

**BC Geological Survey
Assessment Report
29703**

Assessment Report for the

CRAWFORD Property

Soil Sampling

Fort Steele Mining Division

N.T.S. 82 F/ 10E

Latitude 49° 46' 56" N, Longitude 116° 39' 33"W

for

Jasper Mining Corporation
1020, 833 - 4th Avenue S.W.
Calgary, Alberta
T2P 3T5

Submitted by:

Richard T. Walker

of

Dynamic Exploration Ltd.
2601 - 42nd Avenue South
Cranbrook, B.C.
V1C 7H3

Submitted: February 22nd, 2008

SUMMARY

The Crawford property is located approximately 70 kilometres northwest of Cranbrook, east of Kootenay Lake in the Purcell Mountains. The property can be reached by way of an old logging road up Crawford Creek about 15 kilometres northeast of the town of Crawford Bay. Elevation ranges from 1325 to 2320 metres.

The CRAWFORD property is underlain by clastic sediments of the Horsethief Creek Group of Upper Proterozoic age, with Late Proterozoic to lower Cambrian strata of the Hamill Group occurring at the northwest corner of the property. These rocks have been subsequently intruded by a broadly elliptical quartz monzonite stock of Cretaceous(?) age, interpreted to be associated with the Fry Creek Batholith, a large intrusive complex located five kilometres to the north. The Fry Creek Batholith has been correlated to the Bayonne Magmatic Belt (BMB), comprised of Cretaceous age felsic intrusions extending from the Baldy Batholith north of Kamloops to the International Boundary with the United States south of Creston. The Sawyer Creek Stock, located 8 km to the southeast, is similarly correlated to the BMB and may represent an apophyse of the Fry Creek Batholith.

The 2007 program consisted of an initial soil sampling program with which to evaluate the property and as a follow-up to the 2006 Aeroquest airborne geophysical survey, completed between May 15th and September 25th, 2007.

A total of 196 soil samples were recovered along two contour lines and a road traverse, all oriented at a high angle to the structure underlying the property. Samples were taken at a 50 m spacing, all recovered from the "B" Horizon. Holes were dug by hand using a mattock to a depth generally between 10 and 25 cm below surface. Samples were placed in Kraft soil envelopes, air dried to eliminate excess water content and shipped to Acme Analytical Laboratories Ltd in Vancouver, BC for analysis using SS80 preparation and Group 1DX analysis.

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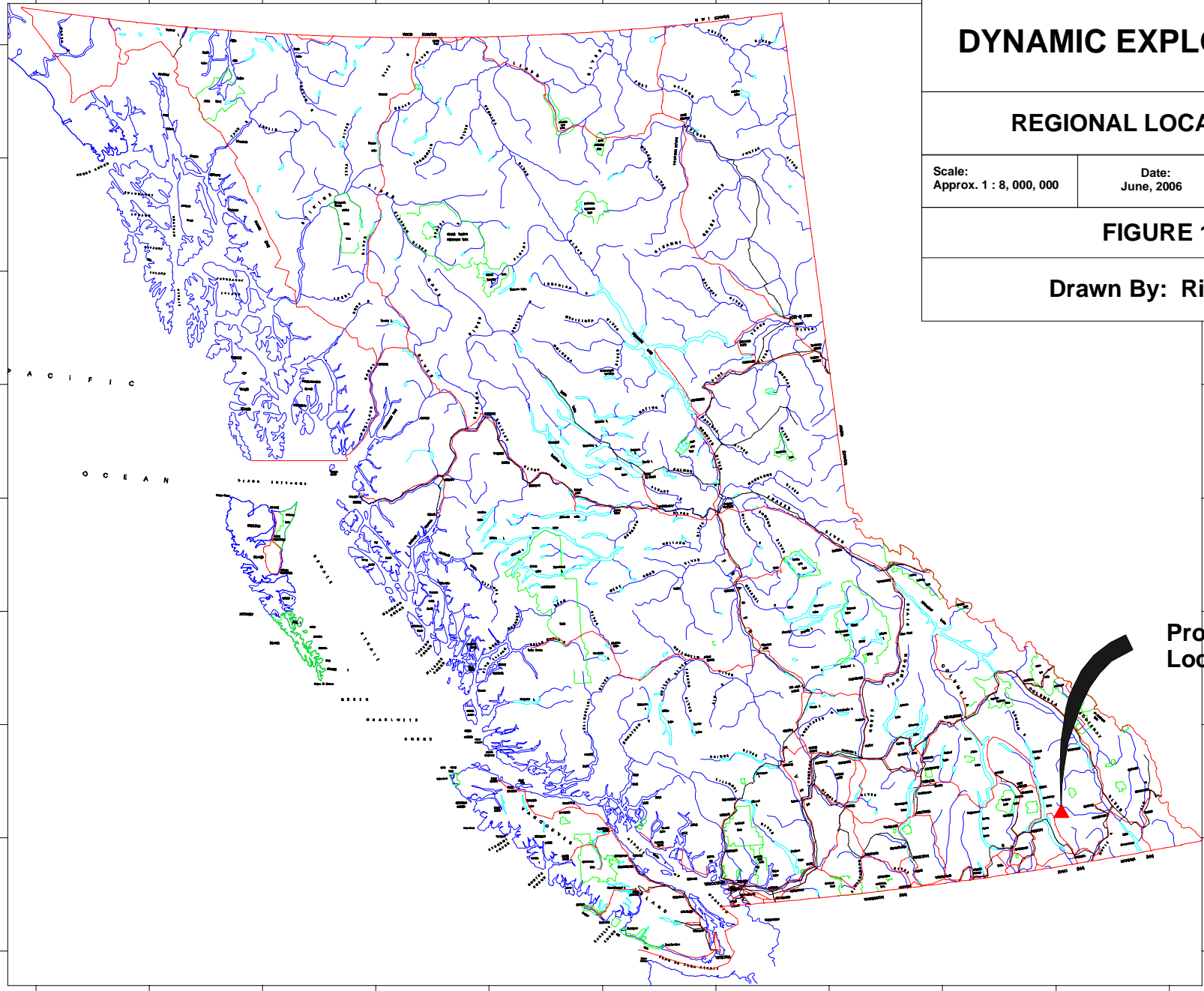
INTRODUCTION

The Crawford property is located approximately 70 kilometres northwest of Cranbrook (Fig. 1 to 3), east of Kootenay Lake in the Purcell Mountains. The property can be reached by way of an old logging road up Crawford Creek about 15 kilometres northeast of the town of Crawford Bay. Elevation ranges from 1325 to 2320 metres.

The CRAWFORD property is underlain by clastic sediments of the Horsethief Creek Group of Upper Proterozoic age, with Late Proterozoic to lower Cambrian strata of the Hamill Group occurring at the northwest corner of the property (Fig. 4). These rocks have been subsequently intruded by a broadly elliptical quartz monzonite stock of Cretaceous(?) age, interpreted to be associated with the Fry Creek Batholith, a large intrusive complex located five kilometres to the north. The Fry Creek Batholith has been correlated to the Bayonne Magmatic Belt (BMB), comprised of Cretaceous age felsic intrusions extending from the Baldy Batholith north of Kamloops to the International Boundary with the United States south of Creston. The Sawyer Creek Stock, located 8 km to the southeast, is similarly correlated to the BMB and may represent an apophyse of the Fry Creek Batholith.

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DYNAMIC EXPLORATION LTD

REGIONAL LOCATION MAP

Scale:
Approx. 1 : 8, 000, 000

Date:
June, 2006

Mapsheet:
N.T.S. 82F / 10E
BCGS: 082F 077

FIGURE 1

Drawn By: Rick Walker

Property
Location

DYNAMIC EXPLORATION LTD

PROPERTY LOCATION MAP

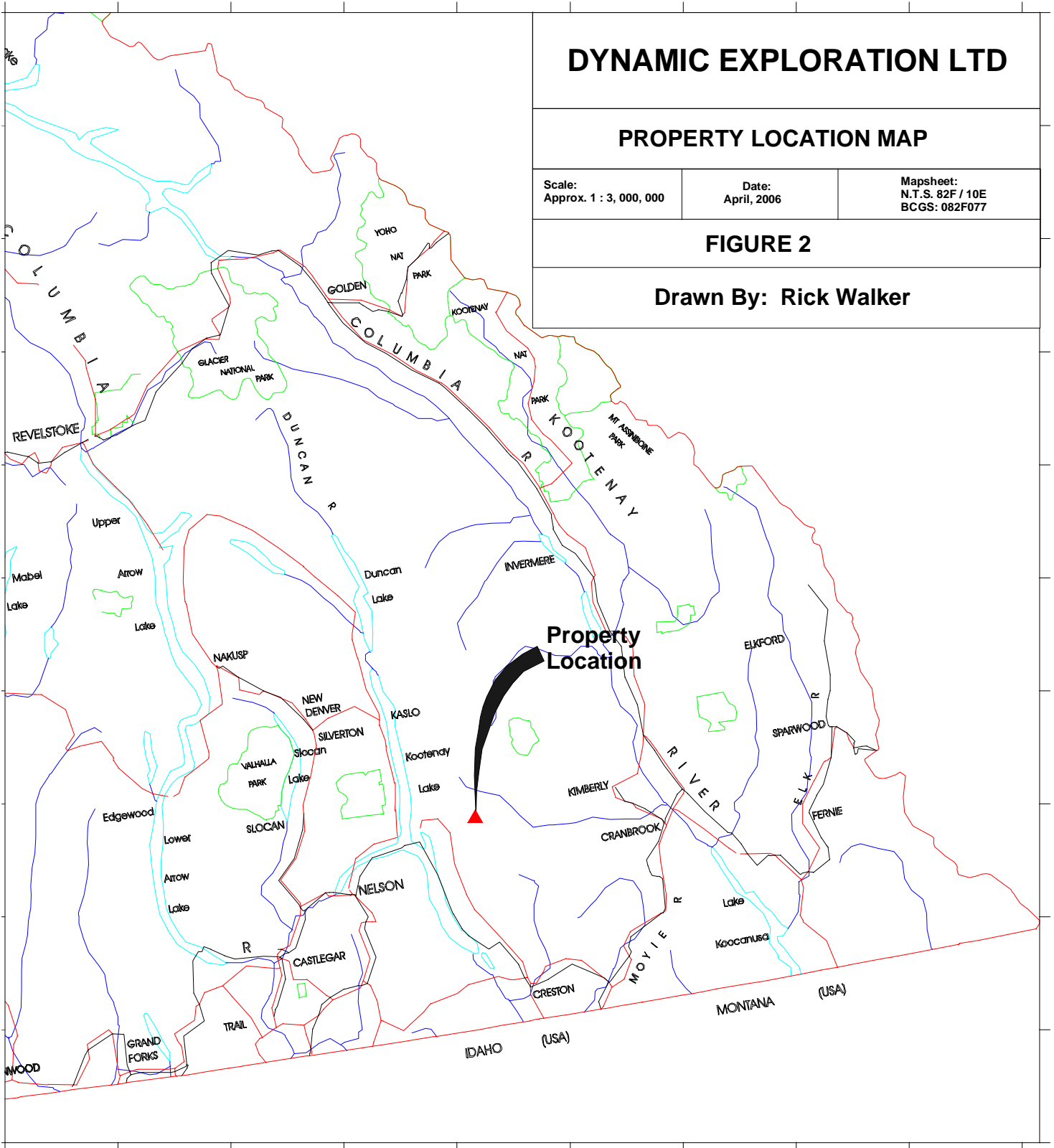
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April, 2006

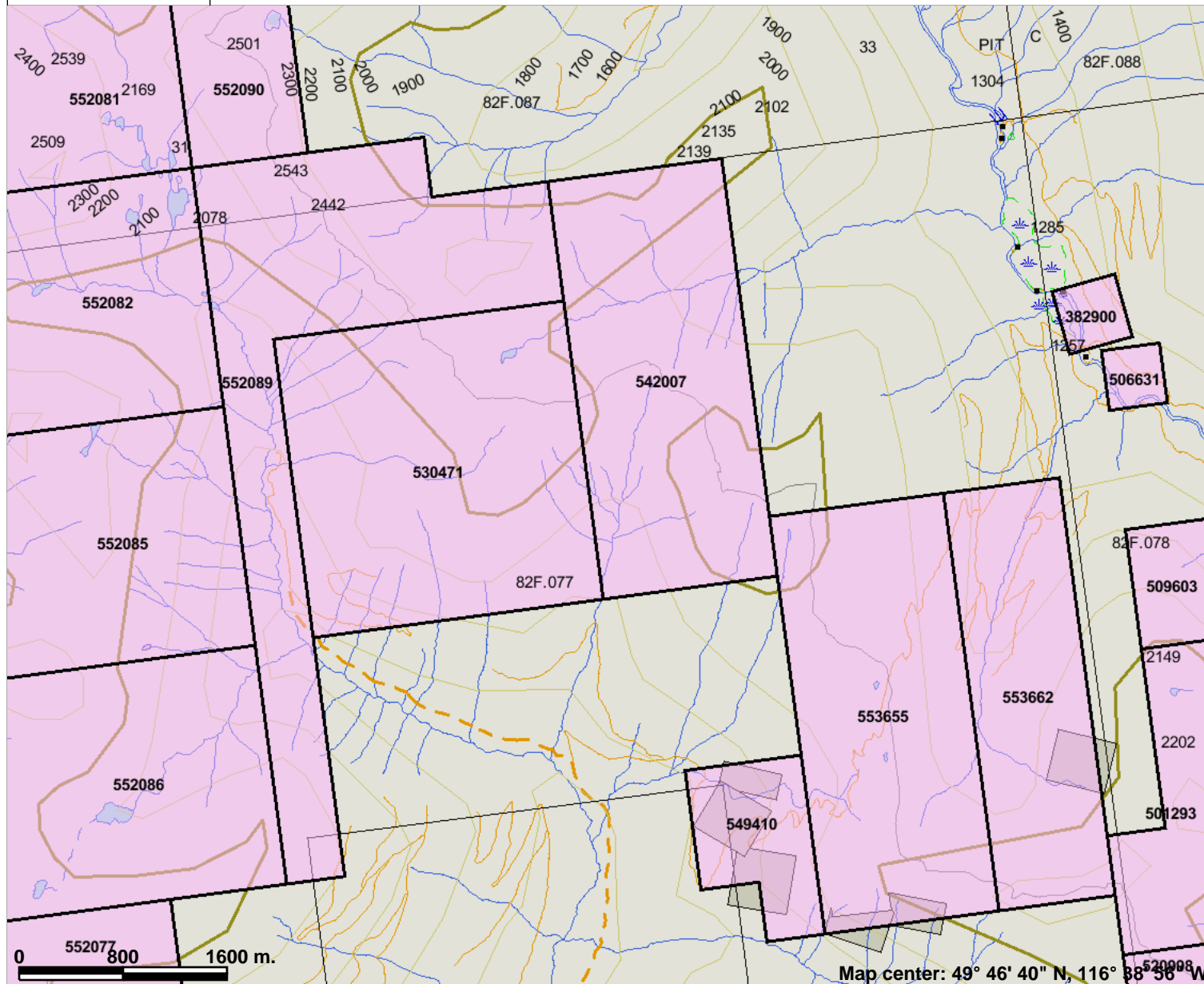
Mapsheet:
N.T.S. 82F / 10E
BCGS: 082F077

FIGURE 2

Drawn By: Rick Walker



Internet Mapping Framework



Legend

- Indian Reserves
- National Parks
- Parks
- Mineral Tenures (Mineral - LRDW)
- Mineral Claim
- Mineral Lease
- Reserves (Mineral - LRDW Sites)**
- Placer Claim Designation
- Placer Lease Designation
- No Staking Reserve
- Conditional Reserve
- Release Required Reserve
- Surface Restriction
- Recreation Area
- Others
- Mining Division (MTO)
- Survey Parcels
- BCGS Grid
- Contours (1:250K)**
- Contour - Index
- Contour - Intermediate
- Area of Exclusion
- Area of Indefinite Contours
- Transportation - Points (TRIM)**
- Helipad
- Transportation - Lines (TRIM)**
- Airfield
- Airport
- Airstrip
- Airport.Abandoned
- Ferry Route
- Road (Gravel Undivided) - 1 Lane
- Road (Gravel Undivided) - 2 Lanes

This map is a user generated static output from an Internet mapping site and is for general reference only. Data layers that appear on this map may or may not be accurate, current, or otherwise reliable. THIS MAP IS NOT TO BE USED FOR NAVIGATION.

Notes: Figure 3: Crawford Property



Scale: 1:46,526

LOCATION AND ACCESS

The CRAWFORD property is located in the western Purcell Mountains (Latitude 49° 46' 56" N, Longitude 116° 39' 33" W), approximately 70 kilometres northwest of Cranbrook, B.C. on N.T.S. mapsheet 82 F/10E (Fig. 1 and 2). The property consists of 2 Mineral Tenures acquired through Mineral Tenures Online (Fig. 3).

The property can be accessed by way of an old logging road up Crawford Creek about 15 kilometres northeast of the town of Crawford Bay. Elevation ranges from 1325 to 2320 metres.

All roads are negotiable using a 2WD vehicle although 4WD is recommended for better clearance.

Helicopter access is recommended for the northeastern?? portions of the property.

PHYSIOGRAPHY AND CLIMATE

The CRAWFORD property is located slightly northwest of Rose Pass (Fig. 2), approximately 70 km northwest of Cranbrook, on the east side of Kootenay Lake (Fig. 1). Relief in the area varies from 1325 metres (4,350 feet) to approximately 2320 metres (7,611 feet).

The claims are generally characterized by moderately to very steep topography, with generally south and west facing slopes along east-west and north-south oriented valleys. Crawford Creek, along the western edge of the property is oriented roughly north-south.

Vegetation in the area consists predominantly coniferous, with deciduous trees preferentially located along the valley bottom. Undergrowth consists largely of small deciduous shrubs, with Devil's Club along watercourses and wet areas.

The claims are located east of Kootenay Lake along a regional topographic high, comprising the local drainage divide, and are therefore subject to heavier precipitation. As a result, the region is characterized by heavy snowfall during the winter months. The property is available for vehicle based, geological exploration from June to late October.

CLAIM STATUS

The property consists of 2 Mineral Tenure On-line (MTO) mineral tenures (Fig. 3), acquired in accordance with existing government claim location regulations. Significant tenure information is summarized below:

Tenure Name	Area (ha)	Tenure #	Expiry Date*
Crawford	521.565	530471	Mar. 24, 2014
Crawford East	<u>438.0745</u>	542007	Mar. 24, 2014
Total:	959.6395		

*After 2007 assessment credit applied.

HISTORY

Prior to 1979 there is no recorded work for the immediate area. The Rose Pass tin occurrence, a quartz vein containing galena, sphalerite, chalcopyrite, and reportedly, stannite, is located about 4 km southeast of the property.

In 1979 two claims comprising 32 units were staked by Cominco Ltd to cover several molybdenite occurrences located by follow-up of regional stream silt anomalies.

In 1980 Cominco carried out a program of geological mapping and soil sampling on the Crawford property (on their Rose property) (Wright 1980). The work consisted of detailed geological mapping and prospecting of the property, and soil sampling on a grid. Sampling consisted of 617 soils and 8 stream silts. All samples were analyzed for Cu, Pb, Zn, Mn, Mo and W. Results showed scattered anomalous values for Cu, Pb, Zn, Mn and W which could not be related to known mineralization. Anomalous molybdenum values were found to correspond to a large quartz monzonite intrusion containing several molybdenite showings. Further work was recommended to determine the significance of the anomalies. No further work appears to have been filed by Cominco on the property so it is not known whether a follow-up program was undertaken.

REGIONAL GEOLOGY

The only previous work undertaken pertaining to the general area of the Crawford claims was that of Reesor (1996, 1993) for the east side of Kootenay Lake. The stratigraphy of the Purcell Supergroup strata has been well described to the east by Höy (1993) and the Purcell and Windermere Supergroup to the north by Pope (1990).

Stratigraphy

Windermere Supergroup

The Windermere Supergroup varies in thickness in the Toby Creek area, from 80 metres to over 3 kilometres and is in sharp contact with the underlying Belt-Purcell Supergroup across an unconformity with considerable topography, interpreted as a result of a local basement high, the "Windermere High" (Reesor 1973). The Windermere Supergroup was deposited above this unconformity and consists of a basal conglomeratic unit, the Toby Formation, and the overlying argillite and pebble conglomerate dominated Horsethief Creek Formation.

Horsethief Creek Group

The Horsethief Creek Group has been subdivided into five lithofacies. These lithofacies define a rudimentary stratigraphy of facies within the Horsethief Creek Formation as individual lithological units are inconsistent due to rapid lateral thickness and facies variations.

The lithofacies identified in the Horsethief Creek Formation are as follows:

- a) siltstone-argillite - dominant in the lower half of the Horsethief Creek Formation and separate the remaining lithofacies throughout the formation. This lithofacies consists of thick sequences of thin bedded (1 to 10 cm), graded siltstone and argillite and finely laminated (1 to 5 mm), black, green and grey argillite.
- b) black carbonate - an easily traced marker used to identify and map the base of the Horsethief Creek Formation consisting of thin bedded (5 to 20 cm), dark grey to black limestone, with variable quartz sand and silt in a calcitic matrix, and thin calcareous quartz-arenite beds.
- c) dolomite - buff weathering dolomite, up to 30 metres thick, dolomite pebble-conglomerate beds and dolomite supported quartzite occur throughout the Horsethief Creek Formation.
- d) quartz feldspar arenites and pebble conglomerates - consist of pebble conglomerates comprised of grain-supported, moderately sorted crystalline quartz and quartz feldspar clasts with variable red jasper, green to grey argillite, quartzite and dolomite clasts in a quartz, feldspar, carbonate, sericite and chlorite matrix. Clasts are generally 1 to 2 centimetres in diameter but may exceed 10 centimetres in length. Coarse arenite beds are similar to the pebble

conglomerates but have a greater proportion of matrix and are generally poorly sorted.

- e) red and varicoloured argillites - are present at the top of the Horsethief Creek Formation and consist of variably coloured argillites with interbedded pink carbonate, and varicoloured impure arenites.

Mesozoic

Granitic Intrusions

Cretaceous intrusives of broadly “granitic” composition are present in a belt extending from the westernmost Rocky Mountains to Kootenay Lake, northward to the Baldy Batholith. Intrusions range from small dykes and sills to larger intrusive complexes such as the Mt. Skelly Batholith and are collectively referred to as the Bayonne Magmatic Belt (or Suite).

“Intrusive rocks ... include a number of small post kinematic mesozonal quartz monzonite, monzonite and syenitic plutons, numerous small quartz monzonite to syenite dikes and sills probably related to these stocks, and late mafic dikes. The Kiakho and Reade Lake stocks, two of the larger of the mesozonal plutons, cut across and apparently seal two prominent east-trending faults that transect the eastern flank of the Purcell anticlinorium, and hence place constraints on the timing of latest movement on these faults.

The Kiakho stock is exposed on the heavily wooded slopes of Kiakho Creek approximately 10 kilometres (west-southwest) ... of Cranbrook ... Exposures consist mainly of large, fresh angular boulders of boulder fields. Although contacts with country rock were not observed, regional mapping indicates that it intrudes clastic rocks of the Aldridge and Creston formations. The distribution of outcrops and a pronounced aeromagnetic anomaly indicate that it cuts the east-trending Cranbrook normal fault with no apparent offset. ...

The Kiakho stock is similar to the Reade Lake stock with the dominant phase being a light grey, medium-grained quartz monzonite. It is generally equigranular but grades into a hypidiomorphic granular porphyritic phase with prominent plagioclase and light grey to flesh-coloured potassic feldspar phenocrysts; both are up to several centimetres in diameter in a granular groundmass of white subhedral plagioclase, light grey potassic feldspar, quartz and black hornblende” (Höy 1993).

The Bayonne Granitic Suite is a composite batholith comprised of a number of smaller Jurassic to Cretaceous age granitoid stocks and plutons which extends from near the International Boundary across Kootenay Lake. On the east side of the Kootenay Lake, the Bayonne Granitic Suite locally includes the Mount Skelly Pluton, a biotite (hornblende) monzogranite with megacrysts of potassium feldspar (Reesor 1996). Rice (1941) grouped these granitoids under the broad heading of the Bayonne Batholith, as described below.

“The Bayonne batholith varies in composition from a granite to a calcic granodiorite; the average composition is that of a fairly alkaline granodiorite. ... Much of the rock has an equigranular texture, but a porphyritic phase occurs in many places, at some of which phenocrysts of potash feldspar 2 or 3 inches long are present. The potash feldspar may be orthoclase or microcline and in some specimens both occur. The plagioclase is oligoclase, generally well twinned and frequently in zoned crystals. Dark brown biotite is the only ferromagnesian mineral abundant, but grains of hornblende occur in rare instances. The usual accessories are present. Sericite and epidote are the commonest secondary minerals, but neither occur in significant amounts except where the rock has been altered.

A marked feature of the Bayonne batholith is its highly variable nature. This is observable not only in the range of composition but in the appearance of the rock. Coarse-grained and fine-grained, porphyritic and non-porphyritic, pink and light or dark grey phases may occur in a single exposure, in some places in streaks and patches. Masses of pegmatite and dykes of pegmatite and aplite occur everywhere. Some of the pegmatite dykes are over 100 feet wide. A few large crystals of blue-green beryl, pink garnet, magnetite, and a little black tourmaline were seen in these pegmatites.

Large inclusions of granitized sediments are locally abundant. ... These inclusions vary in size from a foot to some hundreds of feet. Alteration is severe, but the sedimentary nature of the original rock is, in most cases, still recognizable and the boundary between the granite and the inclusion is generally fairly sharp. Other inclusions or xenoliths (sic.) from a few inches to a foot long also occur, which can readily be distinguished from the first type mentioned. They parallel one another, are darker coloured, their original texture and composition has been more or less completely altered, they are fairly uniform in size, and they usually grade imperceptibly into the granite. They are more widely distributed, indeed very few exposures of any size were examined that did not contain some of these xenoliths (sic.), and in places they are extremely abundant. The xenoliths (sic.) are often most common in the porphyritic phases and scarcer in the non-porphyritic phases of the granite ...“.

Structure

Four major phases of deformation have been identified in the Toby Creek area, Helikian-Devonian extension (D1), Jurassic-Paleocene contraction (D2-D3) and Eocene extension (D4).

The first phase of deformation resulted in unconformities at the base of the Dutch Creek and Mount Nelson Formations (D1a) and the unconformity at the base of the Windermere Supergroup (D1b). Thinning of Paleozoic strata onto the Windermere High is interpreted to reflect the effects of D1c deformation together with the development of small fault-bounded sub-basins.

Contraction during the Columbian (D2) and Laramide (D3) orogenies resulted in a series of northeast vergent thrust faults and the development of a regional foliation (S1). Three major thrust sheets are evident in the Toby Creek area with one, the Mount Nelson thrust sheet, comprised of four smaller fault panels. The three major thrust sheets represent out-of-sequence faults, having propagated toward the hinterland, carried in the hanging wall of the Purcell Thrust.

Contraction during D2 and D3 produced east-vergent imbricate thrust faults and west vergent backthrusts. Many of these faults were subsequently reactivated during the fourth phase (D4) of deformation. High angle brittle faults are also a result of D4.

LOCAL GEOLOGY

Stratigraphy

The CRAWFORD property is underlain by clastic sediments of the Horsethief Creek and Hamill Groups of Upper Proterozoic age. The rocks, which appear to be younging westward have been regionally metamorphosed and altered to light-grey phyllites, muscovite-plagioclase+/-andalusite+/-chlorite schists, quartz pebble conglomerates and light-grey quartzites (Wright 1980). The majority of outcrops, foliation in schists and quartz veins in all rock types generally strike NNE and dip fairly steeply toward the east. These rocks have been intruded by a large elliptical quartz monzonite stock of Cretaceous(?) age.

Given the presence of the large Fry Creek Batholith nearby to the northwest, the felsic intrusive on the property may be a related satellite intrusion or apophyse. Alternatively, it may be a small unrelated Cretaceous intrusion, having been intruded during a regional Mesozoic intrusive event.

Structure

The structure of the CRAWFORD area is dominated by its position on the western flank of the Purcell Anticlinorium, a north plunging fold of regional significance. The Purcell Anticlinorium is allochthonous with respect to North American cratonic basement, having been transported northeastward in the hanging wall of the Purcell Thrust. This major structure has been complicated slightly by a number of regional and local faults, discussed below with reference to the Kootenay Lake mapsheet of Reesor (1996). An early folding event has been proposed for early structures interpreted to have developed in the Late Proterozoic during the Goat River Orogeny (Höy 1993).

The prominent faults in the CRAWFORD area are interpreted to be predominantly the result of the Laramide orogeny, characterized by east-verging, west-dipping thrust faults. The major fault system of the area is the St. Mary / Hall Lake fault system, interpreted to be a long lived fault initiated in the Late Proterozoic as a growth fault and periodically active at least into the Laramide orogeny. Eastward directed movement across the St. Mary / Hall Lake fault resulted in steeply dipping strata on the western limb of the Purcell Anticlinorium being juxtaposed against relatively shallowly to moderately dipping strata closer to the hinge axis.

Significant dip displacement is indicated across the fault east of Sanca Creek where Proterozoic lower

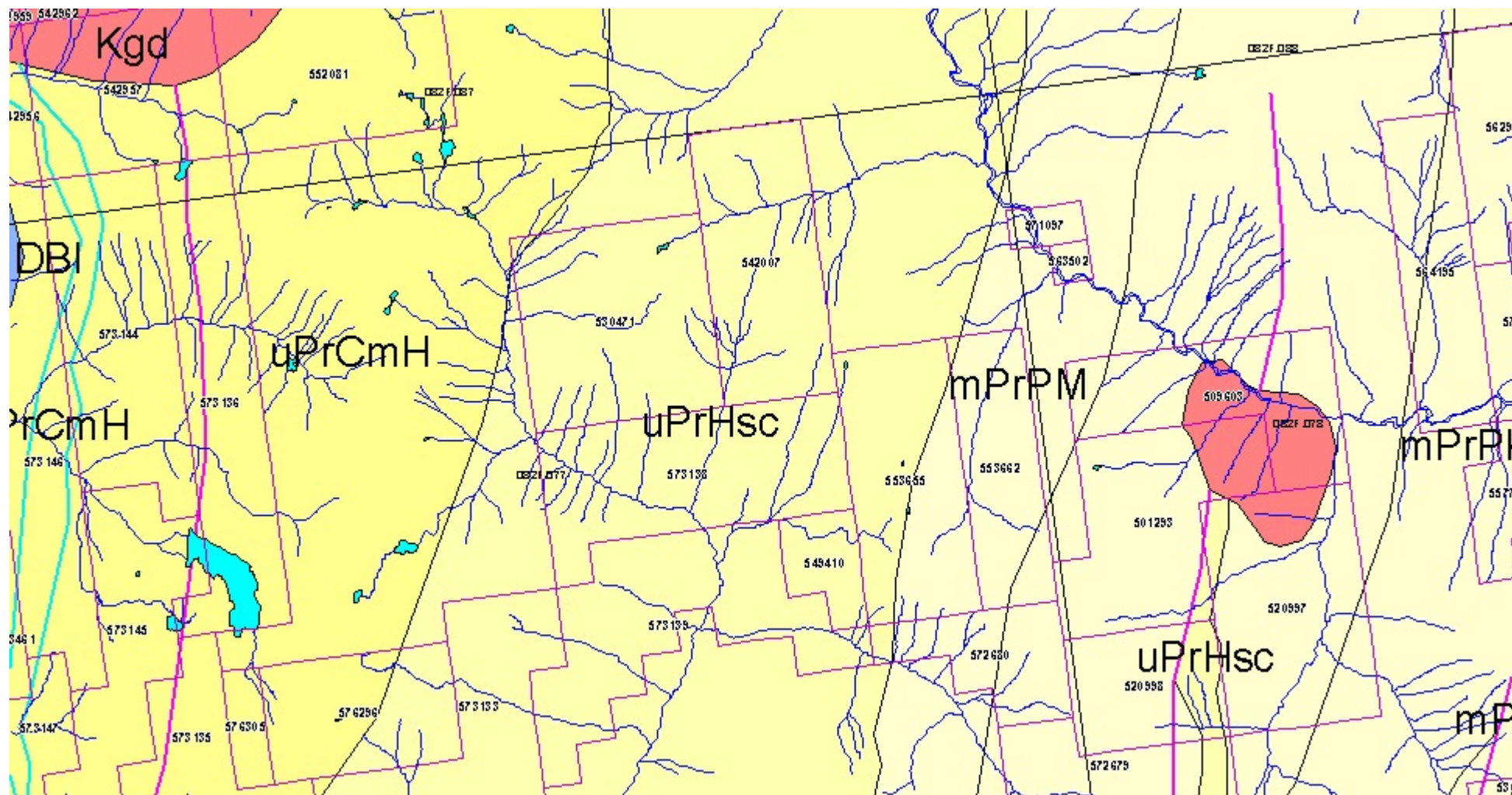


Figure 4 - Geological map of the Crawford area. Property underlain by Upper Proterozoic Horsethief Creek Group (uPrHsc) coarse clastic sedimentary rocks with Mount Nelson Formation to east, Cambrian Hamill Group to west and is located approximately 5 km south of the Fry Creek Batholith (at top of figure) and 8 km west of the Sawyer Stock. Approximate scale 1: 50,000. (Taken from The MapPlace)

Creston strata has been juxtaposed against early Paleozoic Cambrian Eager Formation strata. Later thrust faults are evident in the hanging wall of the St. Mary / Hall Lake fault. The Redding Creek fault is locally significant fault. It is a west dipping, east verging thrust fault that juxtaposes middle Creston strata against the lower member of the Coppery Creek group. A number of smaller, normal faults are indicated in the hanging wall of the Redding Creek Fault, all of which appear to have minor dip (and probably strike-slip) movement. All of the faults in the hanging wall of the St. Mary / Hall Lake fault are interpreted to be older than the Cretaceous Mount Skelly Pluton (Bayonne Magmatic Belt) as all are truncated at the contact of the pluton.

DETAILED/PROPERTY GEOLOGY

The following has been taken from Wright (1980):

“In the NE corner of the property is a 300m wide unit of silvery-grey phyllite, consisting of 90% muscovite and 10% plagioclase, with a NNE strike and steep eastward dip. In the upper part of the unit, the phyllite is interbedded with fine to coarse quartz pebble conglomerate beds up to 50 cm thick. The quartz pebbles which range from 1 to 12 cm in length, averaging 2cm, show a fining westward sequence. The phyllites grade along strike into silvery-grey muscovite-plagioclase schists which are simply a coarser-grained version of the same rock with similar foliation and composition. Fine to coarse quartz pebble conglomerates, as above, are interbedded with the schists. Toward the intrusive contact, the rocks become coarser-grained with the muscovite-plagioclase schists grading into fine to medium-grained and andalusite-muscovite-plagioclase schists. These rocks are typically light to dark grey, weathering medium to dark grey, with 70-80% muscovite, 5-15% plagioclase, 5-7% andalusite, and less than 1% disseminated pyrite.

Adjacent to the quartz monzonite contact, the schists become chloritized, with an average composition of 60% muscovite, 10-20% chlorite, 5-15% plagioclase, 5-15% andalusite and 1% pyrite. These rocks are light to dark greenish-grey, weathering dark reddish-green to grey. Infrequent quartz veins up to 4 cm thick parallel foliation, averaging 022 /60 E.

Within the schists, unit 2a constitutes a 20-25 metre-wide band of medium quartz pebble conglomerate in a coarse-grained quartzite matrix. The milky white quartz pebbles, which range in size from 1-5cm and average 2cm, make up approximately 15-20% of the rock are well-rounded and poorly sorted. Also within the schists, in the eastern half of the property is a 25-50 metre-wide unit of pinkish-white to yellowish-grey, medium-grained quartzite, in places containing less than 1% disseminated pyrite.

Along the western edge of the property is a thick resistant unit of light grey, fine to medium-grained quartzite of the Hamill Group. Manganese and iron staining, and specular hematite are abundant on fracture surfaces. Near the contact with the quartz monzonite intrusion, coatings of pyrolusite and recrystallized quartz are abundant along fractures. In places, small dark red crystals of magnetite make up less than 2% of the rock.

The quartz monzonite, which forms an elliptical stock in the centre of the property is typically white to pinkish-grey, weathering medium to dark grey, fine to coarse-grained, and consists of 30-35% plagioclase, 30-35% K-feldspar, 30% quartz and 5-10% biotite. Aplite veins are common, ranging from 1 to 50 cm in width with an average attitude of 065 /80 S. Milky white quartz veins up to 20 cm thick are less abundant than the aplite veins, and exhibit sharp contacts with random orientations throughout most of the intrusion. Near the contact with the schists, especially along the southern contact, the percentage of biotite in the quartz monzonite rises sharply to approximately 30% giving the rock a dioritic appearance. The contact between the quartz monzonite and the Hamill quartzite is gradational, from unaltered quartz monzonite to kaolinized quartz monzonite, then pinkish quartzite and finally light-grey quartzite.

Alteration is most intense in the quartz monzonite within three hundred metres of the contact. The rock is pervasively sericitized and kaolinized, with silicification, iron and manganese staining along joints. The central portion of the intrusion, where exposed, is relatively fresh.

Mineralization

Molybdenite mineralization, as disseminated flakes and rosettes in quartz veins was noted in several localities within the altered margin of the intrusion. Trace amounts of powellite were also noted from scattered localities”.

2007 PROGRAM

The 2007 program consisted of an initial soil sampling program with which to evaluate the property and as a follow-up to the 2006 Aeroquest airborne geophysical survey, completed between May 15th and September 25th, 2007.

A total of 196 soil samples were recovered along two contour lines and a road traverse, all oriented at a high angle to the structure underlying the property. Samples were taken at a 50 m spacing, all recovered from the "B" Horizon. Holes were dug by hand using a mattock to a depth generally between 10 and 25 cm below surface. Samples were placed in Kraft soil envelopes, air dried to eliminate excess water content and shipped to Acme Analytical Laboratories Ltd in Vancouver, BC for analysis using SS80 preparation and Group 1DX analysis.

A copy of the report is included in Appendix B.

RESULTS

Generally, the results returned from this preliminary soil program on the Cvrawford property returned very little on which to follow up. Only two single sample anomalies having values greater than 8 ppb were returned for gold. Two areas had very weakly anomalous silver (values between 0.5 and 1 ppm). Tungsten values are slightly elevated with regard to a regional background. Lead values are at or below levels interpreted to be regional background. The only element having possible anomalous values is zinc.

Zinc

A number of single or multi-sample samples anomalous for zinc (Fig. 6) were documented from the property. The samples are located in the approximate centre of the property and may define a poorly defined northwest trend. These weak results comprise the only results of any possible significance from this program.

Crawford 2007

DYNAMIC EXPLORATION LTD

SAMPLE LOCATION MAP

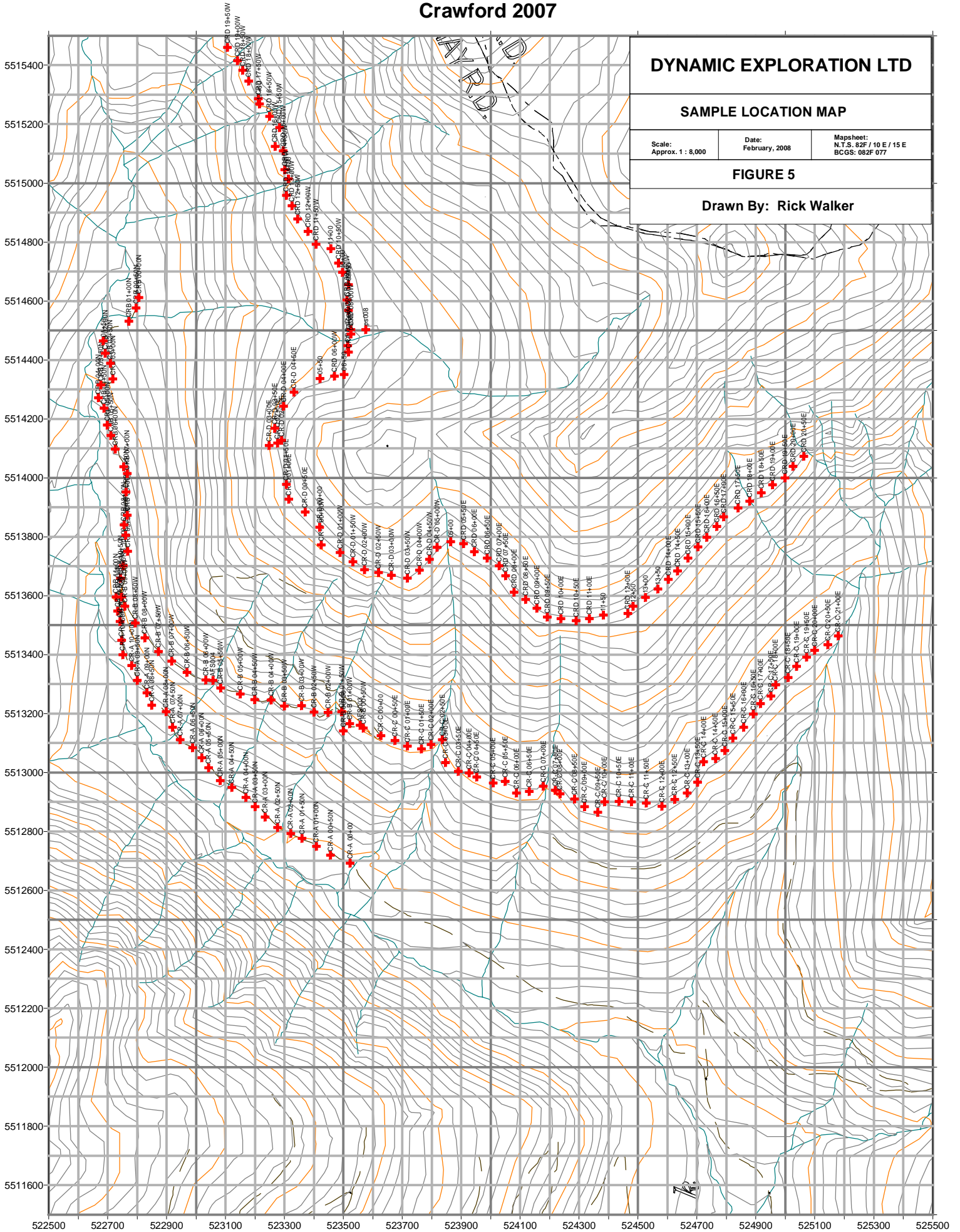
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Approx. 1 : 8,000

Date:
February, 2008

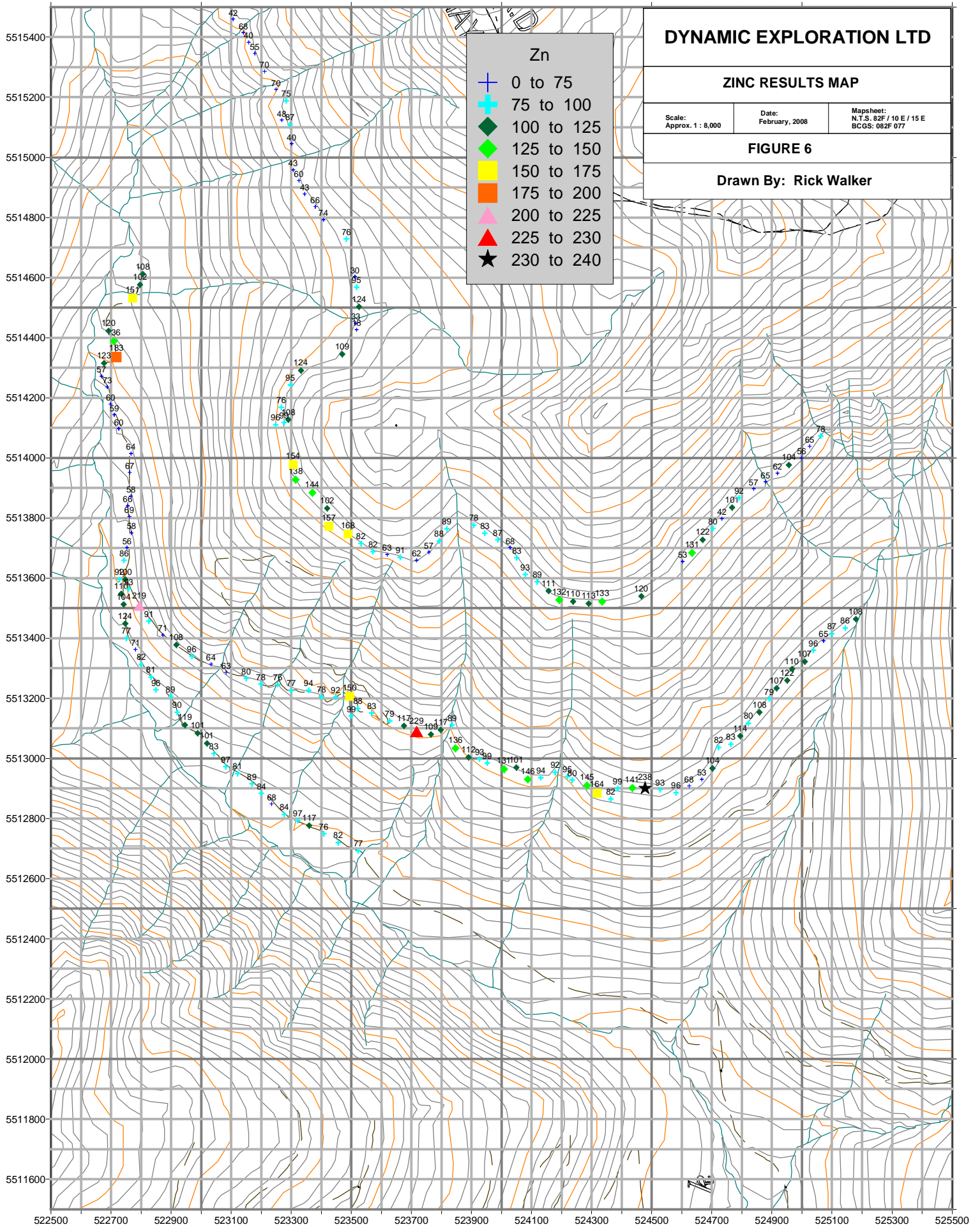
Mapsheet:
M.T.S. 62F / 10 E / 15 E
BCGS: 082F 077

FIGURE 5

Drawn By: Rick Walker



Crawford 2007



DISCUSSION

The property covers a mapped exposure of felsic intrusive material (Reesor 1996) tentatively interpreted to correlate to the Bayonne Magmatic Belt, comprised of Cretaceous age felsic intrusions extending from the Baldy Batholith north of Kamloops to the International Boundary with the United States south of Creston. The felsic intrusive may represent an apophyse of the Fry Creek Batholith, a large intrusive complex located nearby to the northwest. Alternatively, it may represent a separate intrusion similar to the Hall Lake Stock, located approximately 20 kilometres to the southeast.

Previous workers (Wright 1980) carried out a program of geological mapping and soil sampling. The work consisted of detailed geological mapping and prospecting of the property, and soil sampling on a grid. Sampling consisted of 617 soils and 8 stream silts with all samples analyzed for Cu, Pb, Zn, Mn, Mo and W. Results showed scattered anomalous values for Cu, Pb, Zn, Mn and W which could not be related to known mineralization. Anomalous molybdenum values were found to correspond to a large quartz monzonite intrusion containing several molybdenite showings. The molybdenite mineralization, as disseminated flakes and rosettes in quartz veins was noted in several localities within the altered margin of the intrusion. Trace amounts of powellite were also noted from scattered localities.

Based on these results, a decision was made to retain the Crawford property for further evaluation.

CONCLUSIONS

Initial interpretation of the Aeroquest Limited airborne geophysical data, together with limited data from previous programs, strongly suggest further evaluation of the property is warranted.

In 1980 Cominco noted molybdenite mineralization, as disseminated flakes and rosettes in quartz veins in several localities within the altered margin of the intrusion. Trace amounts of powellite were also noted from scattered localities.

Further work was recommended to determine the significance of the anomalies. No further work appears to have been filed by Cominco on the property so it is not known whether a follow-up program was undertaken.

Further evaluation of the Crawford property is recommended particularly in light of the presence of molybdenite mineralization documented by Cominco which may not have been adequately followed-up.

RECOMMENDATIONS

1. Compilation of results from previous programs should be undertaken to build an initial database of with which to continue evaluation of the property;
2. Undertake a soil sampling program. Detailed contour sampling should be undertaken throughout the property to provide better coverage, with smaller grids established to develop better resolution in areas of anomalous results;
3. Geological mapping should be undertaken to:
 - a) identify and/or re-establish mineralized horizons,
 - b) identify and/or confirm the stratigraphy present on the property,
 - c) provide better structural control for the property.
 - d) obtain rock and/or chip samples of mineralized horizons identified on the property;
4. Consider diamond drilling to test surface anomalies identified on the basis of additional soil and rock sampling and/or sub-surface anomalies identified from airborne and/or ground-based geophysical surveys.

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Appendix A

Statement of Qualifications

STATEMENT OF QUALIFICATIONS

I, Richard T. Walker, of 2601 - 42nd Ave South, Cranbrook, B.C., hereby certify that:

- 1) I am a graduate of the University of Calgary of Calgary, Alberta, having obtained a Bachelors of Science in 1986,
- 2) I obtained a Masters of Geology at the University of Calgary of Calgary, Alberta in 1989;
- 3) I am a member in good standing with the Association of Professional Engineers and Geoscientists of the Province of British Columbia;
- 4) I am the Vice President of Exploration for Jasper Mining Corporation, with offices at 1020, 833 - 4th Ave South, Calgary, Alberta;
- 5) I am the author of this report which is based on a soil sampling program completed between May 15th and September 25th, 2007;
- 6) I have a direct interest in Jasper Mining Corporation; and
- 7) I hereby grant my permission to Jasper Mining Corporation to use this report, or any portion of it, for any legal purposes normal to the business of the firm, provided the excerpts used do not materially deviate from the intent of this report as set out in the whole.

Dated at Cranbrook, British Columbia this 22nd day of February, 2008.



Richard T. Walker, P.Geo

APPENDIX B

SOIL SAMPLE RESULTS

ELEMENT	Mo	Cu	Pb	Zn	Ag	Ni	Co	Mn	Fe	As	U	Au	Th	Sr	Cd	Sb	Bi	V	Ca	P	La	Cr	Mg	Ba	Ti	B	Al	Na	K	W	Hg	Sc	Tl	S	Ga	Se	Te	Sn	Zr	
SAMPLES	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	%	ppm	ppm	ppb	ppm	ppm	ppm	ppm	ppm	ppm	%	%	ppm	ppm	%	%	ppm	%	%	%	ppm	ppm	ppm	ppm	%	ppm	ppm	ppm	ppm	ppm	ppm	ppm
G-1	0.5	2.1	3.3	46	<1	5.1	4	494	1.78	<5	2.5	0.6	3.9	47	<1	<1	0.1	35	0.44	0.079	6	35	0.56	207	0.115	1	0.82	0.055	0.46	0.2	1.9	1.9	0.3	<0.5	4	<5	<1	<1	1.1	
CR-A 12+00N	27	36.1	15.1	110	0.1	34.7	18.2	4257	6.54	39.9	16.3	1.4	4.3	64	0.5	0.2	3.7	30	0.27	0.091	24	25	0.52	150	0.086	2	2.29	0.01	0.21	1.1	0.04	3.6	0.3	<0.5	6	0.6	<1	1	2.2	
CR-A 11+50N	42.6	37.3	19	104	0.2	27.2	14.9	3404	6.57	50.1	13	2.1	3	37	0.5	0.4	3.1	34	0.2	0.101	23	24	0.42	111	0.087	2	2.35	0.009	0.21	0.8	0.07	3.2	0.3	<0.5	8	0.7	<1	1	2.7	
CR-A 11+00N	27.2	32.4	13.2	124	0.8	30	9.6	769	3.58	4.4	3.6	1.3	3.6	20	0.3	0.1	4.9	38	0.14	0.125	23	23	0.52	123	0.101	1	2.6	0.005	0.28	1.5	0.07	2.7	0.3	<0.5	11	0.6	<1	1	3.3	
CR-A 10+50N	20.7	30.2	14.8	77	0.1	20.1	9.6	652	2.76	4.8	5.7	0.9	3.9	12	0.3	0.2	6	31	0.06	0.076	29	18	0.42	107	0.091	1	3.01	0.007	0.25	1.8	0.05	2.6	0.3	<0.5	7	0.9	<1	1	4.2	
CR-A 10+00N	5.6	10.1	10.7	71	<1	10.6	4.6	371	2.12	2.9	1.7	5.6	5.3	10	0.2	0.2	6.9	26	0.13	0.104	21	12	0.34	117	0.075	1	1.75	0.005	0.21	3.1	0.05	1.9	0.2	<0.5	6	<5	<1	1	1.8	
CR-A 09+50N	6	16	11.6	82	<1	17.1	5.8	337	2.47	3	1.9	1.4	5.1	10	0.2	0.1	6.1	30	0.09	0.08	21	19	0.47	124	0.083	1	2.56	0.005	0.21	2.6	0.06	2.7	0.3	<0.5	6	0.5	<1	1	2.3	
CR-A 09+00N	3.7	15.9	11.1	81	<1	18.3	5.8	295	2.43	2.7	1.4	1.8	4.7	9	0.2	0.2	4.5	33	0.04	0.054	16	19	0.44	109	0.095	1	2.2	0.005	0.18	1.1	0.04	2.5	0.3	<0.5	7	<5	<1	1	4.9	
CR-A 08+50N	2.8	14.9	13.2	96	<1	8.9	5.5	531	2.13	3.8	1.2	0.9	2.4	10	0.4	0.3	3.3	34	0.06	0.105	10	11	0.2	94	0.114	1	3.72	0.006	0.09	0.8	0.1	1.9	0.2	<0.5	10	0.7	<1	1	15.8	
CR-A 08+00N	3.2	12.6	11.8	89	0.2	8.2	4.9	809	2.16	3.3	1.2	1.9	0.9	12	0.6	0.3	2.6	28	0.08	0.153	10	10	0.2	95	0.174	1	2.49	0.006	0.09	0.9	0.1	1.2	0.2	<0.5	9	0.6	<1	1	4.9	
CR-A 07+50N	5.1	13.2	11.6	90	<1	15.1	6.4	588	2.22	2.9	1.9	1.8	5.1	10	0.2	0.1	7.7	26	0.14	0.118	22	13	0.4	154	0.082	1	1.94	0.006	0.25	2.9	0.03	2	0.3	<0.5	6	<5	<1	1	2.5	
CR-A 07+00N	6	20.7	14.8	119	<1	23.8	7.8	896	2.61	3.7	5.5	1	4.5	16	0.4	0.2	6.5	33	0.1	0.092	26	17	0.41	152	0.106	1	2.63	0.007	0.24	1.7	0.05	2.7	0.3	<0.5	8	0.5	<1	1	4.2	
CR-A 06+50N	6.2	16.8	11	101	<1	22.3	8.5	529	2.45	3.7	3.3	1.1	6.4	19	0.2	0.1	6.8	29	0.2	0.129	27	16	0.46	171	0.087	1	2	0.007	0.3	2.9	0.03	2.5	0.3	<0.5	6	<5	<1	1	1.8	
CR-A 06+00N	4.3	15.5	17.9	101	0.1	15.5	6.8	1077	2.21	2.9	2.4	0.7	3	17	0.5	0.3	5.9	28	0.11	0.092	23	11	0.32	172	0.088	1	2.16	0.006	0.19	1.4	0.06	1.7	0.3	<0.5	8	<5	<1	1	2.8	
CR-A 05+50N	5.1	13.3	14.9	83	<1	15.5	6.8	757	2.14	3.9	2.3	3.7	3.7	19	0.3	0.2	5.1	23	0.14	0.076	21	12	0.38	144	0.076	<1	1.49	0.007	0.2	1.4	0.02	1.7	0.2	<0.5	6	<5	<1	1	1.2	
CR-A 05+00N	7.1	35.5	19.1	97	0.1	22.8	9.9	803	2.64	3.2	11.1	0.8	4.3	14	0.4	0.2	6.4	33	0.06	0.074	51	16	0.43	176	0.108	1	2.46	0.007	0.27	1.7	0.04	3.1	0.3	<0.5	9	0.5	<1	1	3.8	
CR-A 04+50N	5.7	24.3	14.3	81	0.1	20.9	9.7	773	2.67	3.3	3.2	3.8	3.9	14	0.2	0.2	6.3	33	0.09	0.091	25	19	0.44	117	0.093	1	2.07	0.007	0.26	1.8	0.03	2.4	0.3	<0.5	8	<5	<1	1	1.6	
CR-A 04+00N	6.5	21.9	14.4	89	0.1	22	10.2	1051	2.7	3.3	4	1.1	3.9	19	0.3	0.2	6.2	33	0.11	0.077	29	19	0.49	152	0.1	1	2.18	0.008	0.27	1.6	0.03	2.4	0.3	<0.5	8	<5	<1	1	2.1	
CR-A 03+50N	5.8	19	13.9	84	0.1	15.8	7.8	846	2.31	3	3.8	1.2	3.3	19	0.4	0.2	5.4	28	0.15	0.089	28	14	0.39	147	0.085	1	1.87	0.007	0.22	1.6	0.02	2.1	0.2	<0.5	7	<5	<1	1	2	
CR-A 03+00N	4.5	17.7	17	68	0.3	8.5	7.8	1735	2.63	2.7	1.7	1.3	1	22	0.6	0.3	3.6	30	0.12	0.145	12	12	0.15	113	0.081	1	2.16	0.007	0.08	0.9	0.1	1.1	0.2	0.06	12	0.7	<1	1	4.1	
CR-A 02+50N	6.1	17.7	20	84	<1	19.3	8.9	770	2.3	3	2.6	1.1	4.3	24	0.3	0.2	5.3	29	0.15	0.084	28	17	0.43	204	0.087	1	1.78	0.007	0.29	1.9	0.04	2.3	0.3	<0.5	7	<5	<1	1	1.1	
CR-A 02+00N	7.2	26.9	18.8	97	0.2	20.6	11.9	1154	2.86	3.1	5.4	1.5	3.1	21	0.6	0.2	5.2	34	0.13	0.078	33	19	0.45	157	0.111	1	2.06	0.008	0.29	1.5	0.04	2.3	0.3	<0.5	10	0.5	<1	1	3.2	
RE CR-A 02+00N	7.1	26.7	18.8	96	0.2	20.8	11.5	1170	2.87	3.1	5.5	1.4	2.9	20	0.7	0.2	6	33	0.13	0.076	33	19	0.44	155	0.111	1	2.04	0.008	0.29	1.3	0.04	2.3	0.3	<0.5	10	0.5	<1	1	3.3	
CR-A 01+50N	8.4	25.2	18.1	117	0.2	17.1	6.9	799	3.1	3.9	17.8	1.7	6.2	36	0.3	0.2	8.1	36	0.15	0.101	46	11	0.44	359	0.165	1	4.38	0.016	0.34	1.9	0.08	3.4	0.4	0.06	15	0.8	<1	2	20.7	
CR-A 01+00N	4.5	14.8	17.4	76	<1	17	8	657	2.09	2.6	2.4	0.9	4.3	24	0.4	0.2	4.3	26	0.19	0.074	26	16	0.42	177	0.08	1	1.55	0.007	0.27	1.6	0.03	1.9	0.3	<0.5	6	<5	<1	1	0.8	
CR-A 00+50N	7.2	33.5	20.5	82	0.2	22.6	15	1109	2.82	4.5	19.7	1.6	2.8	25	0.5	0.2	4.2	26	0.13	0.091	75	19	0.5	158	0.067	1	3.04	0.008	0.12	1.7	0.06	2.5	0.2	0.07	8	0.8	<1	1	4.1	
CR-A 00+00	5	11.8	15.5	77	<1	11.5	6.3	781	1.91	2.5	2.9	5.1	3.6	36	0.4	0.2	8.8	23	0.2	0.074	32	9	0.33	204	0.07	1	1.62	0.007	0.21	2.1	0.03	1.6	0.2	<0.5	6	<5	<1	1	1	
CR-B 09+50W	10.3	28.6	12.7	100	<1	24.5	11.7	7672	11.48	95.8	23.6	1	5.4	10	0.3	0.2	5.5	31	0.03	0.099	23	24	0.48	174	0.1	1	1.93	0.008	0.32	1.2	0.03	5.3	0.3	<0.5	6	0.7	<1	1	2.6	
CR-B 09+00W	13.3	61.4	12.1	93	<1	51.6	18.3	911	3.3	5.3	8.5	0.8	7.7	13	<1	0.1	3.9	39	0.04	0.064	29	34	0.72	126	0.116	1	3.74	0.008	0.35	1.3	0.03	4.5	0.4	0.06	8	0.6	<1	1	2.1	
CR-B 08+50W	22.6	47.7	13.6	219	0.2	70.4	14.5	816	3.56	4.7	6.5	0.9	4.8	14	0.4	0.2	2.5	42	0.08	0.1	27	41	0.73	121	0.114	1	3	0.008	0.39	0.7	0.05	4.4	0.4	<0.5	8	0.5	<1	1	2.4	
CR-B 08+00W	4.7	28.9	18.2	91	0.2	21.5	10.1	912	2.82	4.7	1.6																													

CR-D 00+50W	5	10.5	21.7	155	<.1	8.2	4.6	573	3.24	7	1.5	2.7	4.5	29	0.3	0.5	8.5	38	0.14	0.08	19	11	0.36	138	0.084	1	1.76	0.007	0.17	3.7	0.04	2	0.3	0.06	12	<.5	<.1	2	0.7
RE CR-D 00+50W	5.1	10.3	21.6	157	<.1	8.5	4.7	560	3.24	7	1.4	5.2	4.5	29	0.4	0.4	7.9	38	0.14	0.081	20	11	0.37	145	0.084	1	1.83	0.007	0.16	3.8	0.04	2	0.3	0.06	11	<.5	<.1	2	0.6
CR-D 00+00	3.3	11.2	16.6	102	<.1	13.6	5.3	278	2.95	5.9	1	1.2	5.7	9	0.2	0.4	4.4	42	0.04	0.1	12	15	0.29	83	0.122	1	1.95	0.008	0.12	2.6	0.03	2.1	0.3	<.05	11	<.5	<.1	2	7.9
CR-D 00+50E	5.2	26.7	19.8	144	0.2	32.8	15	761	3.64	8.1	1.9	1.5	6.7	11	0.2	0.7	4.3	48	0.05	0.074	14	30	0.49	114	0.126	1	3.48	0.008	0.19	3.1	0.04	3.2	0.6	0.06	12	0.5	<.1	1	7
CR-D 01+50E	4.5	24.6	25.8	138	<.1	25.7	13.8	2638	3.33	5.4	1.3	0.8	5.3	13	0.2	0.9	5.2	46	0.07	0.075	15	24	0.44	129	0.126	1	2.4	0.008	0.18	1.4	0.04	2.6	0.4	<.05	12	<.5	<.1	2	3
CR-D 01+50E	4.9	35.2	18	154	0.2	50.9	21.1	761	4.02	3.5	1.6	1.7	6.7	11	0.3	0.3	4.3	50	0.06	0.076	17	37	0.64	109	0.144	3	3.94	0.009	0.2	2.9	0.06	3.6	0.4	<.05	12	0.5	<.1	1	8.8
CR-D 02+50E	4.4	30.2	18.4	108	<.1	38.9	15.4	1050	3.5	4.1	1.3	5.4	6.4	8	0.2	0.6	4.4	46	0.04	0.071	16	35	0.49	98	0.111	1	2.78	0.006	0.18	1.7	0.06	3	0.3	<.05	11	<.5	<.1	1	2.8
CR-D 02+50E	4.6	46.4	17.9	99	0.2	36.1	13.9	338	3.19	5.1	2	3	7.7	8	0.1	0.3	6.1	38	0.04	0.08	15	23	0.49	106	0.127	1	3.68	0.008	0.2	3.2	0.07	3.4	0.4	<.05	10	0.8	<.1	1	21.8
CR-D 03+50E	4.7	15.2	19	96	0.3	15.4	5.5	352	2.93	6.5	1.4	3.1	6.6	7	0.2	0.6	6.4	36	0.03	0.092	12	19	0.35	81	0.106	1	3.43	0.007	0.12	3.4	0.11	2.4	0.2	<.05	10	0.7	<.1	1	12.1
CR-D 03+50E	3.5	11.3	23.6	76	0.1	10.4	4.9	363	3.28	4.8	0.8	2.4	4.7	7	0.2	0.6	5.3	52	0.03	0.074	11	20	0.26	77	0.14	1	2.23	0.008	0.1	2.1	0.06	1.9	0.2	<.05	14	<.5	<.1	2	6.1
CR-D 04+50E	3.1	8.7	19.7	95	<.1	14.7	6.6	1682	2.8	4.1	0.7	1.4	3.9	16	0.3	0.4	4.7	50	0.08	0.066	12	25	0.34	183	0.144	1	1.19	0.008	0.16	1.2	0.02	1.9	0.3	<.05	13	<.5	<.1	2	1
CR-D 04+50E	6.8	15.7	16	124	0.3	16.3	8.7	977	2.97	7.6	5.3	33.8	3.7	12	0.4	0.3	4.7	35	0.07	0.114	17	21	0.44	103	0.103	1	2.81	0.009	0.13	2.4	0.07	2.3	0.2	<.05	10	<.5	<.1	1	4.6
STANDARD DS7	21.2	112.8	72.5	425	0.9	58.4	10.3	636	2.45	52.1	5.4	73.5	4.9	77	6.9	6.6	4.9	89	1.01	0.087	14	212	1.09	387	0.126	41	1.05	0.098	0.45	4.1	0.21	2.9	4.4	0.26	5	3.9	1	6	5.6

Appendix C

STATEMENT OF EXPENDITURES

STATEMENT OF EXPENDITURES

The following expenses were incurred on behalf of the Crawford project between May 15th and September 25th, 2007.

PERSONNEL

Field Manager - 7 days at \$350 / day	\$ 2,450.00
Field Crew - 19 days @ \$250 / day	\$ 4,750.00
Sub-Total	\$ 7,200.00

EQUIPMENT RENTAL

4WD Truck - mileage - 1,994 km @ \$0.75 / km	\$ 1,495.50
Accommodation	\$ 705.92
Mobile radios (Trucks) - 6 days at \$20 / day	\$ 120.00
Hand-held Radios - 17 man-days at \$10 / day	\$ 170.00
Digital Camera - 2 days at \$20 / day	\$ 40.00
Quads - 10 man-days at \$100 / day	\$ 1,000.00
Satellite Phone -	\$ 63.00
Storage Trailer (Equipment) - 7 days at \$20 / day	\$ 140.00
Sub-Total	\$ 3,734.42

FIELD SUPPLIES (Flagging, KRAFT bags, etc.)

26 man-days @ \$20 / day	\$ 520.00
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DISBURSEMENTS

Analyses - 196 soil samples at \$25 / sample	\$ 4,900.00
Fuel	\$ 335.72
Field Supplies	\$ 36.82
Groceries	\$ 128.73
Office Supplies	\$ 177.24
Meals	\$ 172.27
Plotting	\$ 642.40
Shipping	\$ 196.52
Trailer Repairs	\$ 119.71
Sub-Total	\$ 6,709.41

REPORT/REPRODUCTION

R. T. Walker, P.Geo.: 3 days report writing at \$500/day	\$ 1,500.00
1 days analysis / drafting at \$500 / day	\$ 500.00
Sub-Total	\$ 2,000.00

Total **\$ 20,163.83**