 082F078
FIGURE 6		
DYNAMIC EXPLOLRATION LTD		

Mount Rice



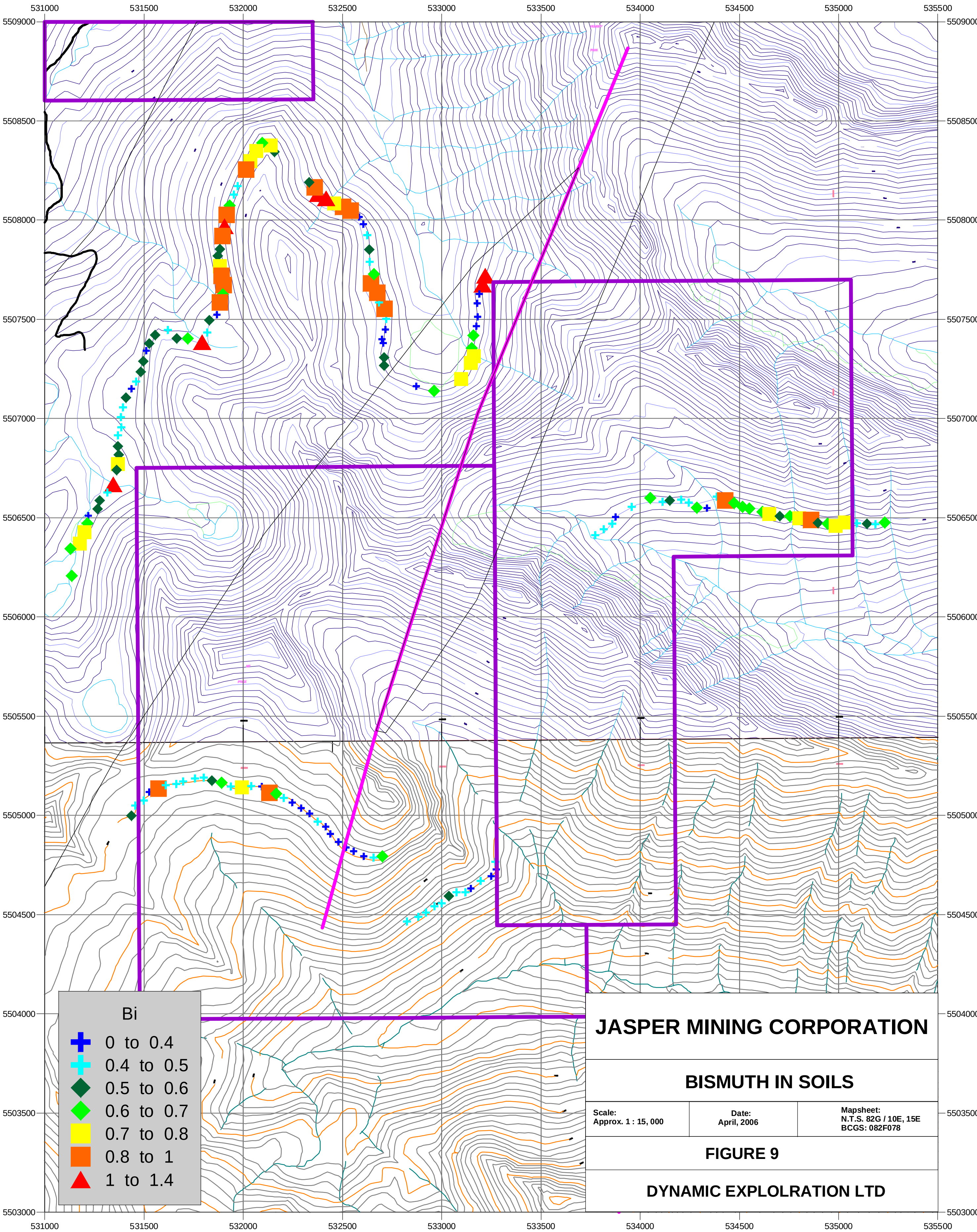
Pb
0 to 21
21 to 25
25 to 29.7
29.7 to 31.9
31.9 to 39.6
39.6 to 45
45 to 67.4
67.4 to 83.9
83.9 to 156.5
JASPER MINING CORPORATION
LEAD IN SOILS
Scale: Approx. 1 : 15, 000
Date: April, 2006
Mapsheet: N.T.S. 82G / 10E, 15E
BCGS: 082F078
FIGURE 7
DYNAMIC EXPLOLRATION LTD

Mount Rice



Zn
0 to 39
39 to 42
42 to 46
46 to 53
53 to 59
59 to 78
78 to 97
97 to 119
119 to 405
JASPER MINING CORPORATION
ZINC IN SOILS
Scale: Approx. 1 : 15, 000
Date: April, 2006
Mapsheet: N.T.S. 82G / 10E, 15E BCGS: 082F078
FIGURE 8
DYNAMIC EXPLOLRATION LTD

Mount Rice



JASPER MINING CORPORATION

BISMUTH IN SOILS

Scale: Approx. 1 : 15, 000	Date: April, 2006	Mapsheet: N.T.S. 82G / 10E, 15E BCGS: 082F078

FIGURE 9

DYNAMIC EXPLOLRATION LTD

The potential for intrusion-related and/or other magmatic related mineralization is suggested by:

1) the general association of molybdenum with Cretaceous intrusions of the Bayonne Magmatic Belt,
2) spatial association between silver-bearing to silver-rich base metal veins and documented intrusions (i.e. Perry Creek - Moyie River area, Rose Pass area (Welcome-Enterprise) and, in particular, the Sanca - Akokli Creek area), and
3) the documented presence of relatively small felsic intrusions in the general area (i.e. Hall Lake Stock, Sawyer Stock, Ailsa Lake, Mount Skelly Complex, Fry Creek Batholith, etc), and

On the basis of mapping by Reesor (1996), the Toby Formation is present in the footwall of the Redding Creek Fault south of the Sawyer Creek Stock (Sawyer property 6.5 km northwest) and has been cut-out along trend to the north (i.e. north of the Sawyer Creek Stock). Bapty and MacLachlan (1996) described a pod and/or lens of massive galena, sphalerite and tetrahedrite in one of the chutes on the south side of a small tributary into Sawyer Creek. Previous mapping (Bapty and Walker 1994) located evidence of similar mineralization on a ridge top traverse and descent through the north facing cliffs.

Geochemical evidence for mineralization adjacent to the Sawyer Creek Stock, located in the footwall of the Redding Creek Fault, suggests additional pods, or perhaps larger bodies of galena+sphalerite +tetrahedrite mineralization may be present. The mineralization was described as being "... a small limestone-replacement deposit with a sedimentary-exhalative origin" (Bapty and MacLachlan 1996). The author believes a replacement (manto) style deposit is possible, given theabundance of carbonate-bearing intervals (limestone and dolomite) in the Kitchener Formation on the Mount Rice property. Intrusion of the Sawyer Creek Stock along the Redding Creek Fault (given that Reesor (1996) does not extend the fault trace through the stock, indicating the intrusion post dates the fault), would provide a proximal heat source, both to dissolve carbonate-bearing intervals and provide possible metalliferous fluids to precipitate within the resulting void space along, and adjacent to, the Redding Creek Fault.

The prominent anomaly on the First Vertical Derivative map (Fig. 10) is localized along the surface trace of both the geological contact between the Creston and Kitchener formations and the small fault evident on the property. The First Vertical Derivative (FMD) map is a filter of the magnetic data intended to emphasize magnetic changes in the magnetic data at "shallow" levels. The Creston Formation has a strong regional magnetic signature and, therefore, the FMD map is interpreted to reflect the contact between the siliciclastic dominated Creston Formation and the carbonate-dominant Kitchener Formation. However, there are a series of three elongate to elliptical anomalies evident within the Kitchener Formation between the Sawyer and Mount Rice properties which may represent near surface (blind) Cretaceous instrusions similar to the Sawyer Stock.

This interpretation may be supported by the geochemical results returned from the Mount Rice

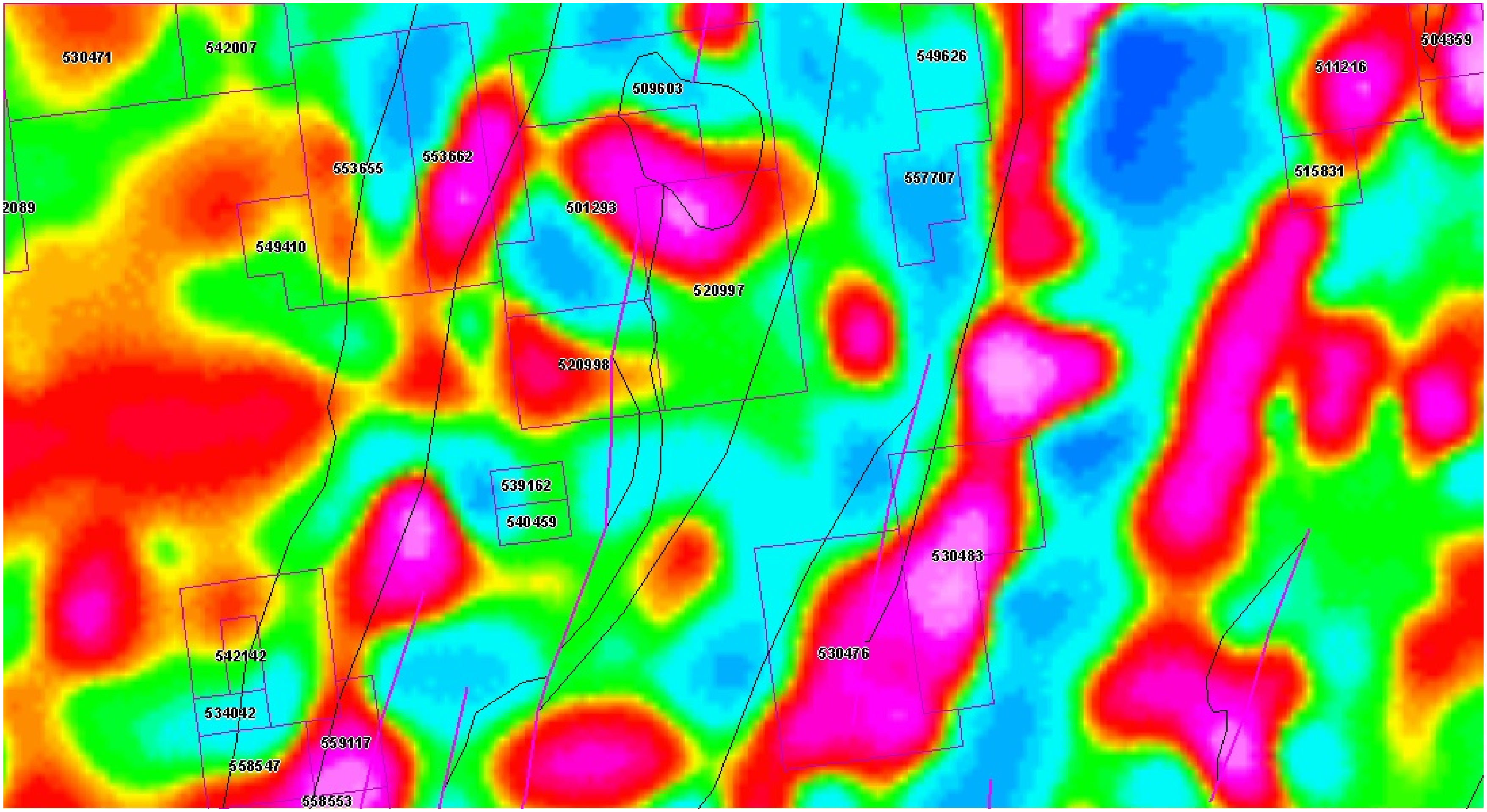


Figure 10: First Vertical Derivative Geophysical Map of the area surrounding the Mount Rice property (claims in lower right). Note: association of anomaly associated with Sawyer Stock (Sawyer claims north-central portion of figure). Taken from MapPlace, approximate scale 1:75,000)

program in that copper (generally a higher temperature metal) is moderately to highly anomalous along the northern line, in the highest class having values to 81.5 ppm. In addition, bismuth is "highly" anomalous along the northern line, having values to 1.4 ppm, relative to the other two lines. This line is approximately 1.28 km from the centre of the prominent, though small, anomaly on the FMD map immediately north of the property.

Other metals that might provide insight regarding intrusion-related gold potential (Ag, Au, Sn and W) did not return any apparently meaningful anomalies to pursue with the exception of arsenic. Arsenic data returned strongly anomalous data on the northern line and at the eastern end of the eastern line. The anomalous arsenic data (though relatively low; maximum of 49.21 ppm) supports the anomalous copper on the northern line and may, therefore, indicate a magmatic component to the anomalous metal results returned.

The data may also indicate a spatial association with stratigraphy in that the metal data appears to have a relative low associated with the contact between the Kitchener and Creston formations. This is difficult to evaluate as the stratigraphic contact between the Creston and Kitchener formations, as well as the fault contact, is plotted as a steeply dipping contact with no apparent influence from topography. Further mapping would be required to confirm the nature of this contact, with has been projected from adjacent areas, and therefore allow an evaluation of any spatial association between the soil results and either contact.

CONCLUSIONS

The only realistic conclusion that can arise from such a small program is that further work needs to be done in order to evaluate the property. Additional soils need to be collected from throughout the property so as to provide meaningful coverage of the property and a reasonably sized data-set from which to identify, and quantify, potentially anomalous results.

Further work should be undertaken to assess the possibility for a magmatic contribution from possible blind Cretaceous intrusions north and west of the property, as proposed herein.

RECOMMENDATIONS

1. Continue the soil sampling program. Additional sampling should include acquisition of further contour samples. Samples should be taken along a series of contours throughout the property to provide coarse coverage of the property, with smaller grids established to develop better resolution in areas of anomalous results, possible in the area of the northern and western soil lines;

2. Creeks draining the property should be silt sampled.

3. Prospecting to identify any mineralization exposed on the property should be considered, with geological mapping of any outcrop identified completed so as to address correlation of the strata exposed and identify any intrusive lithologies.

4. A short duration helicopter-supported camp should be considered so as to undertake high elevation soil sampling, prospecting, geological mapping and, if warranted, rock sampling. In particular, an attempt should be made to map the surface trace of the stratigraphic contact between the Creston / Kitchener formations as well as the fault coring the property.

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Appendix A

Statement of Qualifications

STATEMENT OF QUALIFICATIONS

I, Richard T. Walker, of 656 Brookview Crescent, Cranbrook, BC, hereby certify that:

1) I am a graduate of the University of Calgary of Calgary, Alberta, having obtained a Bachelors of Science in 1986.

2) I obtained a Masters of Geology at the University of Calgary of Calgary, Alberta in 1989.

3) I am a member of good standing with the Association of Professional Engineers and Geoscientists of the Province of British Columbia.

4) I am the Vice President - Exploration for Jasper Mining Corporation, with an office at 2601 42nd Avenue, Crescent, Cranbrook, British Columbia.

5) I am the author of this report which is based on work completed under my supervision between November 9 and 11, 2007.

6) I was personally involved in the acquisition of the claims described herein.

Dated at Cranbrook, British Columbia this 22nd day of June, 2007

Richard T. Walker, P.Geo.

Appendix B

Soil Sample Results

DISTANCE	EASTING	NORTHING	Mo	Cu	Pb	Zn	Ag	Ni	Co	Mn	Fe	As	U	Au	Th	Sr	Cd	Sb	Bi	V	Ca	P	La	Cr	Mg	Ba	Ti	B	Al	Na	K	W	Hg	Sc	Tl	S	Ga	Se	Te	Sn	Zr	Sample
MR-8 48+50N	533219	5507720	1.8	36.4	42.4	65	0.1	11.3	5	137	2.59	11.2	2.4	0.7	3.6	6	0.3	0.7	1.3	14	0.04	0.067	34	10	0.41	20	0.012	<1	1.23	0.005	0.06	0.2	0.04	0.9	0.2	0.08	4	0.7	<1	<1	0.1	15
MR-8 48+00N	533206	5507674	2.8	21.8	32.5	97	0.2	23.4	26.4	902	3.02	31.2	2.3	0.7	2	7	0.4	1	1.3	22	0.07	0.094	21	20	0.41	25	0.025	1	1.72	0.007	0.08	0.4	0.05	0.8	0.2	0.1	7	0.8	<1	1	1.5	15
MR-8 47+50N	533189	5507627	1	14.4	13.8	23	<.1	5.1	3.8	563	1.46	2.6	1.7	0.5	2.5	2	0.1	0.2	0.3	10	0.02	0.034	15	6	0.15	49	0.013	<1	1.09	0.003	0.06	0.1	0.04	0.7	0.1	<.05	4	<.5	<1	<1	0.4	15
MR-8 47+00N	533178	5507580	0.9	14.4	14.8	24	0.1	5.9	3.9	364	1.48	2.8	1.7	1.8	2.6	2	0.1	0.2	0.3	12	0.03	0.04	16	6	0.16	48	0.02	2	1.1	0.004	0.07	0.1	0.04	0.7	0.1	0.06	4	<.5	<1	<1	0.7	7.5
MR-8 46+50N	533181	5507512	0.8	14.2	14.9	23	<.1	5.3	4	340	1.4	2.7	1.6	0.8	3	2	0.1	0.2	0.3	10	0.03	0.038	13	6	0.16	42	0.011	1	1.02	0.002	0.06	0.1	0.03	0.6	0.1	<.05	3	<.5	<1	<1	0.8	15
MR-8 46+00N	533175	5507467	0.2	1.6	5.2	5	<.1	1.1	0.6	40	0.2	0.9	0.2	<.5	2.3	1	<.1	0.2	0.2	5	0.01	0.019	16	2	0.03	15	0.004	1	0.46	0.003	0.03	0.1	0.01	0.3	0.1	<.05	3	<.5	<1	<1	0.1	15
MR-8 45+50N	533160	5507418	1	7.3	17	24	<.1	3.7	3.2	1029	1.06	4.2	0.5	<.5	0.5	5	0.1	0.4	0.6	16	0.13	0.044	14	5	0.09	49	0.018	1	0.45	0.005	0.06	0.1	0.02	0.4	0.1	<.05	4	<.5	<1	<1	0.2	15
MR-8 45+00N	533152	5507355	0.6	6.7	67.4	37	0.1	5.7	10.8	2228	1.31	6	2.5	0.5	0.2	7	0.6	1.2	0.6	18	0.15	0.115	7	6	0.16	89	0.014	2	0.78	0.007	0.09	0.1	0.07	0.3	0.2	0.13	4	<.5	<1	1	0.2	15
MR-8 44+50N	533162	5507315	0.8	5.6	29.8	49	<.1	7	3.7	631	1.45	4.7	0.9	1.4	0.6	7	0.4	0.6	0.7	21	0.12	0.065	11	7	0.22	85	0.032	1	0.96	0.005	0.1	0.1	0.06	0.5	0.2	<.05	6	<.5	<1	1	1.1	15
MR-8 44+00N	533147	5507280	0.6	8.5	83.9	40	0.1	8.4	17.3	1801	1.33	5.7	1.9	0.8	1.1	5	0.3	1.2	0.7	11	0.08	0.09	11	6	0.25	69	0.016	1	1.04	0.004	0.09	0.2	0.08	0.5	0.2	0.08	2	<.5	<1	1	0.2	15
43+50	533130	5507243																																								
MR-8 43+00N	533099	5507200	1.4	16.3	111.9	100	0.2	15.8	7.5	1109	2.17	11.3	5.2	1.7	2.4	9	1	1	0.7	14	0.26	0.119	12	14	0.44	49	0.014	1	0.83	0.005	0.06	0.3	0.06	1.1	0.1	<.05	3	0.5	<1	1	0.8	15
42+50	533060	5507144																																								
42+00																																										
MR-8 41+50N	532962	5507139	0.5	4.2	44.3	24	<.1	5.7	2.3	32	1.16	4.7	0.6	<.5	6	2	0.3	0.6	0.6	15	0.02	0.024	18	5	0.16	30	0.03	1	0.86	0.006	0.05	0.1	0.05	0.7	0.1	<.05	5	<.5	<1	1	4	15
41+00	532912	5507147																																								
MR-8 40+50N	532871	5507164	0.1	1.4	9.8	11	0.2	0.9	0.5	15	0.16	1.1	0.3	<.5	1.7	3	0.3	0.2	0.2	5	0.02	0.028	21	2	0.02	18	0.008	1	0.35	0.007	0.05	<.1	0.03	0.4	0.1	<.05	3	<.5	<1	<1	0.7	7.5
40+00	532821	5507182																																								
39+50	532768	5507196																																								
39+00	532740	5507204																																								
MR-8 38+50N	532710	5507268	1.1	9.7	23.3	43	<.1	7.6	6.6	457	2.08	3.2	1	0.7	1.9	3	0.2	0.4	0.5	18	0.07	0.05	10	7	0.32	67	0.029	1	1.54	0.007	0.05	0.2	0.05	0.9	0.1	<.05	6	<.5	<1	1	2.6	15
MR-8 38+00N	532710	5507308	0.6	8.8	39.6	52	<.1	7.2	7.5	1786	1.39	3	0.6	0.5	0.7	7	0.5	0.4	0.5	15	0.17	0.055	13	7	0.47	129	0.015	1	1.02	0.005	0.07	0.1	0.03	0.6	0.1	<.05	5	<.5	<1	1	0.2	15
MR-8 37+50N	532705	5507381	0.4	3.8	10.1	29	<.1	6.2	2.2	321	1.21	1.7	0.3	0.9	1.7	2	0.1	0.2	0.2	14	0.05	0.034	11	8	1.08	59	0.006	<1	1.18	0.003	0.04	0.1	0.02	0.7	0.1	<.05	5	<.5	<1	<1	0.1	15
MR-8 37+00N	532700	5507400	0.7	6	36	33	0.1	8.1	2.3	180	2.62	2.2	0.8	<.5	3.3	3	0.2	0.2	0.3	26	0.03	0.041	7	12	1.68	42	0.044	1	2.31	0.006	0.03	0.2	0.05	1.5	0.1	<.05	10	<.5	<1	1	3.6	15
MR-8 36+50N	532716	5507448	0.5	6.3	30.9	33	<.1	6.2	2.8	520	1.62	3.4	0.6	<.5	1.3	2	0.1	0.4	0.3	18	0.02	0.034	9	8	0.89	45	0.02	<1	1.5	0.004	0.03	0.1	0.03	0.8	0.1	<.05	7	<.5	<1	1	0.7	15
MR-8 36+00N	532720	5507501	0.6	6.7	22.3	54	<.1	5.5	4.4	1830	1.63	4	0.4	<.5	0.7	4	0.3	0.4	0.4	16	0.06	0.041	13	6	0.33	169	0.016	1	0.86	0.005	0.06	0.1	0.02	0.6	0.1	<.05	5	<.5	<1	<1	0.2	15
MR-8 35+50N	532711	5507554	0.8	10	29.9	56	0.1	5.6	6.8	665	2.08	6.2	0.7	<.5	1.2	4	0.3	0.3	0.8	21	0.06	0.039	11	6	0.13	111	0.026	<1	1.71	0.007	0.05	0.2	0.06	0.8	0.1	<.05	8	<.5	<1	1	1.7	15
MR-8 35+00N	532684	5507585	0.6	6.5	18.1	35	<.1	3.9	3.2	387	1.21	5	0.3	<.5	1.9	4	0.2	0.2	0.4	19	0.08	0.033	19	4	0.08	74	0.035	1	0.51	0.006	0.04	0.1	0.01	0.5	0.1	<.05	6	<.5	<1	1	0.5	15
MR-8 34+50N	532676	5507635	0.9	14.9	74.4	78	0.1	5.5	7.1	1060	2.27	3.4	1.2	0.5	0.8	7	0.6	0.3	0.9	18	0.17	0.063	12	6	0.12	250	0.012	1	1.24	0.005	0.06	0.1	0.03	0.7	0.1	<.05	5	<.5	<1	1	1.1	15
MR-8 34+00N	532643	5507682	1.2	14.6	68.2	116	0.2	7.1	8.3	2353	2.12	6.2	1.5	<.5	0.7	10	1.6	0.8	0.8	20	0.27	0.057	7	6	0.13	430	0.03	1	1.03	0.007	0.05	0.2	0.06	0.8	0.1	<.05	6	<.5	<1	1	1.9	15
MR-8 33+50N	532658	5507729	0.6	7.4	33.3	41	0.3	4.3	2.7	170	1.4	4.9	0.5	<.5	1.1	2	0.2	0.7	0.6	17	0.01	0.043	13	5	0.12	48	0.032	1	0.75	0.005	0.06	0.1	0.04	0.6	0.2	<.05	6	<.5	<1	1	1.1	15
MR-8 33+00N	532638	5507790	0.7	10.9	39.9	55	1	5.5	5.1	1134	1.73	5.1	0.7	2.2	1.8	3	0.4	1	0.4	19	0.02	0.108	7	6	0.13	91	0.059	1	2.19	0.009	0.05	0.2	0.1	1.1	0.1	<.05	7	<.5	<1	1	8.2	15
MR-8 32+50N	532635	5507851	0.4	6	39.9	38	0.2	3.4	2.2	1090	1.05	5.4	0.3	<.5	1.7	4	0.2	0.5	0.5	18	0.03	0.031	11	5	0.08	111	0.04	1	0.75	0.008	0.04	0.1	0.03	0.8	0.2	<.05	7	<.5	<1	1	1.3	15
32+00	532635	5507879																																								
MR-8 31+50N	532625	5507924	0.7	15.9	30.2	57	<.1	10.6	6.8	322	2.82	6.9	0.6	<.5	3.2	2	0.1	0.7	0.4	14	0.02	0.066	12	7	0.27	67	0.028	1	1.2	0.004	0.08	0.1	0.04	1.2	0.2	<.05	4	<.5	<1	<1	3.3	15
MR-8 31+00N	532605	5507979	0.2	3.8	7	14	<.1	1.6	1.6	32	0.37	0.8	0.2	<.5	3.1	2	<.1	0.2	0.2	8	0.01	0.009	13	2	0.05	93	0.017	1	0.5	0.004	0.04	0.1	<.01	0.6	0.3	<.05	5	<.5	<1	<1	0.6	15
MR-8 30+50N	532583	5508017	0.6	7.3	25.8	43	<.1	7.7	5	591	2.59	3.5	0.5	<.5	1.6	4	0.2	0.9	0.3	20	0.06	0.043	11	6	0.22	185	0.036	1	0.89	0.005	0.1	0.1	0.04	1	0.2	<.05	5	<.5	<1	1	1.3	15
MR-8 30+00N	532541	5508047	0.8	11.7	31.9	79	0.1	8.7	6.9	2486	2.37	7.3	0.4	<.5	1.2	6	0.5	1	0.8	23	0.1	0.053	8	8	0.27	204	0.067	1	1.12	0.009	0.1	0.2	0.06	0.9	0.3	<.05	7	<.5	<1	1	1.9	15
MR-8 29+50N	532502	5508068	1.6	11	29.7	86	0.1	8.6	9.1	2020	3.39	4.9	0.5	<.5	2.7	4	0.3	1.1	0.8	27	0.07	0.033	11	8	0.19	126	0.048	1	0.76	0.005	0.06	0.2	0.03	1.1	0.2	<.05	7	<.5	<1	1	1.5	15
MR-8 29+00N	532458	5508085	1	18.3	76.6	119	0.3	16.1	11.4	860	3.06	15.3	0.8	<.5	2.7	3	0.6	3.3	0.7	17	0.08	0.048	12	5	0.13	141	0.015	1	0.62	0.005	0.09	0.3	0.02	1	0.1	<.05	3	<.5	<1	1	0.5	15
MR-8 28+50N	532417	5508110	2.2	20.5	56.4	88	0.2	17.2	16.4	1196	2.92	37.3	12.3	1.2	2	12	1.2	0.7	1.1	15	0.53	0.071	17	7	0.26	47	0.019	2	1.05	0.008	0.07	0.3	0.06	1	0.1	0.07	4	1.1	<1	1	1.2	15
MR-8 28+00N	532377	5508131	3.5	13.6	43.9	93	<.1	8.1	5.8	849	2.2	15.2	1.1	0.9	2.7	6	0.6	1	1.1	20	0.14	0.05	19	7	0.19	99	0.038	2	0.86	0.007	0.06	0.2	0.07	0.9	0.2	<.05	7	<.5	<1	1	0.6	7.5
MR-8 27+50N	532360	5508165	1.8	32.6	89.1	404	0.4	30.4	15.2	1326	3.29	31.2	8.6	1.2	5.2	11	1.9	1.6	0.8	21	0.38	0.062	36	18	0.36	285	0.046	1	2.21	0.008	0.07	0.3	0.1	2.5	0.2	0.06	5	0.6	<1	1	4.2	15
MR-8 27+00N	532332	5508192	1.1	11.3	25.3	48	<.1	12.2	9.2	233	2.89	3.8	0.8	0.9	7.7	4	0.1	0.4	0.5	30	0.03	0.047	10	13	0.74	68	0.084	1	3.02	0.007	0.05	0.2	0.04	1.9	0.2	<.05	8	<.5	<1	1	28.2	15
MR-8 26+50N	532157	5508345	0.8	9.2	35.9	26	0.4	4.7	3.3	551	1.31	7.7	0.5	1	1.1	3	0.2	0.4	0.5	19	0.01	0.037	16	5	0.12	33	0.032	2	0.94	0.007	0.05	0.2	0.03	0.5	0.2	<.05	7	<.5	<1	1	0.8	15
MR-8 26+00N	532138	5508377	1	10.7	52.1	34	0.2	5.4	2.3	62	1.86	9.5	0.5	<.5	3.2	4	0.1	0.3	0.7	28	0.02	0.052	15	5	0.11	17	0.048	2	0.84	0.01	0.05	0.2	0.02	0.5	0.1	<.05	8	<.5	<1	1	1.6	15
MR-8 25+50N	532095	5508389	0.8	11.7	9.1	21	0.1	6.3	2.7	104	1.41	25.8	0.6	2.1	2	3	<.1	0.2	0.6	18	0.01	0.026	23	7	0.13	29	0.015	1	1.04	0.005	0.05	0.1	0.02	0.5	0.2	<.05	6	<.5	<1	1	0.5	15
MR-8 25+00N	532065	5508349	1.5	20.8	13.2	47	0.4	11.6	5.4	344	3.18	21.5	1.2	1.1	5.3	4	0.1	0.8	0.7	23	0.01	0.061	18	9	0.31	31	0.021	<1	1.52	0.005	0.08	0.2	0.05	1.1	0.2	<.05	6	<.5	<1	1	2.4	15
MR-8 24+50N	532037	5508298	1.1	13.3	16.6	37	<.1	8.6	3.8	140	2.56	12.9	0.6	1	5.7	4	0.1	0.6	0.7	30	0.01	0.034	18	8	0.24	32	0.051	1	1.12	0.007	0.06	0.2	0.03	1	0.2	<.05	8	<.5	<1	1	2.9	15
MR-8 24+00N	532015	5508254	1	14.5	21.1	51	0.4	8.1	5.4	334	2.3	16.4	0.8	2.2	4.6	4	0.1	0.6	0.8	25	0.03	0.074	11	8	0.22	48	0.069	1	2.21	0.008	0.07	0.3	0.05	1.4	0.2	<.05	9	<.5	<1	1	15.1	15
23+50	531999	5508208																																								
MR-8 23+00N	531972	5508172	0.6	7	16.8	33	0.2	4.8	2.9	225	2.06	8.6	0.4	1.1	1.6	3	0.1	0.5	0.4	23	0.02	0.058	8	6	0.15	60	0.035	1	1.96	0.008	0.04	0.2	0.06	1.1	0.1	<.05	7	<.5	<1	1	5.5	15
MR-8 22+50N	531952	5508129	0.9	10	22.9	32	0.3	5.8	3.8	175	2.01	12.9	0.8	3.1	3.7	7	0.1	0.3	0.4	24	0.05	0.067	8	6	0.13	39	0.062	1	2.38	0.011	0.04	0.2	0.07	1.3	0.1	<.05	8	<.5	<1	1	16.8	15
MR-8 22+00N	531932	5508073	0.5	7	8.9	13	0.2	2.5	1.1	95	1.05	6.9	0.5	1.3	0.9	4	0.2	0.2	0.6	18	0.04	0.048	15	4	0.06	42	0.021	<1	0.88	0.009	0.04	0.1	0.05	0.5	0.1	<.05	6	<.5	<1	1	0.6	15
MR-8 21+50N	531917	5508027	0.8	10	15.8	40	0.1	9.9	8.1	309	2.95	22.3	0.3	<.5	2.6	6	0.2	0.6	0.8	78	0.05	0.066	14	11	0.36	38	0.111	1	1.05	0.008	0.08	0.2	0.03	2.3	0.2	<.05	12	<.5	<1	1	2.5	15
MR-8 21+00N	531906	5507969	1.4	9.9	15.5	27	0.2	6.8	4.5	246	2.53	15.9	0.6	0.9	5.7	3	<.1	0.5	1.1	25	0.01	0.036	18	6	0.15	31	0.024	<1	0.98	0.004	0.05	0.2	0.03	0.7	0.1	<.05	7	<.5	<1	1	2.2	15
MR-8 20+50N	531895	5507920	1.3	14.6	23.9	54	0.1	11.4	6.8	245	3.05	15.2	0.8	1.2	6.3	3	0.1	0.5	0.8	46	0.02	0.032	21	11	0.36	37	0.039	1	1.43	0.004	0.08	0.2	0.03	1.9	0.2	<.05	8	<.5	<1	1	1.6	15
MR-8 20+00N	531882	5507853	1.3	42.3	32.5	99	<.1	21.6	19.2	345	6.04	13.9	0.6	1.1	5.2	6	0.1	0.7	0.5	111	0.1	0.072	12	22	1.06	55	0.13	1	2.47	0.005	0.22	0.2	0.03	5.9	0.3	<.05	10	<.5	<1	1	3.1	15
MR-8 19+50N	531872	5507819	1.1	7.4	12	28	<.1	6.8	3.5	98	1.86	10.7	0.5	1.2	4.8	4	0.1	0.5	0.5	37	0.03	0.025	23	7	0.22	34	0.046	1	1.08	0.005	0.05	0.2	0.02	1	0.1	<.05	10	<.5	<1	1	1.1	15
MR-8 19+00N	531883	5507769	0.9	7.1	14.8	29	0.1	6.7	3	127	2.13	12.4	0.6	2.9	2.7	3	<.1	0.3	0.7	25	0.03	0.029	19	7	0.21	40	0.019	<1	1.33	0.005	0.08	0.2	0.02	0.9	0.1	<.05	8	<.5	<1	1	0.6	15
MR-8 18+50N	531890	5507720	1.2	15.3	28.5	88	0.2	10.7	6.9	366	2.6	22.9	0.7	1.5	4.3	5	0.2	0.4	0.9	21	0.06	0.04	20	7	0.15	70	0.024	<1	1.27	0.006	0.06	0.2	0.03	0.8	0.1	<.05	7	<.5	<1	1	1.1	15
MR-8 18+00N	531902	5507672	2	27.4	30.6	99	0.5	17.2	11.5	680	3.12	29.8	1	0.5	3.8	8	0.5	0.5	0.8	23	0.1	0.043	18	9	0.29	184	0.019	1	1.21	0.006	0.08	0.1	0.03	1	0.1	<.05	7	<.5	<1	1	0.6	15
MR-8 17+50N	531898	5507625	1.1	15.9	26.5	40	0.2	9.6	5.7	120	2.13	8.1	0.7	0.6	4.1	2	0.1	0.5	0.6	20	0.02	0.027	20	6	0.18	38	0.018	1	0.95	0.004	0.05	0.1	0.02	0.8	0.2	<.05	5	<.5	<1	<1	0.4	15
MR-8 17+00N	531883	5507585	1.5	26.9	37.2	82	0.3	15	6.9	145	3.05	10.7	2.1	1.7	5.3	6	0.3	0.6	0.8	27	0.08	0.034	17	12	0.46	144	0.063	1	2.23	0.008	0.06	0.2	0.07	1.8	0.1	<.05	10	<.5	<1	1	7.3	15
MR-8 16+50N	531867	5507524	0.6	7.7	29.8	28	<.1	3.8	2.2	72	0.84	3	0.4	0.6	0.6	4	0.4	0.6	0.3	22	0.03	0.019	14	6	0.06	39	0.024	1	0.58	0.008	0.03	0.1	0.03	0.6	0.2	<.05	6	<.5	<1	1	0.2	15
MR-8 16+00N	531830	5507496	0.6	7.2	15.5	15	0.4	3.8	1.6	56	1.69	5	0.6	1.9	2.4	3	0.2	0.5	0.5	21	0.01	0.031	9	7	0.06	37	0.036	1	1.47	0.009	0.03	0.1	0.08	0.8	0.1	<.05	8	<.5	<1	1	8.1	15
MR-8 15+50N	531819	5507433	0.3	3.1	7.9	10	<.1	3.4	1.4	58	0.49	3.1	0.2	0.6	3	3	0.1	0.4	0.4	8	0.04	0.012	22	4	0.08	29	0.013	1	0.46	0.006	0.04	<.1	0.01	0.4	0.2	<.05	5	<.5	<1	<1	0.1	15
MR-8 15+00N	531794	5507386	1.4	14.2	57.2	27	0.3	6	24.5	1066	1.89	11.1	1	1.4	0.5	4	0.8	0.6	1.3	27	0.04	0.052	15	7	0.15	72	0.027	1	0.89	0.007	0.05	0.2	0.05	0.5	0.2	<.05	6	<.5	<1	1	0.2	15
MR-8 14+50N	531720	5507405	3	15.7	47.5	81	0.4	9.8	8.3	1624	2.66	7.8	2	1.6	2.5	9	0.4	1.1	0.6	31	0.15	0.052	15	10	0.22	122	0.118	1	1.96	0.013	0.07	0.2	0.12	1.4	0.1	<.05	14	<.5	<1	1	9.5	15
MR-8 14+00N	531666	5507405	0.7	12.8	15.5	30	0.3	6	3.5	119	1.99	5.8	0.6	1.7	0.8	2	0.2	0.6	0.5	15	0.01	0.051	12	7	0.22	34	0.029	<1	1.27	0.008	0.04	0.1	0.07	0.6	0.1	<.05	7	<.5	<1	1	2.1	15
MR-8 13+50N	531621	5507447	0.3	6.4	16.3	19	0.2	3.1	2.1	415	0.83	4.2	0.3	0.7	0.4	2	0.2	0.5	0.4	11	0.03	0.05	17	4	0.12	38	0.021	<1	0.51	0.006	0.06	0.1	0.03	0.3	0.2	<.05	4	<.5	<1	<1	0.3	15
MR-8 13+00N	531557	5507421	0.5	17	21	43	0.1	12.7	7.5	417	2.49	6.7	0.5	0.6	4.9	2	0.1	1	0.5	19	0.04	0.044	17	9	0.79	76	0.049	1	1.23	0.006	0.11	0.1	0.04	1.4	0.2	<.05	6	<.5	<1	1	1.3	15
MR-8 12+50N	531527	5507378	0.5	19.3	24.9	47	0.2	12.4	6.2	185	2.62	9.1	0.6	1	5.4	2	0.2	1.3	0.5	20	0.03	0.043	16	10	0.63	75	0.037	1	1.34	0.005	0.09	0.2	0.03	1.3	0.2	<.05	6	<.5	<1	1	1.9	15
MR-8 12+00N	531511	5507342	0.1	2	9.4	10	0.1	1.1	0.6	118	0.18	1.2	0.2	0.5	0.8	2	0.1	0.2	0.3	6	0.03	0.013	19	2	0.05	33	0.012	1	0.38	0.005	0.04	<.1	0.03	0.4	0.2	<.05	5	<.5	<1	<1	<.1	15
MR-8 11+50N	531496	5507288	0.7	8.1	14.4	38	0.2	4.5	3.5	259	1.78	3.6	0.5	8.6	1.6	3	0.3	0.4	0.5	21	0.03	0.037	11	7	0.14	76	0.046	1	1.7	0.009	0.04	0.2	0.05	0.9	0.2	<.05	7	<.5	<1	1	3.5	15
MR-8 11+00N	531484	5507235	0.7	6.9	30.1	46	0.2	4	2.8	1253	1.38	4.9	0.5	<.5	1.3	5	0.9	0.9	0.5	21	0.13	0.036	12	5	0.11	152	0.062	1	0.74	0.01	0.06	0.1	0.07	0.7	0.2	<.05	8	<.5	<1	1	1.8	15
MR-8 10+50N	531461	5507187	2.1	29.4	30.6	77	0.4	12.5	6.4	4530	1.9	5.6	10.8	<.5	0.6	25	1.3	0.9	0.4	27	0.92	0.125	12	12	0.33	415	0.063	3	2.6	0.019	0.07	0.1	0.13	1.4	0.2	0.12	10	0.9	<1	1	6.7	15
MR-8 10+00N	531436	5507151	1.7	81.4	38.4	56	0.7	11.2	5.3	2942	0.92	3.8	77.1	2.3	0.4	51	1.2	0.8	0.3	9	2.28	0.199	20	9	0.32	254	0.014	7	1.79	0.013	0.04	<.1	0.21	1.1	0.3	0.27	4	3.1	<1	<1	5.9	7.5
MR-8 9+50N	531410	5507105	1	12.2	38.9	45	0.2	4.4	3	128	2.49	4.6	0.8	1.1	2.4	2	0.3	0.6	0.5	21	0.02	0.031	10	6	0.1	55	0.029	<1	1.48	0.006	0.03	0.2	0.05	0.9	0.1	0.06	6	<.5	<1	1	2.5	15
MR-8 9+00N	531395	5507057	1	30.4	45	97	0.6	10.3	4.7	2473	1.1	3.7	12.1	0.9	0.6	37	1.8	0.9	0.4	7	1.85	0.185	12	5	0.34	239	0.01	4	1.6	0.009	0.05	0.1	0.19	1	0.2	0.21	2	0.8	<1	<1	4.7	0.5
MR-8 8+50N	531383	5507007	0.9	15.4	52.4	137	0.9	10.8	8.5	1878	4.86	4.5	1.6	4.9	3	10	1	0.9	0.4	22	0.36	0.086	10	7	0.16	162	0.057	1	3.5	0.007	0.02	0.2	0.12	2.1	0.1	0.07	6	<.5	<1	1	7.7	15
MR-8 8+00N	531386	5506956	1.5	19.5	15.4	46	0.2	5.4	4.2	184	3.72	4.5	1.2	1.6	2.1	4	0.5	0.5	0.4	29	0.05	0.071	6	8	0.13	78	0.104	1	3.9	0.009	0.04	0.3	0.14	1.6	0.1	0.07	14	0.8	<1	1	21.2	15
MR-8 7+50N	531368	5506916	0.8	9.6	14.1	48	0.1	8.9	4.4	831	2.97	3.3	0.5	1.1	3.6	2	0.1	0.6	0.4	20	0.04	0.026	15	8	0.3	92	0.033	<1	1.14	0.004	0.09	0.2	0.04	1.1	0.2	<.05	5	<.5	<1	1	0.7	15
MR-8 7+00N	531368	5506861	0.6	22.2	24	39	0.1	11.7	7.2	232	2.58	11	0.8	0.7	5	1	0.1	0.8	0.5	8	0.04	0.041	13	5	0.34	42	0.008	<1	0.84	0.001	0.06	0.1	0.04	0.8	0.1	<.05	2	<.5	<1	<1	0.8	15
MR-8 6+50N	531373	5506818	0.7	19.1	17	51	<.1	11.3	7.3	158	2.44	11.9	0.7	1.1	7.1	1	0.1	0.7	0.5	9	0.03	0.024	18	7	0.38	72	0.004	<1	1.53	0.002	0.06	0.1	0.03	1	0.1	<.05	3	<.5	<1	<1	5	15
MR-8 6+00N	531369	5506770	0.9	16.4	42.3	57	<.1	13.5	6.6	165	2.75	14.5	1.5	0.9	3.3	6	0.4	0.7	0.7	14	0.24	0.039	15	8	0.39	169	0.013	1	1.33	0.005	0.07	0.1	0.05	0.9	0.1	<.05	4	<.5	<1	<1	0.7	15
MR-8 5+50N	531363	5506740	1.3	35	40.2	150	0.3	13.3	9.7	2909	1.96	9.7	23.6	0.8	0.9	22	1.3	0.9	0.5	17	1.32	0.183	13	8	0.37	1089	0.017	3	1.48	0.01	0.06	0.1	0.16	1	0.2	0.18	4	1	<1	<1	3.4	15
MR-8 5+00N	531346	5506669	4.6	31.2	117.2	148	0.3	22.8	24.5	8228	5.88	49.2	27.4	1.7	2.8	11	1.9	2.1	1	9	0.43	0.085	10	7	0.41	712	0.011	1	0.99	0.005	0.06	0.1	0.09	1.7	0.2	0.09	3	2	<1	<1	1.3	15
MR-8 4+50N	531316	5506627	1.8	62.9	27.7	35	1.4	11.7	3.7	324	1.47	5.8	11.2	1.8	0.5	12	0.2	0.7	0.4	15	0.3	0.185	43	7	0.24	66	0.011	1	2.13	0.012	0.05	0.1	0.13	0.8	0.1	0.15	4	1.3	<1	<1	1.5	7.5
MR-8 4+00N	531276	5506587	0.9	8.2	15.4	27	<.1	6.5	3.7	121	1.78	4.2	0.5	0.6	2.6	5	0.2	0.5	0.5	38	0.06	0.024	13	8	0.37	45	0.084	<1	0.75	0.006	0.06	0.1	0.06	1.2	0.1	<.05	6	<.5	<1	1	0.9	15
MR-8 3+50N	531266	5506544	0.9	50.6	16.4	99	0.2	23.2	19.2	685	5.08	2.9	0.9	<.5	1.8	13	0.4	0.5	0.5	151	0.33	0.079	10	22	1.46	134	0.253	1	2.45	0.009	0.33	0.1	0.03	5.5	0.2	<.05	14	<.5	<1	1	2	15
MR-8 3+00N	531220	5506510	0.7	8.6	14.3	43	0.3	6.7	3.8	634	2.26	2.7	0.6	1.3	1.3	4	0.2	0.3	0.3	20	0.07	0.053	11	8	0.35	65	0.046	1	1.7	0.01	0.05	0.1	0.08	1.3	0.2	0.06	8	<.5	<1	1	1.7	15
MR-8 2+50N	531216	5506473	1	6	28.7	15	0.1	3.4	1.5	44	1.25	5.3	0.8	1.3	1.2	5	0.6	0.8	0.6	18	0.06	0.041	6	5	0.09	40	0.065	<1	1.37	0.011	0.03	0.1	0.1	0.9	0.1	<.05	8	<.5	<1	1	7.1	15
MR-8 2+00N	531199	5506428	0.7	11	14.7	19	0.2	3.9	2.1	56	1.42	8.5	0.4	1.1	1.6	3	0.2	0.5	0.7	17	0.04	0.035	14	4	0.07	49	0.02	<1	0.55	0.005	0.03	0.1	0.06	0.4	0.1	<.05	6	<.5	<1	1	0.3	15
MR-8 1+50N	531177	5506369	1.3	13.1	19.1	28	<.1	5.2	2.5	83	1.66	13.7	0.6	0.5	2.5	2	0.2	0.7	0.7	17	0.02	0.033	18	6	0.12	37	0.014	<1	0.8	0.004	0.06	0.2	0.03	0.6	0.1	<.05	5	<.5	<1	1	0.2	15
MR-8 1+00N	531130	5506344	3	7.9	41.7	40	<.1	9.8	5.2	304	2.4	6.3	9.4	0.9	4.7	6	0.2	1.4	0.6	29	0.1	0.026	14	12	0.42	637	0.029	<1	2.14	0.009	0.06	0.3	0.06	2.2	0.2	<.05	10	<.5	<1	1	3.8	15
00+50																																										
MR-8 0+0	531136	5506208	3.4	32.7	80	212	0.3	33.9	18.7	520	5.76	18.2	11.9	1.9	4	7	1.9	5.5	0.6	18	0.33	0.137	16	14	0.87	320	0.028	1	1.26	0.007	0.1	0.4	0.07	3.3	0.2	0.1	3	1.1	<1	<1	1.4	7.5
MR-10 0+0	533774	5506412	1.1	16	19.4	59	0.1	10.6	6.2	1376	1.85	3.1	1.8	1.3	5.3	3	0.2	0.5	0.4	20	0.03	0.037	12	11	0.32	46	0.07	1	1.99	0.006	0.06	0.2	0.08	1.6	0.2	<.05	7	0.8	<1	1	7.2	15
MR-10 0+50E	533817	5506442	0.8	14.8	31.5	59	0.3	9.4	6	514	2.21	2.2	3.5	<.5	6.3	3	0.1	0.2	0.4	19	0.04	0.031	17	10	0.58	49	0.063	1	1.59	0.006	0.06	0.1	0.06	1.4	0.2	<.05	7	<.5	<1	1	2.8	15
MR-10 1+0E	533860	5506470	0.7	13.7	29.2	64	0.2	7.4	5.7	1293	1.82	2.8	0.7	<.5	1.9	2	0.3	0.4	0.4	13	0.03	0.052	13	8	0.46	45	0.027	<1	1.2	0.004	0.06	0.1	0.06	0.6	0.2	0.06	4	<.5	<1	<1	0.6	15
MR-10 1+50E	533876	5506504	0.7	8.8	22.2	42	0.2	14.1	5.6	930	1.67	2.1	0.8	<.5	2.8	2	0.1	0.2	0.3	14	0.01	0.031	14	12	0.33	43	0.023	<1	0.81	0.003	0.06	0.1	0.03	0.6	0.1	<.05	4	<.5	<1	<1	0.4	15
02+00	533923	5506530																																								
MR-10 2+50E	533958	5506556	0.9	8.4	15.6	32	0.1	4.9	4.8	1840	1.44	2.1	0.7	<.5	1.7	3	0.3	0.4	0.4	16	0.02	0.032	14	6	0.11	56	0.029	<1	0.62	0.004	0.06	0.2	0.04	0.4	0.1	<.05	5	<.5	<1	<1	0.5	15
03+00	534016	5506565																																								
MR-10 3+50E	534051	5506599	2	9.1	100.9	53	0.1	9.9	8.8	2577	1.79	3	1.4	0.6	1.2	5	0.5	0.3	0.6	21	0.05	0.048	16	27	0.19	68	0.019	1	0.75	0.005	0.06	0.1	0.03	0.6	0.1	<.05	5	<.5	<1	1	0.3	15
MR-10 4+00E	534112	5506580	0.8	4.8	24.4	40	<.1	5.4	2.7	95	1.73	3.3	1.3	<.5	5.3	3	0.2	0.3	0.4	17	0.02	0.018	17	10	0.14	34	0.025	<1	0.68	0.004	0.06	0.1	0.04	0.6	0.1	<.05	5	<.5	<1	<1	0.8	15
MR-10 4+50E	534149	5506587	4.5	12.9	16.7	38	0.2	8.7	5	469	2.61	3	7.4	<.5	3.8	10	0.3	0.5	0.5	28	0.11	0.043	16	10	0.24	56	0.087	2	1.76	0.009	0.06	0.2	0.06	1.2	0.1	0.08	12	<.5	<1	1	6.1	15

MR-10 5+00E	534207	5506591	0.3	4.3	15.5	18	<.1	3.6	2.2	88	0.74	2.7	0.8	<.5	0.8	4	0.2	0.4	0.4	12	0.04	0.027	15	5	0.1	23	0.011	1	0.62	0.006	0.04	0.1	0.02	0.5	0.1	<.05	3	<.5	<1	<1	0.2	15
MR-10 5+50E	534246	5506577	0.9	4.4	7.5	23	<.1	5.5	2.9	157	1.64	1.9	0.8	<.5	4	3	0.2	0.4	0.4	19	0.02	0.015	18	6	0.13	27	0.038	<1	0.54	0.004	0.04	0.2	0.02	0.5	0.1	<.05	6	<.5	<1	1	0.8	15
MR-10 6+00E	534286	5506552	1.1	9.9	34.8	51	<.1	11.7	7	185	2.06	5.1	3.5	<.5	6.3	7	0.3	0.2	0.6	13	0.11	0.028	18	15	0.33	65	0.007	<1	1.25	0.004	0.06	0.1	0.03	1	0.1	<.05	4	<.5	<1	<1	0.5	15
MR-10 6+50E	534336	5506548	1.5	2.9	6.4	19	<.1	4.3	2.3	54	1.21	1.4	0.8	<.5	4.4	3	0.1	0.2	0.3	15	0.03	0.012	21	6	0.08	40	0.015	<1	0.45	0.003	0.05	0.1	0.01	0.5	0.1	<.05	3	<.5	<1	<1	0.3	15
MR-10 7+00E	534386	5506606	4.7	13.4	28.2	42	0.1	7.4	6.5	1424	1.5	5.3	31	0.6	0.8	15	0.6	0.5	0.4	15	0.26	0.098	50	12	0.17	42	0.008	1	1.06	0.007	0.06	0.1	0.04	0.6	0.1	0.1	4	0.6	<1	<1	0.4	7.5
MR-10 7+50E	534429	5506587	0.8	10.3	20.5	78	<.1	17.2	8.6	1100	2.42	2.1	1.5	1.4	0.4	12	0.3	0.4	0.8	29	0.16	0.107	16	32	0.27	148	0.017	2	0.79	0.005	0.09	0.1	0.02	0.7	0.1	0.11	5	<.5	<1	1	0.3	15
MR-10 8+00E	534473	5506575	0.8	22.8	26.6	69	<.1	22.7	12.5	1353	2.31	2.3	5.9	0.5	1.8	10	0.3	0.4	0.6	29	0.14	0.086	16	38	0.45	80	0.015	1	0.96	0.007	0.07	0.2	0.02	1.2	0.1	0.08	3	<.5	<1	<1	0.1	15
MR-10 8+50E	534516	5506556	0.7	12.2	36.3	74	<.1	11.7	12.5	2570	1.73	2.7	1.3	1.3	0.7	10	0.4	0.3	0.6	21	0.17	0.087	15	21	0.26	87	0.008	1	0.62	0.004	0.07	0.1	0.02	0.6	0.1	0.08	2	<.5	<1	<1	0.2	15
MR-10 9+00E	534551	5506546	0.8	7.9	14.4	24	<.1	8.8	6.6	413	2.15	2.5	1.3	0.5	5.1	3	0.1	0.4	0.6	22	0.01	0.031	22	15	0.14	24	0.019	<1	0.82	0.003	0.04	0.2	0.03	0.9	0.1	<.05	4	0.6	<1	<1	0.4	15
MR-10 9+50E	534616	5506529	1	10.9	20	43	<.1	11.9	9.5	1384	2.76	1.8	2	0.7	3.5	5	0.1	0.3	0.6	30	0.06	0.048	23	17	0.23	58	0.03	1	0.93	0.004	0.06	0.1	0.02	1.2	0.1	<.05	5	<.5	<1	1	0.3	15
MR-10 10+00E	534651	5506520	0.9	9.6	10.2	29	<.1	10.9	7.4	139	2.74	2.6	1.1	0.9	9.2	3	0.1	0.3	0.7	22	0.01	0.024	24	19	0.19	19	0.026	<1	1.04	0.003	0.04	0.1	0.04	1.1	0.1	<.05	4	<.5	<1	<1	1.4	15
MR-10 10+50E	534704	5506508	2.3	8	7.4	33	<.1	10.9	6.6	390	2.03	2.1	0.9	<.5	7.8	3	0.1	0.4	0.5	16	0.02	0.015	25	15	0.19	41	0.012	<1	0.63	0.002	0.05	0.1	0.01	0.8	0.1	<.05	3	<.5	<1	<1	0.4	15
MR-10 11+00E	534754	5506508	0.7	7.8	42.4	46	<.1	12.5	6.3	116	2.06	10.5	0.8	0.8	7.1	2	0.2	0.6	0.6	11	0.02	0.028	22	9	0.13	15	0.007	<1	0.56	0.002	0.03	0.1	0.02	0.5	<.1	<.05	3	<.5	<1	<1	0.3	15
MR-10 11+50E	534802	5506497	1.8	28.2	156.2	153	0.1	18.6	15.5	1394	2.58	28.1	1.4	<.5	2.3	7	1.5	0.6	0.7	34	0.19	0.083	16	20	0.46	53	0.01	<1	1.05	0.004	0.05	0.3	0.02	2	0.1	0.06	3	<.5	<1	<1	0.2	15
MR-10 12+00E	534862	5506490	6	16.9	21.1	66	<.1	13.1	9.6	541	3.26	13.8	1	0.9	7.5	3	0.2	0.7	0.8	33	0.02	0.024	20	15	0.23	36	0.055	1	1.06	0.005	0.06	0.4	0.03	1.3	0.1	<.05	8	<.5	<1	1	3.2	15
MR-10 12+50E	534896	5506475	2.6	9.6	22.2	40	<.1	9	5.9	300	1.96	38.6	2.1	1	6.4	3	<.1	0.3	0.5	19	0.04	0.018	19	8	0.13	34	0.013	<1	0.93	0.003	0.04	0.2	0.02	0.7	0.1	<.05	4	<.5	<1	<1	0.4	15
MR-10 13+00E	534945	5506470	0.7	9.2	10.9	33	<.1	9.2	5.5	150	1.88	3.7	0.8	<.5	7.3	3	0.1	0.3	0.6	19	0.02	0.019	19	8	0.13	42	0.014	<1	0.92	0.004	0.05	0.2	0.02	0.7	0.1	<.05	3	<.5	<1	<1	1.1	15
MR-10 13+50E	534985	5506460	2.1	11.9	19.5	37	<.1	9.2	6	395	1.95	12	0.7	<.5	6.9	5	0.2	0.9	0.7	26	0.08	0.021	22	9	0.15	46	0.036	1	0.74	0.005	0.05	0.2	0.03	0.8	0.1	<.05	5	<.5	<1	1	0.7	15
MR-10 14+00E	535033	5506477	5.6	32.9	28.6	57	0.2	12.9	14.3	4131	2.27	18.2	8.5	0.5	2.7	17	1.3	0.4	0.7	17	0.33	0.062	28	14	0.27	68	0.013	1	1.43	0.007	0.05	0.1	0.06	1.1	0.2	<.05	4	<.5	<1	<1	0.5	15
MR-10 14+50E	535092	5506472	0.9	7.2	7.7	23	<.1	7.7	6.4	228	1.93	3.4	0.6	<.5	5.7	4	0.1	0.4	0.4	37	0.04	0.019	21	9	0.46	54	0.037	<1	0.95	0.004	0.04	0.2	0.01	1.3	0.1	<.05	6	<.5	<1	1	0.5	15
MR-10 15+00E	535142	5506469	0.8	7.6	21	23	<.1	7.4	7.7	255	1.92	4.9	0.6	0.9	4.8	3	0.1	0.5	0.5	33	0.03	0.019	15	9	0.34	41	0.023	1	0.95	0.004	0.04	0.2	0.02	1.2	0.1	<.05	5	<.5	<1	1	0.8	15
MR-10 15+50E	535185	5506467	1.3	4.9	10.8	18	<.1	7.5	5.5	53	2.38	5.6	1	<.5	9.7	3	0.1	0.2	0.4	22	0.02	0.015	23	13	0.31	39	0.009	<1	1.14	0.002	0.03	0.1	0.02	0.9	<.1	<.05	4	<.5	<1	<1	1.5	15
MR-10 16+00E	535232	5506477	1.7	7.6	14.7	39	<.1	10	6.2	351	2.55	6	0.7	<.5	7.9	5	0.2	0.3	0.6	26	0.04	0.021	23	13	0.29	57	0.048	<1	1.05	0.004	0.06	0.2	0.03	1	0.1	<.05	8	<.5	<1	1	1.7	15
BA-A 00+00	531437	5504998	1	8.7	20.4	30	<.1	9.2	3	97	2.29	8	0.7	1.1	4.2	4	0.2	0.7	0.5	33	0.02	0.037	19	14	0.22	22	0.052	1	1	0.006	0.05	0.3	0.04	0.9	0.1	<.05	7	<.5	<1	1	2	
BA-A 00+50E	531455	5505050	0.9	5.2	16.2	20	<.1	4	1.6	38	1.6	3.4	0.7	1.1	4.1	4	0.1	0.5	0.4	27	0.02	0.029	20	8	0.15	28	0.041	1	1.14	0.006	0.04	0.1	0.04	0.8	0.1	<.05	8	<.5	<1	1	1.9	
BA-A 01+00E	531498	5505075	0.9	6	13.3	22	<.1	5.7	2.1	62	1.73	2.2	0.9	<.5	3.9	3	0.1	0.4	0.4	23	0.02	0.023	16	8	0.17	25	0.039	1	1.17	0.004	0.05	0.2	0.05	0.8	0.2	<.05	6	<.5	<1	1	1.8	
BA-A 01+50E	531527	5505117	0.5	6	16.7	33	0.1	7.5	3.3	166	1.01	1.4	4.2	0.8	0.3	5	0.2	0.3	0.3	16	0.08	0.082	10	8	0.25	50	0.023	2	1.28	0.006	0.06	0.1	0.04	0.4	0.1	0.07	6	<.5	<1	1	1.6	
BA-A 02+00E	531574	5505136	0.9	3	42.1	46	<.1	4.4	2.6	85	0.9	2.4	4.1	0.6	1.6	7	0.5	0.5	0.9	15	0.1	0.025	11	8	0.22	166	0.024	2	1.19	0.009	0.05	0.1	0.03	1	0.1	<.05	8	<.5	<1	1	0.6	
BA-A 02+50E	531609	5505152	1.1	6	14.4	43	<.1	6.1	2.8	58	2.36	3.3	0.8	1.1	4.3	4	0.4	0.4	0.4	33	0.03	0.017	12	9	0.16	98	0.047	1	1.56	0.006	0.04	0.2	0.05	1.1	0.1	<.05	11	<.5	<1	1	6.8	
BA-A 03+00E	531662	5505159	0.8	7.2	44.5	65	0.2	8.2	3.9	414	1.22	4.7	3.5	1.2	0.2	9	0.6	0.6	0.4	18	0.16	0.114	9	9	0.25	110	0.013	2	1.29	0.008	0.07	0.1	0.05	0.4	0.1	0.09	5	<.5	<1	1	0.8	
BA-A 03+50E	531697	5505171	1	7.8	15.2	43	<.1	9.3	4.5	292	1.71	2.3	1.8	<.5	1.8	5	0.1	0.3	0.4	24	0.06	0.051	12	11	0.25	71	0.046	1	2.04	0.006	0.06	0.1	0.03	1.1	0.1	<.05	6	<.5	<1	1	3.9	
BA-A 04+00E	531757	5505186	0.6	3.1	17.7	26	<.1	3.7	1.7	77	0.88	2.2	1.9	1.2	3.5	5	0.3	0.3	0.4	17	0.06	0.021	18	6	0.12	67	0.029	1	0.85	0.004	0.05	0.1	0.03	0.7	0.1	<.05	6	<.5	<1	<1	0.8	
BA-A 04+50E	531801	5505191	2.4	9.3	74.8	69	0.4	13.8	6	382	1.78	1.9	15.5	0.8	0.6	8	0.4	0.3	0.4	29	0.07	0.066	12	14	0.33	98	0.029	2	1.55	0.007	0.07	0.1	0.05	0.8	0.1	0.08	7	<.5	<1	1	1.1	
BA-A 05+00E	531842	5505175	2	6.9	57.4	59	0.1	7.8	6.9	823	1.33	10.5	10.7	7.2	0.7	12	1	0.5	0.5	19	0.15	0.074	12	9	0.28	150	0.017	2	1.1	0.008	0.08	0.1	0.07	0.7	0.2	0.08	5	<.5	<1	1	0.5	
BA-A 05+50E	531892	5505166	1.4	5.8	107.7	72	<.1	7.3	9	1327	1.55	9.3	6.3	0.7	2	14	1.5	0.6	0.6	19	0.18	0.042	16	8	0.21	171	0.029	2	1.16	0.009	0.08	0.1	0.04	1	0.2	<.05	5	<.5	<1	1	0.8	
BA-A 06+00E	531938	5505145	0.7	8.6	30	53	0.4	7.8	5	1053	1.43	4.1	0.9	0.9	0.3	4	0.4	0.7	0.4	18	0.04	0.11	15	8	0.18	39	0.013	2	1.07	0.005	0.08	0.1	0.05	0.4	0.1	<.05	5	<.5	<1	<1	1.7	
BA-A 06+50E	531993	5505142	0.9	8.8	67.1	35	0.2	7.7	4.5	302	1.38	10.5	1	1.4	0.4	4	0.4	1	0.7	16	0.02	0.078	15	8	0.15	24	0.012	1	0.84	0.004	0.06	0.1	0.07	0.4	0.1	0.06	3	0.5	<1	1	0.4	
BA-A 07+00E	532041	5505148	1.5	7.2	12.3	57	<.1	8.7	5.8	642	2.01	2.7	2.4	0.8	1.8	5	0.3	0.5	0.4	26	0.03	0.043	18	11	0.23	70	0.031	1	1.22	0.005	0.07	0.2	0.03	0.9	0.1	<.05	6	<.5	<1	1	0.7	
BA-A 07+50E	532094	5505146	0.5	1.6	5.4	29	<.1	5.6	3.5	77	0.99	1.1	1.2	1	1.8	2	0.1	0.3	0.2	5	0.01	0.035	21	3	0.09	41	0.003	<1	0.55	0.003	0.07	<.1	0.02	0.3	0.1	<.05	2	<.5	<1	<1	0.5	
BA-A 08+00E	532132	5505113	0.8	8.3	25	84	0.1	5.9	3.6	105	1.31	2.1	1.4	3.8	2.3	4	0.4	0.3	0.8	17	0.02	0.03	19	7	0.16	50	0.03	1	0.93	0.004	0.05	0.1	0.03	0.7	0.1	<.05	5	<.5	<1	<1	0.9	
BA-A 08+50E	532167	5505110	1.1	10.3	83.9	116	0.4	10.6	8.5	1235	1.52	2.8	3.4	1.3	0.7	6	0.5	0.3	0.6	20	0.05	0.099	11	10	0.21	57	0.022	2	1.94	0.007	0.06	0.1	0.08	0.6	0.1	0.11	6	0.6	<1	<1	1.8	
BA-A 09+00E	532205	5505089	1.7	5.8	17.6	36	<.1	5.2	3.1	253	1.39	1.7	1	1.2	3.7	5	0.1	0.3	0.4	23	0.03	0.021	15	7	0.1	48	0.053	1	1.08	0.007	0.06	0.1	0.04	0.9	0.1	<.05	7	0.5	<1	1	2.7	
BA-A 09+50E	532247	5505064	1	8.4	13	25	0.1	4.7	2.7	314	1.78	2.7	1	2.4	6.4	3	0.2	0.4	0.3	26	0.02	0.038	10	8	0.12	34	0.089	1	2.65	0.009	0.04	0.1	0.08	1.7	0.1	<.05	10	0.5	<1	1	14.3	
BA-A 10+00E	532292	5505038	1.2	10.1	10.4	39	<.1	5.1	3.1	882	1.88	2.5	1	1.2	5.6	3	0.2	0.3	0.3	29	0.03	0.047	8	8	0.11	44	0.089	1	3	0.009	0.03	0.2	0.09	1.5	0.1	<.05	10	0.6	<1	1	13	
BA-A 10+50E	532335	5505009	1.2	9.2	11.4	34	0.1	4.8	3.3	648	1.61	2.7	0.9	1.4	3.3	4	0.2	0.6	0.3	24	0.03	0.042	8	7	0.1	46	0.084	1	2.46	0.009	0.04	0.2	0.08	1.4	0.1	<.05	9	<.5	<1	1	12.2	
BA-A 11+00E	532375	5504968	1	7.1	13.2	28	<.1	5.4	2.6	512	1.53	2.6	0.7	1.4	4.5	4	0.2	0.7	0.4	32	0.02	0.025	18	10	0.13	41	0.068	1	1.26	0.006	0.04	0.1	0.04	1	0.1	<.05	8	<.5	<1	1	4.8	
BA-A 11+50E	532415	5504943	0.8	7.7	11.3	13	<.1	1.8	1.4	304	1.67	1.9	1.2	0.9	3.4	3	0.1	0.4	0.3	24	0.02	0.051	6	6	0.03	32	0.091	1	3.27	0.013	0.02	0.1	0.07	1.8	0.1	<.05	11	<.5	<1	1	27.7	
BA-A 12+00E	532440	5504906	0.5	5.8	10.1	33	<.1	7.1	4	154	1.23	2.4	1.4	1.1	10.6	4	0.2	0.4	0.3	13	0.02	0.029	29	6	0.18	62	0.024	1	1.48	0.004	0.07	0.1	0.03	1	0.1	<.05	3	<.5	<1	<1	3.1	
BA-A 12+50E	532479	5504866	0.8	13.6	11.8	25	0.2	5.4	2.2	108	1.85	3.9	1.3	2.2	6.5	4	0.2	0.4	0.3	25	0.03	0.059	10	9	0.1	45	0.077	1	3.49	0.012	0.04	0.2	0.06	2.1	0.1	<.05	8	<.5	<1	1	39.2	
BA-A 13+00E	532517	5504839	1	14.5	11.7	24	0.3	4.8	2	102	1.81	3.5	1	2.2	3.7	4	0.2	0.4	0.3	33	0.03	0.065	6	8	0.07	43	0.112	<1	3.06	0.015	0.03	0.1	0.05	1.9	0.1	<.05	11	0.6	<1	1	37.7	
BA-A 13+50E	532557	5504820	0.7	5.8	10.2	20	<.1	3.9	1.9	80	1.44	2.8	0.8	1.3	4.7	3	0.2	0.7	0.3	21	0.01	0.026	18	6	0.09	24	0.047	<1	1.08	0.006	0.04	0.1	0.04	1	0.1	<.05	7	<.5	<1	1	3.9	
BA-A 14+00E	532608	5504794	0.7	9.7	10.6	23	0.3	4.4	2.3	58	1.54	2.3	1.4	2.5	5.7	4	0.1	0.2	0.3	21	0.02	0.067	21	6	0.09	37	0.053	1	1.48	0.007	0.05	0.1	0.04	1.2	0.1	<.05	7	0.5	<1	1	7.5	
BA-A 14+50E	532656	5504788	0.9	60.8	9.8	29	0.2	6.7	3.7	117	1.86	2.8	1.4	1.9	8	3	0.1	0.3	0.4	18	0.02	0.053	23	8	0.15	44	0.033	1	2.22	0.004	0.07	0.2	0.05	1.5	0.2	<.05	6	0.6	<1	1	5.5	
BA-A 15+00E	532701	5504794	1	10.2	45.9	46	<.1	9.2	5.5	821	1.89	6.5	1.5	1.6	7.9	7	0.4	1	0.6	29	0.06	0.053	23	10	0.14	121	0.058	1	2.26	0.006	0.07	0.2	0.08	1.8	0.1	<.05	8	<.5	<1	1	9.5	
BA-B 00+00	532824	5504465	0.7	3.1	9.4	20	0.1	5	2.6	92	1.08	2.8	1.2	0.8	5.8	2	0.1	0.4	0.4	10	0.01	0.022	26	5	0.15	28	0.007	<1	0.88	0.003	0.06	0.1	0.02	0.7	0.1	<.05	3	<.5	<1	<1	0.2	
BA-B 00+50E	532883	5504488	0.6	13	61.2	80	0.2	9	6.4	185	1.51	3.7	1.9	0.6	11.1	3	0.2	0.4	0.4	16	0.02	0.026	22	8	0.16	58	0.032	1	2.02	0.005	0.07	0.1	0.05	1.6	0.1	<.05	4	0.5	<1	<1	7.4	
BA-B 01+00E	532921	5504511	0.5	9.2	34.2	36	0.2	6	4.1	110	1.39	4.1	1.6	0.8	8.1	4	0.1	0.4	0.4	17	0.02	0.014	22	7	0.1	49	0.021	<1	1.15	0.005	0.07	0.1	0.02	1.2	0.1	0.06	4	<.5	<1	1	3.1	
BA-B 01+50E	532963	5504542	0.4	5	15.5	27	<.1	5.2	3.1	408	1.14	1.8	1.1	0.8	5.8	4	0.1	0.4	0.4	20	0.02	0.014	24	7	0.11	62	0.031	1	1.09	0.005	0.08	0.1	0.03	1.1	0.2	<.05	5	<.5	<1	1	1.7	
BA-B 02+00E	533000	5504557	0.5	13.2	18.2	33	0.1	9.1	5.7	100	1.6	3.2	3	1.3	9.7	4	0.1	0.4	0.4	14	0.02	0.026	28	7	0.16	65	0.028	<1	1.39	0.003	0.07	0.2	0.04	1.1	0.1	<.05	3	<.5	<1	<1	5.9	
BA-B 02+50E	533036	5504593	0.3	23.5	48	67	0.2	10.3	8.3	910	1.43	7.2	2.9	1.1	13.3	5	0.4	0.8	0.5	10	0.05	0.035	32	6	0.22	79	0.009	<1	1.02	0.003	0.09	0.1	0.02	0.9	0.1	<.05	2	<.5	<1	<1	0.4	
BA-B 03+00E	533074	5504613	0.3	3.8	8.3	25	<.1	7.1	4	59	1.68	1.9	1	0.5	10.3	3	0.1	0.2	0.4	14	0.01	0.017	31	6	0.24	46	0.008	<1	1.12	0.003	0.08	0.1	0.02	1.2	0.1	<.05	4	<.5	<1	<1	0.7	
BA-B 03+50E	533119	5504613	0.4	9.3	12.6	17	0.1	5.8	3.3	65	1.48	1.9	1.6	1.3	9.3	3	0.1	0.3	0.4	16	0.02	0.019	23	7	0.12	47	0.025	1	1.59	0.004	0.07	0.1	0.03	1.3	0.1	<.05	4	<.5	<1	1	8.1	
BA-B 04+00E	533146	5504631	0.5	5.8	11.5	24	<.1	6	3.4	343	1.3	2.2	1.1	1	6.3	3	0.2	0.3	0.3	16	0.02	0.036	20	7	0.09	66	0.026	1	1.63	0.005	0.06	0.1	0.04	1.3	0.1	<.05	5	<.5	<1	<1	2.7	
BA-B 04+50E	533196	5504671	0.5	12.7	10.2	30	0.2	8.8	4.2	107	1.55	1.9	1.3	1.5	8.1	4	0.2	0.3	0.4	18	0.04	0.027	20	8	0.15	55	0.04	1	1.68	0.006	0.06	0.2	0.08	1.4	0.1	<.05	5	<.5	<1	<1	8.2	
BA-B 05+00E	533250	5504694	0.4	5.2	10.5	24	0.1	5.3	2.9	60	1.31	1.9	0.9	0.5	5.7	3	0.1	0.3	0.3	16	0.02	0.02	19	6	0.11	44	0.031	<1	1.17	0.008	0.06	0.1	0.04	1	0.1	<.05	5	<.5	<1	1	2.7	
BA-B 05+50E	533276	5504727	0.3	3.8	5.9	16	<.1	7.3	3.5	52	1.85	1.3	1.7	1	7.9	2	<.1	0.2	0.3	13	0.01	0.031	22	8	0.12	29	0.018	1	1.09	0.003	0.05	0.1	0.04	0.8	<.1	<.05	2	<.5	<1	<1	2.3	
BA-B 06+00E	533270	5504766	0.7	4.7	10.3	23	0.2	3.1	2	97	2.33	4.3	1	1	5.7	3	0.2	0.5	0.4	25	0.02	0.052	15	9	0.08	29	0.04	1	1.61	0.006	0.05	0.2	0.06	1.1	0.1	<.05	7	<.5	<1	1	4.3	

Appendix C

Statement of Expenditures

STATEMENT OF EXPENDITURES

The following expenses were incurred on the Mount Rice property for the purpose of geological exploration between November 9 and 11, 2007; comprised of three helicopter supported traverses.

PERSONNEL	
8 man-days at $250 / day	$ 2,000.00
EQUIPMENT RENTAL	
Hand-held radios - 8 man-days at $10 / day	$ 80.00
FIELD SUPPLIES	
8 man-days @ $15 / day	$ 120.00
HELICOPTER - 3 days	$ 6,000.00
ANALYSES	
164 Soil Samples at $25 / sample	$ 4,100.00
SHIPPING	$ 90.00
REPORT/REPRODUCTION	
R. T. Walker, P.Geo.: 3.0 days @ $500/day	$ 1,500.00
Total:	**$13,890.00**

Appendix D

Program-Related Documents

BRITISH COLUMBIA

Contact Us ▸ Help

B.C. HOME

Mineral Titles

Mineral Titles Online

Mineral Claim Exploration and Development Work/Expiry Date Change

Confirmation

Mineral Claim Exploration and Development Work/Expiry Date Change

- Select Input Method
- Select/Input Tenures
- Input Lots
- Data Input Form
- Review Form Data
- Process Payment
- Confirmation

Recorder: MOUNTAIN STAR RESOURCES LTD (139398)

Submitter: MOUNTAIN STAR RESOURCES LTD (139398)

Recorded: 2007/MAR/22

Effective: 2007/MAR/22

D/E Date: 2007/MAR/22

Your report is due in 90 days. Please attach a copy of this confirmation page to the front of your report.

Event Number: 4139255

Work Start Date: 2006/JUL/27
Work Stop Date: 2006/NOV/12

Total Value of Work: $ 14372.38
Mine Permit No:

Work Type: Technical Work
Technical Items: Geochemical, PAC Withdrawal (up to 30% of technical work performed)

Summary of the work value:

Tenure #	Claim Name/Property	Issue Date	Good To Date	New Good To Date	# of Days For-ward	Area in Ha	Work Value Due	Sub-mission Fee
530476	RICE WEST	2006/mar/24	2007/mar/24	2011/mar/24	1461	522.41	$ 10448.16	$ 836.42
530483	RICE EAST	2006/mar/24	2007/mar/24	2011/mar/24	1461	417.86	$ 8357.10	$ 669.03

- Main Menu
- Search for Mineral / Placer / Coal Titles
- View Mineral Tenures
- View Placer Tenures
- View Coal Tenures
- MTO Help Tips

Exit this e-service

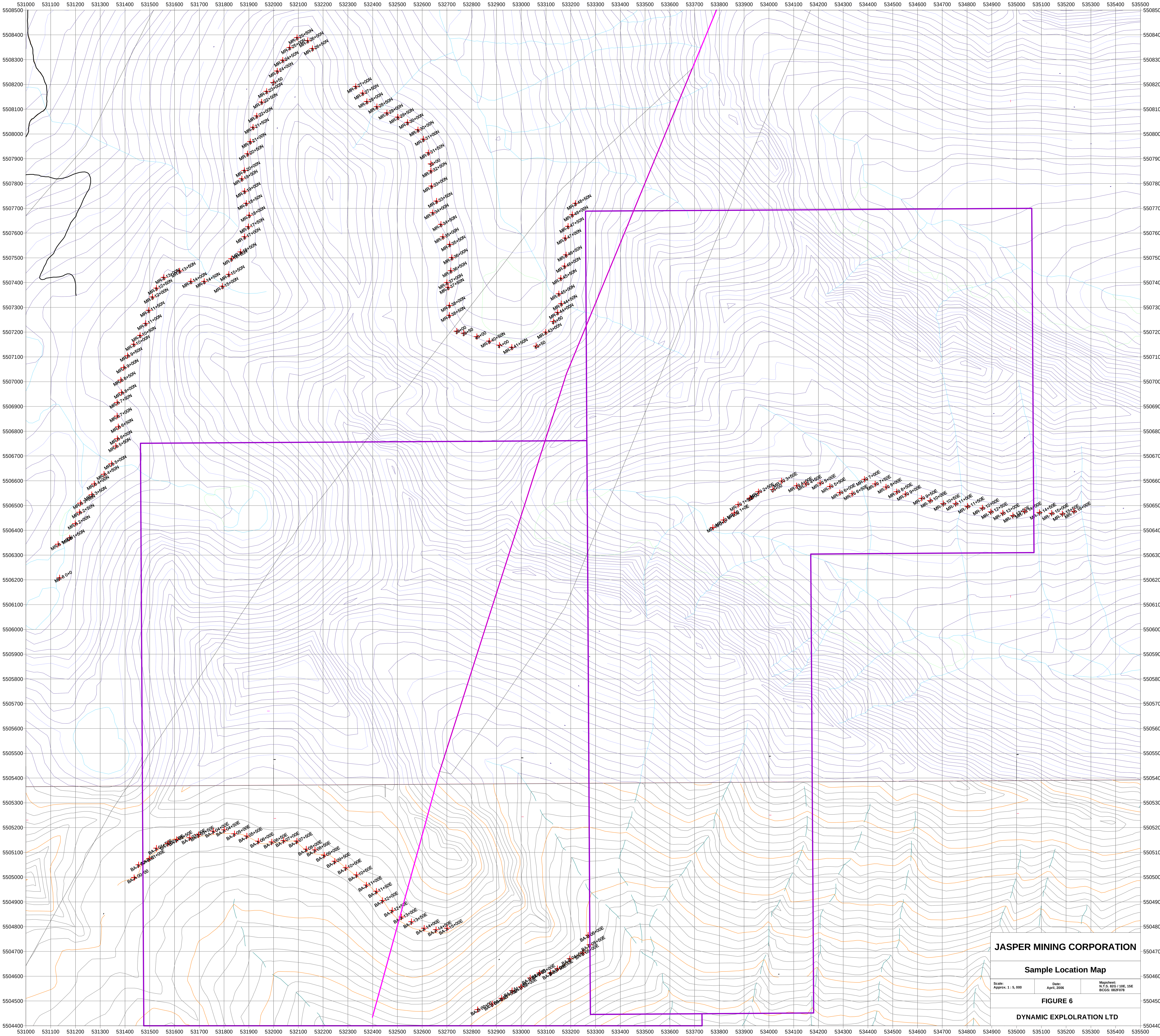
Total required work value:	$	18805.26
PAC name:		Mountain Star Resources Ltd
Debited PAC amount:	$	4432.88
Credited PAC amount:	$	0.00
Total Submission Fees:	$	1505.45
Total Paid:	$	1505.45

The event was successfully saved.

Please use ***Back*** button to go back to event confirmation index.

Back

Mount Rice


JASPER MINING CORPORATION
Sample Location Map
Scale: Approx. 1 : 5, 000
Date: April, 2006
Mapsheet: N.T.S. 82G / 10E, 15E BCGS: 082F078
FIGURE 6
DYNAMIC EXPLOLRATION LTD