

**BC Geological Survey
Assessment Report
30527**

ASSESSMENT REPORT

FOR

808718(BC) LTD.

GEOPHYSICAL SURVEY

ON

NECHAKO (OOTS A LAKE) PROPERTY

*LAT: 53° 44' N LONG: 125° 48' W (NAD83)
UTM 315000E, 5958000N (NAD83, Zone 10)*

Location: Chelaslie Arm, Nechako Basin area, British Columbia, Canada

Mining Zone: Omineca

NTS mapsheet: 093F/5, 093F/6, 093F/11, 093F/12, 093F/13

*BCGS TRIM MAPSHEET: 093F042, 093F043, 093F052, 093F053, 093F061,
093F062, 093F071, 093F072, 093F081, 093F082*

**ASSESSMENT REPORT BY
BRIAN CHEN**

**S.J.V. CONSULTANTS LTD.
JANUARY 2009**

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1. Introduction

This Assessment Report describes exploration work of a 3D Induced Polarization (3D IP) survey carried out in 2008 for 808718(BC) Ltd. on the Nechako (Ootsa Lake) property, which is located in the Omineca mining zone, approximately 120km southwest of Vanderhoof, BC, Canada. The geophysical survey, consisting of 32 lines totaling 76 line-kilometers, was undertaken between August and October and covered a block of 4 claims: 539037, 539041, 539042 and 539054. The work is being applied to the entire Nechako (Ootsa Lake) property.

The report is based on the geophysical report, a compilation of published and unpublished data, maps and reports made by cited persons. The report is written under the supervision of a “qualified person”. Sections 2 – 8.2 of this report are taken with permission from *Technical Report on the Geology and Geophysics of the South Block Property, Golden Dragon Explorations Inc.* (2008). Statement of qualifications for the authors are listed in Appendices B and C.

2. Legal description of property

The Nechako (Ootsa Lake) group of claims consist of 121 mineral claims (83 claims in Meteor claim block, 27 claims in South claim block, 5 claims in Chel claim block and 6 claims in Link Claim block) using the “cell system” of Mineral Titles Online (BC) totaling approximately 54989.45 hectares in surface area. The center of the property is situated approximately 90 kilometres by road from Burns Lake and is bounded to the north by Intata Reach of the Nechako Reservoir and Chelaslie Arm to the south(See Figure 1). The property is situated on National Topographic System 1:50,000 map sheets 93F/5, 93F/6, 93F/11, 93F/12, 93F/13 and BC Provincial 1:20,000 map sheets 093F042, 093F043, 093F052, 093F053, 093F061, 093F062, 093F071, 093F072, 093F081 and 093F082 respectively. The geographical center of the property is 53°44'N and 125°48'W with UTM coordinates of 315000mE, and 5958000mN, NAD 83, Zone 10.

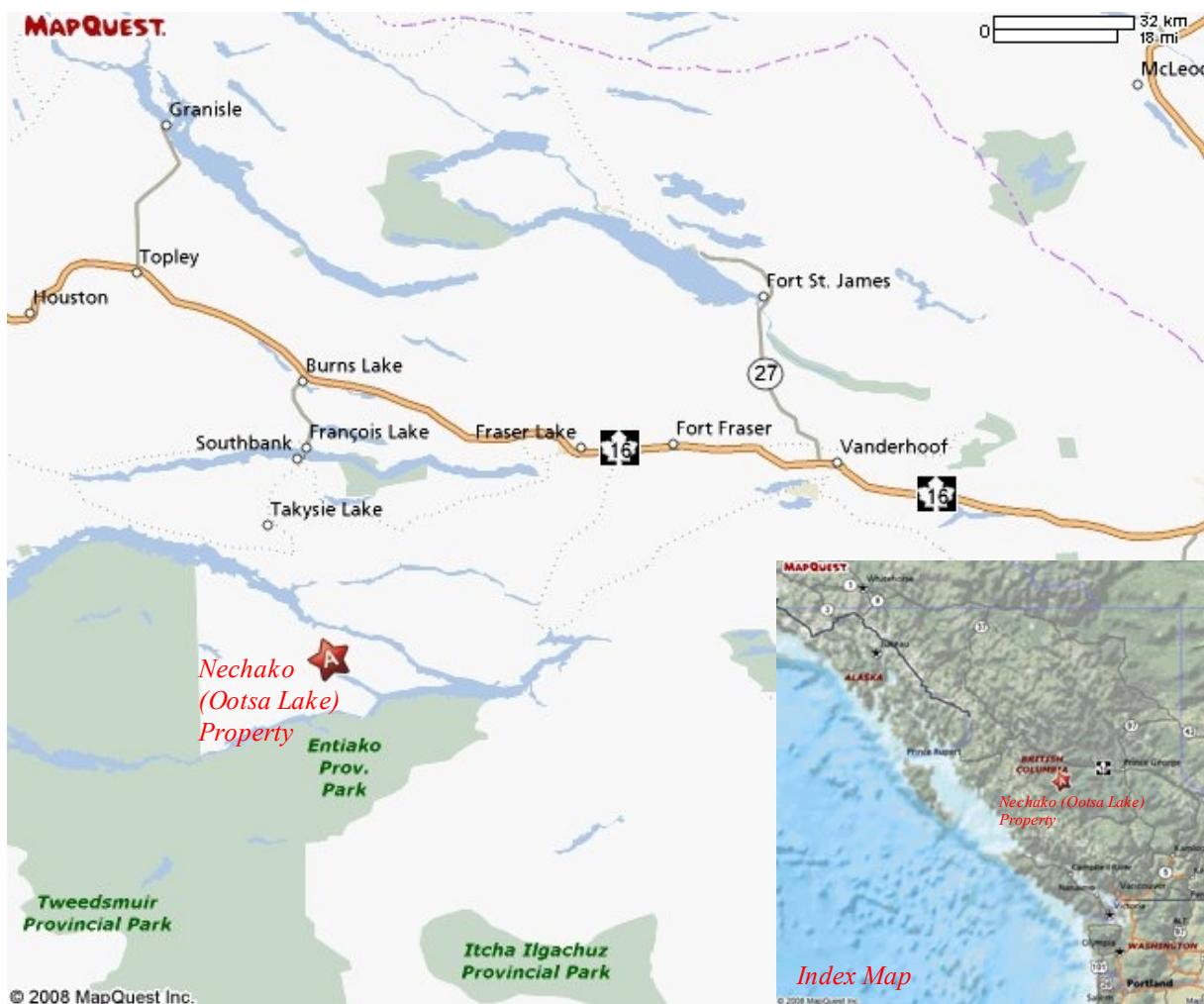


Figure 1: Location of Nechako (Ootsa Lake) Property
(Base Map Derived From MapQuest, www.mapquest.com)

The property shape and boundary are displayed in Figure 2; details of the claims are shown in Table 1. There are no known environmental concerns or parks designated for any area contained within the claims. Of note is that along the Nechako Reservoir any area below 300 metres ASL is potentially liable to be flooded with no compensation. The property has no encumbrances. The claims have not been legally surveyed.

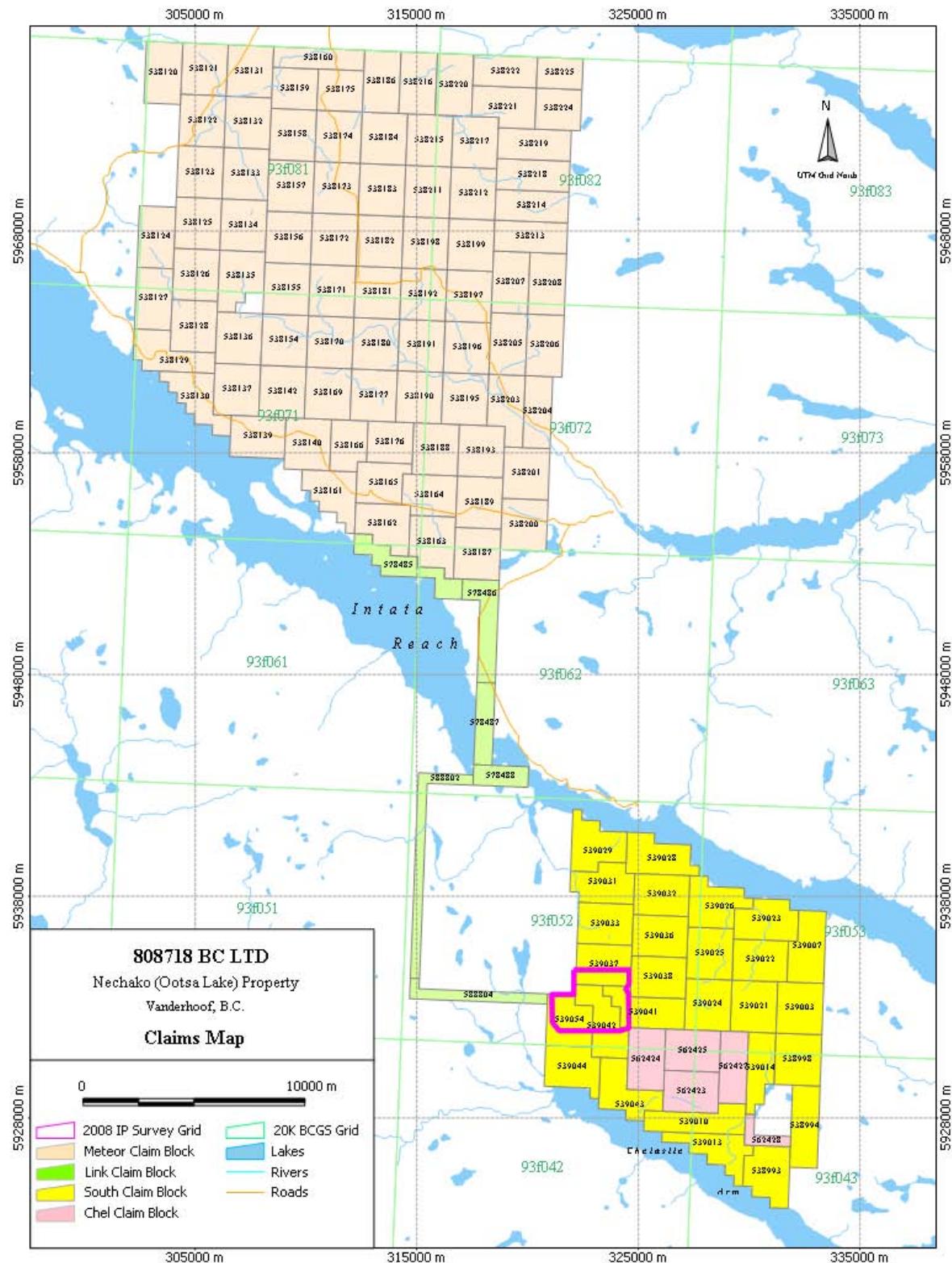


Figure 2: Claim map of Nechako (Ootsa lake) property.

| TENURE # | NAME | OWNER FMC | MAP | GOOD TO DATE | STATUS | AREA (ha) |
|----------|-----------|---------------|------|--------------|--------|-----------|
| 538120 | METEOR 1 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 457.429 |
| 538121 | METEOR 2 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 476.469 |
| 538122 | METEOR 3 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 476.703 |
| 538123 | METEOR 4 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 476.937 |
| 538124 | METEOR 5 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 458.147 |
| 538125 | METEOR 6 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 477.168 |
| 538126 | METEOR 7 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 477.402 |
| 538127 | METEOR 8 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 458.418 |
| 538128 | METEOR 9 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 477.637 |
| 538129 | METEOR 10 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 477.77 |
| 538130 | METEOR 11 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 458.846 |
| 538131 | METEOR 12 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 476.472 |
| 538132 | METEOR 13 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 476.707 |
| 538133 | METEOR 14 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 476.941 |
| 538134 | METEOR 15 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 477.169 |
| 538135 | METEOR 16 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 477.407 |
| 538136 | METEOR 17 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 477.685 |
| 538137 | METEOR 18 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 477.92 |
| 538139 | METEOR 19 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 458.999 |
| 538140 | METEOR 20 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 478.145 |
| 538142 | METEOR 21 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 477.922 |
| 538154 | METEOR 22 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 477.688 |
| 538155 | METEOR 23 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 477.453 |
| 538156 | METEOR 24 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 477.219 |
| 538157 | METEOR 25 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 476.99 |
| 538158 | METEOR 26 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 476.754 |
| 538159 | METEOR 27 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 381.233 |
| 538160 | METEOR 28 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 381.119 |
| 538161 | METEOR 29 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 459.231 |
| 538162 | METEOR 30 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 459.363 |
| 538163 | METEOR 31 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 478.572 |
| 538164 | METEOR 32 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 440.084 |
| 538165 | METEOR 33 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 401.777 |
| 538166 | METEOR 34 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 363.385 |
| 538169 | METEOR 35 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 477.926 |
| 538170 | METEOR 36 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 477.691 |
| 538171 | METEOR 37 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 477.457 |
| 538172 | METEOR 38 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 477.223 |

| | | | | | | |
|--------|-----------|---------------|------|-------------|------|---------|
| 538173 | METEOR 39 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 476.989 |
| 538174 | METEOR 40 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 476.753 |
| 538175 | METEOR 41 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 381.233 |
| 538176 | METEOR 42 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 382.495 |
| 538177 | METEOR 43 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 477.929 |
| 538180 | METEOR 44 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 477.694 |
| 538181 | METEOR 45 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 477.46 |
| 538182 | METEOR 46 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 477.225 |
| 538183 | METEOR 47 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 476.987 |
| 538184 | METEOR 48 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 476.751 |
| 538186 | METEOR 49 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 457.432 |
| 538187 | METEOR 50 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 478.615 |
| 538188 | METEOR 51 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 478.141 |
| 538189 | METEOR 52 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 478.378 |
| 538190 | METEOR 53 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 477.929 |
| 538191 | METEOR 54 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 477.695 |
| 538192 | METEOR 55 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 477.46 |
| 538193 | METEOR 56 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 478.14 |
| 538195 | METEOR 57 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 477.929 |
| 538196 | METEOR 58 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 477.694 |
| 538197 | METEOR 59 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 477.46 |
| 538198 | METEOR 60 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 477.225 |
| 538199 | METEOR 61 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 477.225 |
| 538200 | METEOR 62 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 478.473 |
| 538201 | METEOR 63 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 478.236 |
| 538203 | METEOR 64 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 458.834 |
| 538204 | METEOR 65 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 401.493 |
| 538205 | METEOR 66 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 458.563 |
| 538206 | METEOR 67 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 458.56 |
| 538207 | METEOR 68 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 458.292 |
| 538208 | METEOR 69 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 458.288 |
| 538211 | METEOR 70 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 476.985 |
| 538212 | METEOR 71 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 476.981 |
| 538213 | METEOR 72 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 458.086 |
| 538214 | METEOR 73 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 457.943 |
| 538215 | METEOR 74 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 476.749 |
| 538216 | METEOR 75 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 457.43 |
| 538217 | METEOR 76 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 476.745 |
| 538218 | METEOR 77 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 457.808 |

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|--------|-----------|---------------|------|-------------|------|----------|
| 538219 | METEOR 78 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 457.673 |
| 538220 | METEOR 79 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 457.427 |
| 538221 | METEOR 80 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 476.57 |
| 538222 | METEOR 81 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 400.187 |
| 538224 | METEOR 82 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 381.262 |
| 538225 | METEOR 83 | 146491 (100%) | 093F | 2009/may/19 | GOOD | 285.848 |
| 538993 | SOUTH 16 | 204672 (100%) | 093F | 2009/may/19 | GOOD | 481.3979 |
| 538994 | SOUTH 17 | 204672 (100%) | 093F | 2009/may/19 | GOOD | 461.9291 |
| 538998 | SOUTH 18 | 204672 (100%) | 093F | 2009/may/19 | GOOD | 480.87 |
| 539003 | SOUTH 19 | 204672 (100%) | 093F | 2009/may/19 | GOOD | 480.64 |
| 539007 | SOUTH 20 | 204672 (100%) | 093F | 2009/may/19 | GOOD | 480.37 |
| 539010 | SOUTH 21 | 204672 (100%) | 093F | 2009/may/19 | GOOD | 481.167 |
| 539013 | SOUTH 22 | 204672 (100%) | 093F | 2009/may/19 | GOOD | 442.7927 |
| 539014 | SOUTH 23 | 204672 (100%) | 093F | 2009/may/19 | GOOD | 384.7345 |
| 539021 | SOUTH 28 | 204672 (100%) | 093F | 2009/may/19 | GOOD | 480.64 |
| 539022 | SOUTH 29 | 204672 (100%) | 093F | 2009/may/19 | GOOD | 461.2117 |
| 539023 | SOUTH 30 | 204672 (100%) | 093F | 2009/may/19 | GOOD | 480.25 |
| 539024 | SOUTH 31 | 204672 (100%) | 093F | 2009/may/19 | GOOD | 480.64 |
| 539025 | SOUTH 32 | 204672 (100%) | 093F | 2009/may/19 | GOOD | 480.41 |
| 539026 | SOUTH 32 | 204672 (100%) | 093F | 2009/may/19 | GOOD | 480.18 |
| 539028 | SOUTH 33 | 204672 (100%) | 093F | 2009/may/19 | GOOD | 479.9736 |
| 539029 | SOUTH 34 | 204672 (100%) | 093F | 2009/may/19 | GOOD | 479.9278 |
| 539031 | SOUTH 35 | 204672 (100%) | 093F | 2009/may/19 | GOOD | 441.7 |
| 539032 | SOUTH 36 | 204672 (100%) | 093F | 2009/may/19 | GOOD | 460.94 |
| 539033 | SOUTH 37 | 204672 (100%) | 093F | 2009/may/19 | GOOD | 461.08 |
| 539036 | SOUTH 38 | 204672 (100%) | 093F | 2009/may/19 | GOOD | 461.1215 |
| 539037 | SOUTH 39 | 204672 (100%) | 093F | 2009/may/19 | GOOD | 461.2555 |
| 539038 | SOUTH 40 | 204672 (100%) | 093F | 2009/may/19 | GOOD | 461.3 |
| 539041 | SOUTH 41 | 204672 (100%) | 093F | 2009/may/19 | GOOD | 480.67 |
| 539042 | SOUTH 42 | 204672 (100%) | 093F | 2009/may/19 | GOOD | 480.74 |
| 539043 | SOUTH 43 | 204672 (100%) | 093F | 2009/may/19 | GOOD | 481.05 |
| 539044 | SOUTH 44 | 204672 (100%) | 093F | 2009/may/19 | GOOD | 461.715 |
| 539054 | SOUTH 45 | 204672 (100%) | 093F | 2009/may/19 | GOOD | 442.2789 |
| 562423 | CHEL | 204672 (100%) | 093F | 2009/may/19 | GOOD | 461.7939 |
| 562424 | CHEL1 | 204672 (100%) | 093F | 2009/may/19 | GOOD | 461.66 |
| 562425 | | 204672 (100%) | 093F | 2009/may/19 | GOOD | 461.6156 |
| 562427 | CHEL3 | 204672 (100%) | 093F | 2009/may/19 | GOOD | 403.97 |
| 562428 | CHEL4 | 204672 (100%) | 093F | 2009/may/19 | GOOD | 134.7433 |
| 578485 | LINK 1 | 146491 (100%) | 093F | 2010/mar/14 | GOOD | 459.545 |

| | | | | | | |
|------------|--------|---------------|------|-------------|------|----------|
| 578486 | LINK 2 | 146491 (100%) | 093F | 2010/mar/14 | GOOD | 459.7796 |
| 578487 | LINK 3 | 146491 (100%) | 093F | 2010/mar/14 | GOOD | 306.8052 |
| 578488 | LINK 4 | 146491 (100%) | 093F | 2010/mar/14 | GOOD | 230.216 |
| 588802 | LINK 5 | 146491 (100%) | 093F | 2010/jul/23 | GOOD | 480.002 |
| 588804 | LINK 6 | 146491 (100%) | 093F | 2010/jul/23 | GOOD | 307.6052 |
| 121 Claims | | | | | | 54989.45 |

Data from Mineral Titles Online, January 2009

Table 1: Claim data for Nechako (Ootsa Lake) property.

808718(BC) Ltd. entered into a property purchase agreement with Golden Dragon Explorations Inc. (GDE) on November 30th, 2007 to acquire a 100% percent interest in the 48 claims (South claim block) for \$1.00 and assumed all of the outstanding debt obligations as follow:

- Chris Bass Management fees \$30000 (monthly instalments of \$7500)
- Aeroquest \$61259.70
- G&A obligation \$25000

Also, 10868000 GDE shares would be returned to treasury for cancellation.

A check of the BC Mineral Titles Online website indicates that all 83 claims in Meteor claim block are owned 100% by United Exploration Management Inc. (UEMI). CM Meteor Resources Ltd. had entered into a purchase agreement with UEMI (Assessment Report for Geophysical Surveys on the Meteor Property, Wesley Raven, P.Geo., November 9, 2007). Recently, 808718(BC) Ltd. has a property purchase agreement with CM Meteor Resources Ltd. to acquire all of the rights, title and interest in the Meteor claim block and assumed all of the outstanding debt obligations.

3. Physiography and access

The Nechako (Ootsa Lake) property is located on the Nechako plateau which maintains a fairly constant overall elevation, but can be quite dissected at the local scale in a distinctive basin and range (horst and graben) topography producing areas of abundant outcrops. Elevations in the area vary from 1417 m at the top of Deerhorn Hill to 715m on François Lake. To the west, the area abuts on the Quanchus Range with a chain of peaks in the 2100 to 2300 m range.

Access to and throughout the property is good. Major highways border the Nechako Plateau: to the north (Highway 16) and the east (Highway 97). More locally, access to the property is via Vanderhoof, then south on the Kenney Dam Road to the Ice Bridge Ferry which crosses Intata Reach then by logging road around the property.

The main economic activity in the area is logging. There are a few ranches along the lower Nechako River and some farming northwest of Cheslatta Lake in the Takysie-Grassy Plains area. Tourism is a minor activity and consists mostly of fishing and, in the fall, hunting. Vegetation is dominated by evergreens (pine and spruce) with poplar and cottonwood in low-lying areas. The climate is typical of central British Columbia with below freezing temperatures (0° C to -40° C) from November to April and periods of hot weather in the summer ranging from 20° to 40° C. Precipitation averages 427.8 millimetres a year, with a substantial portion in the form of snow averaging 90.5 centimetres per year.

The region has been severely damaged by infestations of the Rocky Mountain Pine beetle. Vast areas have been affected by this insect which has killed large stands of commercial timber. Because of these infestations forest fires may pose a threat to exploration activities during the summer months. Along the Nechako Reservoir, any area below 300 metres ASL is potentially liable to be flooded with no compensation.

4. History

The assessment of the exploration work carried out in 2008, was the geophysical survey conducted in South claim block of the Nechako (Ootsa Lake) property which is situated immediately south of Intata Reach. Therefore, the description of the history is limited to the area around the geophysical survey grid. The South claim block of the Nechako (Ootsa Lake) property has been investigated by several regional exploration programs dating back to the 1960's which concentrated on small showings (MR, Bull and WT, now held by others). Early on most of the work was concentrated on exploring for copper-molybdenum mineralization. By the 1980's, the interior plateau region of central British Columbia was recognized to have comparable structural and lithological characteristics to gold-producing regions in the basin and range structural province in Nevada. Some preliminary geochemical sampling and claim staking was conducted on parts of the property by JMT Services Corp. for other clients.

Exploration intensified in the area during the late 1980's and early to mid 1990's. Several major mining companies including Hudson's Bay Mining and Smelting, Cogema Resources, Noranda and Phelps Dodge Corporation of Canada conducted gold exploration programs in the region. Exploration was greatly aided by regional studies conducted by both the Geological Survey of Canada's aeromagnetic surveys and the Province of British Columbia's regional lake sediment and water geochemistry surveys completed in 1994.

To the author's knowledge, there are no geological, geochemical or other exploration data specific to the subject property. Adjacent claims to the west and to the north are held by Goldmember Ventures Corp.

A brief description of the exploration programs conducted in the area of the Nechako (Ootsa Lake) claims is summarized below. The information below is provided as background material for the reader. The writer has derived the description from the BC Government Minfile and has no reason to doubt the accuracy of the descriptions. The information may not necessarily be indicative of the mineralization on the properties that are subject of this technical report. The source of the information is without exception publicly available documents. The writer has no affiliation with any of the properties. Assessment reports are available for the properties and the reader is encouraged to read them if additional information is needed.

Bull 4, Minfile Number: 093F063

The Bull 4 showing is located about 90 kilometres south of Burns Lake on the north shore of Chelaslie Arm. Stream and lake sediment surveys were conducted in this area in the 1960s by several different companies. In 1973, Placer Developments Ltd. staked their MR claims (now the Bull 1 claim) and completed reconnaissance mapping and sampling. In 1980, Prism Resources staked the Precious Metal claims to cover the MR claim and conducted reconnaissance mapping and sampling. In 1992, Sleeping Gold Ltd. conducted geological mapping, prospecting and sampling on the Bull 1-4 claims.

The region in which the showing occurs is within the Intermontane Belt, underlain dominantly by Lower to Middle Jurassic volcanic and sedimentary rocks of the Hazelton Group. These assemblages are overlain by the Upper Cretaceous to Lower Tertiary Ootsa Lake Group and Miocene plateau basalt. Intruding Lower Jurassic rocks of the Hazelton Group in the northeastern part of the map sheet is a belt of granodiorite, diorite and quartz diorite plutons of the Lower Jurassic Topley intrusive suite. Felsic plutons of probable Cretaceous age intrude both Lower and Middle Jurassic Hazelton strata.

The Bull claims are underlain by Lower to Middle Jurassic Hazelton Group rhyolitic and andesitic volcanics with minor epiclastic sediments. These are intruded by Eocene Ootsa Lake Group rhyolite dikes which are in turn cut by diabase dikes. A quartz vein stockwork and breccia zone, hosted in mafic tuff and breccia, contains up to 5 per cent pyrite, 3 per cent galena and 1 per cent sphalerite. The veining is exposed over a strike length of 20 metres and the width of the zone is between 2 and 4 metres. The vein system is open along strike to the west but to the east mineralization appears to be cut off by rhyolite dikes. There appears to be at least 2 vein structures striking east-west with near vertical dips.

The best sample was a grab sample taken across 30 centimetres of a quartz vein stockwork with 1 per cent galena and 5 per cent pyrite which assayed 2.1 grams per tonne gold, 22 grams per tonne silver and 0.273 per cent lead (Assessment Report 22535, sample #463762).

WT, Minfile Number: 093F031

The region in which the showing occurs is within the Intermontane Belt, underlain dominantly by Lower to Middle Jurassic volcanic and sedimentary rocks of the Hazelton Group. These assemblages are overlain by the Upper Cretaceous to Lower Tertiary Ootsa Lake Group and Miocene plateau basalt. Intruding Lower Jurassic rocks of the Hazelton Group in the northeastern part of the map sheet is a belt of granodiorite, diorite and quartz diorite plutons of the Lower Jurassic Topley intrusive suite. Felsic

plutons of probable Cretaceous age intrude both Lower and Middle Jurassic Hazelton strata.

The WT showing is underlain mainly by a Cretaceous felsic pluton comprising hornblende diorite to quartz diorite to the north and nonporphyritic to porphyritic latite to the south. At the contact of the intrusion with the enclosing rocks of the Hazelton Group, hornfels and skarn have developed. Pyrite, chalcopyrite, molybdenite, minor bornite, pyrrhotite and magnetite are variably dispersed throughout the plutonic rocks and the hornfels-skarn unit. Chalcopyrite occurs mainly as disseminations along fracture planes along with epidote and chlorite, and as veinlets and stringers associated with magnetite, in both biotite-rich diorite and the hornfels-skarn unit. Molybdenite occurs in association with very fine-grained dark coloured biotite-rich inclusions within the biotite diorite.

Government programs

The first recorded work done in the area was a Geological Survey of Canada mapping program, led by H. W. Tipper in 1949. The results of this program were published in GSC Memoir 324 (Tipper, 1963). The government has been active in the area, mapping bedrock and surficial deposits of the NTS 93F/3 and portions covering the 93F/2 and 92F/3 map sheets. A lake sediment geochemical survey provided good coverage of map sheets 93F/11, 12, 13, and 14. The BC Geological Survey also did miscellaneous detailed surveys of showings and geochemical anomalies within the area. The Geological Survey of Canada flew an airborne magnetic survey covering most or all of the gap from 53°15' to 51°15' north latitude and from the Fraser River to the Coast Range.

1971-1973 Noranda

Beginning in 1971, Noranda conducted a geochemical survey covering the WT claims which hosts the WT showing now held by John C. Bot as the Shellie and Shellie 2 claims. A total of 14.8 line miles of grid were established with soil samples taken every 200 feet. In 1972, Noranda continued exploring the claims by conducting geological mapping and prospecting and hand trenching over the area gridded in the previous year. A total of 16 rocks samples were collected with best values of 0.47 oz/ton silver and 0.58% copper. In 1973, Noranda conducted an Induced Polarization survey over the gridded area and reported a weak IP response trending to the west.

1980 Prism Resources

Prism staked their Precious Metal claims in 1980 to cover Placer's MR lead-zinc anomaly. They conducted reconnaissance geological mapping and sampling that year, taking a total of 218 soil, silt and rock samples from the property. Samples were analyzed for copper, lead, zinc and silver, defining a 600

by 700 meter lead-zinc soil geochemical anomaly with maximum values of 1950 ppm lead and 2760 ppm zinc. Silver values were erratic and generally low (Harivel and Livingston, 1981). No fieldwork was carried out in 1982 on the Precious Metal claims, but the 1980 sample pulps were analyzed for gold, arsenic and molybdenum. Gold and molybdenum values were generally low in both soils and rocks; high arsenic values in soils, to a maximum of 145 ppm, coincided with the previously-defined lead-zinc anomaly (Harivel and Livingston, 1982).

1992 Sleeping Gold Ltd.

During May and June of 1991, Sleeping Gold Ltd. carried out a preliminary exploration program on the bull property, (now the claim held by Iqbal Boga) consisting of geological mapping, prospecting and soil sampling. This program was designed to verify the reported lead-zinc-arsenic soil anomalies, locate their source and evaluate the property's potential for epithermal and volcanogenic massive sulphide mineralization. A total of 24 rock samples and 152 soil samples were taken. A quartz vein stockwork and breccia at this location is accompanied by pervasive silicification and up to 5% pyrite, 3% galena and 1% sphalerite. There appears to be at least two vein structures striking east-west with near vertical dips. The veining is exposed over a strike length of 20 meters, and the width of the zone is difficult to estimate due to lack of exposure, but individual vein outcrops exceed 2.0 meters in width and the positions of these outcrops relative to each other would indicate a total width of at least 4.0 meters. The vein system remains open along strike to the west and probably continues to the east as far as the baseline, but appears to be cut off by rhyolite dykes further east.

5. Geologic setting

5.1. Regional geology

The Tertiary geologic elements of the Nechako Plateau are part of a regional extensional system that extends from the Republic area of northern Washington State, northwesterly for some 1000 kilometres into the Babine district of north-central British Columbia. This belt trends northwest with the approximate dimensions of 1000 by 200 kilometres. It crosses all major terrane boundaries and underlies the Quesnel, Kootenay and Omineca Terranes in the south and the Stikine Terrane in the north, crossing the oceanic Cache Creek Group. It overlaps the southern margin of the Bowser Basin where it continues northward as a thin strip along the eastern margin of the Coast Range.

Stratigraphic and intrusive rocks in the Stikine Terrane range in age from Paleozoic to Pleistocene. With respect to the Eocene mineral setting, the geologic elements of the Stikine Terrane may be divided into three separate packages: basement rocks, later Upper Cretaceous-Eocene rocks associated with mineralization and cover rocks (Table 2).

Basement Rocks - Lower Upper Cretaceous and Older

Basement rocks to the Tertiary in the Stikine Terrane comprise Upper Paleozoic to lower Upper Cretaceous strata grouped into two major time-stratigraphic assemblages. The oldest assemblage consists of arc volcanics of Upper Paleozoic to Middle Jurassic age which includes limestone, volcanics and sediments of the Upper Paleozoic Cache Creek Assemblage, submarine and marine island arc volcanics and sediments of the Carnian to Norian subalkaline, basaltic Stuhini (Takla) Group, and the Sinemurian to Bajocian calc-alkaline Hazelton Group.

| Stratified Rocks | Intrusive and Metamorphic Rocks |
|--|---|
| 11. Anahim Volcanics (Pliocene-Pleistocene) | |
| 10. Chilcotin Volcanics (Miocene) | |
| 9. Endako Group (Eocene-Oligocene) | |
| 8. Ootsa Lake Group (Eocene and Paleocene) | G. Eocene (stocks, plugs, dykes, rhyolite, felsite, porphyry, diorite, gabbro) |
| 7. Kasalka-Kingsvale Groups (Upper Cretaceous) | F. Upper Cretaceous-Paleocene (Quanchus Intrusions: stocks and batholiths, diorite to quartz monzonite) |
| 6. Skeena-Jackass Mountain Groups (Lower Cretaceous) | E. Mid-Cretaceous (mainly tonalite to quartz monzonite of Coast Range complex) |
| 5. Gambier Group (Upper Jurassic-Lower Cretaceous) | D. Jurassic-Cretaceous (François Lake Batholith; quartz diorite to granite, includes quartz-feldspar porphyry) |
| 4. Relay Mountain-Bowser Groups (Upper Jurassic-Lower Cretaceous) | |
| 3. Hazelton Group (Lower and Middle Jurassic) | C. Middle Jurassic (locally foliated granodiorite and quartz monzonite) |
| 2. Stuhini Group (Upper Triassic) | B. Permian (mainly granodiorite in lower Chilcotin River) |
| 1. Cache Creek Group (Upper Palaeozoic) | A. Metamorphic Rocks (gneiss, schist, metavolcanics, cataclasites) |

Table 2: Geological elements of Stikine Terrane.

The arc volcanic assemblages are overlain by two sedimentary assemblages, the Middle Jurassic to Lower Cretaceous Bowser Lake Group and the Lower and Upper Cretaceous Skeena Group. Deltaic assemblages of the Bowser Lake Group were deposited mainly in the Bowser Basin to the north of the Nechako reconnaissance area, except for its basal beds. These basal beds belong to the Ashman Formation and represent a black clastic-chert pebble conglomerate unit that covers much of the Stikine Terrane. Marine and nonmarine sediments of the Neocomian to Cenomanian Skeena and Jackass Mountain Groups blanketed much of the Stikine Terrane and sourced from the east, off the Cache Creek, Quesnel and Omineca Terranes. The blanket of Skeena Group clastics across Stikinia outlines a

regional datum to which deformation and deposition of younger strata may be related. This surface represents one of three main erosional surfaces in central BC.

The basement rocks have been affected by regional compressive tectonics. Westerly verging compression along the east margin of the Stikine Terrane, associated with the amalgamation of Stikinia, Quesnellia and the Cache Creek Terranes to the North American Craton, affects rocks as young as Upper Jurassic. Easterly verging compression along the west margin of the Stikine Terrane, associated with the amalgamation of the Wrangellia with Stikinia affects rocks as young as Late Cretaceous.

Intrusive rocks associated with the basement strata include the Upper Jurassic-Lower Cretaceous François Lake intrusions to the northeast of the reconnaissance area, and mid-Cretaceous plutons of the Coast Crystalline Complex.

Many of the northwest and northeast trending fault zones that control the distribution of the Tertiary geologic elements are fault zones whose activity can be traced back to the Upper Triassic and Lower Jurassic.

Upper Cretaceous to Miocene

The Upper Cretaceous to Eocene metallogenic event is associated with three stratigraphic assemblages, the late Upper Cretaceous andesitic Kasalka Group, the felsic Eocene Ootsa Lake Group and the basaltic Eocene to Oligocene Endako Group. These assemblages represent a generalized cycle of early andesitic volcanism, explosive felsic volcanism, bimodal felsite-basic volcanism and later basic volcanism. The early andesitic Kasalka Group, and the felsic Ootsa Lake Group strata were deposited in calderas and caldera complexes. The distribution of the older facies of the Endako Group are in part controlled by the felsic calderas. The felsic calderas are large, composite features that may measure more than 50 kilometres in diameter and are nested caldera complexes. The volcanic assemblages are associated with a fault array whose main expression is extensional. This sequence of caldera associated volcanism and extensional faulting is a common sequence through the length of the extensional belt, from the Mexican border to Babine Lake and is associated with a vast array of significant mineral deposits.

The Kasalka Group volcanics (McIntrye, 1985) occur as a number of caldera basins throughout west-central British Columbia, on the Stikine Terrane between the Blackwater Linear zone and the north flank of the Skeena Arch. They are mainly feldspathic andesitic volcanics but local basins include explosive and passive felsic volcanism. They are associated with granodioritic stocks and plugs of the

Quanchus and Bulkley Intrusions. In a number of locations in central BC, red and green polylithic volcanic and granitic cobble conglomerate underlies basal Kasalka strata. Age of Kasalka volcanics and associated intrusives range from 85My to 60My and fall mainly in the 72 to 67My interval.

The Ootsa Lake Group volcanics (Duffel, 1959) are typified by light coloured felsic volcanics. They underlie broad areas of the southern Stikine Terrane from Babine Lake to the Chilcotin River and include a variety of depositional types. They occur in structurally controlled basins and in large caldera complexes. Two caldera complexes underlie the Necho Reconnaissance area, the Mt. Dent Caldera Complex in the south (Alexis property area) and the Cheslatta Caldera Complex in the north, (Nechako property area). Subvolcanic intrusives are common; coeval plutonic rocks are rare within the caldera complexes but common in the basement. The Ootsa Lake Group ranges in age from 58 to 47My with the interval of 52 to 48My representing timing of the main felsic eruptive events.

The Endako Group (Armstrong, 1949) is a wide ranging assemblage of mainly basaltic rocks. In a general sense, the Endako Group overlies and is younger than the Ootsa Lake Group. Basaltic and andesitic rocks are commonly associated with felsic rocks in the calderas. Ages of the Endako Group show a range from 50 to 37My. The early basaltic rocks of the Endako Group overlap in both ages and depositional sites with the felsites of the Ootsa Lake Group. Although the Ootsa Lake Group and the early Endako Group are mapped as separate entities, the interval of their coincidence in space and time infers a genetic relationship.

Post-Ootsa Lake Group basaltic volcanism occurred intermittently throughout the area, from 45My to recent (Mathews, 1984 and 1989; Rouse, 1988). Basaltic volcanics younger than 35My are correlated with the Chilcotin Group. Felsic volcanics are known to be locally associated with intervals of this basalt event but no significant centre has yet been recognized.

Pliocene-Pleistocene

Outcrops of the Anahim Group peralkaline basalts have been observed in two locations of the south area: west of Nazko, a 3-km wide cinder cone overlies glacial till, and a few outcrops were found in the Moore Creek area.

"During the Pleistocene all of Central British Columbia was covered by glacier ice that molded a multitude of features from which the glacial events can be interpreted" (Tipper, 1971). The bulk of glacial features in Central British Columbia have been produced by the Fraser Glaciation, the last major advance. Minor late re-advances are observed around the Anahim volcanoes and along the Coast

Ranges.

Within the study area glacial transport direction varies from N0° to 30°, south of the Blackwater lineament, to N60° to 90° north of it. Glacial deposits consist mostly of lodgement till with some areas of ablation till, esker systems, and fluvio-glacial material. A thin veneer of ablation till may occasionally overlie lodgement till. There are no extensive glacial lake deposits (sands and clays). Evidence of multiple glaciation has been observed in a few localities in the form of lodgement till overlying fluvio-glacial deposits.

5.2. Regional Structure

The Nechako Basin is within the Intermontane Belt of the Canadian Cordillera, mainly on the Stikinia Terrane, but overlapping onto the Cache Creek Terrane. "A regional dextral transcurrent strain regime appears to have been important in the evolution of early Cenozoic structures in the southern part of the Intermontane Belt [...] These structures have been related to right lateral transform motions and to regional extension" (Gabrielse *et al.*, 1992). This regime resulted in alternating basins and arches along the Intermontane Belt: Nechako Basin, Skeena Arch, Bowser Basin and Stikine Arch. The Nechako Basin can be assimilated to a pull-apart basin formed between the Fraser River Fault System and the Coast Range Megalineament or one of its parallel structures extending north from the Yalakom Fault. The internal structure of the Nechako Basin reflects the same structural regime.

The Nechako Arch is composed mainly of pre-Tertiary rocks, and is itself cut by two northwest trending ranges, the Nechako and the Fawnie Horst; the south of the Nechako horst shows several dextral offsets along N70° degree faults, the major one being the Top Lake Fault; at the southwest end of the Nechako Arch the Wolf graben is filled mainly by Ootsa Lake group rhyolites; a similar graben extends south of the Cheslatta Caldera across an area of older rocks which may be an extension of the Nechako arch.

5.3. Property Geology

The Nechako (Ootsa Lake) property is located on the southwest flank of the Cheslatta Caldera Complex which is a broad, circular area of some 60 kilometres across enclosing about 3000 square kilometers. The southern border of the property abuts against the Nechako Arch Arch and the contact closely follows the linear trace of Knewstubb and Natalkuz Lakes. This contact is suggested to be fault

controlled. The western border of the property is bounded by the Snag graben.

The caldera is situated on the obtuse side of a major kink on the north flank of the Nechako Arch. This kink zone is defined by the area of intersection of regional northwest and northeast trending faults. The caldera overlaps this zone and the margin of the caldera is outlined by the disappearance of the trace of these faults into the caldera.

The caldera complex is underlain dominantly by Early Tertiary felsic volcanics of Ootsa Lake Group and basic volcanics of the Endako Group. A number of different facies assemblages of Tertiary volcanic and sedimentary rocks are distributed throughout the caldera complex. This suggests the presence of a number of separate volcanic centres and indicates that the Cheslatta Caldera Complex consists of a nested array of smaller scale calderas within the larger structure.

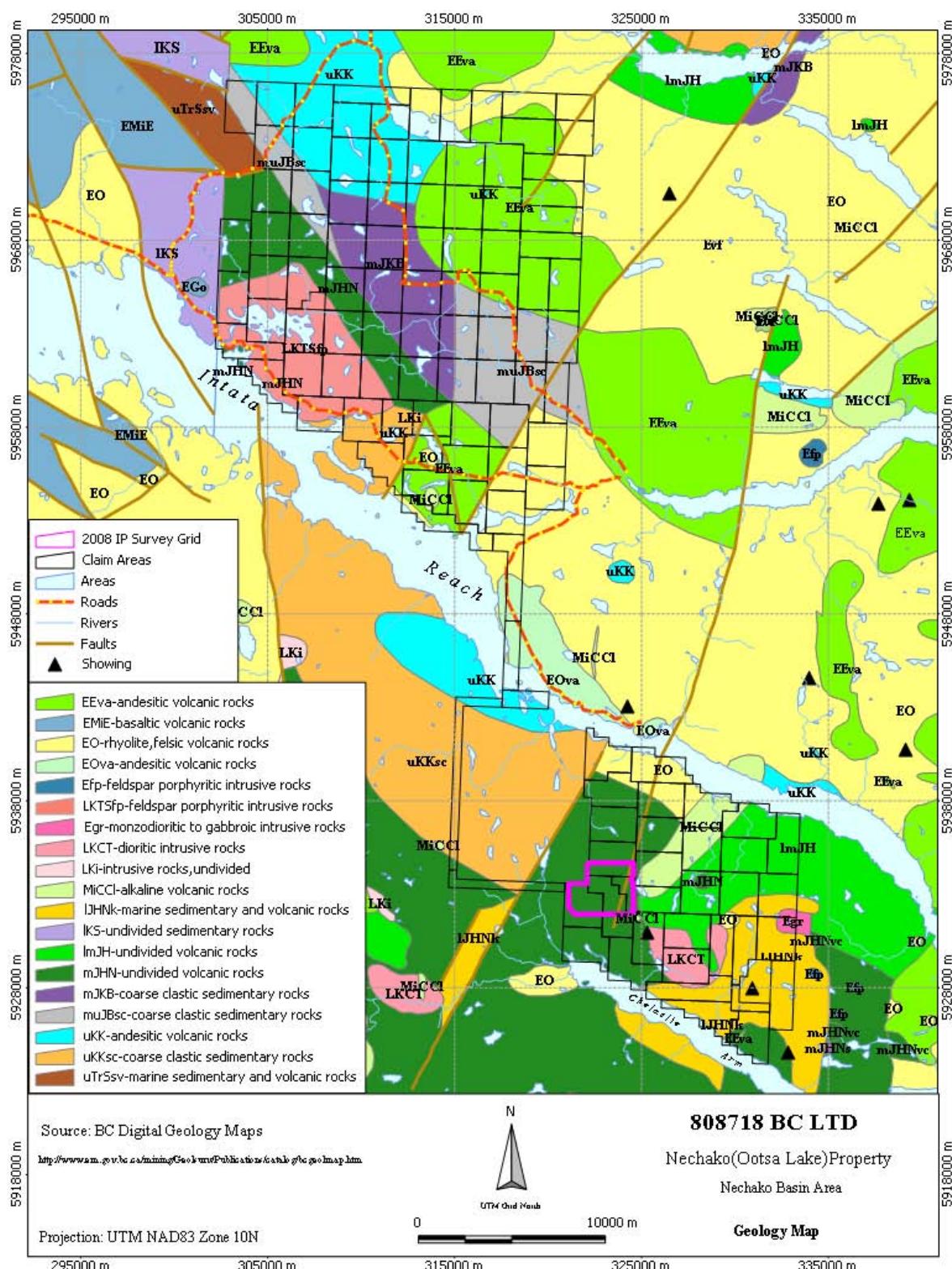


Figure 3: Regional Geology Map of the Nechako (Ootsa Lake) Property.

The Cheslatta Caldera Complex is rimmed by a suite of intermediate volcanics correlative with volcanics of the late Upper Cretaceous Kasalka Group (McIntyre, 1985). These mainly andesitic volcanics occur in small discrete areas 5 to 10 kilometres across that are interpreted to represent separate volcanic basins or calderas. These volcanics are included with the Ootsa Lake and Endako Groups as part of the Cheslatta Caldera Complex and the andesites represent the beginning stages of the evolution of the caldera complex.

The basement to the Upper Cretaceous and Tertiary volcanics of the Cheslatta Complex is similar to the Nchako Arch and includes volcanics and sediments of the Upper Triassic Stuhini Group, the Lower and Middle Jurassic Hazelton Group, the Middle and Upper Jurassic Bowser Lake Group and the Lower Cretaceous Skeena Group. A unit found in the area comprises red and green volcanic-polylithic conglomerate, fanglomerate, sandstone and mudstone that appear to be younger than the Skeena Group. These rocks have been found in scattered remnants throughout west central British Columbia and are informally known as "basal Kasalka beds" (Woodsworth, 1979).

The Quanchus Intrusions, Late Cretaceous batholiths of granitic to granodiorotic composition cut Hazelton Group rocks in the south west part of the property. These batholiths are generally coarse-grained, equigranular and light coloured. Potassium-argon dating indicates an age of 64.3 ± 2.4 Ma for the Capoose batholith, twenty kilometers southeast of the South Block property (Andrew, 1988).

6. Deposit types

The Nchako (Ootsa Lake) property has the potential to host mineralization from a varied number of deposit models. The property has the potential to host:

- Epithermal gold-silver deposits (example: Wolf Property to the south)
- Cu-Mo porphyry mineralization (some Mo mineralization was noted at the Chelaslie showing)
- Skarn Cu-Mo-W deposits (Exo property to the south).

Epithermal deposits

The following description on low sulphidation epithermal gold-silver mineralization is adapted from A. Panteleyev, 1996 from the British Columbia Geological Survey.

The depositional environment/geological setting of this type of deposit is generally in high-level hydrothermal systems from depths of ~1km to surficial hot spring settings. The deposits are hosted in

regional-scale fracture systems related to grabens, (resurgent) calderas, flow-dome complexes and rarely, maar diatremes. Extensional structures in volcanic fields (normal faults, fault splays, ladder veins and cymoid loops, etc.) are common and locally graben or caldera-fill clastic rocks are present. High-level (subvolcanic) stocks and/or dikes and pebble breccia diatremes occur in some areas. Locally resurgent or domal structures are related to underlying intrusive bodies.

Host lithologies can include most types of volcanic rocks, but calcalkaline andesitic compositions predominate. Some deposits occur in areas with bimodal volcanism and extensive subaerial ashflow deposits. A less common association is with alkalic intrusive rocks and shoshonitic volcanics.

The mineralized zones are typically localized in structures, but may occur in permeable lithologies. Upward-flaring mineralized zones centred on structurally controlled hydrothermal conduits are typical. Large (>1m wide and hundreds of metres in strike length) to small veins and stockworks are common with lesser disseminations and replacements. Vein systems can be laterally extensive but ore shoots have relatively restricted vertical extent. High-grade ores are commonly found in dilational zones in faults at flexures, splays and in cymoid loops. Typically the veins display textures including open-space filling, symmetrical and other layering, crustification, comb structure, colloform banding and multiple brecciation. The veins generally consist of quartz, amethyst, chalcedony, quartz pseudomorphs after calcite, calcite, adularia, sericite, barite, fluorite, and Ca/Mg/Mn/Fe carbonate minerals such as rhodochrosite, hematite and chlorite.

The predominant minerals in these types of deposits include pyrite, electrum, gold, silver, and argentite with lesser amount of chalcopyrite, sphalerite, galena, tetrahedrite, silver sulphosalt and/or selenide minerals. Deposits can be strongly zoned along strike and vertically, and are commonly zoned vertically over 250 to 350m from a base metal poor, Au/Ag-rich top to a relatively Ag-rich base metal zone and an underlying base metal rich zone grading at depth into a sparse base metal, pyritic zone. From surface to depth metal zones contain: Au/Ag/As/Sb/Hg, Au/Ag/Pb/Zn/Cu and Ag/Pb/Zn. In alkalic host rocks tellurides, V/mica (roscoelite) and fluorite may be abundant, with lesser molybdenite.

Silicification is extensive in ores as multiple generations of quartz and chalcedony are commonly accompanied by adularia and calcite. Pervasive silicification in vein envelopes is flanked by sericite-illite-kaolinite assemblages. Intermediate argillic alteration (kaolinite-illite-montmorillonite, or smectite) forms adjacent to some veins; advanced argillic alteration (kaolinite-alunite) may form along the tops of mineralized zones. Propylitic alteration dominates at depth and peripherally.

In some districts the epithermal mineralization is tied to a specific metallogenic event, either structural, magmatic, or both. The veins are emplaced within a restricted stratigraphic interval generally within 1km of the paleosurface. Mineralization near surface takes place in hotspring systems, or the deeper underlying hydrothermal conduits. At greater depth it can be postulated to occur above, or peripheral to, porphyry and possibly skarn mineralization. Normal faults, margins of grabens, coarse clastic caldera moat-fill units, radial and ring dike fracture sets and both hydrothermal and tectonic breccias are all ore fluid channeling structures. Through-going, branching, bifurcating, anastamosing and intersecting fracture systems are commonly mineralized. Ore shoots form where dilational openings and cymoid loops develop, typically where the strike or dip of veins change. Hanging wall fractures in mineralized structures are particularly favourable for high-grade ore.

These deposits form in both subaerial, predominantly felsic, volcanic fields in extensional and strike-slip structural regimes and island arc or continental andesitic stratovolcanoes above active subduction zones. Near-surface hydrothermal systems, ranging from hotspring at surface to deeper, structurally and permeability focused fluid flow zones are the sites of mineralization. The ore fluids are relatively dilute and cool solutions that are mixtures of magmatic and meteoric fluids. Mineral deposition takes place as the solutions undergo cooling and degassing by fluid mixing, boiling and decompression.

Porphyry deposits

The flowing description is taken almost verbatim from a paper entitled Porphyry Deposits by W.D. Sinclair for the geological survey of Canada.

The most applicable genetic model for porphyry deposits is a magmatic-hydrothermal one, or variations thereupon, in which the ore metals were derived from temporally and genetically related intrusions. Large polyphase hydrothermal systems developed within and above genetically-related intrusions and commonly interacted with meteoric fluids (and possibly seawater) on their tops and peripheries. During the waning stages of hydrothermal activity, the magmatic-hydrothermal systems collapsed inward upon themselves and were replaced by waters of dominantly meteoric origin. Redistribution, and possibly further concentration of metals, occurred in some deposits during these waning stages. Porphyry deposits occur in a variety of tectonic settings.

Porphyry Cu deposits typically occur in the root zones of andesitic stratovolcanoes in subduction-related, continental and island-arc settings (Mitchell and Garson, 1972; Sillitoe, 1973, 1988a; Sillitoe and Bonham, 1984). Porphyry Cu/Au deposits, such as those associated with Triassic and Lower

Jurassic silica-saturated alkaline intrusions in British Columbia, formed in an island-arc setting, although possibly during periods of extension.

7. Mineralization

There is only one known historical mineralized showing within the claims, the Chelaslie Arm occurrence (Minfile 2008). The following description is summarized from the BC Minfile database.

Chelaslie Arm 93F 025

The region in which the showing occurs is within the Inter-montane Belt, underlain dominantly by Lower to Middle Jurassic volcanic and sedimentary rocks of the Hazelton Group. These assemblages are overlain by the Upper Cretaceous to Lower Tertiary Ootsa Lake Group and Miocene plateau basalt. Intruding Lower Jurassic rocks of the Hazelton Group in the northeastern part of the map sheet is a belt of granodiorite, diorite and quartz diorite plutons of the Lower Jurassic Topley intrusive suite. Felsic plutons of probable Cretaceous age intrude both Lower and Middle Jurassic Hazelton strata.

The Chelaslie Arm showing comprises molybdenite in a small quartz vein cutting granitic rocks of probable Cretaceous age which has intruded Lower Jurassic Hazelton Group volcanics.

However, there are two mineral showings surrounded by the South Block Claims. These are Bull 4 (Minfile number 93F063) and WT (Minfile number 93F031) which are described in more detail under the heading "History".

8. Exploration

The summary of the exploration work is limited to the South claim block of the Nechako (Ootsa Lake) property within which the 2008 geophysical survey was carried out.

8.1. 2006 Work

In 2006, the exploration program initially undertaken by the company consisted of a regional study and reprocessing of the existing Government geophysical (gravity and magnetic) data. The company employed consultants from S.J.V. Consultants Ltd. of Delta, BC, to reprocess airborne magnetic, gravity and geochemical data covering the South block project (part of Nechako property) and to produce a series of 3 dimensional inversion maps and to offer their conclusions and recommendations to aid the company in its exploration programs. The following is a summary of the reports.

The government gravity map tends to display very large, deep structures due to the data being

gathered at 2000m grid intervals. The South Block property is located over a pivot point where a North-South trending regional contact to the north changes direction to Northeast-Southwest trending contact.

The government regional magnetic data was gathered on 800m spaced survey lines and is therefore denser than the gravity data. The magnetic map displays good correlation with the gravity data. The “bouguer” high along the western half of the area coincides with a weak magnetic low. NE-SW trending magnetic lineations loosely follow the gravity lineaments at the corners of the block. However, the magnetic map is significantly more complex than the gravity map and reveals more detailed structures and bodies closer to the ground surface.

The magnetic map delineates regional trends in addition to those mapped by the gravity survey. It also delineates NW-SE trends that follow the edges of the study area. These trends may be due in part to the coincidence of major bodies of water (Chelaslie Arm to the south and Intata Reach to the north) but they are also believed to be related to the underlying geology. The pattern resulting from these interesting trends shows the study area to be roughly covering the gap formed at the intersection of four regional lineations, possible faults.

A series of 3D magnetic inversion maps were generated displaying the cutoff surface of 0.012 SI. The shapes generated may approximate rocks containing a higher magnetic susceptibility than surrounding rocks or may represent the surface trace of intrusive rocks.

A series of thematic maps displaying the elements silver, arsenic, gold, molybdenum, lead and antimony were generated from industry and government data. The plots show predominately lake sediment results but some regional rock samples returned weakly anomalous values in silver. Lead, molybdenum and gold. The lake sediments tended to show weakly to moderately anomalous values in the elements reported.

A compilation map is shown in Figure 4.

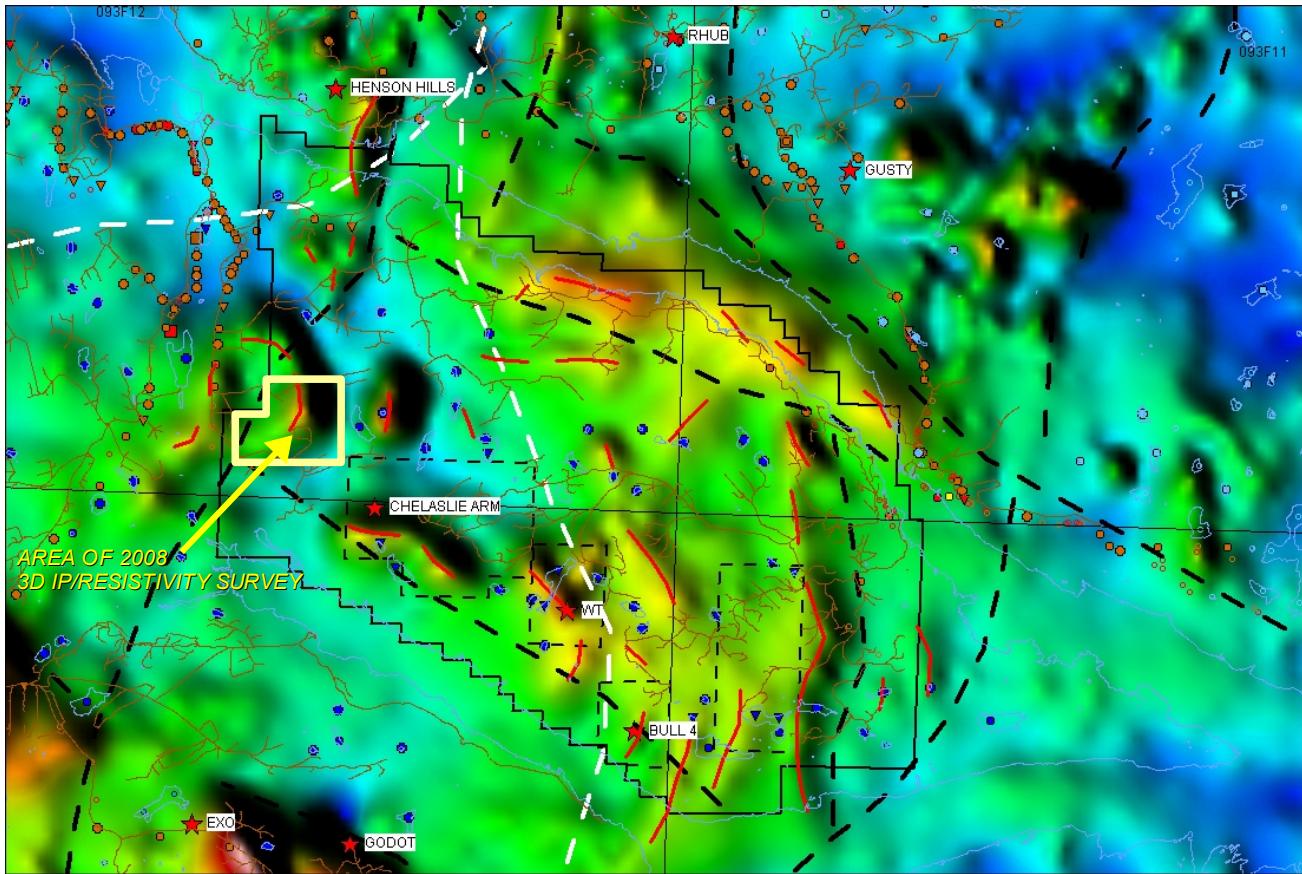


Figure 4: South Block Compilation Map, August, 2006.

Background Magnetic Color Contour Map – Regional Gravity Trends (White Dashes) – Regional Magnetic Trends (Black Dashes) – Localized Magnetic High Trends (Red Lines) – MinFile Occurrences (Red Stars) – Thematic Geochemistry Symbols (Au) – Roads (Brown Lines) – Lakes (Blue Lines)

From Technical Report, South Claims Property, Golden Dragon Explorations Inc. with addition of area of 2008 3D IP/Resistivity Survey

8.2. 2007 Airborne Geophysical Survey

A helicopter-borne geophysical survey was carried out on behalf of Golden Dragon Explorations Inc. on the South Claims project, near Burns lake, British Columbia by Aeroquest International (www.aeroquest.ca). The following is derived from a summary of the technical report by Warren D.Robb (P.Geo.) and Barry Price (P.Geo.) dated in February 2008.

The AeroTEM II time domain EM and magnetic helicopter-borne survey took place from August 4th – 6th, 2007. The survey was flown with a line spacing of 100 meters. The control (tie) lines were flown perpendicular to the survey lines with a spacing of 1000 meters. The total survey coverage was 985.8

line kilometers, all of which was within the defined project area. The claims and survey grid are shown in Figure 5.

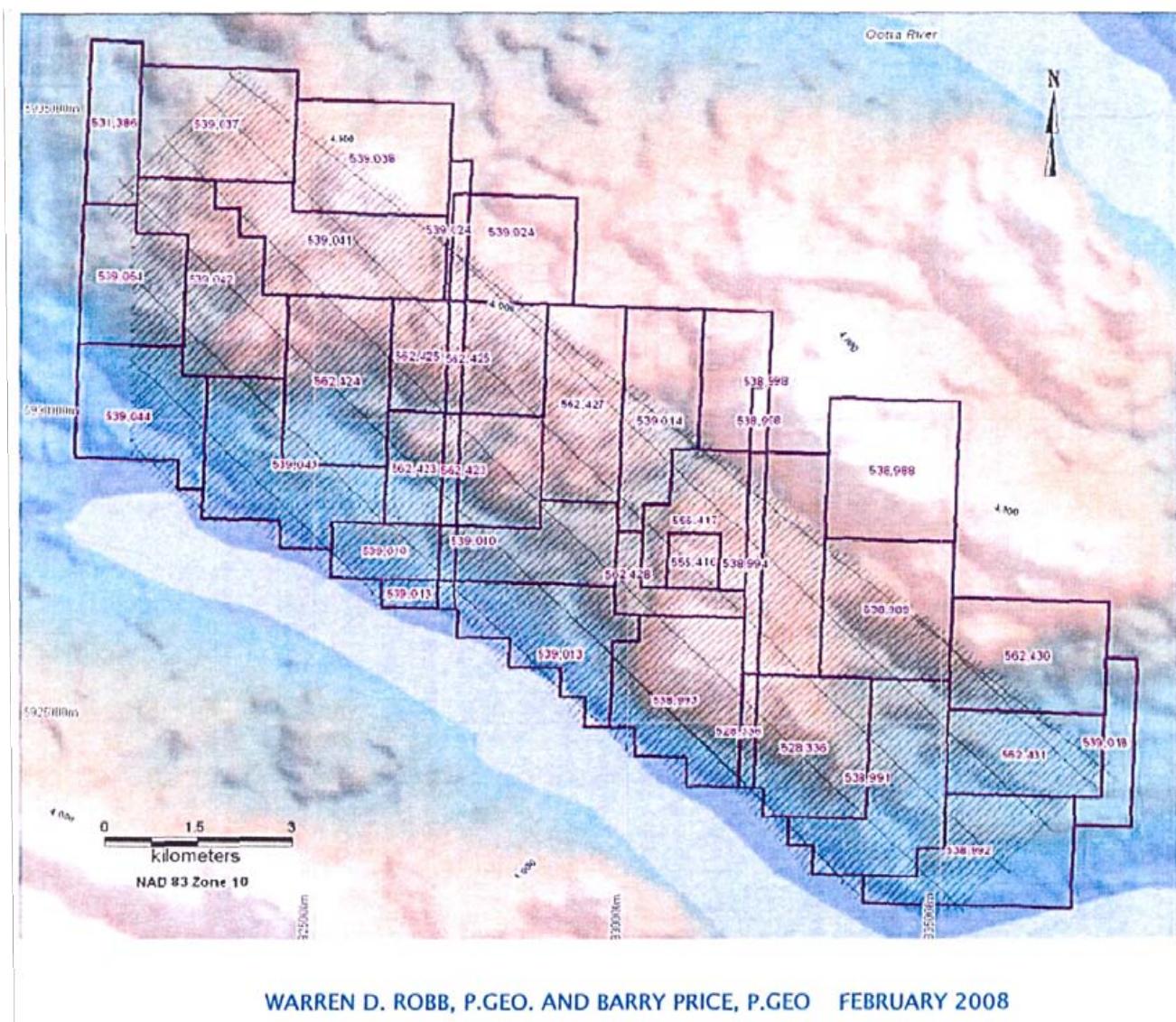


Figure 5: Claims and Survey Grid.

2007 Airborne Geophysical Survey, from Technical Report, South Claims Property, Golden Dragon Explorations Inc.

The survey area is covered by three map plates and three geophysical data products are delivered as listed below:

- TMI – Coloured Total Magnetic Intensity (TMI) with contours and EM anomaly symbols;
 - Zoff3 – AeroTEM Z3 off-time with contours and EM anomaly symbols;

- EM – AeroTEM off-time profiles Z5 – Z 15 with EM anomaly symbols.

The coordinate/projection system for the maps is NAD83 – UTM Zone 10N. For reference, the latitude and longitude in WGS84 are also noted on the maps. All the maps show flight path trace, skeletal topography and conductor picks represented by an anomaly symbol classified according to calculated off-time conductance, a thick or thin classification and an anomaly identifier label. The anomaly symbol legend and survey specifications are displayed on the left margin of the maps. Warren D.Robb (P.Geo.) and Barry Price (P.Geo.) have reviewed the plans of magnetic and EM surveys and in brief the following area are of interest and need to be prospected, mapped and if warranted, subjected to further ground-based surveys:

1. The northwestern part of grid (Plate 1) is very active in terms of electromagnetic anomalies;
2. Electromagnetic activity is subdued in the central area (Plate 2) and fewer anomalies exist in the southeastern area (Plate 3);
3. Magnetic anomalies in Plates 1 and 2 area are due to either magnetic intrusions or magnetic rich volcanics. In Plate 3 area magnetic anomalies are more linear trends, possibly reflecting dyking or magnetic and non magnetic flows;
4. Both magnetic highs area and low area should be mapped and prospected, as magnetic lows may indicate altered area, and magnetic areas could indicate skarn areas;
5. All strong electromagnetic (EM) anomalies deserve to be located on the ground and prospected.

A compilation map is presented in Figure 6.

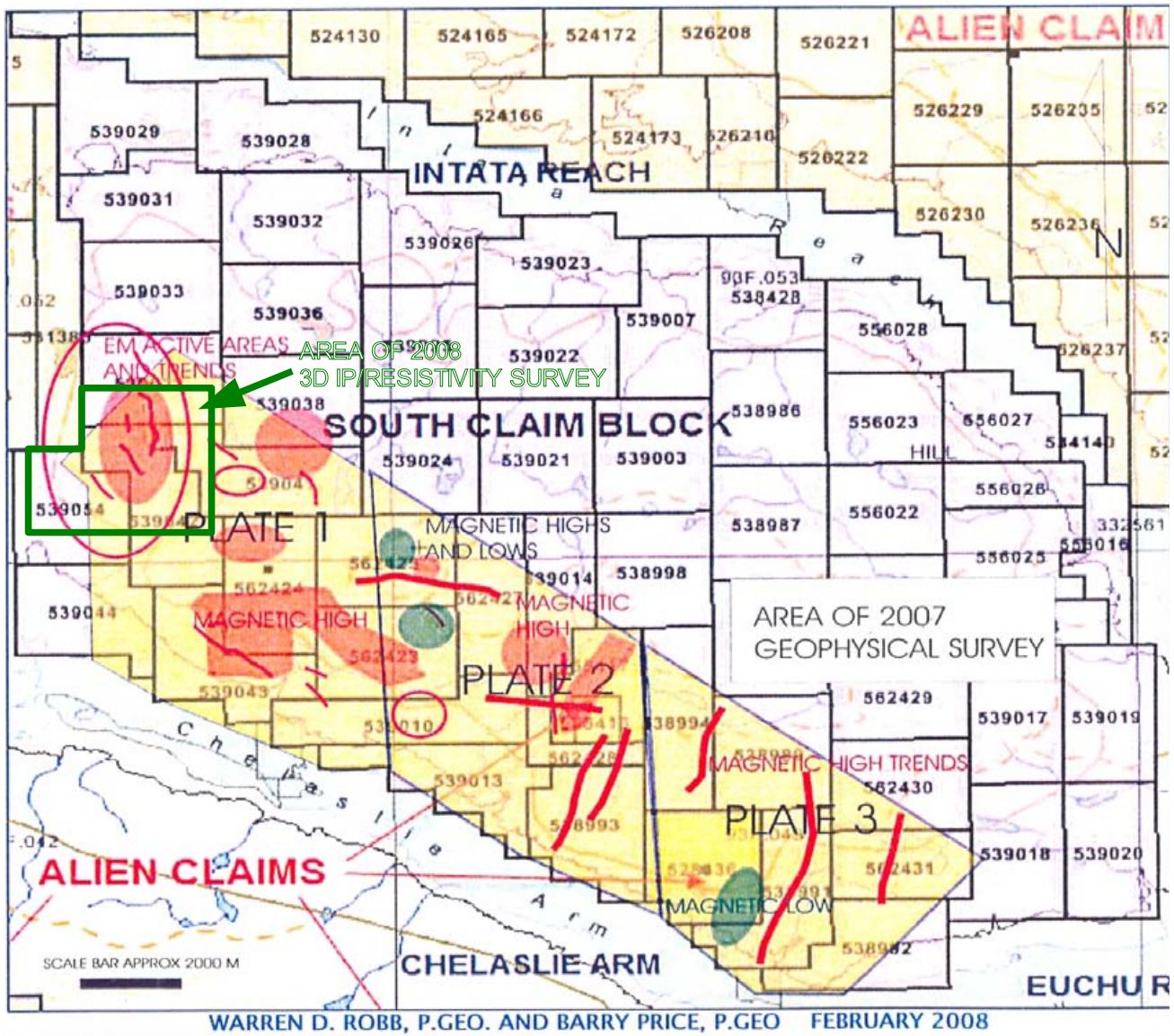


Figure 6: Regional Geophysical Compilation Map.

From Technical Report, South Claims Property, Golden Dragon Explorations Inc. with addition of area of 2008 3D IP/Resistivity Survey

8.3. 2008 3D IP/Resistivity Survey

A 3-D IP/Resistivity survey was carried out on behalf of 808718 (BC) Ltd. on its Nechako (Ootsa Lake) property between August and October, 2008. The survey was contracted to SJ Geophysics Ltd. of Delta, BC. The survey was a follow-up ground exploration program recommended by Warren D. Robb (P.Geo.) and Barry Price (P.Geo.) in their technical report, to cover anomalies identified by the airborne surveys and to help map geology and search for evidence of a epithermal alteration system.

A detailed description of the IP/Resistivity survey is provided in a report by Brian Chen, which includes descriptions of the equipment, arrays, processing and 3D IP methodology as well as an interpretation of the results. The following is the summary from the report.

8.3.1. Location and Line information

The geophysical grid on Nechako (Ootsa Lake) property is located 120km southwest of Vanderhoof, British Columbia, on the south shore of the Ootsa Lake (Figure 1). To access the south shore of Ootsa Lake, a private barge must be contracted at Intata Reach on the north shore. To access the Intata Reach from Vanderhoof, follow the Kenney Dam Road to the dam then turn west and drive along the Deerhorn Forest Service Road until it joins the Ootsa Lake Road. The drive to the barge is approximately two and a half hours. On the south shore, drive east along Swamp Road until reaching the camp just east of Table Hill. The camp was situated on a barge contracted from Cheslatta Forest Products and moored to the shore. To access the grid from the camp, the crew drove along Swamp Road until kilometer marker 23 then turned left and followed a forest service road to the grid. Figure 7 shows the survey grid within the Nechako (Ootsa Lake) property. Accommodation was provided by 808718(BC) Ltd. See Appendix C for a summary of costs relating to the IP survey and the accommodations.

The design grid consisted of 37 east-west trending lines. Local line and station coordinates were the final four digits of the UTM coordinates. Lines were labeled from L2000N to L5600N and spaced at 100m intervals with stations at 50m intervals. Lines L2000N to L3500N started at station 1150E; Lines L3600N to L4600N started at station 2150E. Western end of lines L2000N to L2300N were cut short because of bear sightings in the area. All lines, except for lines 2000N, 3700N, 3800N and 3900N,

ended at station 4500E. Line 2000N was cut short 350m on its east due to dense bush, Lines 3700N, 3800N and 3900N were cut short 150m to 250m on their east to avoid the swamp. There were three north-south trending tie lines along stations 2150E, 3250E and 4500E. Survey lines are shown in Figure 7. Appendix D summarizes the lines and stations surveyed.

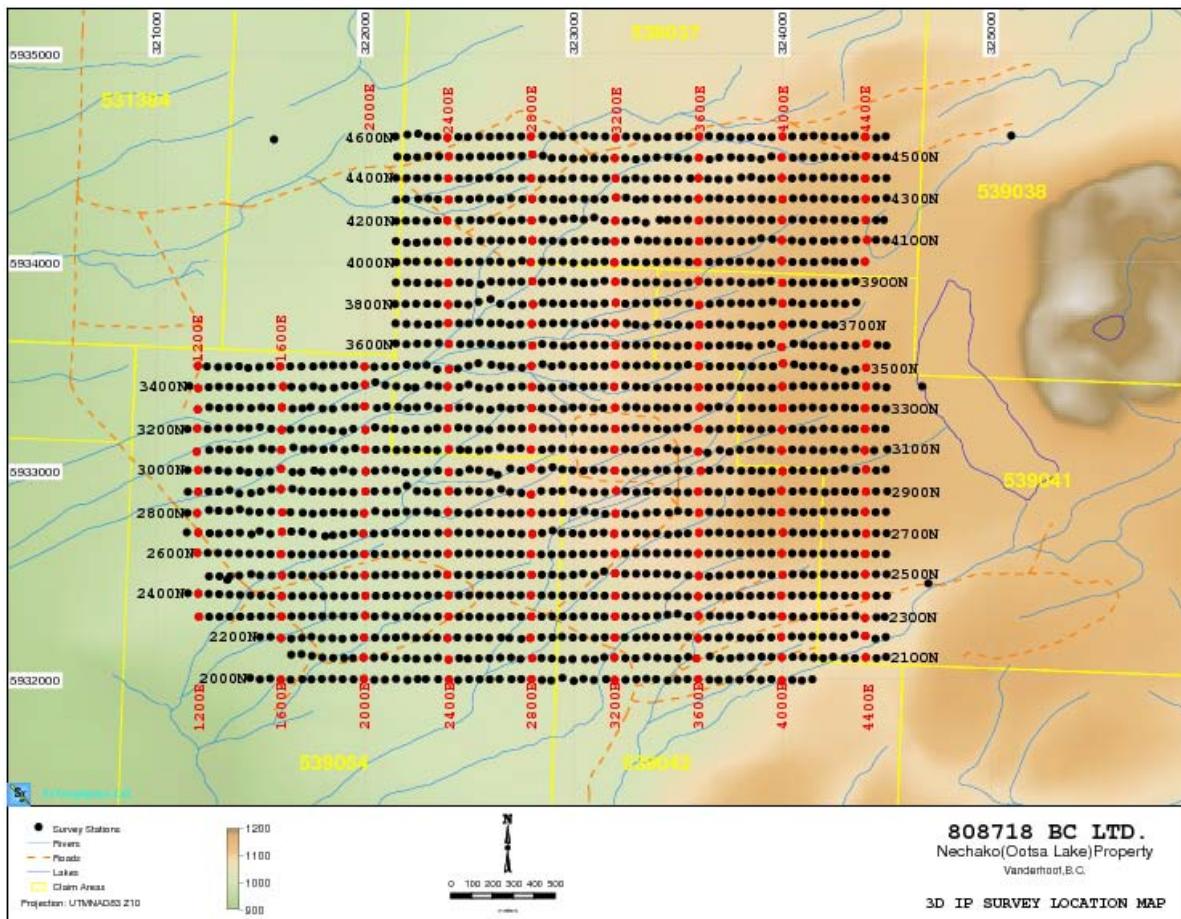


Figure 7: 3D IP/Resistivity Survey Grid.

8.3.2. Line Cutting

Prior to commencing the geophysical survey, between July to October 2008, crew from 0806827(BC) Ltd. started line cutting the grid.

8.3.3. Field work and instrumentation

The 3D IP survey on the Nechako (Ootsa Lake) property was conducted from August to October, 2008. A total of 76 line-kilometers of 32 lines was surveyed in 37 days by a crew of SJ Geophysics Ltd. consisting of between 4 to 7 workers. A modified pole-dipole configuration array was used with 11-16 potential dipoles at 50-100m separations. The IP data was collected using SJ Geophysics Full Wave Form receivers. A GDD transmitter was used to inject current on a 2 seconds on, 2 seconds off duty cycle for the duration of the program. Complete specification of the instruments employed can be found in Appendix E. For the production phase, the 3D configuration consisted of two current lines being recorded into the receiver line. The two current injection locations were on the two adjacent survey lines 100m away from the receiver line. Six remote stations, consisting of 3-4 long electrodes, were employed for the entire IP survey. To achieve better depth penetration and lower noise, the East remote was used when surveying the West half of the grid; vice versa for the West remote.

The potential array was implemented using specialized 8 conductor IP cables configured with 50m takeouts for the potential rods. At each current station, the electrodes used consisted of 15mm stainless steel rods of approximately 1m in length. For the potential line, the electrodes consisted of 10mm stainless steel rods 50cm in length. The exact location of the remote current is used in the geophysical calculations.

Location data was collected using Garmin handheld GPS units to an accuracy of 5m to 6m and Sunto Inclinometers to an accuracy of 1 to 2°. The location data was integrated with BC Trim DEM for the inversion process. The projection for the location data is UTM NAD83, Zone 10.

3D IP surveys are designed to take advantage of the interpretational functionality offered by 3D inversion techniques. Unlike conventional IP, the electrode arrays are no longer restricted to an in-line geometry. Typically, current electrodes and receiver electrodes are located on adjacent lines. Under these conditions, multiple current locations can be applied to a single receiver electrode array and data acquisition rates can be significantly improved over conventional surveys.

In a common 3D IP configuration, a receiver array is established end-to-end on a survey line while current electrodes are located on two adjacent lines. The survey typically starts at one end of the line and proceeds to the other end. A 12 dipole array normally consists of eight 50m dipoles and four 100m dipoles, which are located at the far end(s) of the array to achieve greater signal strength. Current

electrodes are advanced along the adjacent lines, starting at one end and advanced approximately 500m through the array at 50m increments. At this point, the receiver array is advanced 400m and the process is repeated down the line. Receiver arrays are typically established on every second line (200m apart) thereby providing subsurface coverage at 100m.

8.3.4. Data presentation

Inversion models

The inversion models presented in the following section were computed using data gathered during the 2008 survey. The software used is summarized in Appendix F.

Given the size of the grid and the high data quality, the inversion models were expected to show relatively complex geophysical features. The topography for the inversion model were extracted from the BC TRIM DEM as the elevations provided correlated well with the survey stations elevations provided by the digital GPS measurements. Moreover, the BC TRIM DEM provided good topographic coverage between the lines to assist in creating more accurate inversion models.

A single inversion with a mesh composed of 20m easting cells and 25m northing cells was calculated.

Visualization of the inversion models

False color contour maps of the inverted resistivity and chargeability results were produced for selected depths. Data was positioned using UTM coordinates gathered during the field work. These maps display the regional distribution of the geophysical trends, outlining strike orientation and possible fault offsets.

The topography variations add a level of complexity to the interpretation, especially with the use of plan maps. Plan maps can be displayed in two ways: depth below topography or as horizontal slices in terms of elevation. For the purposes of this report, the plan maps produced were created at depth below the surface.

Plan maps plotted for both resistivity and chargeability at depths of 25m, 50m, 75m, 100m, 150m, 200m, 250m and 300m are provided in Appendix G (with the section maps). Vertical slices of the resistivity and chargeability models are also plotted as false color sections for each survey line (Tx and Rx). This allows the direct comparison of the resistivity and chargeability variations.

With the computer technology that exists today, the 3D inversion results can also be viewed using a

3D visualization program such as UBCGIF's MeshTools3D program or open-source software such as Paraview. These programs allow one to plot contour and thresholds of the resistivity and chargeability models simultaneously. It enhances the interpretation process by illustrating the direct association between the different parameters.

8.3.5. Discussion of the IP/Resistivity models

The geophysical survey grid is underlain dominantly by undivided volcanic rocks of Hazelton Group with one north-northeast trending regional fault running through its eastern portion. The results of 2007 helicopter-borne AeroTEM EM and magnetic surveys revealed that the 3D IP/Resistivity survey area is "very active in terms of electromagnetic anomalies". Also, there are magnetic highs in the central part and east corner of the survey grid.

The resistivity inversion result is demonstrated in Figures 8 and 9. Figure 8 shows a 3D view of the relatively resistivity highs of the inversion model with isosurfaces. Figure 9 shows the inverted resistivity false color contour plan map at a depth of 100m below surface.

The resistivity result reveals a relatively low background resistivity value of approximately 80 Ohm-m. The relatively high resistivity features could be roughly grouped into two zones (resistivity zone 1 and 2) from east to west (see Figure 9). Several resistivity contacts are also sketched in Figure 9 and denoted by bold dashed lines in white. The west portion of the survey area is dominated by relatively high resistivity feature (resistive zone 1). Two northeast and east-northeast trending resistivity contacts are identified in this zone. The resistivity contact in south portion of resistive zone 1 related to a low resistivity layer which dips northwards at a low angle. The 3D view in upper panel of Figure 10 shows the low resistivity layer between the two resistive bodies (>500Ohm-m). Resistive zone 2 is situated in the north central portion of the survey grid, the relatively resistivity high feature dips eastwards at an angle of approximate 45 degree (see the 3D view in lower panel of Figure 10).

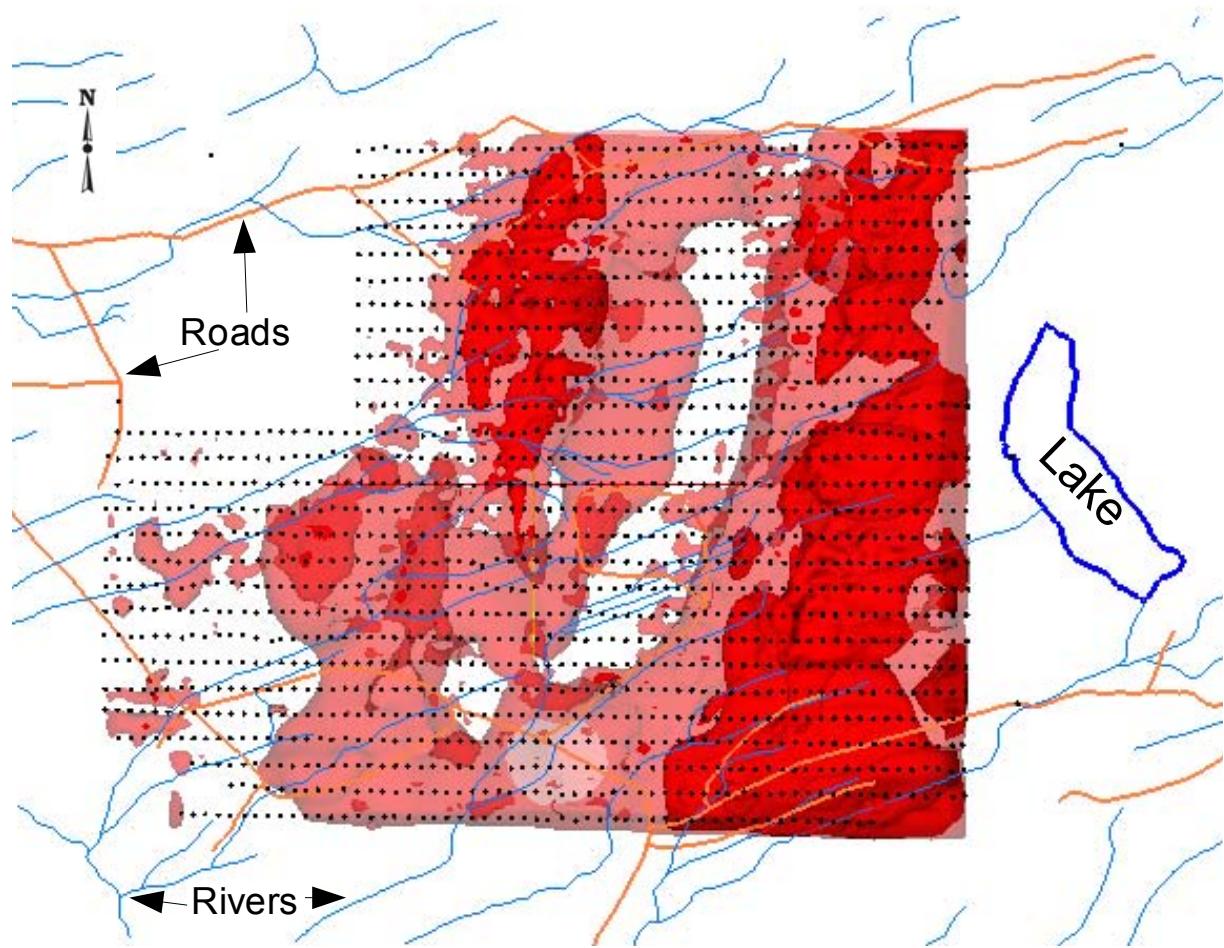


Figure 8: 3D Perspective Plots of Simplified Resistivity Inversion Model

Relatively high resistivity features ($>500\text{Ohm}\cdot\text{m}$ and $>100\text{Ohm}\cdot\text{m}$) are shown with red and light red isosurfaces respectively.

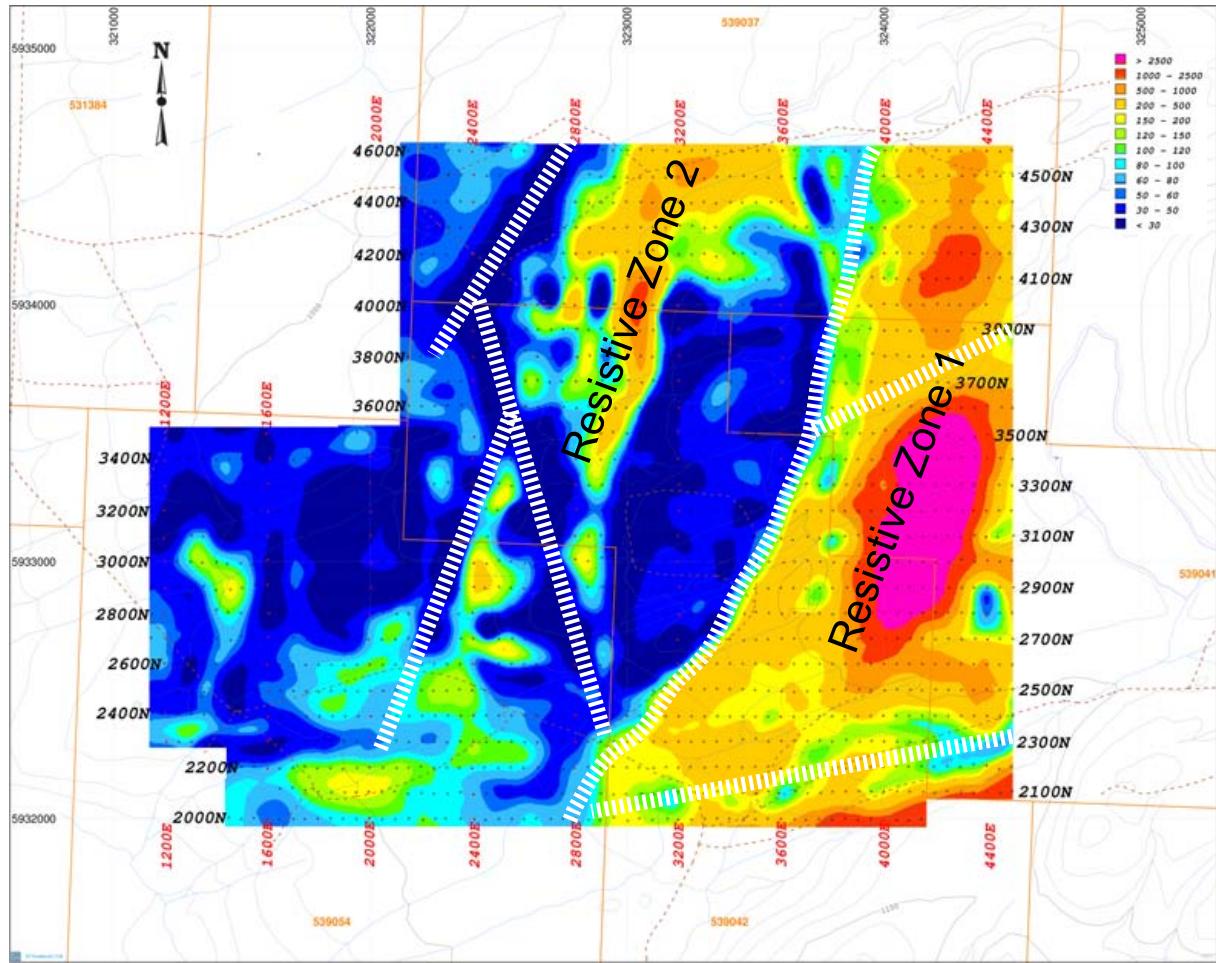
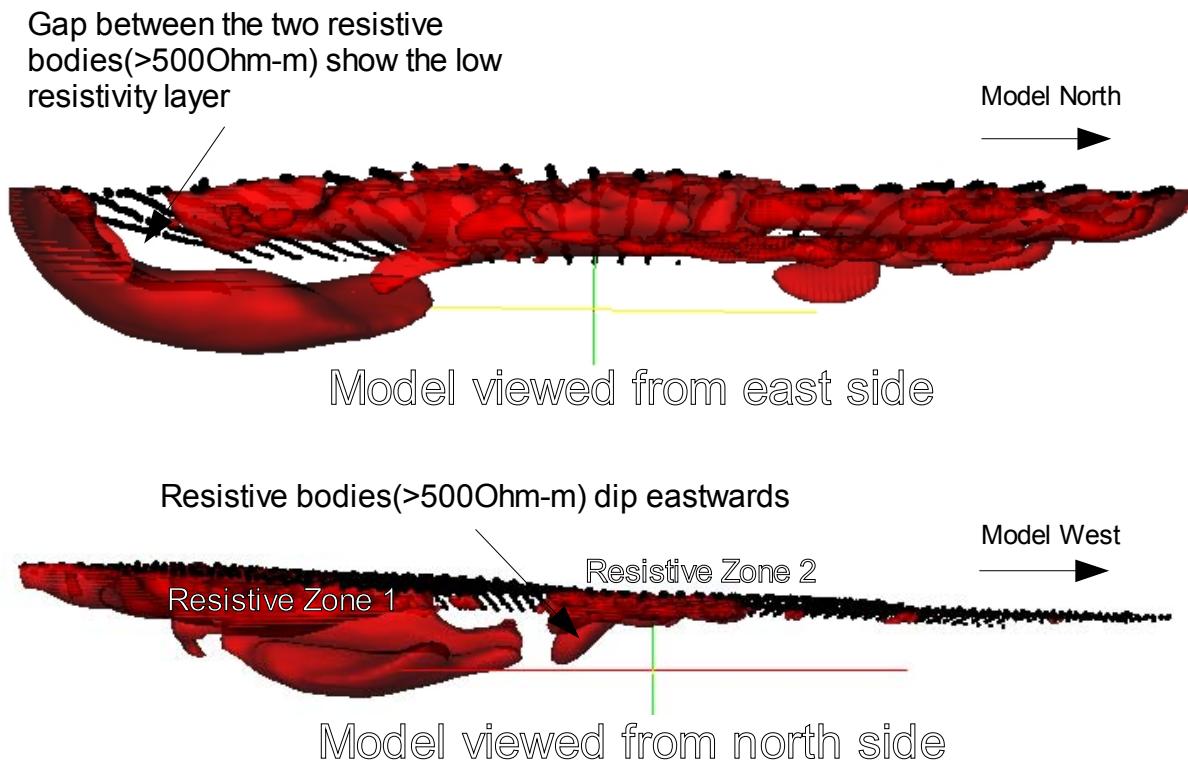


Figure 9: Inverted Resistivity False Color Contour Plan Map

Depth slice at 100m below surface. Bold dashed line in white shows the resistivity contact.



*Figure 10: 3D Perspective Plot of Simplified Resistivity Inversion Model
Relatively high resistivity features (>500Ohm-m) is shown with red isosurfaces.*

The chargeability result is shown in Figures 11 and 12. Figure 11 shows a 3D view of the relatively chargeability highs of the inversion model with isosurfaces. Chargeability highs ($> 5\text{ms}$) are shown with dark green isosurfaces, weak responses ($> 3.5\text{ms}$) are displayed with light green isosurfaces. Figure 12 shows the chargeability false color contour plan maps at a depth of 100m below surface with resistivity contacts. High chargeability features are shown in hot colors.

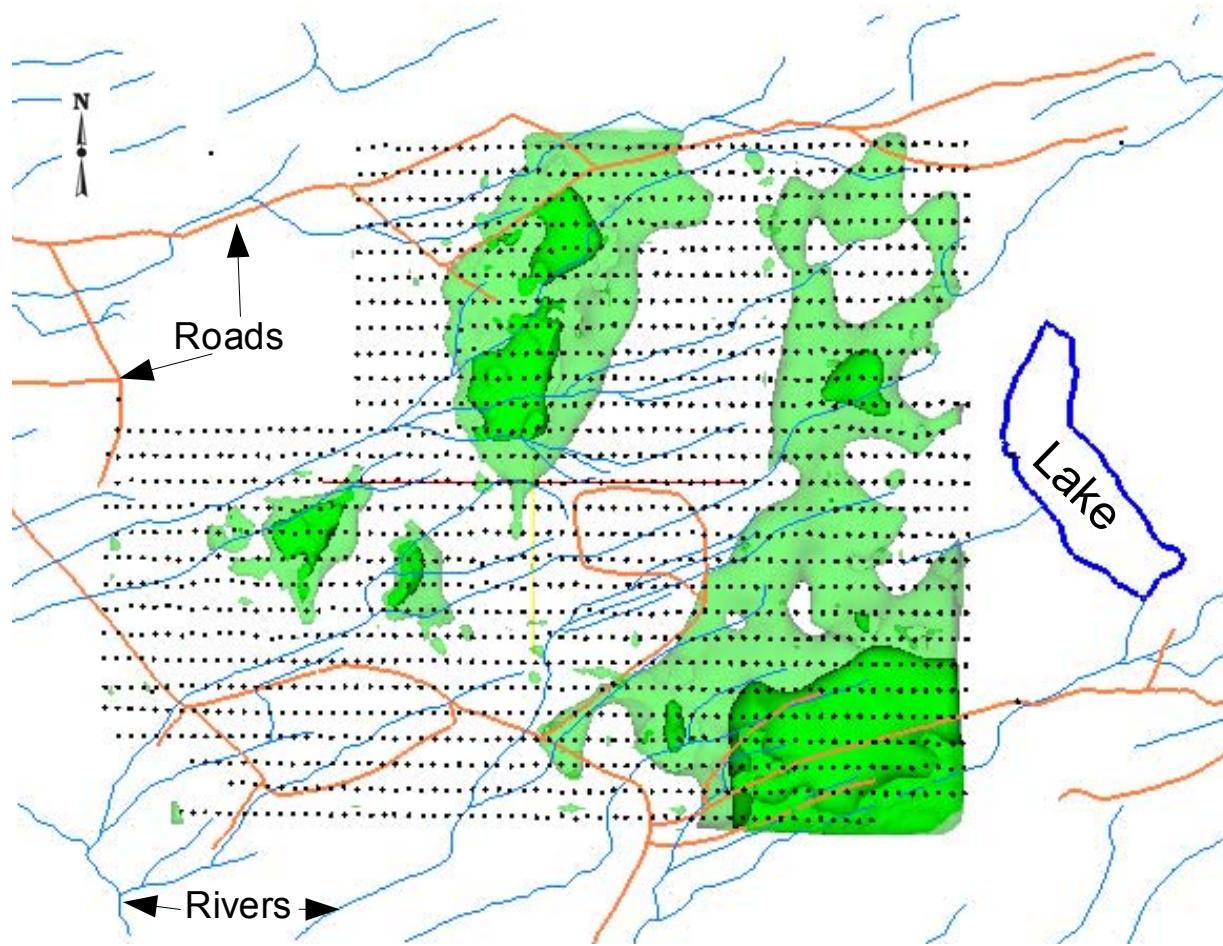


Figure 11: 3D Perspective Plots of Simplified Chargeability Inversion Model

Relatively chargeability highs (>5ms) are shown with green isosurfaces, weak chargeability responses(>3.5ms)are in light green.

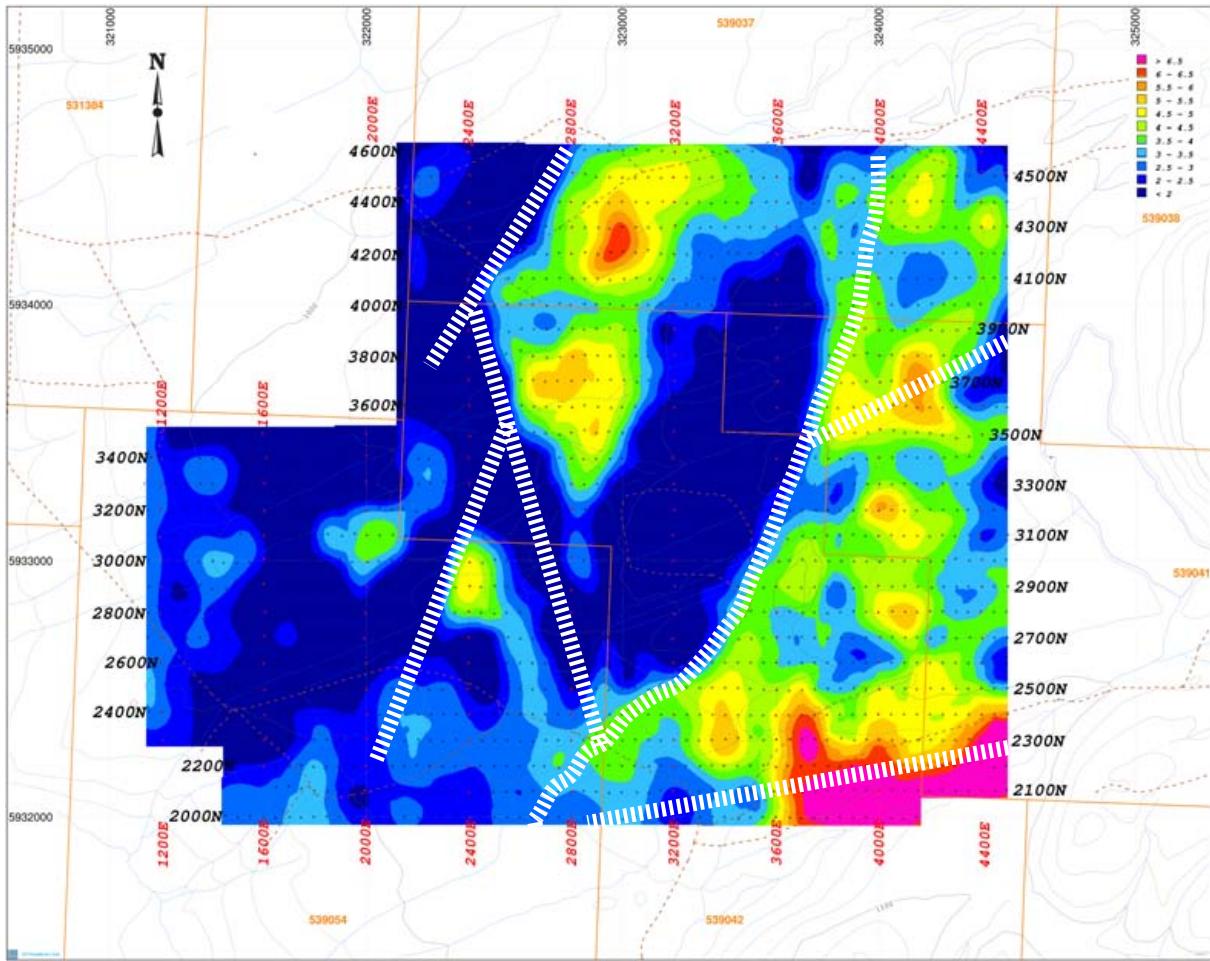


Figure 12: Inverted Chargeability False Color Contour Plan Map

Depth slice at 200m below surface. High chargeability values are presented in hot colors. The bold dashed lines in white show the resistivity contact.

The chargeability model is characterized by a relatively low background chargeability value of 3ms. In resistive zone 1, the resistive features seem to have weak chargeability response in general. The two resistivity contacts seem to associate with chargeability highs (Figure 12). The main chargeability highs (>5ms) are located in the southeastern corner of the grid, between lines 2000N and 2700N, stations 3600E to 4500E. It coincides with the resistivity contact associated with a low resistivity layer dipping northwards at a low angle. The regional geology (see Figure 3) shows a small dioritic intrusive showing near the southeast corner of the grid. From the helicopter-borne magnetic survey (2007), the diorite intrusive body is characterized by magnetic high. The chargeability highs might be related to the intrusive. Figure 13, the IP section of line 2300N, demonstrates the relation between the low

resistivity layer and the chargeability highs. Although, the regional geology shows that a north-northeast trending regional fault runs across the grid in the resistive zone 1, there is no strong evidence to demonstrate the correlation between the fault and the weak chargeability high responses in terms of the shape of the chargeability response.

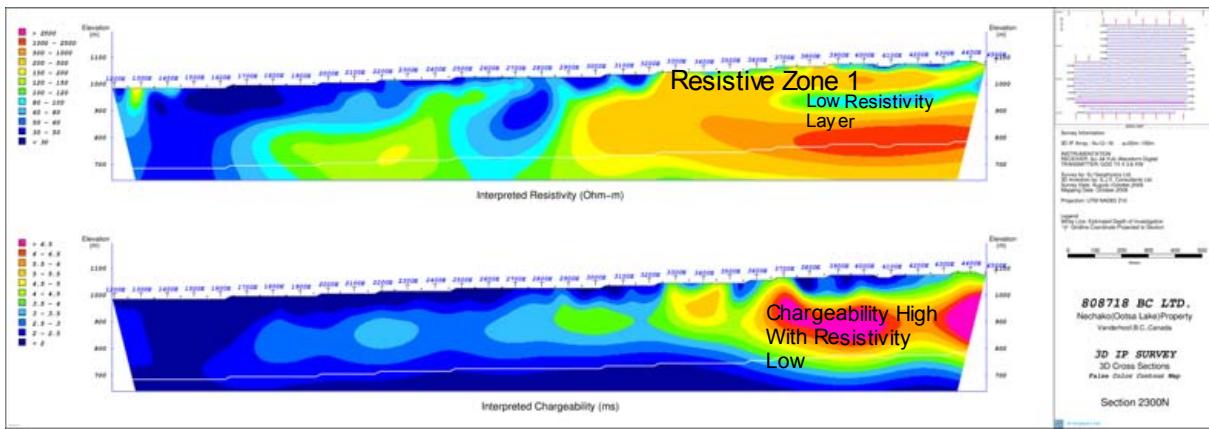


Figure 13: 3D Resistivity and Chargeability Sections, Line 2300N.

In resistive zone 2, the weak chargeability high responses ($>3.5\text{ms}$) seem to correlate with the resistive features. Figure 14 exhibits a 3D view of the integrated inversion models of both resistivity and chargeability. Figure 15, the IP section of line 3800N, shows the correlation of the relatively resistivity highs and the weak chargeability high responses at depth on this section. The bold dashed line in red denotes the trend of the dipping of the resistive feature. From this point of view, the weak chargeability response might be related to the relatively resistive volcanic rock (see the regional geology map). The low resistivity lineaments/resistivity contacts (denoted as bold dashed lines in white in Figures 9 and 12) may have a correlation with the EM anomalies discovered by the AeroTEM survey. The magnetic high in the central part of the grid, outlined by the helicopter-borne survey and the regional magnetic result, could be roughly related to the units in resistive zone 2, which is situated in the central part of the grid, due to airborne survey resolution. Detail geology and geochemical surveys are recommended to map and prospect the outlined features by the geophysical survey.

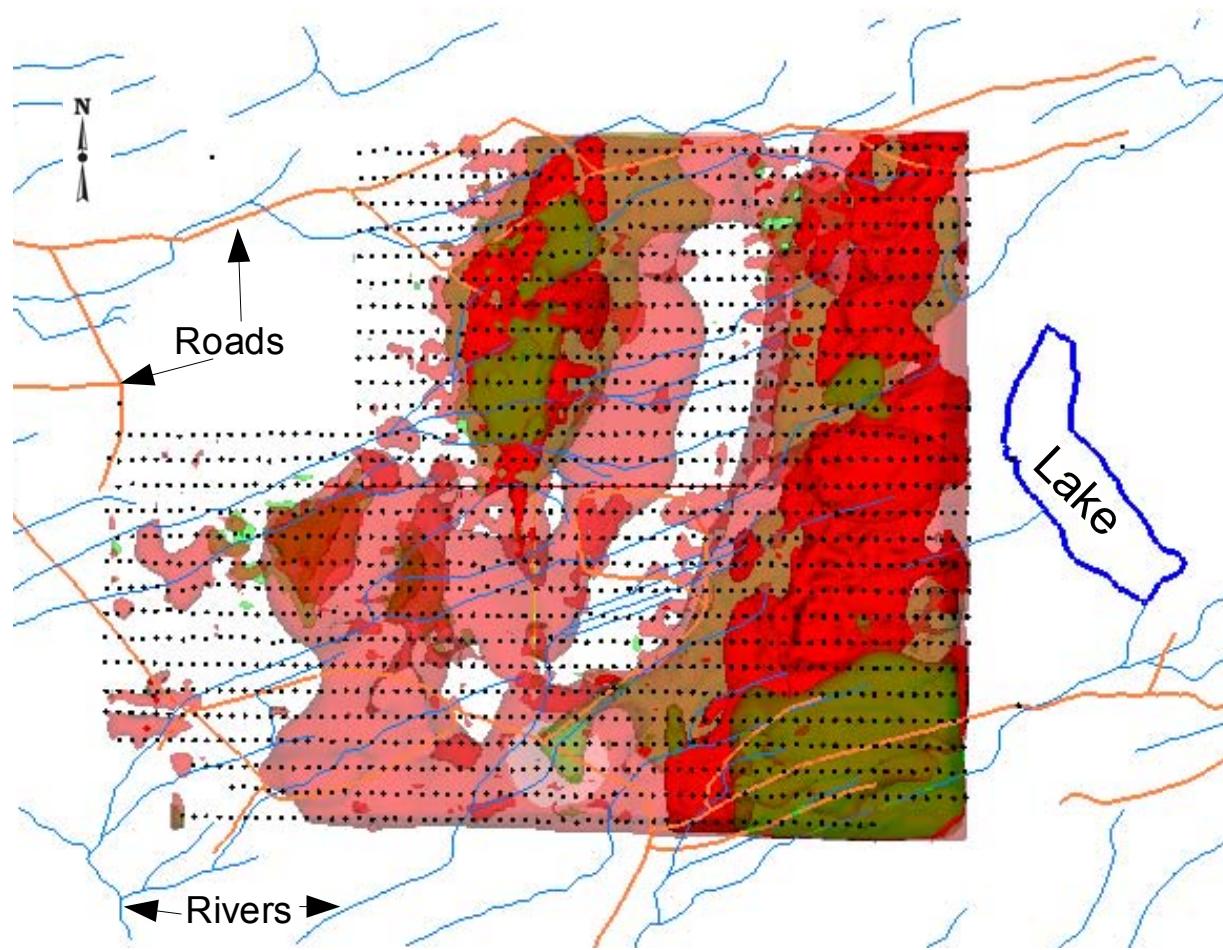


Figure 14: 3D Perspective Plots of Simplified Chargeability Inversion Model

High chargeability features are shown in green, values > 8ms are in dark green, values > 6ms are in light green.

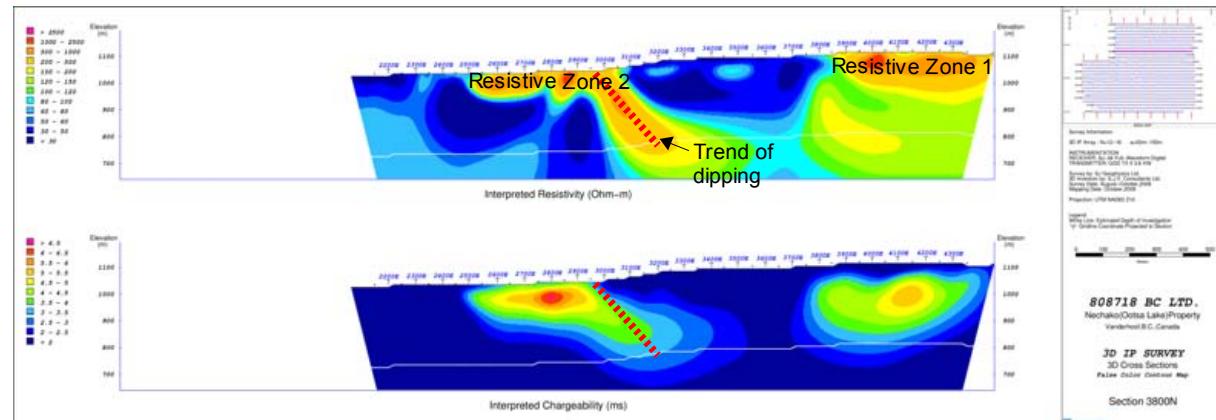


Figure 15: 3D Resistivity and Chargeability Sections, Line 3800N.

9. Conclusions and recommendations

On the 2008 3D IP/Resistivity survey grid, the relatively high resistivity features in the grid could be roughly grouped into two zones. The main chargeability high is located in the southeast corner of the grid and associated with a low resistivity layer dipping northwards at a low angle. It could be related to the dioritic intrusion. More geophysical survey to chase the main chargeability high is recommended if it shows evidence of relating to mineralization sought. The weak chargeability highs occur in the two resistive zones and show correlation with the relatively high resistivity features. They could be the responses arose from resistive volcanic rocks. There is no direct evidence to establish the relation between the weak chargeability highs and the north-northeast trending regional fault that runs across the grid in the resistive zone.

A ground magnetic survey over the same geophysical grid is recommended to help identify faults/alteration zones or intrusions. Exploration work including geological, geochemical to follow up the main chargeability high and the weak chargeability highs are also suggested.

Respectfully Submitted,
per S.J.V. Consultants Ltd.

Brian Chen, M.Sc.

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- Schimann, K (1993) Assessment report Yellow Moose Property AR#23099
- Schimann, K (1993) Assessment report Quartz Lake Property AR#23386
- Schimann, K (1993) Assessment report Yellow Moose Property AR#23387
- Schimann, K (1993) Assessment report Brewster Lake Property AR#23388
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- Schimann, K (1994) Assessment report Lucas West Property AR#23744
- Schimann, K (1994) Assessment report Lucas Property AR#23745
- Schimann, K (1994) Assessment report Tam Property AR#23746
- Schimann, K (1994) Assessment report Saunders Property AR#23747
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Appendix B: Statement of qualifications (Brian Chen)

I, Brian Chen, of the city of Delta, British Columbia, hereby certify that:

- 1.** I graduated from the University of Science and Technology of China in 1989 with a Bachelor of Science degree in geophysics and from South China Sea Inst. Of Oceanology, CAS in 1992 with a Master of Science degree in Mathematical geology.
- 2.** I have been working in geophysics since 1992.
- 3.** I have no interest in 808718(BC) Ltd. or in any property within the scope of this report, nor do I expect to receive any.
- 4.** My work is regularly reviewed by a registered professional geoscientist registered within the province of British Columbia.

Signed by: _____

Brian Chen, M.Sc., B.Sc.

Date: _____

Appendix C: Statement of costs

808718(BC) Ltd., Nechako (Ootsa Lake) property statement of expenditures

From submission 1, dated 2008/Sep/18 event numbers 4237153 and 4237156:

Costs provided by 808718(BC) Ltd.

| | |
|---|--------------------------|
| Equipment/Vehicle (Transport and Rental)/Accommodation (Camp Rental, Food, Fuel etc) | 107253.21\$ |
| Line cutting | 79435.13\$ |
| Support/Management | 9032.73\$ |
| 3D Induced Polarization survey | 221882.06\$ |
| Total (\$CDN) | <hr/> 417603.13\$ |

Appendix D: Survey line summary

| <i>Line(N)</i> | <i>Bol Station(E)</i> | <i>EOL Station(E)</i> | <i>Type</i> | <i>Remote Used</i> | <i>Surveyed Length</i> |
|----------------|-----------------------|-----------------------|-------------------|---|------------------------|
| 2000 | 1450 | 4150 | CX | 2501N/1350E, 2501N/4700E | 2700 |
| 2100 | 1650 | 4400 | RX | | 2750 |
| 2200 | 1500 | 4500 | CX | 2501N/1350E, 2501N/4700E | 3000 |
| 2300 | 1200 | 4400 | RX | | 3200 |
| 2400 | 1150 | 4500 | CX | 2501N/1350E, 2501N/4700E | 3350 |
| 2500 | 1200 | 4400 | RX | | 3200 |
| 2600 | 1150 | 4500 | CX | 2501N/1350E, 2501N/4700E | 3350 |
| 2700 | 1200 | 4400 | RX | | 3200 |
| 2800 | 1150 | 4500 | CX | 2501N/1350E, 2501N/4700E | 3350 |
| 2900 | 1200 | 4400 | RX | | 3200 |
| 3000 | 1150 | 4500 | CX | 2501N/1350E, 2501N/4700E 3401N/4700E,3601N/1201E | 3350 |
| 3100 | 1200 | 4400 | RX | | 3200 |
| 3200 | 1150 | 4500 | CX | 3401N/4700E,3601N/1201E | 3350 |
| 3300 | 1200 | 4400 | RX | | 3200 |
| 3400 | 1150 | 4500 | CX | 3401N/4700E,3601N/1201E | 3350 |
| 3500 | 1200 | 4400 | RX | | 3200 |
| 3600 | 2150 | 4500 | CX | 3401N/4700E,3601N/1201E 4601N/1600E,4601N/5100E | 2350 |
| 3700 | 2150 | 4250 | RX | | 2100 |
| 3800 | 2150 | 4350 | CX | 4601N/1600E,4601N/5100E | 2200 |
| 3900 | 2150 | 4350 | RX | | 2200 |
| 4000 | 2150 | 4400 | CX | 4601N/1600E,4601N/5100E | 2250 |
| 4100 | 2150 | 4500 | RX | | 2350 |
| 4200 | 2150 | 4500 | CX | 4601N/1600E,4601N/5100E | 2350 |
| 4300 | 2150 | 4500 | RX | | 2350 |
| 4400 | 2150 | 4500 | CX | 4601N/1600E,4601N/5100E | 2350 |
| 4500 | 2150 | 4500 | RX | | 2350 |
| 4600 | 2150 | 4500 | CX | 4601N/1600E,4601N/5100E | 2350 |
| | | | Total line metres | | 76150 |

Appendix E: Instrument specifications

SJ-24 FULL WAVEFORM DIGITAL IP RECEIVER

| Technical: | |
|---------------------------------|---|
| Input impedance: | 10 Mohm |
| Input overvoltage protection: | up to 1000V |
| External memory: | Unlimited readings |
| Number of dipoles: | 4 to 16+, expandable |
| Synchronization: | Software signal post-processing user selectable |
| Common mode rejection: | More than 100 dB (for Rs=0) |
| Self potential (Sp): | Range:-5V to + 5V Resolution: 0.1mV Proprietary intelligent stacking process rejecting strong non-linear SP drifts. |
| Primary voltage: | Range: 1µV – 10V (24bit) Resolution: 1µV Accuracy: typ. <1.0% |
| Chargeability: | Resolution: 1µV/V Accuracy: typ. <1.0% |
| General (4 dipole unit): | |
| Dimensions (HWD): | 18 x 16 x 9cm |
| Weight: | 1.1kg |
| Battery: | 12V external |
| Operating temperature range: | -20° to 40°C |

GDD Tx II IP TRANSMITTER

| | |
|------------------------|---|
| Input voltage: | 120V / 60 Hz, or 240V / 50Hz (optional) |
| Output power: | 3.6kW maximum |
| Output voltage: | 150 to 2200V |
| Output current: | 5mA to 10A |
| Time domain: | 1, 2, 4, 8 second on/off cycle |
| Operating temp. range: | -40° to +65°C |
| Display: | Digital LCD read to 0.001A |
| Dimensions (HWD): | 34 x 21 x 39 m |
| Weight: | 20kg |

Appendix F: Software programs

3D IP inversion program

“Inversion” programs have recently become available that allow a more definitive interpretation, although the process remains subjective.

The purpose of the inversion process is to convert surface IP/Resistivity measurements into a realistic “Interpreted Depth Section.” However, note that the term is left in quotation marks. The use of the inversion routine is a subjective one because the input into the inversion routine calls for a number of user selectable variables whose adjustment can greatly influence the output. The output from the inversion routines do assist in providing a more reliable interpretation of IP/Resistivity data, however, they are relatively new to the exploration industry and are, to some degree, still in the experimental stage.

The inversion programs are generally applied iteratively to evaluate the output with regard to what is geologically known, to estimate the depth of detection, and to determine the viability of specific measurements.

The Inversion Program (DCINV3D) used by the SJ Geophysical Group was developed by a consortium of major mining companies under the auspices of the UBC-Geophysical Inversion Facility. It solves two inverse problems. The DC potentials are first inverted to recover the spatial distribution of electrical resistivity, and, secondly, the chargeability data (IP) are inverted to recover the spatial distribution of IP polarizable particles in the rocks.

The interpreted depth section maps represent the cross sectional distribution of polarizable materials, in the case of IP effect, and the cross sectional distribution of the resistivity, in the case of the resistivity parameter.

GRASS

Geographic Resources Analysis Support System, commonly referred to as GRASS GIS, is a Geographic Information System (GIS) used for data management, image processing, graphics production, spatial modelling, and visualization of many types of data. It is free software (open source).

Originally developed as a tool for land management and environmental planning by the military, GRASS has evolved into a powerful utility with a wide range of applications in many different areas of scientific research. GRASS is currently used in academic and commercial settings around the world, as

well as many governmental agencies including NASA, NOAA, USDA, DLR, CSIRO, the National Park Service, the U.S. Census Bureau, USGS, and many environmental consulting companies.

The new GRASS6 release introduces a new topological 2D/3D vector engine and support for vector network analysis. Attributes are now managed in a SQL-based DBMS. A new display manager has been implemented. The NVIZ visualization tool was enhanced to display 3D vector data and voxel volumes. GRASS is integrated with GDAL/OGR libraries to support an extensive range of raster and vector formats, including OGC-conformal Simple Features.

For more information, visit <http://grass.itc.it/>.

LOGISTICAL REPORT

FOR

808718 (BC) LTD.

3D INDUCED POLARIZATION SURVEY

ON THE

NECHAKO (OOTSA LAKE) PROPERTY

LAT: 53° 31' 7"N LONG: 125° 40' 36"W (NAD83)

Location: Vanderhoof, British Columbia, Canada

Mining Zone: Omineca

NTS mapsheet: 093F12

BCGS TRIM MAPSHEET: 93F052

SURVEY CONDUCTED BY
SJ GEOPHYSICS LTD.
AUGUST – OCTOBER 2008

REPORT WRITTEN BY
JOHN LINDNER, JAY WATT
SJ GEOPHYSICS LTD.
OCTOBER 2008

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1. Introduction

A three-dimensional induced polarization (3D IP) survey was conducted on the Ootsa Lake property for 808718 (BC) Ltd. by SJ Geophysics Ltd. The property was located 120 km southwest of Vanderhoof, BC, Canada. The ground geophysical program, consisting of 32 lines totaling 76 line-kilometers, was surveyed between August and October, 2008. Initial data processing and quality control was performed in the field; final quality control and interpretation were completed by S.J.V. Consultants in our Delta, BC, office.

This logistical report summarizes the operational aspects of the survey and the survey methodologies used; it does not discuss any interpretation of the results of the geophysical survey.

2. Location and Line Information

The Ootsa Lake property is located 120km southwest of Vanderhoof, British Columbia, on the south shore of the Ootsa Lake. To access the south shore of Ootsa Lake, a private barge must be contracted at Intata Reach on the north shore. To access the Intata Reach from Vanderhoof, follow the Kenney Dam Road to the dam then turn west and drive along the Deerhorn Forest Service Road until it joins the Ootsa Lake Road. The drive to the barge is approximately two and a half hours. On the south shore, drive east along Swamp Road until reaching the camp just east of Table Hill. The camp was a barge contracted from Cheslatta Forest Products and moored to the shore. To access the grid from the camp, the crew drove along Swamp Road until kilometer marker 23 then turned left and followed a forest service road to the grid. Figure 1 shows the grid location.

The design grid consisted of 37 east-west trending lines. Local line and station coordinates were the final four digits of the UTM coordinates. Lines labeled from L2000N to L5600N. Lines spaced at 100m intervals with stations at 50m intervals. Lines L2000N to L3500N started at station 1150E; Lines L3600N to L4600N started at station 2150E. Western end of lines L2000N to L2300N were cut short because of bear sightings in the area. All lines, except for lines 2000N, 3700N, 3800N and 3900N, ended at station 4500E. Line 2000N was cut short 350m on its east due to dense bush, Lines 3700N, 3800N and 3900N were cut short 150m to 250m on their east to avoid the swarm. There were three north-south trending tie lines along stations 2150E, 3250E and 4500E. See table on Appendix C for detailed survey line information.

Lines were put in by GPS and were cut with machete and axe by a line cutting crew contracted by the client. Lines were marked with pink flagging tape and stations were marked with a combination of pink and blue flagging tape. The grid was a combination of dense forest, swamp and young growth (on old clear cut blocks). Numerous roads and tracks provided access across the grid (Figure 2). The elevation across the grid ranges from 970 to 1150m.

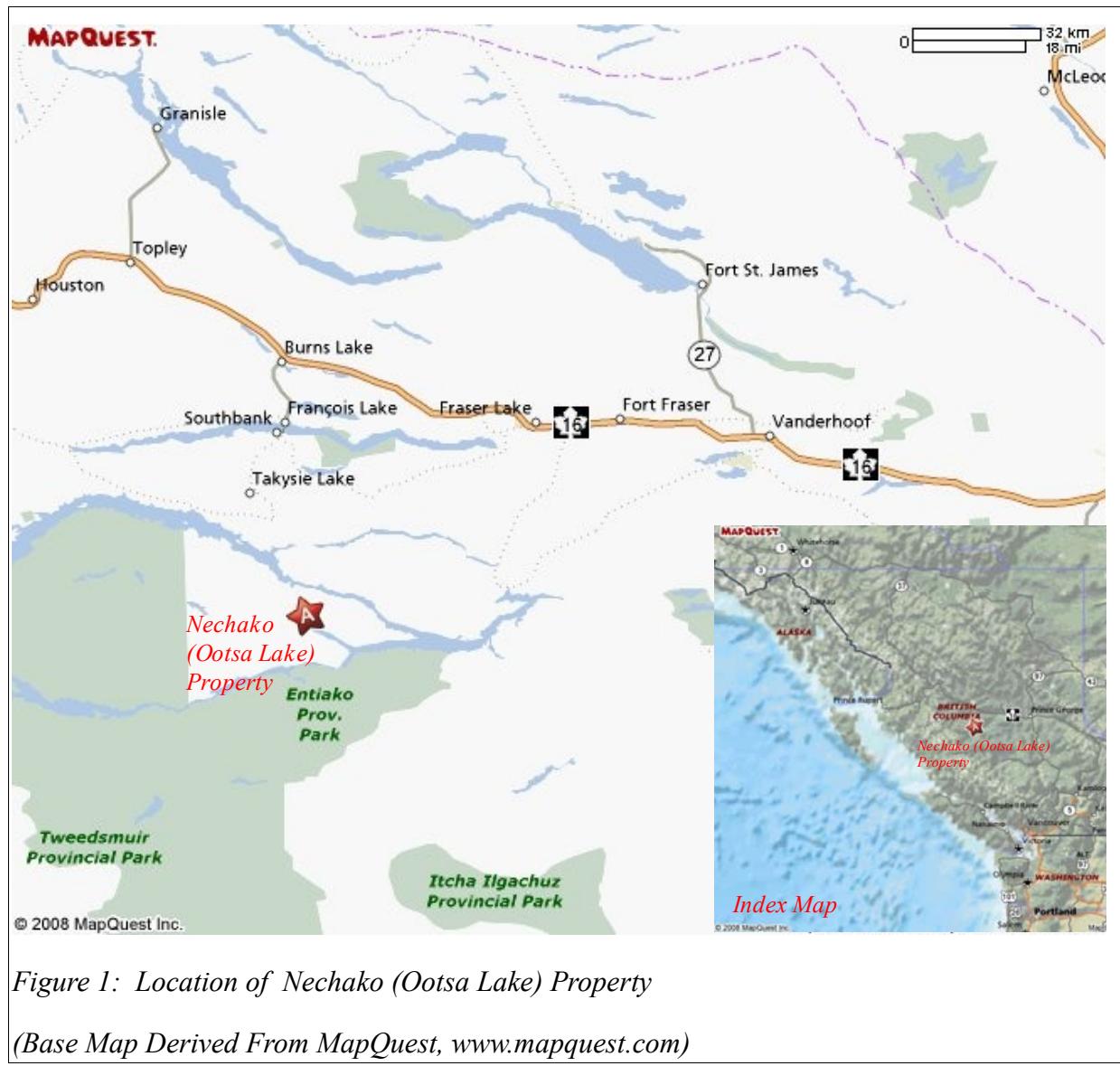
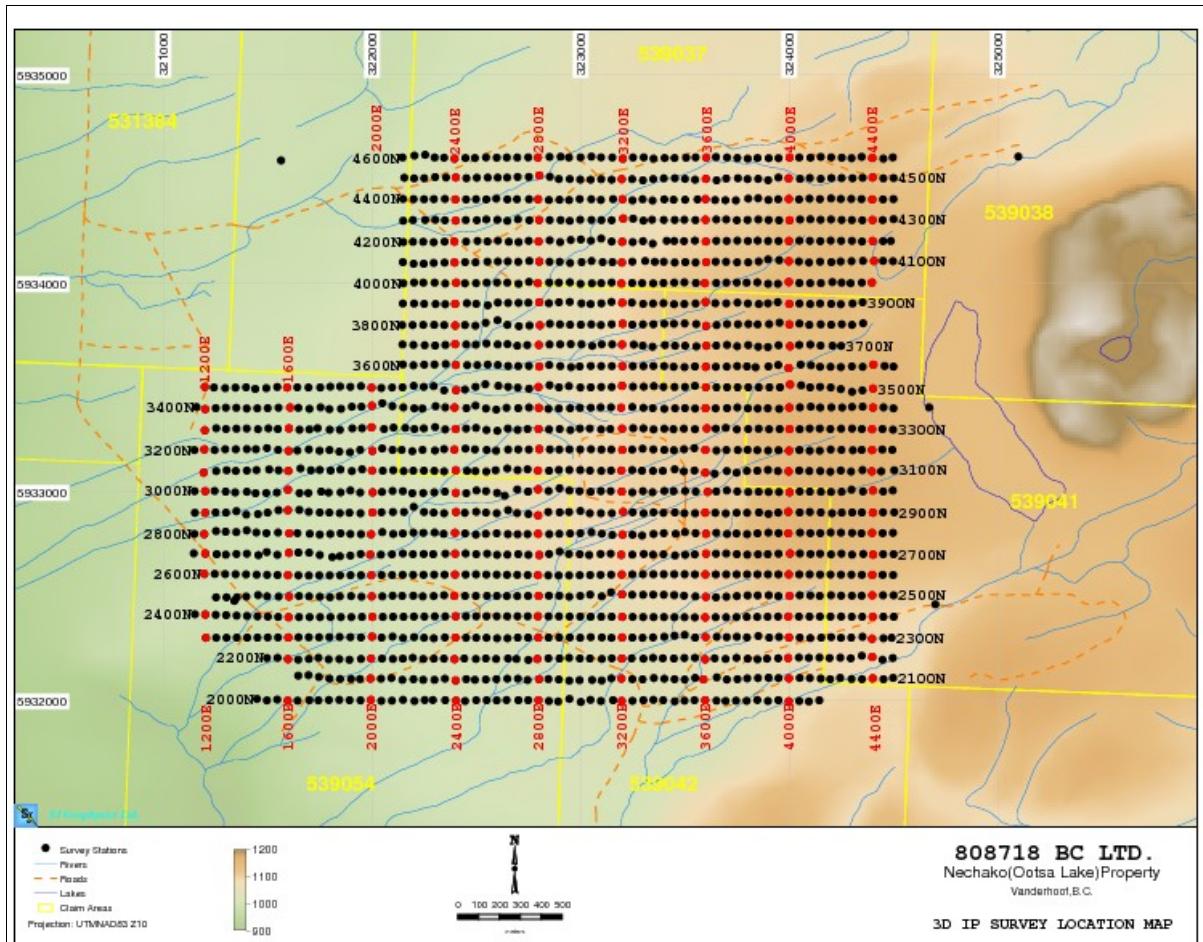


Figure 1: Location of Nechako (Ootsa Lake) Property

(Base Map Derived From MapQuest, www.mapquest.com)



3. Field Work and Instrumentation

3.1. Field Logistics

During the survey, the SJ Geophysics Ltd. crew consisted of between 4 and 7 workers. Initially, the crew consisted of John Lindner (geophysicist), John Wilkinson (logistics), Liam Fowlie and Dustin Walcer with assistance from Jason Prince (a helper hired by the client). Jay Watt (operator), Rene Poulin and Morgan Bezembinder joined the crew when John Wilkinson and John Lindner went on break. Adam Pearce and Casey Vandenberg mobilized to join the crew on September 9th and 19th respectively, Dustin Walcer and Morgan Bezembinder left on break on September 19th. Butch Houlind (expeditor) and Jason Gordon (representative from the chestalla forest products) provided logistical support. The crew stayed on a camp barge moored to the south shore of Ootsa Lake on the Intata Reach just east of Table Hill.

John L. and John W. mobilized with truck and survey equipment from Terrace, B.C., on August 27, 2008, picking up Liam and Dustin in Smithers. After a night in Vanderhoof, the crew met with Butch and Mike Mulberry (client's field representative), then drove to pick up camp supplies that day.

On August 29, the crew laid out remotes, mother lines and made some test measurements. IP data acquisition began on August 30 from Line 2000N and proceeded northwards. IP data collection stopped on October 1st at the client's request. The ten northern lines, Lines 4700N to 5600N, were not surveyed. The crew cleaned up the grid and demobilized on October 2nd.

During the field survey, bear activity and a few poor weather days slowed down the production, other than that, the whole survey ran smoothly.

3.2. Survey Parameters and Instrumentation

A modified pole-dipole configuration was used with 11-16 potential dipoles at 50-100m separations. The potential array was connected using special 8-conductor cables with 50m takeouts spliced to short (50 cm) stainless steel electrodes hammered into the ground. Data were collected using a SJ-24 full waveform receiver. A GDD transmitter was used to inject current on a 2 seconds on, 2 seconds off duty cycle.

Six remote stations, consisting of 3-4 long electrodes, were employed for the entire IP survey. To achieve better depth penetration and lower noise, the East remote was used when surveying the

West half of the grid; vice versa for the West remote.

Location data was collected using Garmin handheld GPS units to an accuracy of 5m to 6m and Sunto Inclinometers to an accuracy of 1 to 2°. The location data was integrated with BC Trim DEM for the inversion process. The projection for the location data is UTM NAD83, Zone 10.

4. Geophysical Techniques

4.1. IP Method

The time domain IP technique energizes the ground with an alternating square wave pulse via a pair of current electrodes. During current injection, the apparent (bulk) resistivity of the ground is calculated from the measured primary voltage and the input current. Following current injection, a time decaying voltage is also measured at the receiver electrodes. This IP effect measures the amount of polarizable (or “chargeable”) materials in the subsurface rock.

Under ideal circumstances, high chargeability corresponds to disseminated metallic sulfides. Unfortunately, IP responses are rarely uniquely interpretable as other rock materials are also chargeable, including some graphitic rocks, clays and some metamorphic rocks (e.g., serpentinite). Therefore, it is prudent from a geological perspective to incorporate other data sets to assist in interpretation.

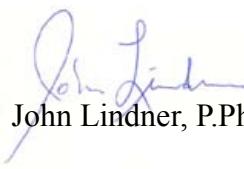
IP and resistivity measurements are generally considered repeatable to within about five percent. However, changing field conditions, such as variable water content or electrode contact, reduce the overall repeatability. These measurements are influenced to a large degree by the rock materials near the surface (or, more precisely, near the measuring electrodes). In the past, interpretation of a traditional IP pseudosection was often uncertain because strong responses located near the surface could mask a weaker one at depth.

4.2. 3DIP Method

Three dimensional IP surveys were designed to take advantage of the interpretative functionality offered by 3D inversion techniques. Unlike conventional 2DIP, the electrode arrays are no longer restricted to an in-line geometry. In the standard 3DIP configuration, a receiver array is established along a survey line while current electrodes are located on two adjacent lines. Current electrodes are advanced along the adjacent lines at fixed increments (25, 50, 100 or 200 m). A typical receiver array

consists of 12 to 16 dipoles separated by the same interval as the current lines or by some multiple of that interval. These spacings are sometimes modified to compensate for local conditions, such as inaccessible sites and streams, or the overall conductivity of ground. Receiver arrays are typically established on every second line. By injecting multiple current locations to a single receiver electrode array, data acquisition rates are significantly improved over conventional surveys.

Respectfully submitted,
As per S.J.V. Consultants Ltd.


John Lindner, P.Phys., M.Sc. (Physics), B.Sc. (Physics and astronomy)

Appendix A: Statement of Qualifications

I, John Lindner, of the city of Vancouver, British Columbia, hereby certify that:

1. I graduated from the University of Lethbridge in 2006 with a Masters of Science in physics and from the University of Victoria in 2003 with a Bachelors of Science in physics and astronomy.
2. I have been working in the mineral exploration industry since 2007.
3. I have no interest in 808718 BC Ltd. or in any property within the scope of this report, nor do I expect to receive any.

Signed by: John Lindner on

John Lindner, P.Phys., M.Sc. (Physics), B.Sc. (Physics and astronomy)

Appendix B: Instrument Specifications

SJ-24 Full Waveform Digital IP Receiver

Technical:

| | |
|-------------------------------|---|
| Input impedance: | 10 MΩ |
| Input overvoltage protection: | Up to 1000V |
| External memory: | Unlimited readings |
| Number of dipoles: | 4 to 16+, expandable |
| Synchronization: | Software signal post-processing user selectable |
| Common mode rejection: | More than 100 dB (for Rs =0) |
| Self potential (Sp): | Range:-5V to + 5V Resolution: 0.1 mV Proprietary intelligent stacking process rejects strong non-linear SP drifts |
| Primary voltage: | Range: 1µV – 10V (24 bit) Resolution: 1µV Accuracy: typically <1.0% |
| Chargeability: | Resolution: 1µV/V Accuracy: typically <1.0% |

Four-dipole digitizer:

| | |
|-------------------|----------------|
| Dimensions (HWD): | 18 x 16 x 9 cm |
| Weight: | 1.1 kg |
| Battery: | 12V external |
| Operating range: | -20 to 40°C |

GDD Tx II IP Transmitter

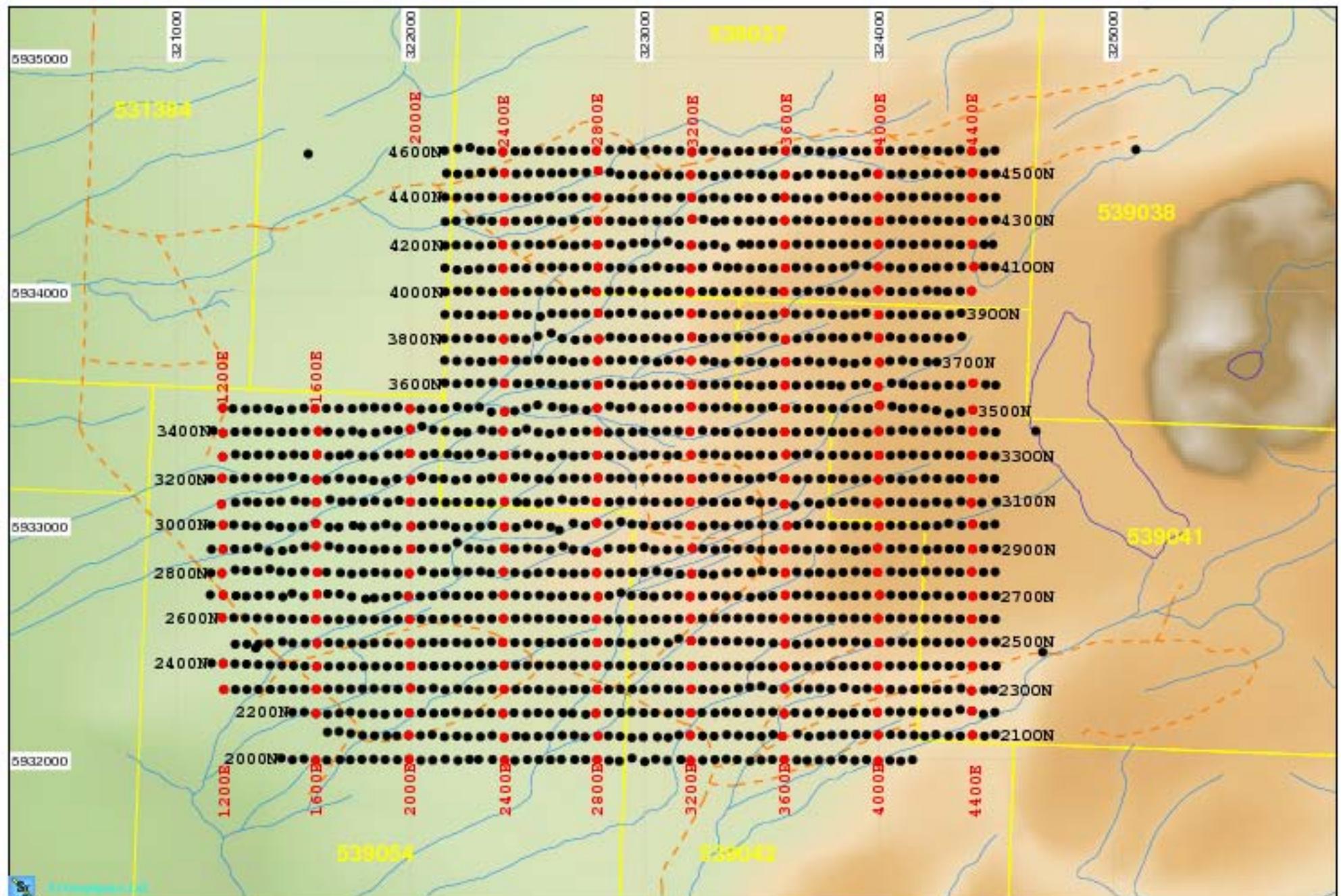
| | |
|------------------------|--|
| Input voltage: | 120V / 60 Hz or 240V / 50Hz (optional) |
| Output power: | 3.6 kW maximum |
| Output voltage: | 150 to 2200 V |
| Output current: | 5 mA to 10 A |
| Time domain: | 1, 2, 4, 8 second on/off cycle |
| Operating temp. range: | -40° to +65° C |
| Display: | Digital LCD read to 0.001 A |
| Dimensions (h w d): | 34 x 21 x 39 cm |
| Weight: | 20 kg |

Appendix C: 3D IP Summary Table

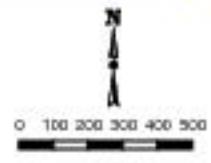
| <i>Line(N)</i> | <i>Bol Station(E)</i> | <i>EOL Station(E)</i> | <i>Type</i> | <i>Remote Used</i> | <i>Surveyed Length</i> |
|----------------|-----------------------|-----------------------|-------------|---|------------------------|
| 2000 | 1450 | 4150 | CX | 2501N/1350E, 2501N/4700E | 2700 |
| 2100 | 1650 | 4400 | RX | | 2750 |
| 2200 | 1500 | 4500 | CX | 2501N/1350E, 2501N/4700E | 3000 |
| 2300 | 1200 | 4400 | RX | | 3200 |
| 2400 | 1150 | 4500 | CX | 2501N/1350E, 2501N/4700E | 3350 |
| 2500 | 1200 | 4400 | RX | | 3200 |
| 2600 | 1150 | 4500 | CX | 2501N/1350E, 2501N/4700E | 3350 |
| 2700 | 1200 | 4400 | RX | | 3200 |
| 2800 | 1150 | 4500 | CX | 2501N/1350E, 2501N/4700E | 3350 |
| 2900 | 1200 | 4400 | RX | | 3200 |
| 3000 | 1150 | 4500 | CX | 2501N/1350E, 2501N/4700E 3401N/4700E,3601N/1201E | 3350 |
| 3100 | 1200 | 4400 | RX | | 3200 |
| 3200 | 1150 | 4500 | CX | 3401N/4700E,3601N/1201E | 3350 |
| 3300 | 1200 | 4400 | RX | | 3200 |
| 3400 | 1150 | 4500 | CX | 3401N/4700E,3601N/1201E | 3350 |
| 3500 | 1200 | 4400 | RX | | 3200 |
| 3600 | 2150 | 4500 | CX | 3401N/4700E,3601N/1201E 4601N/1600E,4601N/5100E | 2350 |
| 3700 | 2150 | 4250 | RX | | 2100 |
| 3800 | 2150 | 4350 | CX | 4601N/1600E,4601N/5100E | 2200 |
| 3900 | 2150 | 4350 | RX | | 2200 |
| 4000 | 2150 | 4400 | CX | 4601N/1600E,4601N/5100E | 2250 |
| 4100 | 2150 | 4500 | RX | | 2350 |
| 4200 | 2150 | 4500 | CX | 4601N/1600E,4601N/5100E | 2350 |
| 4300 | 2150 | 4500 | RX | | 2350 |
| 4400 | 2150 | 4500 | CX | 4601N/1600E,4601N/5100E | 2350 |

| <i>Line(N)</i> | <i>Bol Station(E)</i> | <i>EOL Station(E)</i> | <i>Type</i> | <i>Remote Used</i> | <i>Surveyed Length</i> |
|-----------------------|------------------------------|------------------------------|-------------------------|---------------------------|-------------------------------|
| 4500 | 2150 | 4500 | RX | | 2350 |
| 4600 | 2150 | 4500 | CX | 4601N/1600E,4601N/5100E | 2350 |
| | | | Total line metres | | 76150 |

Appendix G: Plan and section maps





● Survey Stations
 — Rivers
 - - Roads
 — Lakes
 Yellow Boxes: Claim Areas
 Projection: UTM NAD83 Z10





808718 BC LTD.
 Nechako(Ootsa Lake)Property
 Vanderhoof, B.C.


3D IP SURVEY LOCATION MAP


































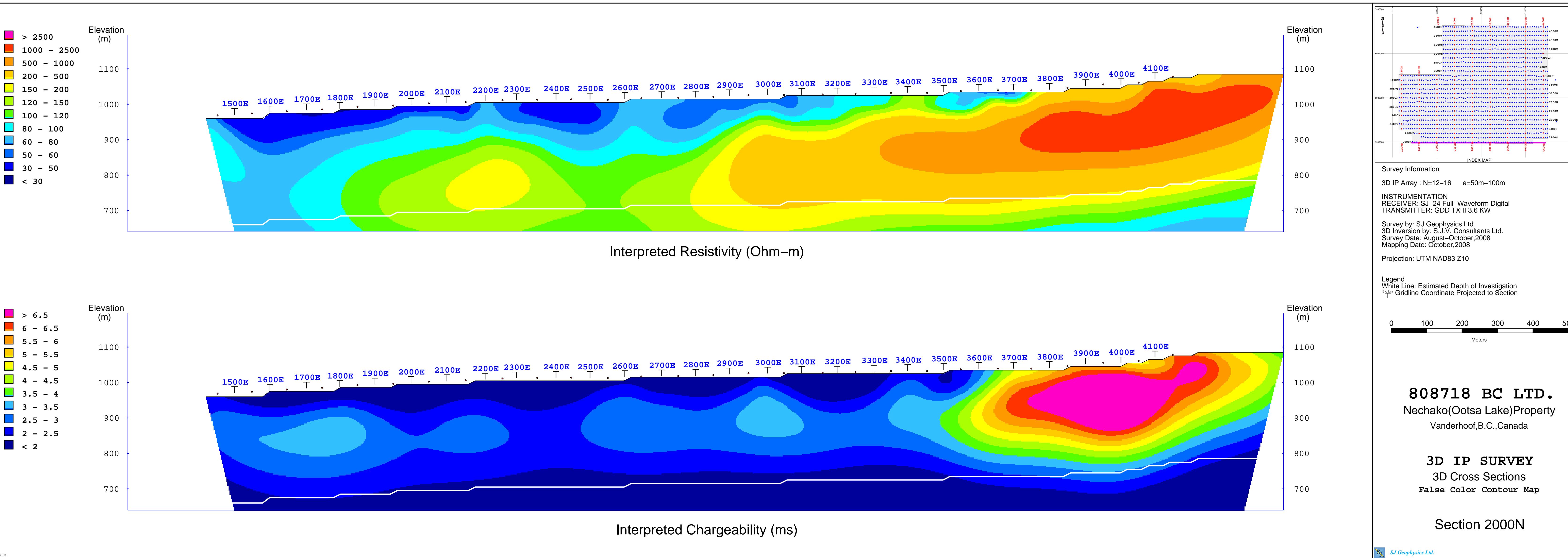


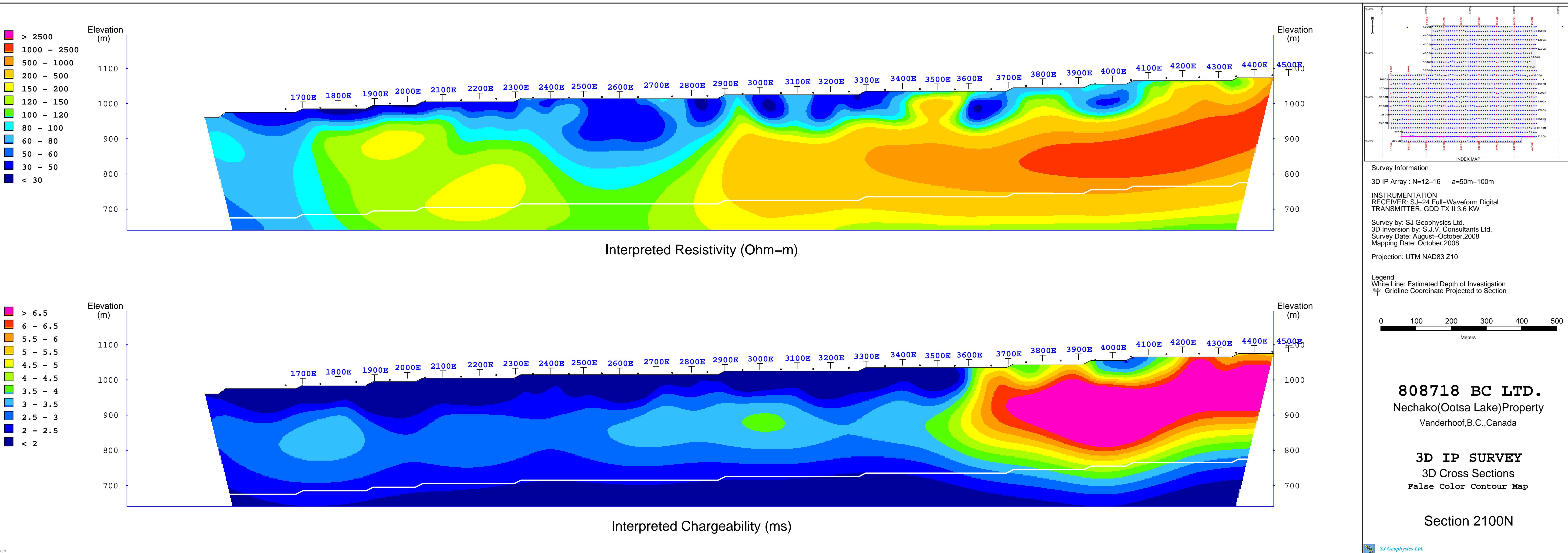


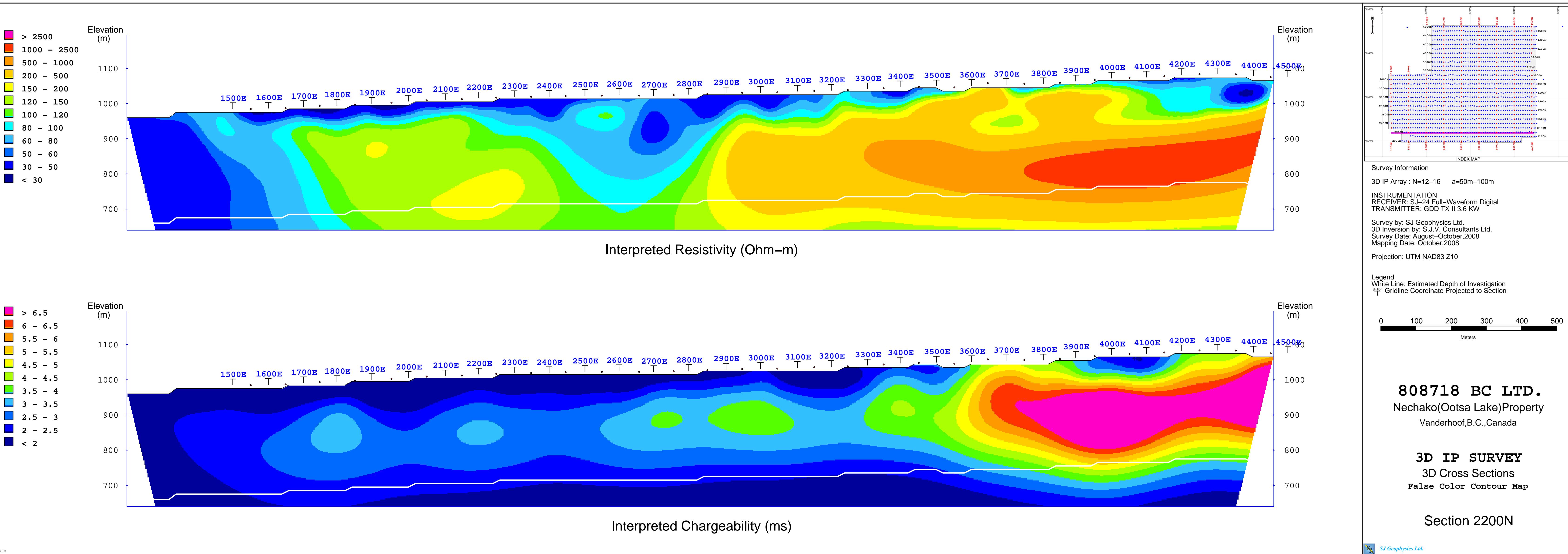


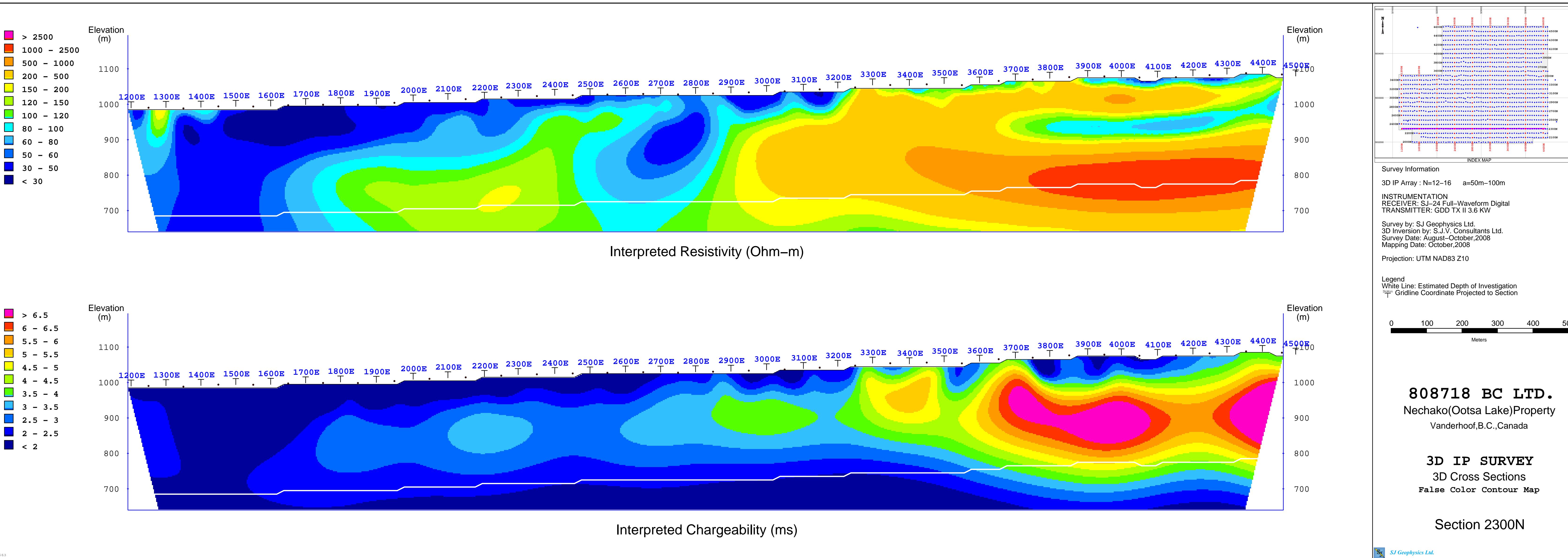


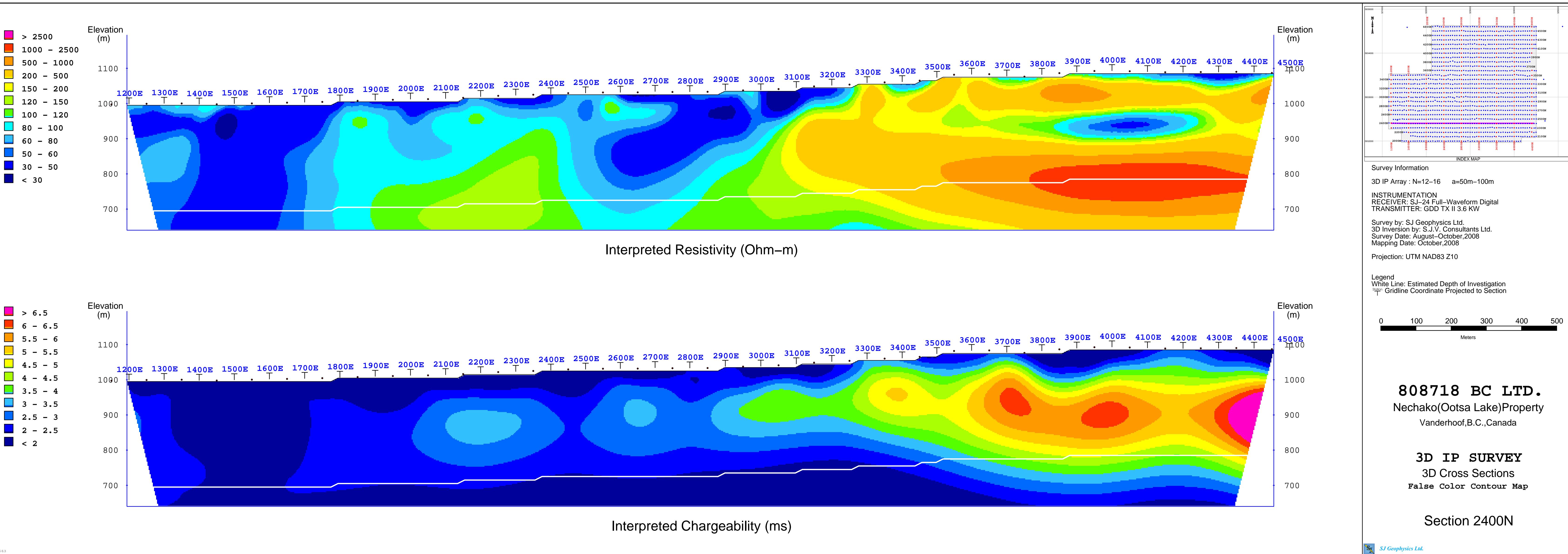


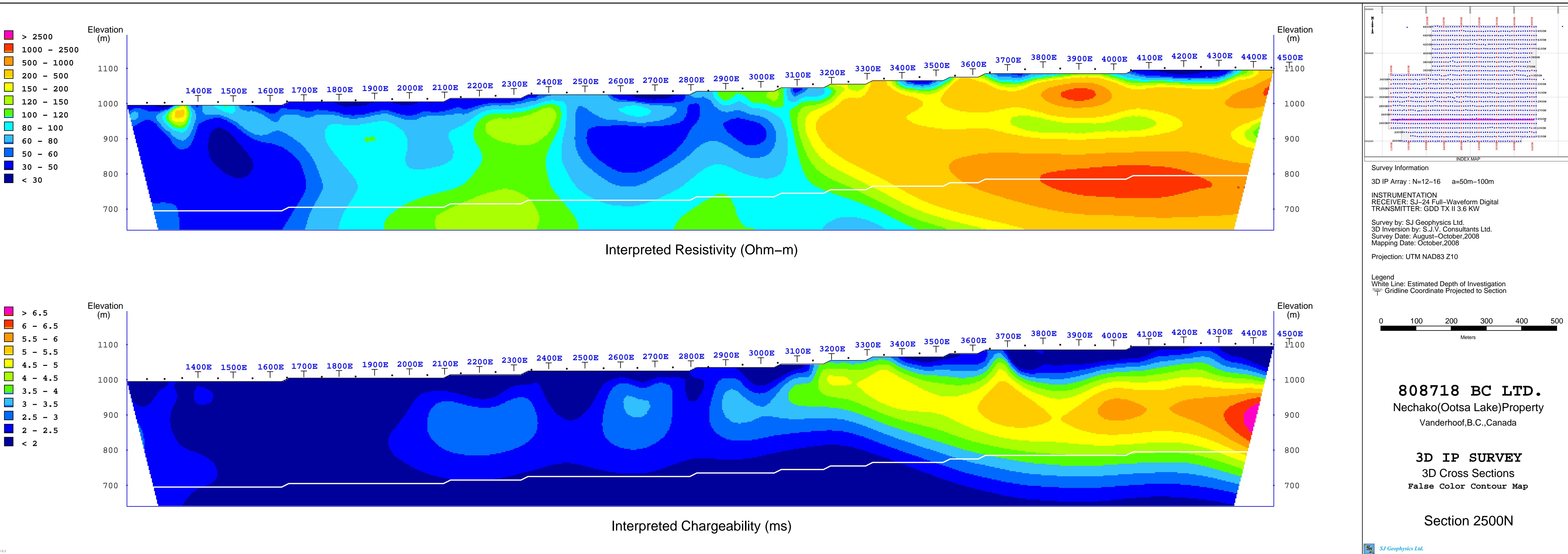


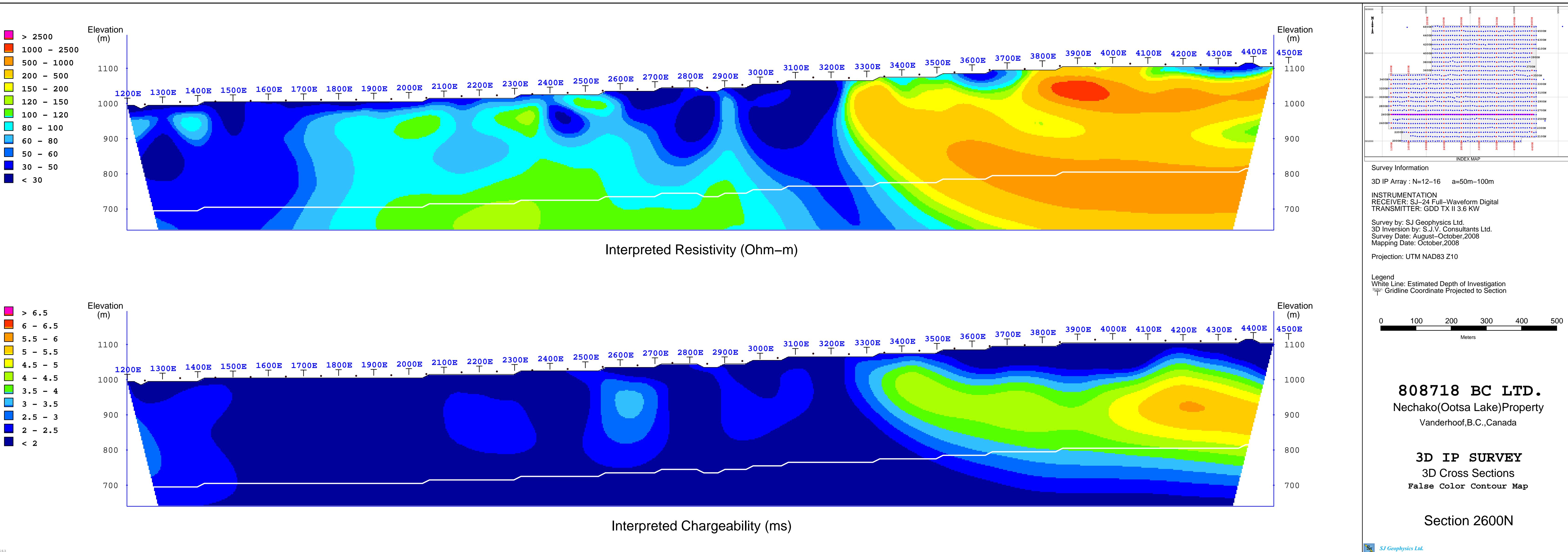


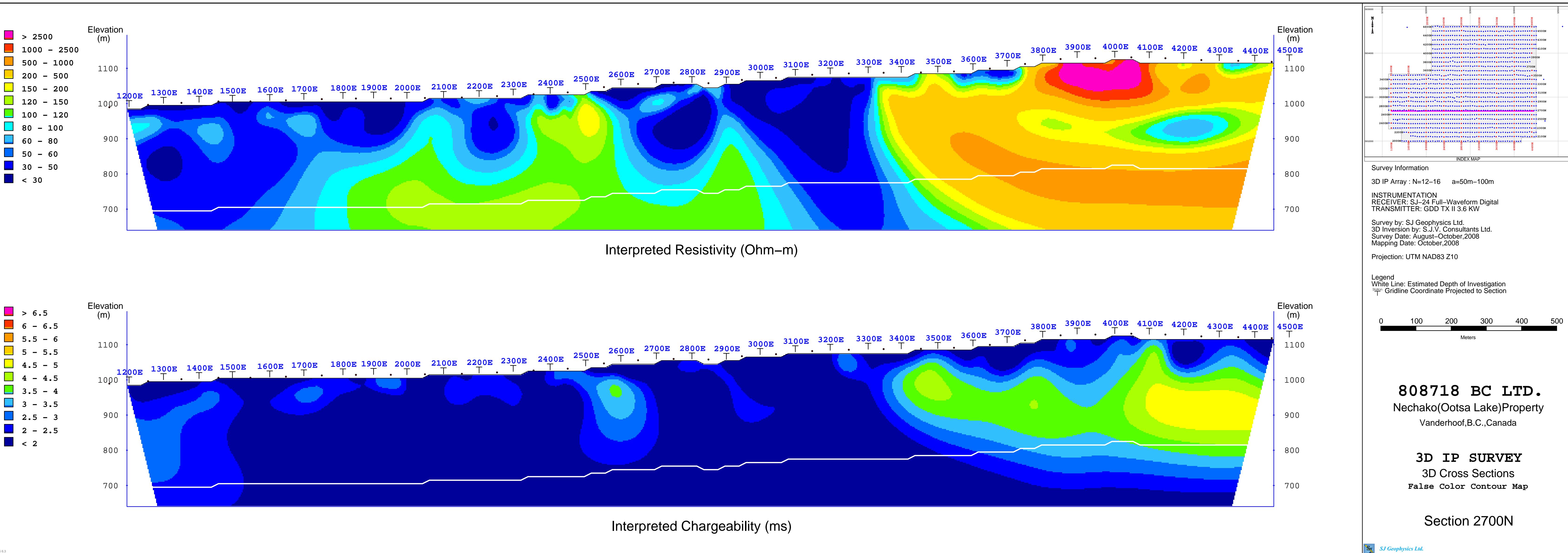


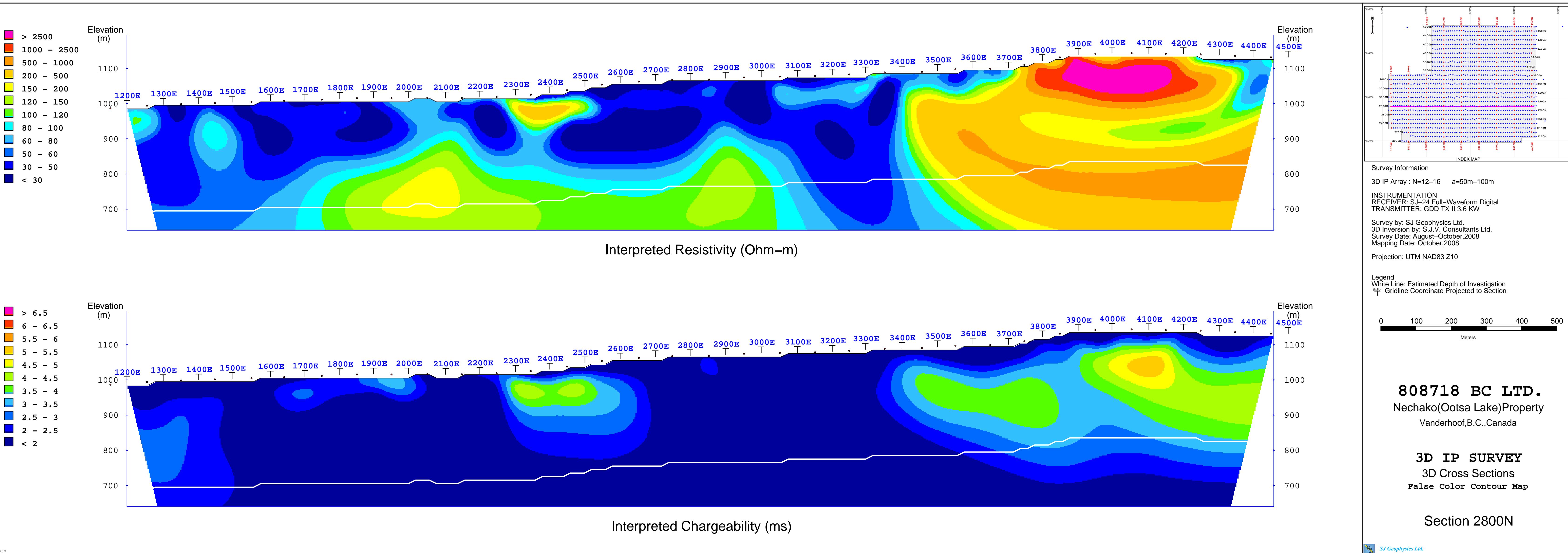


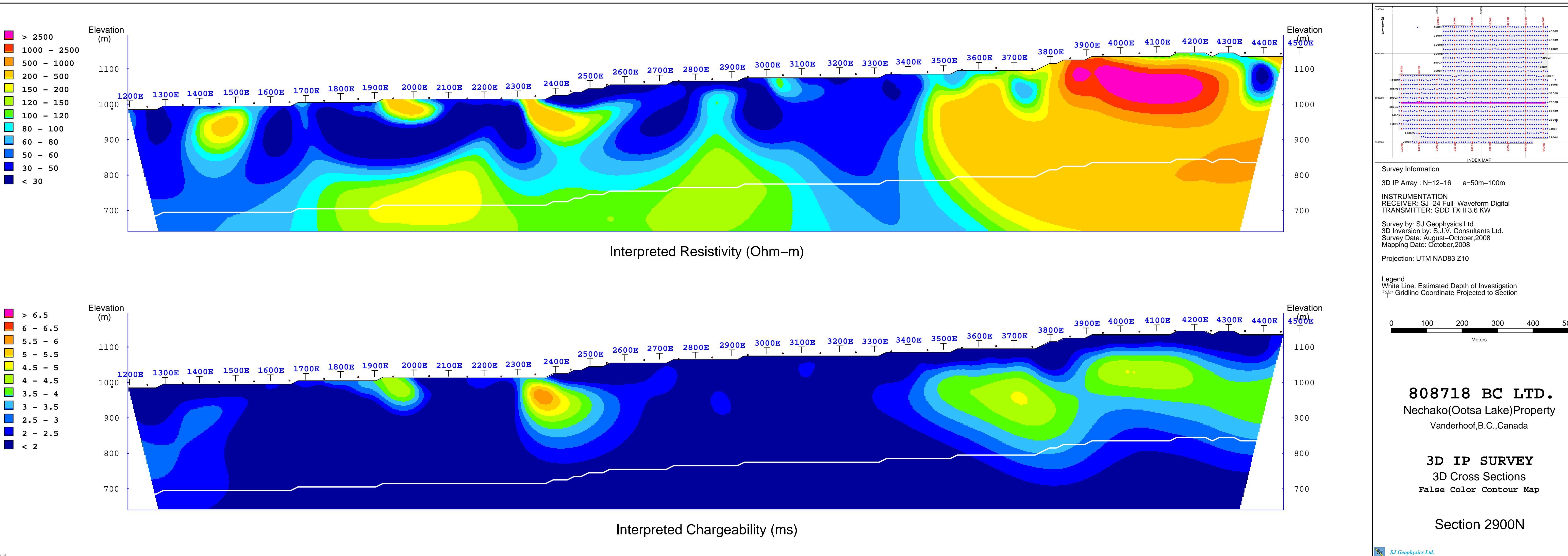


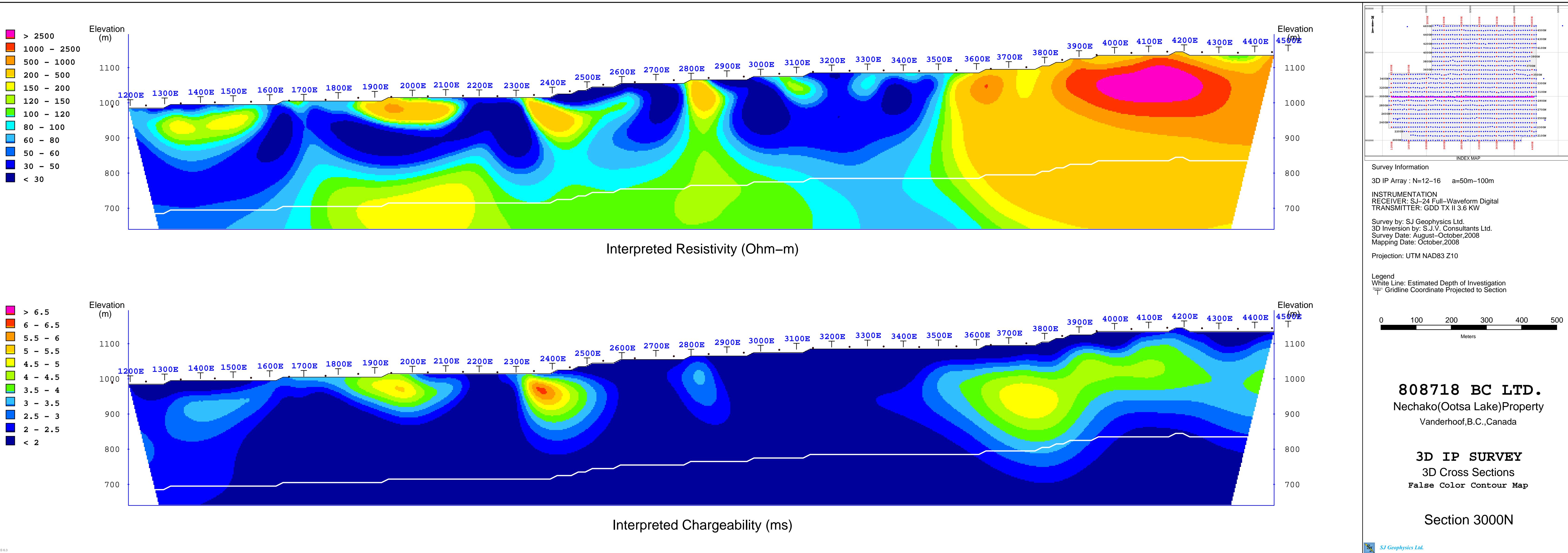


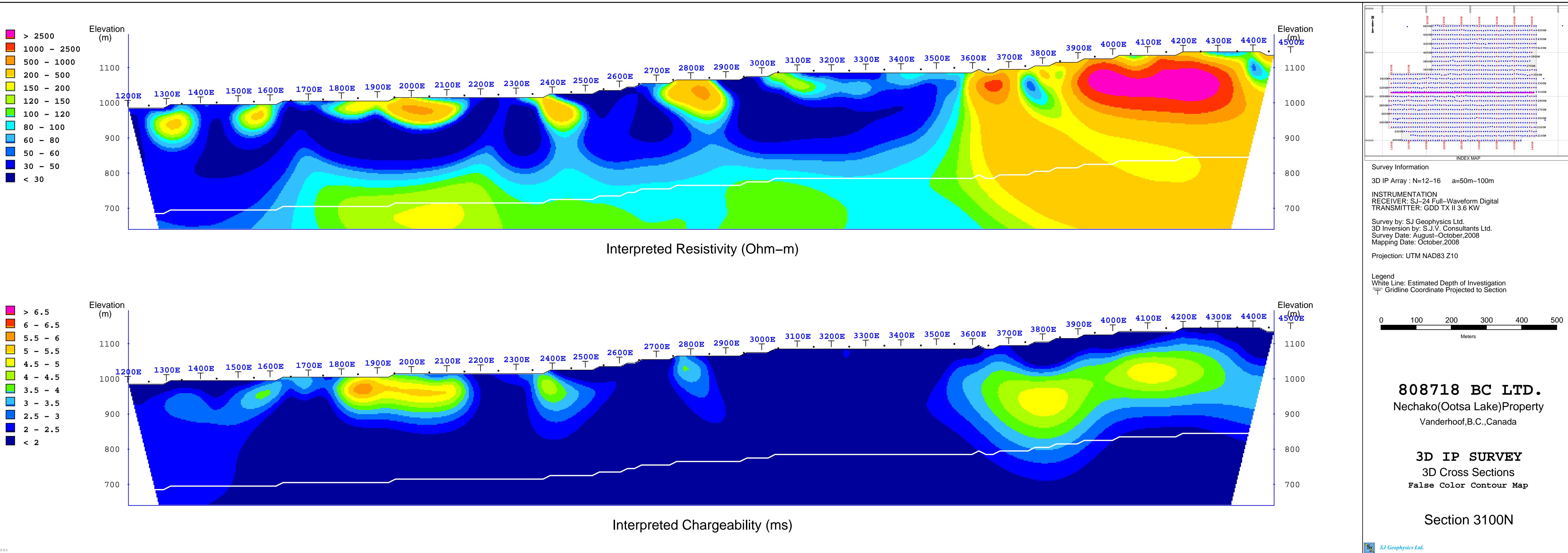


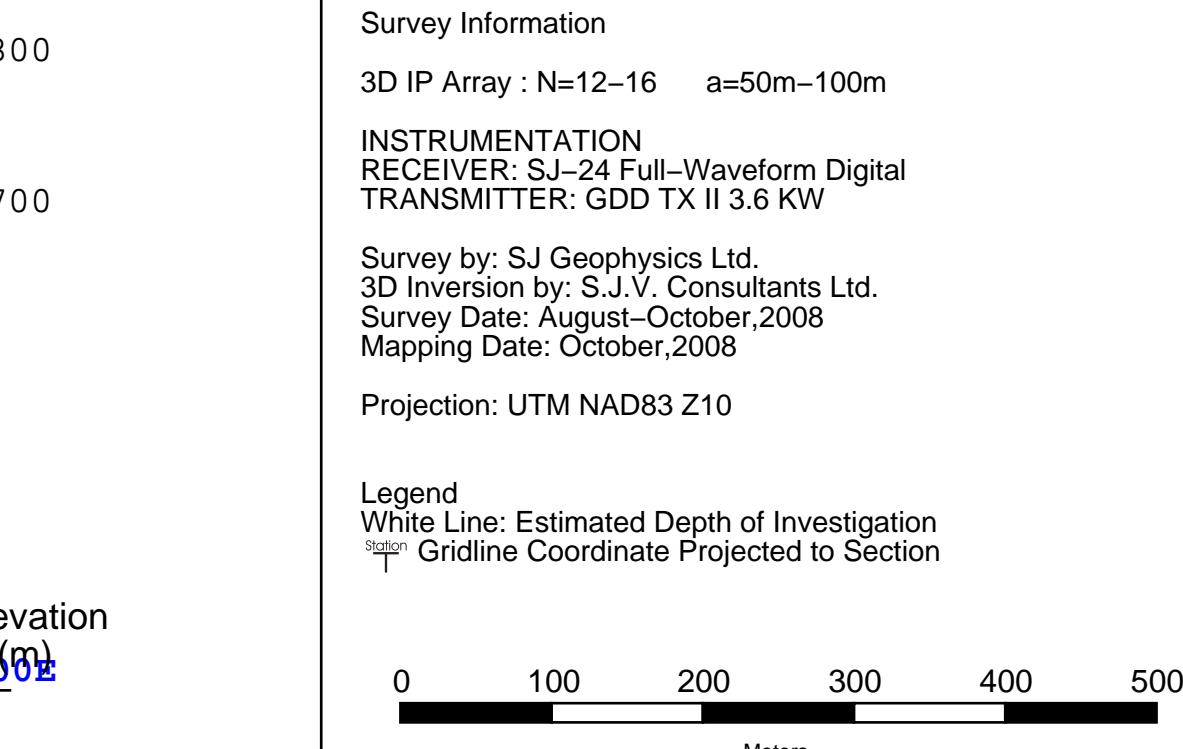
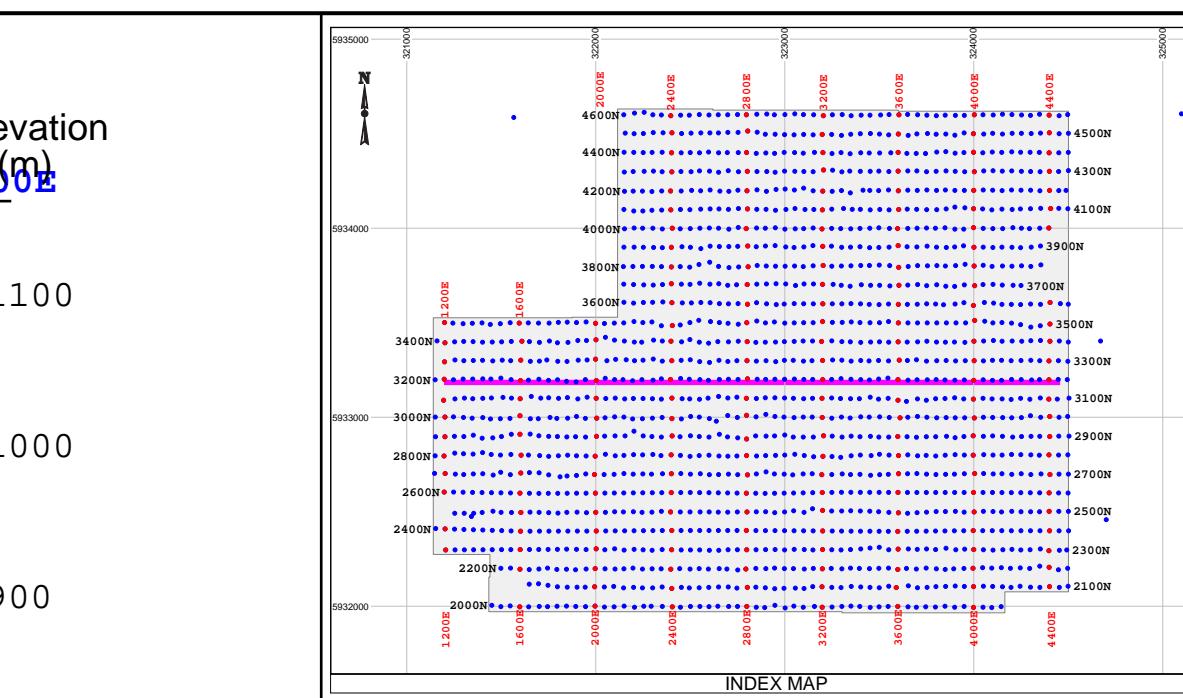
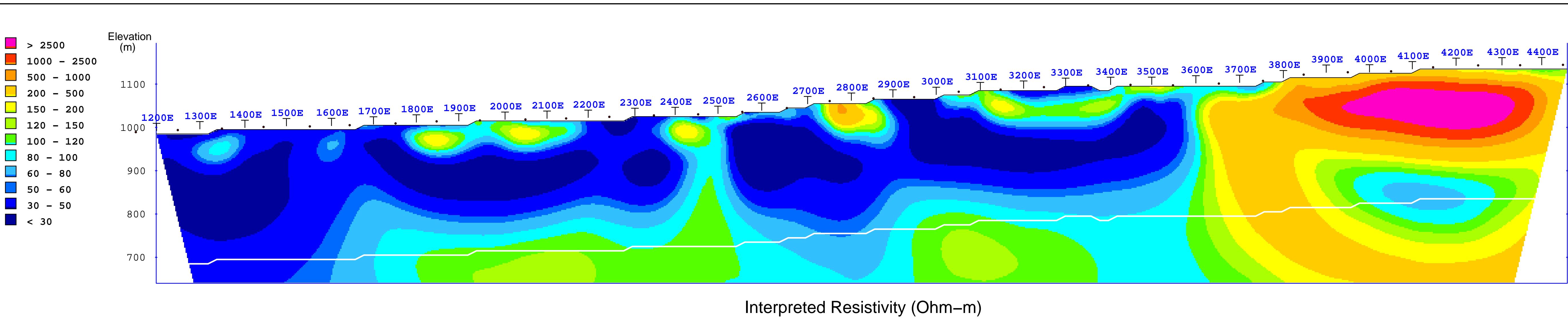












3D IP SURVEY
3D Cross Sections
False Color Contour Map

Section 3200N

