

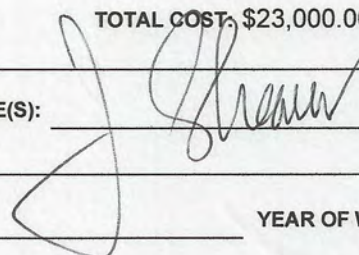
Ministry of Energy, Mines & Petroleum Resources
Mining & Minerals Division
BC Geological Survey

**Assessment Report
Title Page and Summary**

TYPE OF REPORT [type of survey(s)]: Geological Assessment
~~XXXXXXXX~~

TOTAL COST: \$23,000.00

AUTHOR(S): J. T. Shearer, M.Sc., P.Geo.

SIGNATURE(S): 

NOTICE OF WORK PERMIT NUMBER(S)/DATE(S): _____

YEAR OF WORK: 2014

STATEMENT OF WORK - CASH PAYMENTS EVENT NUMBER(S)/DATE(S): 5539864

PROPERTY NAME: Le Mare

CLAIM NAME(S) (on which the work was done): _____

COMMODITIES SOUGHT: Cu/Au

MINERAL INVENTORY MINFILE NUMBER(S), IF KNOWN: _____

MINING DIVISION: Nanaimo

NTS/BCGS: 92L/5 (92L.031)

LATITUDE: 50 ° 25 ' 06 " LONGITUDE: 127 ° 53 ' 10 " (at centre of work)

OWNER(S):

1) J. T. Shearer

2) _____

MAILING ADDRESS:

Unit 5 - 2330 Tyner Street

Port Coquitlam, BC

OPERATOR(S) [who paid for the work]:

1) _____

2) _____

MAILING ADDRESS:

PROPERTY GEOLOGY KEYWORDS (lithology, age, stratigraphy, structure, alteration, mineralization, size and attitude):

Mafic to felsic volcanics of the Jurassic Bonanza Group rocks underlie the Property which have been altered

to clay minerals and mineralized with copper minerals over a wide area, 0.2%Cu along one road over 180m

REFERENCES TO PREVIOUS ASSESSMENT WORK AND ASSESSMENT REPORT NUMBERS: _____

Assessment Reports 30608; 29686

TYPE OF WORK IN THIS REPORT	EXTENT OF WORK (IN METRIC UNITS)	ON WHICH CLAIMS	PROJECT COSTS APPORTIONED (incl. support)
GEOLOGICAL (scale, area)			
Ground, mapping	1:5,000	563795, 546563, 546689, 546543	\$23,000.00
Photo interpretation			
GEOPHYSICAL (line-kilometres)			
Ground			
Magnetic			
Electromagnetic			
Induced Polarization			
Radiometric			
Seismic			
Other			
Airborne			
GEOCHEMICAL (number of samples analysed for...)			
Soil			
Silt			
Rock			
Other			
DRILLING (total metres; number of holes, size)			
Core			
Non-core			
RELATED TECHNICAL			
Sampling/assaying			
Petrographic			
Mineralographic			
Metallurgic			
PROSPECTING (scale, area)			
PREPARATORY / PHYSICAL			
Line/grid (kilometres)			
Topographic/Photogrammetric (scale, area)			
Legal surveys (scale, area)			
Road, local access (kilometres)/trail			
Trench (metres)			
Underground dev. (metres)			
Other			
TOTAL COST:			\$ 23,000.00

GEOCHEMICAL ASSESSMENT REPORT

on the LE MARE COPPER-GOLD PROPERTY

Nanaimo Mining Division
N.T.S.: 92 L/5 (092L.031 and .041)
50° 25' 06"N., 127° 53' 10" W.
U.T.M.: 5585732 N., 579137 E.
Event # 5539864

for

Homegold Resources Ltd.
5-2330 Tyner Street,
Port Coquitlam, British Columbia, V3C 2Z1

by

J. T. Shearer; M.Sc., P.Geo. (BC & Ont.)

January 2, 2015

Fieldwork completed between January 6, 2014 and January 1, 2015

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SUMMARY

The Le Mare property comprises 17 map-staked claims covering 6,198.061 hectares (15,309.21 acres) in the Nanaimo Mining Division and in the Rupert Land District of western British Columbia (Figures 1 and 3). It is located on N.T.S. map sheet 92 L/5, as well as on B.C. map sheets: 092L 031 and 041. All claims comprising the property are in good standing until January 5, 2012. (additional Assessment credits are available to extend the expiry date for several years.)

The mineral claims comprising the Le Mare property are owned by J.T. Shearer; M.Sc., P.Geo. On October 17, 2009, J.T. Shearer and New Destiny Mining Corp. entered into an agreement whereby New Destiny could obtain 100% interest in and to the claims, and to all mineral rights secured by those claims with the exception of "industrial minerals in particular but not limited to chalky geysirite and geysirite". New Destiny's potential interest is subject to a 2% net smelter return royalty payable to J.T. Shearer, of which New Destiny may purchase 50% at any time for \$250,000. The option is exercisable upon payment of a total of \$55,000 (\$15,000 paid) and completion of a total \$500,000 worth of work on the Le Mare property (more than \$300,000 completed which includes the current program).

The Le Mare hydrothermal system is located on crown land in the southwestern part of the property area. The Mah-te-nicht No. 8 Indian Reserve is located adjacent with the northeastern property boundary, about 4.5 km (2.75 mi) north-northeast of, and in a different drainage from the Le Mare hydrothermal system. However, if ocean-going barge loading facilities were to be developed on the south shore of Quatsino Sound, the Quatsino Band would become involved in the design and construction of those facilities. J.T. Shearer has been consulting with the Quatsino Band Council since February, 2007 with regard to exploration of the Le Mare property. There is no plant or equipment, inventory, mine or mill structure on these claims. Currently, an environmental bond of \$4,000 is posted under Permit No. MX-8-253 for road renovation, the development of potential drill sites and diamond drilling.

The Le Mare property is located near the northwestern end of Vancouver Island. It is bounded in part to the west by the Pacific Ocean and to the north by Quatsino Sound. A massif in the northwestern part of the property culminates in the peak of Mount Bury at an elevation of about 610 m (2,000 ft). Another massif that hosts the Le Mare hydrothermal system occupies the property's southwestern part. Le Mare Peak is a 762-m (2,500-ft) high promontory located near the massif's centre. These steep-sided massifs are separated by the relatively flat Mahatta and Culleet creek valleys. The surface of Le Mare Lake, located in the Culleet Creek valley near the property centre, is at an elevation of about 25 m (82 ft).

About 85% of the original west-coast rain forest in the property-area has been clear-cut during the past 40 years. Most of the slopes underlain by the Le Mare hydrothermal system are either bare, or covered with dense juvenile secondary forest growth. Little timber suitable for mining is left on the property.

The northern end of Vancouver Island is accessible by boat, barge, and by road via the Island Highway (B.C. Highway 19) which transects the town of Port McNeill on the island's northeastern coast. B.C. Highway 25, a secondary paved road, connects Port McNeill with Port Alice located near the head of Neroutsos Inlet. Access from Port Alice to the Le Mare property area is via a series of well-maintained logging roads passable by 2-wheel drive vehicles during all times of the year. Most of the property-area is covered by a system of logging roads in various states of repair. Barge and ship loading facilities to support a large scale mine could be developed on the sheltered southern shore of Quatsino Sound near the property's northern boundary.

Port McNeill and Port Alice are the nearest towns with sufficient supply and service capacity to support an exploration or drilling program. The industrialized areas of southwestern British Columbia are readily accessible via water, road, and air from Port McNeill and Port Hardy. Accommodations and basic supplies for an exploration field crew are available at Port Alice and Mahatta Camp, located 8km east of the claims.

The Quatsino Sound area experiences cool wet winters and cool, moderately wet summers. Snow falls in the property-area by December and stays on the ground very briefly at higher elevations. The current exploration target, the Le Mare hydrothermal system, is on crown land with no special restrictions on development thereon.

The Le Mare property is west of the major electrical power source at Port Alice power transmission line. Ocean-going barge transport to the property area would reduce the cost of fuel and supplies. Creeks south and east of the property area could be dammed in order to generate power for a mine-mill complex. Water for milling could be drawn from Culleet or Gooding creeks, or from the outflow from a nearby generating station. An acceptable mill site and tailings storage areas could be constructed in the floors of the Gooding Creek and upper Culleet Creek valleys.

The Port McNeill-Port Hardy area has already demonstrated that it was able to attract personnel to work at the Island Copper mine located between the two towns between 1970 to 1996. That area has sufficient amenities to attract the people needed to operate a new mine near to it.

The Le Mare property hosts mostly mafic volcanic rocks of the Early to Middle Jurassic-age Bonanza Supergroup, including auto-breccias, lahars, and minor amounts of tuff and other pyroclastic beds. Rhyolitic rocks comprise a major amount of the stratigraphy in the property-area. These volcanic rocks are intruded by felsic dykes that may be equivalent to the rhyodacitic porphyries that are associated with mineralization at the Island Copper Cluster deposits located about 32 km (19.3mi) east-northeast of the Le Mare hydrothermal system. The volcanic rocks at the Le Mare hydrothermal system have deformed into a series of open to close outcrop-scale drape-folds related to local intrusion. Regional and contact metamorphism do not exceed lower the greenschist facies.

The Le Mare hydrothermal system appears to have been only relatively shallow unroofing by erosion. The top of the potassic alteration zone is exposed along the crests of Le Mare and Gooding ridges, located between Le Mare Lake and Gooding Cove in the southwestern part of the property. Local magnetic field gradient indicates that this system occupies a 5 X 3 km (3.05 X 1.83 mi) or 15km² (5.6 mi²) oval-shaped area, that may be hosted by a dilational jog in a regional right-lateral fault system. The proposed fault system is similar to the one that hosts the Island Copper Cluster deposits near Port McNeill and Port Hardy, British Columbia.

At surface, copper mineralization occurs in discrete showings-areas, located preferentially in the central parts of sub-vertical hydrothermal plumes. These plumes have core-zones of orthoclase-quartz-biotite (potassic) alteration, enveloped in siliceous exteriors. Orthoclase-quartz-biotite alteration is succeeded by quartz-jasper alteration; both phases are mineralized with chalcopyrite, and minor amounts of bornite. This potassic alteration is accompanied by co-incident soil-copper and magnetic anomalies. Discovering economically viable concentrations of copper mineralization within the Le Mare hydrothermal system depends on the successful identification of zones where these hydrothermal plumes and copper occurrences coincide.

Molybdenum enrichment occurs in areas flanking phyllic alteration in a 600-m (1,968.5-ft) diameter alteration plume, covering a 0.28 km² (0.1 mi²) area in the eastern part of system in the South Gossan zone. Another, much less extensive plume of argillic-phylic alteration, is exposed between the Culleet Creek zone and Culleet Lake in the system's northwestern part. These two plumes cover less than 2% of the total exposure area of the Le Mare hydrothermal system. Argillic-phylic alteration post-dates and overprints potassic alteration.

Both sample results and the distribution of soil-copper and molybdenum anomalies, demonstrate that copper and molybdenum mineralization are associated with early potassic and subsequent argillic-phylic-veinalteration events respectively. They occur together in significant amounts only where molybdenum enrichment has overprinted that of copper.

Highly anomalous gold values were discovered in the central part of the Le Mare hydrothermal system mostly west and southwest of the New Destiny Showing in soil samples. Values range up to 947ppb gold on Claim 657343. The New Destiny showing was trenched with a tracked excavator and returned >0.2% copper over 200 metres (Figure 20).

Most aspects of the Le Mare hydrothermal system are similar with those of the Island Copper Cluster deposits. Geology, alteration, and mineralization at surface at the Le Mare hydrothermal system correspond with those attributes at the Island Copper mine above the main deposit. These similarities indicate that the Le Mare hydrothermal system may host a calc-alkalic porphyry copper-molybdenum deposit of the Island Copper Cluster type.

The Early Jurassic-age land surface above the Le Mare hydrothermal system and whatever near surface hot-spring environment that it may have hosted, has been lost to erosion. Only a few narrow fault controlled, advanced argillic alteration occur in the argillic-phylic alteration plume in the South Gossan zone. They attest to the former existence of acid leaching with the alteration system.

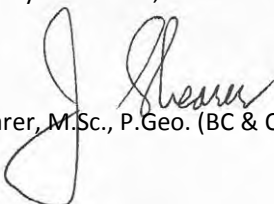
Most exploration has been conducted in the northeastern part of the Le Mare hydrothermal system; its southeastern part remains sparsely explored to unexplored. Six BQ diamond drill holes penetrated the northeastern margin of the Le Mare system in 1992. One hole that penetrated the Culleet Creek potassic alteration plume, intersected five 2-m (6.56-ft) and one 4.7-m (15.42-ft) long intersections containing from 500 to 959 ppm copper, which is similar to the tenor of copper mineralization in nearby trenches. Copper mineralization at surface is locally quite variable. Such variability should be expected in mineralization located near the top of the potassic alteration zone of a porphyry copper-molybdenum deposit. Less than 1% of the surface area of the Le Mare hydrothermal system has been drilled.

Trenching in 2011, followed by 3m wide chip sampling on the New Destiny Copper Showing (discovered in late 2009) returned a 180m continuous copper values averaging 0.25% Copper.

A second-phase exploration program is recommended. The first phase comprises geological mapping, prospecting, soil survey. If reasonable encouragement is generated by the results of the first-phase program, it should be followed by a second-phase program of 3-dimensional induced polarization and ground-magnetic surveys. The estimated costs of the second recommended phase of exploration is as follows:

Program	Estimated Cost inc. H.S.T. + Contingency
A 1st Phase: geological mapping, prospecting, and soil survey was completed between Feb. 14 and May 15, 2011 As documented in this report: It is recommended that a 2nd Phase program of 3-dimensional induced polarization and ground magnetic surveys be completed at an estimated cost of:	\$ 598,074

Respectfully submitted,



J. T. Shearer, M.Sc., P.Geo. (BC & Ont.)

INTRODUCTION

The author, J. T. Shearer; M.Sc., P.Ge. (BC & Ontario) was commissioned by Homegold Resources Ltd. to complete an Assessment Work program on the Le Mare property.

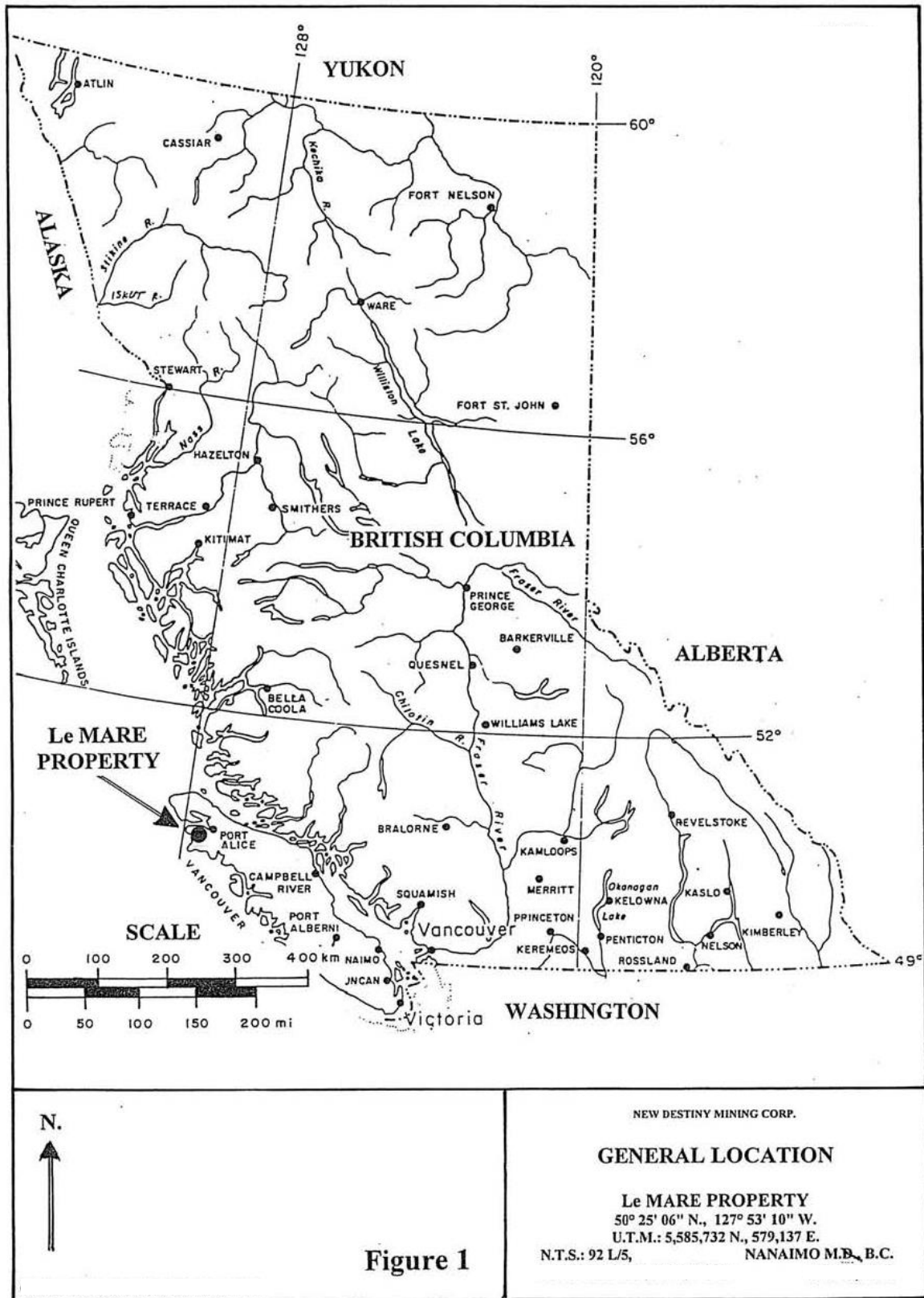
This report was written to document the results of the Phase I Program and recommend a suitable follow-up work.

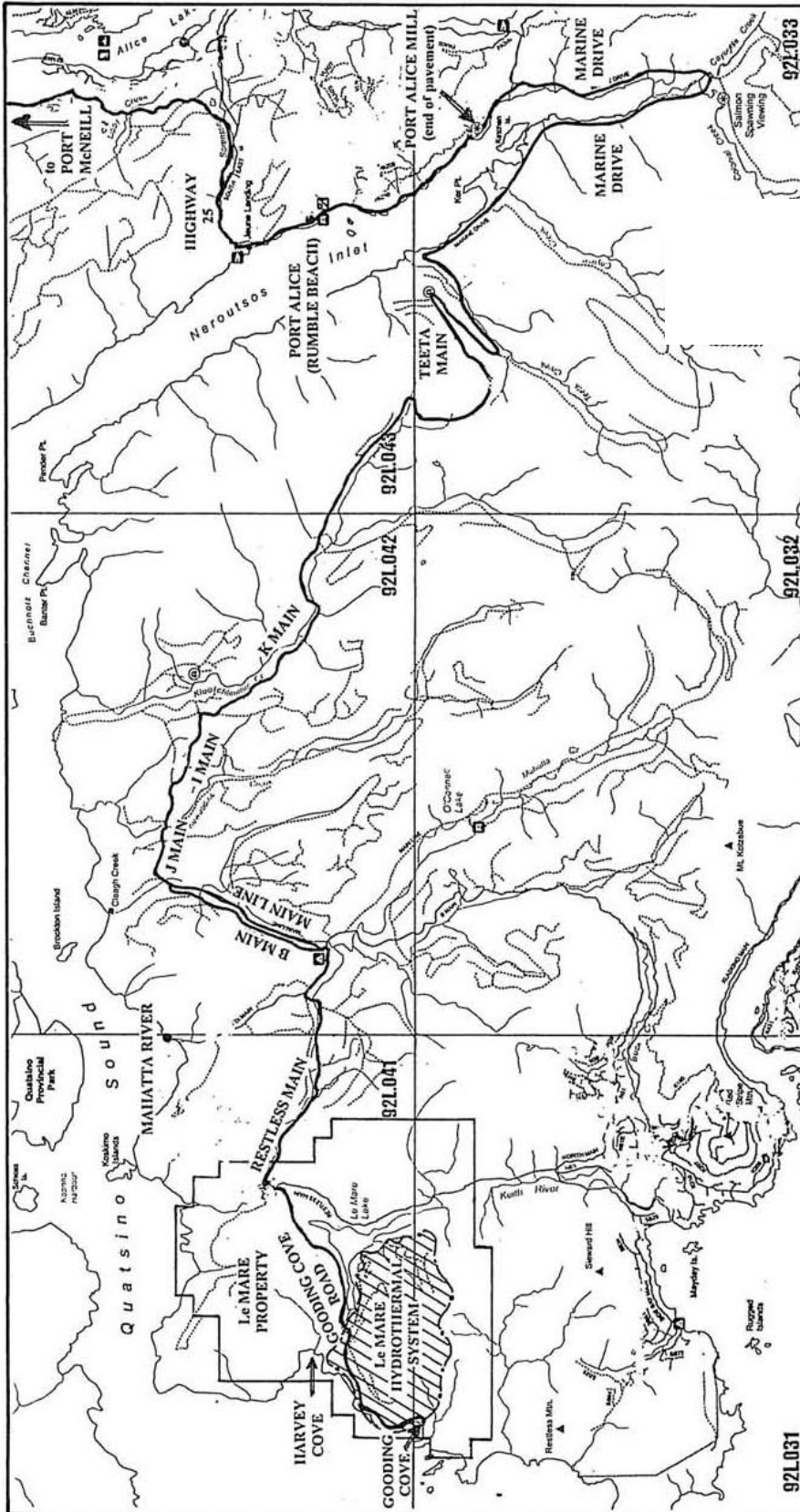
This report is based upon published records of the results of previous exploration in the Le Mare property-area, of property examinations and regional geological mapping conducted by geologists of the British Columbia Geological Survey and of the Geological Survey of Canada, the results of the 2009, and recent (2011), exploration programs. Citations of that work are in standard format (section 10.0, this report). The current (2014) exploration program of geological mapping, prospecting and examination of workings was conducted or supervised by J.T. Shearer; M.Sc., P.Ge., the property owner.

PROPERTY DESCRIPTION AND LOCATION

The Le Mare property comprises 17 map-staked claims covering 6,198.061 hectares (15,309.21 acres) in the Nanaimo Mining Division and in the Rupert Land District of western British Columbia (Figures 1 and 3). It is located on N.T.S. map sheet 92 L/5 as well as on B.C. map sheets: 092L 031 and 041.

The mineral claims comprising the Le Mare property are owned by J.T. Shearer; M.Sc., P.Geo. (Table 1).

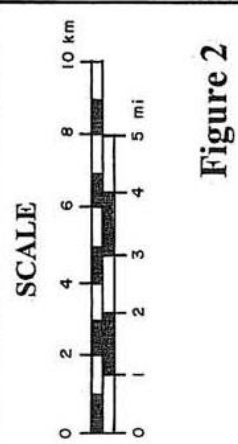




NEW DESTINY MINING CORP.

REGIONAL ACCESS

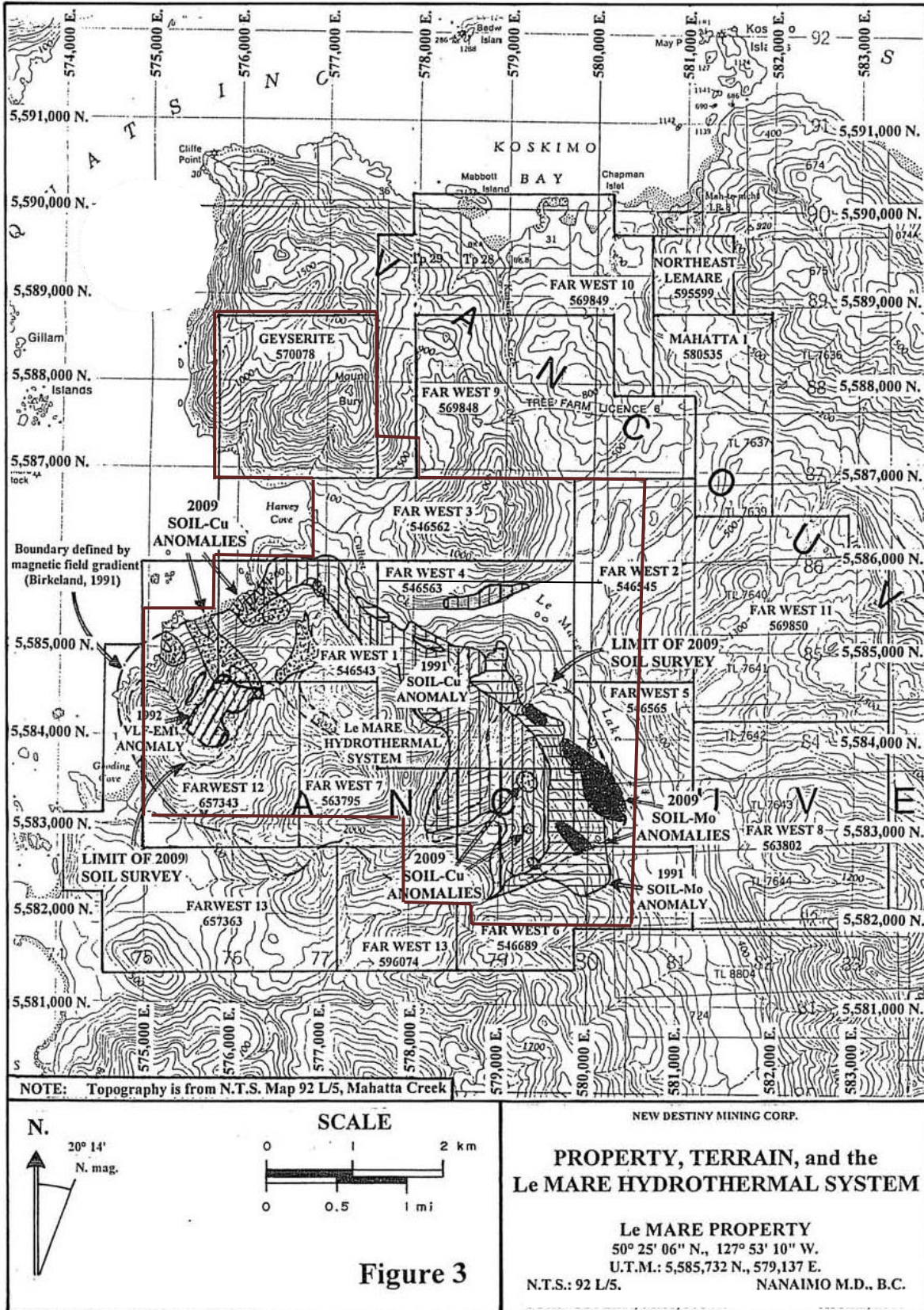
Le MARE PROPERTY
 50° 25' 06" N., 127° 53' 10" W.
 U.T.M.: 5,585,732 N., 579,137 E.
 N.T.S.: 92 L/5,
 NANAIMO M.D., B.C.



SCALE

NOTE: The base map for this Figure was produced by West Fraser Timber Co. Ltd.

Figure 2



CLAIM STATUS

The tenures of the eleven (11) claims comprising the Le Mare property (Figure 3) are as follow:

Table 1
Map-staked Claims Comprising the Le Mare Property

Tenure Number	Claim Name	Area: hectares	Record Date	Expiry Date	Owner
546543	Far West 1	247.09	December 4, 2006	February 5, 2017	J.T. Shearer
546562	Far West 3	185.29	December 5, 2006	February 5, 2017	J.T. Shearer
546545	Far West 2	205.90	December 4, 2006	February 5, 2017	J.T. Shearer
546563	Far West 4	514.83	December 5, 2006	February 5, 2017	J.T. Shearer
546689	Far West 6	391.44	December 6, 2006	February 5, 2017	J.T. Shearer
563795	Far West 7	247.18	July 29, 2007	February 5, 2017	J.T. Shearer
569849	Far West 10	20.58	November 10, 2007	February 5, 2017	J.T. Shearer
567343	Far West 12	453.10	October 22, 2009	February 5, 2017	J.T. Shearer
546565	Far West 5	164.78	December 5, 2006	February 5, 2017	J.T. Shearer
596074	Far West 13	41.20	December 14, 2008	February 5, 2017	J.T. Shearer
570078	Geyselite	123.50	November 14, 2007	February 5, 2017	J.T. Shearer

Total 2,594.89 ha

Cash may be paid in lieu if no work is performed. Following revisions to the Mineral Tenures Act on July 1, 2012, claims bear the burden of \$5 per hectare for the initial two years, \$10 per hectare for year three and four, \$15 per hectare for year five and six and \$20 per hectare each year thereafter.

These are map-staked claims that are located on the computer-generated provincial mineral tenure grid (www.mtonline.bc.ca). No posts or lines exist on the ground; thus, there is no uncertainty regarding the area covered by the claims.

Exploration damage bonds are required if exploration programs such as, line cutting for grid establishment, road building, trenching, and drilling that result in significant surficial disturbance are conducted. Currently, a bond of \$4,000 is posted under Permit No. MX-8-253 for road renovation and the development of potential drill sites. It is anticipated that the cost of bonds for the second phases of the recommended program will be about \$10,000 respectively. Applications for revisions to permit No. MX-8-253 for those phases of work have not been made yet.

ACCESSIBILITY

The Le Mare property is located near the northwestern end of Vancouver Island. It is bounded in part to the west by the Pacific Ocean and to the north by Quatsino Sound. A massif in the northwestern part of the property culminates in the peak of Mount Bury at an elevation of about 610 m (2,000 ft). Another massif that hosts the Le Mare hydrothermal system occupies the property's southwestern part. Le Mare Peak is a 762-m (2,500-ft) high promontory located near the massif's centre. These steep-sided massifs are separated by the relatively flat Mahatta and Culleet Creek valleys. The surface of Le Mare Lake, located in the Culleet Creek valley near the property centre, is at an elevation of about 25 m (82 ft) (Figure 3).

About 70% of the original west-coast rain forest in the property-area has been clear-cut during the past 20 years. Most of the slopes underlain by the Le Mare hydrothermal system are either bare or covered with dense juvenile secondary forest growth. Little timber suitable for mining is left on the property.

The northern end of Vancouver Island is accessible by boat, barge, and by road via the Island Highway (B.C. Highway 19) which transects the town of Port McNeill on the island's northeastern coast. B.C. Highway 25, a secondary paved road, connects Port McNeill with Port Alice located near the head of Neroutsos Inlet (Figure 2). Access from Port Alice to the Le Mare property area is via: Marine Drive, Teeta Main, K Main, I Main, J Main, B Main, and Restless Main roads. These logging roads are well-maintained and passable by 2-wheel drive vehicles during drier times during the year. The trip takes from 1.5 to 2 hours depending on road conditions. Most of the property-area is covered by a system of logging roads in various states of repair. Barge loading facilities to support an open-pit mine could be developed on the sheltered southern shore of Quatsino Sound near the property's northern boundary.

Port McNeill is the nearest town with sufficient supply and service capacity to support an exploration or drilling program. Accommodations and basic supplies to support an exploration field crew are available at Port Alice and Winter Harbour, located northwest of Quatsino Sound. During the 2009 exploration program, the crew stayed in the camp at Mahatta River (Figure 2).

The Quatsino Sound area experiences cool wet winters and cool, moderately wet summers. Snow falls in the property-area by December and stays on the ground very briefly at higher elevations.

The current exploration target, the Le Mare hydrothermal system, is on crown land with no special restrictions on development thereon (Figure 3). Upon development permitting, one normally is able to secure surface rights necessary to conduct a permitted mining operation. The writer knows of no legal impediment to New Destiny Mining Corp. being able to secure such surface rights as part of the permitting process.

Creeks south and east of the property area could be dammed in order to generate power for a mine-mill complex. Water for milling could be drawn from Culleet or Gooding creeks, or from the outflow from a nearby generating station. An acceptable mill site and tailings storage areas could be constructed in the floors of the Gooding Creek and upper Culleet Creek valleys (Figure 3).

Both the mining business and the pool of professionals and skilled tradesmen who serve it are international and mobile. The Port McNeill-Port Hardy area has already demonstrated that it was able to attract personnel to work at the Island Copper mine located between the two towns. That area has sufficient amenities to attract the people needed to operate a new mine near to it.

HISTORY

Chronology of Exploration of Claims in the Le Mare Property-area from 1979 to Present

A.O. Birkeland (1991) briefly described the pre-1979 exploration throughout the area covered by the current Le Mare property as follows:

During the late 1960s and early 1970s, exploration for porphyry Cu-Mo-Au deposits similar to the Island Copper Mine operated by BHP Utah was conducted by several companies on the western portion of Vancouver Island. The earliest reference to claim staking activity in the LeMare area was during 1970 when the Cam claims were recorded along the north shore of LeMare Lake. No assessment work was filed at that time.

Birkeland, A.O.; 1991: p. 4.

- 1979: The Le Mare 1 (477) and Le Mare 2 (496) claims comprising 4 units each were staked along the northwestern shore of Le Mare Lake and along the shore road southwest of Harvey cove respectively (Figure 4). The claims were recorded on November 9 and 13, 1979.
- 1980: D.G. Leighton and Associates Ltd. conducted a prospecting program on the Le Mare claims for British Newfoundland Exploration Ltd. (BRINCO) (Bilquist, 1980) (Figure 4). A two-man crew spent four days prospecting road exposures, taking a total of 28 rock samples. Finely disseminated vein pyrite with sporadic chalcopyrite, bornite, and malachite were found in roadside exposures of felsic volcanic rocks along the northwestern shore of Le Mare Lake (Figure 4) on the Le Mare 1 (477) claim. Chip samples from the Le Mare Lake section contained from 0.13 to 0.14% copper. Grab samples contained up to 0.49% copper. Secondary potassium feldspar was noted. On the Le Mare 2 (496) claim, andesitic flows and dacitic pyroclastic rocks along the road southwest of Harvey Cove was found to contain fracture-related pyrite, chalcopyrite, azurite, and sphalerite. Samples from there contained from 0.2 to 1.4% copper (Figures 4 and 18). BRINCO was unimpressed with the find and financed no further work on it.
- 1981 to 1990:
No exploration work from the Le Mare property-area has been recorded.
- 1991: Research by Keewatin Engineering Inc. during March, 1991, revealed that a belt similar to the Island Copper Belt was located between Kyuquot Sound and Quatsino Sound. It was named the Mahatta-Kashutl belt. Upon findings from re-manipulation of regional aeromagnetic data (section 3.2.2, this report) (Figure 10), and a field examination of the Le Mare Lake area, the 216-unit LeMare property was staked by Keewatin (Birkeland, 1991). The May, 1991 LeMare property was owned by Stow Resources Ltd. of Vancouver, B.C. It covered an area similar to that covered by the current Le Mare property. Moss-mat and stream-sediment sampling conducted over the whole current Le Mare property-area resulted in definition of a primary target that extended for 6 km (3.7 mi) southeastward from Harvey Cove to east of Le Mare Lake (Figure 4). Subsequently, geological mapping, and soil sampling was conducted along the logging roads on the slopes southwest of Le Mare Lake. Geological and alteration mapping was conducted over a total area of 2.44 km² (0.91 mi²) (sections 3.4, 4.2, and 5.3.2, this report) (Figures 4, 14, 17, and 21E to 22W).

Soil samples were collected at mostly 25-m (82-ft) intervals along the roads. A total of: 136 moss mat and silt, 855 soil, and 316 rock samples were collected during the 1991 program. Birkeland (1991) defined anomalous thresholds from the second positive standard deviation levels in the distributions of 1991 soil-metal concentrations and lowered them somewhat to make them more representative of hypothetical regional sampling as follows:

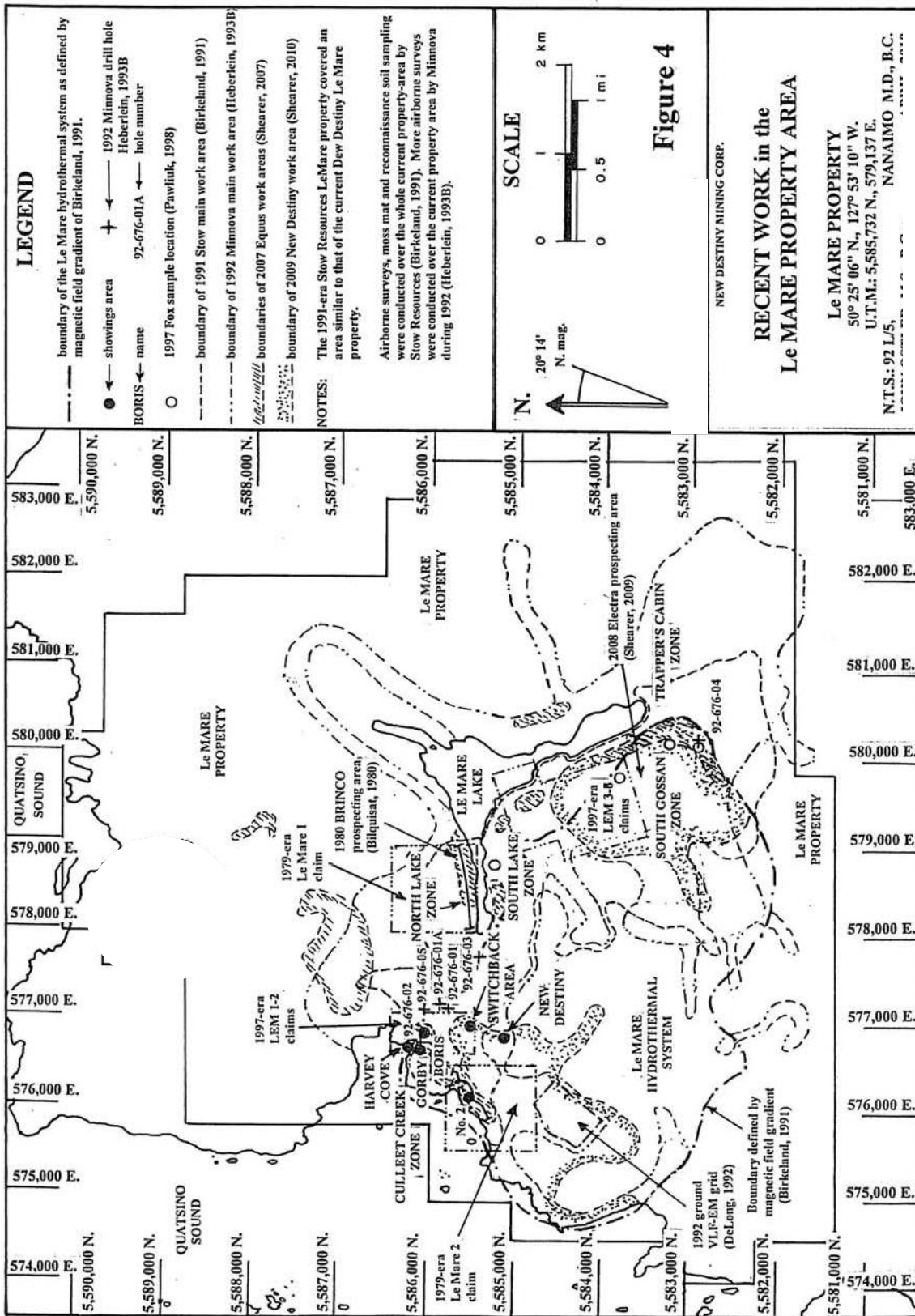


FIGURE 4

Table 2
Birkeland's 1991 Soil-metal Threshold Concentrations

Soil-metal	Copper	Molybdenum	Gold	Silver	Zinc
Anomalous threshold 2nd. Positive Standard. D.	138.6 ppm	4.56 ppm	17 ppb	200 ppb	190.6 ppm
Selected threshold	90 ppm	4 ppm	20 ppb	200 ppb	250 ppm

The 1991 Stow soil survey resulted in the identification of 4.5-km (2.75-mi) long anomalous area along the slopes southwest of Le Mare Lake (Figures 3 and 5). Birkeland (1991) concluded that alteration and mineralization was exposed as a 6-km (3.66-mi) long linear belt of copper enrichment flanked to the northeast by belts of gold, molybdenum, and zinc enrichment, and to the southwest by a belt of zinc enrichment. That belt was postulated to have extended from Harvey Cove in the northwest to southeast of the southeastern shore of Le Mare Lake. Upon examining the property, the writer disagreed with that interpretation (sections 3.2.2, 4.2, and 5.3.2, this report).

Mineralization of several showings areas near Le Mare Lake were examined, including: the South Gossan zone, Trapper's Cabin area, Culleet Creek zone, South Lake zone, Le Mare No. 2 showing, and the North and South Lake zones (Figures 4 and 6). Roadside grab and chip samples were taken throughout the 1991 study area where disseminated and vein-hosted copper and molybdenum mineralization were encountered (Birkeland, 1991). Trenching and composite chip sampling was conducted at the Culleet Creek zone (Figures 4, 6, and 17). There, disseminated and vein-hosted copper mineralization, mostly chalcopyrite and bornite, was found to be associated with silicification and "apple green" alteration. Weighted averages of the results of the 1991 chip sampling of those trench-areas were tabulated by the writer (section 4.2.2, this report) (Table 8).

An "orientation" ground magnetometer survey comprising one line of unspecified length and location was conducted in the South Gossan zone. Readings were taken at 25-m (82-ft) intervals along the line. The results of that survey were reported as follow:

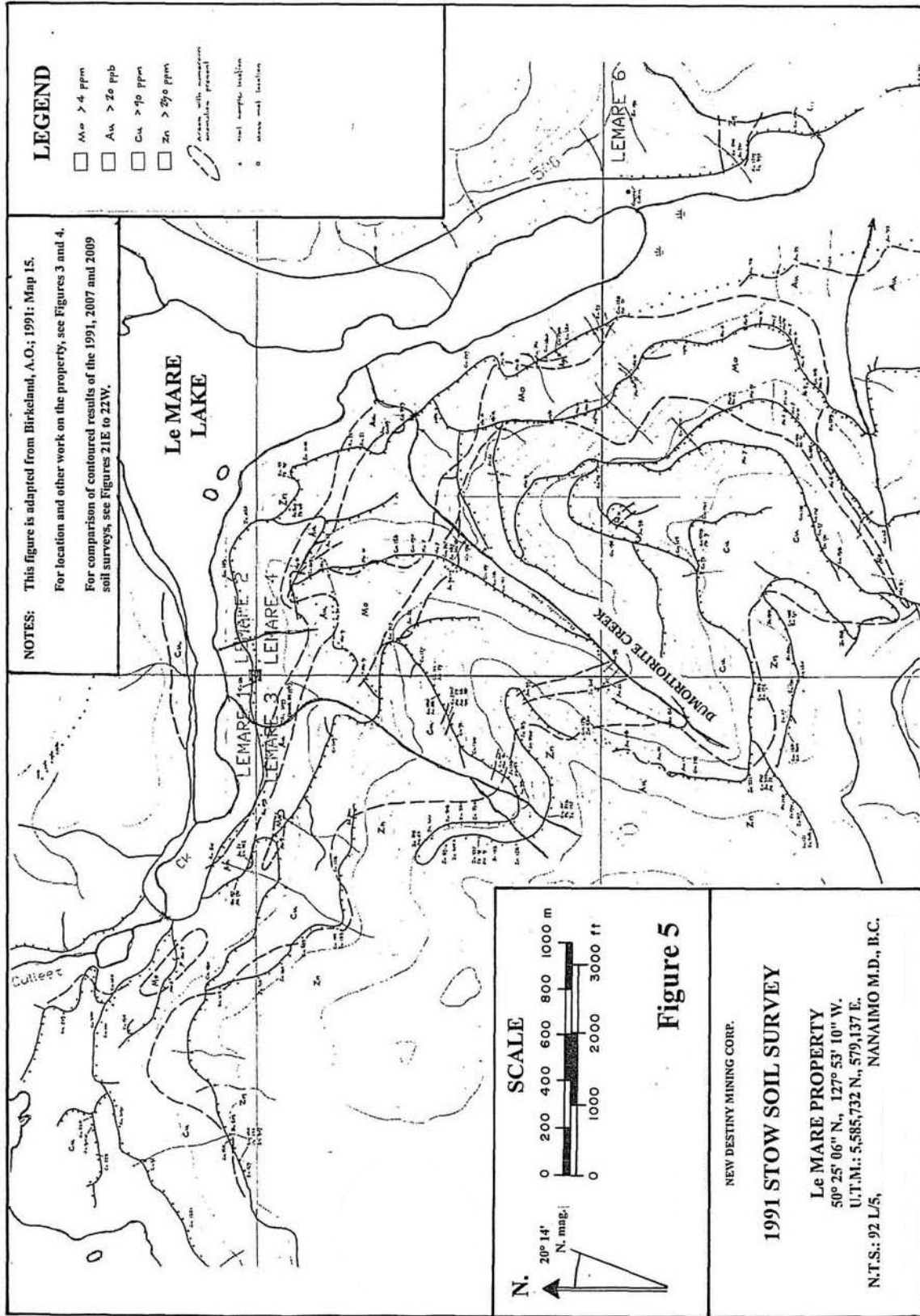
Results indicated that values within the South Gossan alteration zone were relatively constant with values ranging between 56,150 and 56,250 gammas (nanoteslas). At the alteration contact, a 7 station high to 56,650 followed by a 7 station low to 55,800 gammas encountered a magnetic cross-over of approximately 850 gammas. Within the wallrock volcanics, spiky readings fluctuating 600 to 700 gammas with means at approximately 56,200 gammas occurred.

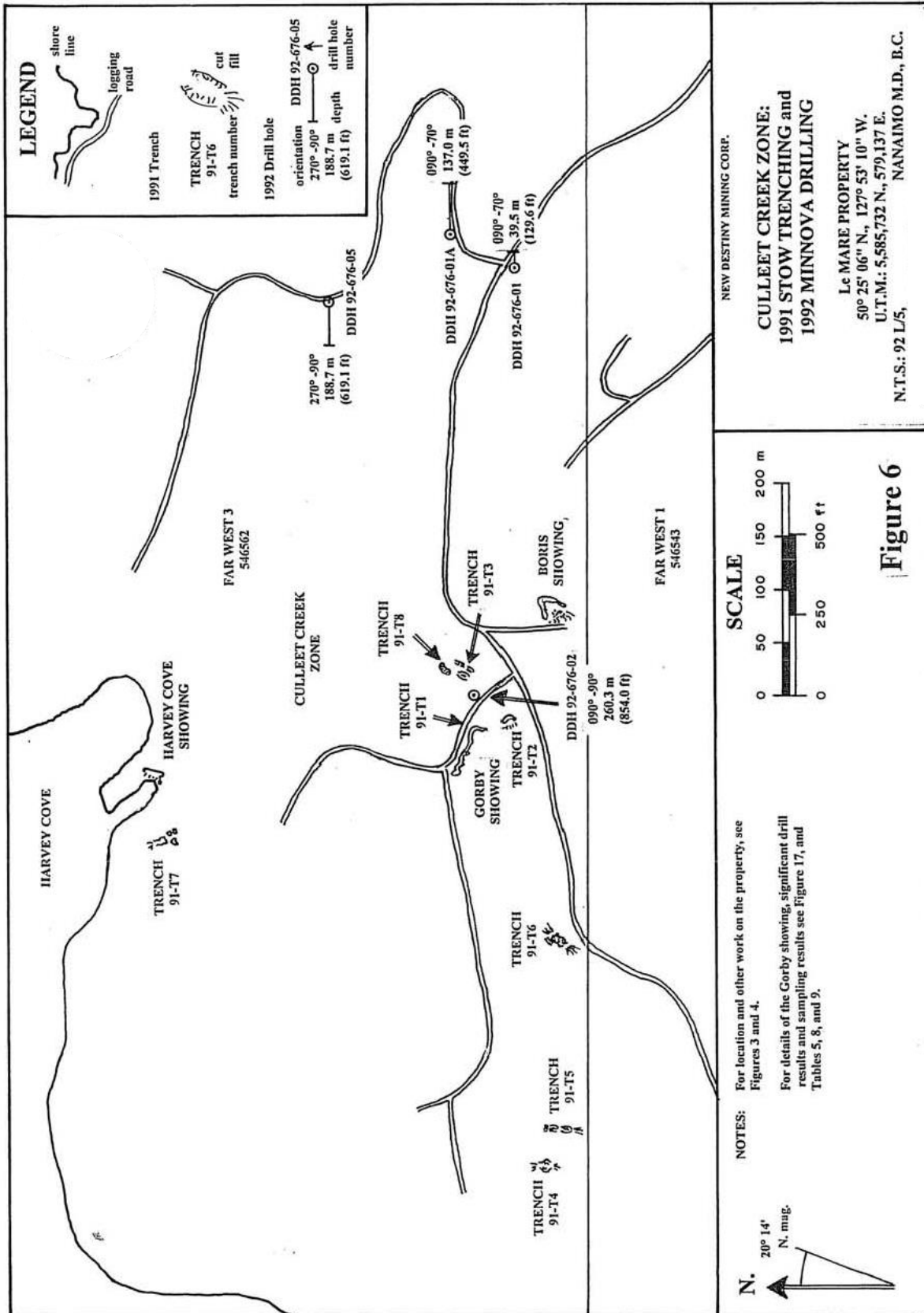
Birkeland, A.O.; 1991: p. 20.

The writer assumes that this line was run east-west into the argillic-phyllitic alteration plume and the coincident aeromagnetic low south of Dumortiorite Creek (Figures 4, 9, and 15) in the South Gossan zone.

Craig Leitch (1991) (Appendix VIII in Birkeland, 1991) conducted a petrographic study of 26 rock specimens from southwest of Le Mare Lake. Alteration types found included: potassic, propylitic, argillic, phyllic, and silicic (section 4.2.1, this report) (Figures 14 and 15).

1992: Stow Resources's Lemare property was enlarged by staking from September, 1991 to January, 1992 when Minnova Inc. optioned it from Stow. Immediately upon securing its option, Minnova commissioned Aerodat to fly airborne: magnetic, electromagnetic and gamma-ray spectrometer surveys along a total of 435 km (265.4 mi) of flight line according to Dave Heberlein (1993B). D. J. Pawliuk (1998) mentioned that a report of the survey was written by ?.?. Woolham in 1992. That report was unavailable to the writer.





Normally, results from airborne radiometric surveys are most definitive from surveys that are conducted during warm, dry weather during the late summer months. Minnova's 1992 gamma-ray survey was conducted during a period of heavy rains and possible snow during late winter over British Columbia's westcoast rain forest. The writer suspects that the results of that survey were not very useful.

1992 Continued

The summer, 1992 program comprised at least 5 km² (1.9 mi²) of geological mapping at 1:5,000 and 1:10,000 scales (not all was reported) (section 3.4, this report) and geochemical sampling: 1,154 rock, 39 soil, 72 moss mat and 55 silt samples were collected (Heberlein, 1993A). Moss-mat samples were collected from all of the significant drainages in the current Le Mare property-area. The focus of the 1992 soil and rock sampling program was in the northwestern part of the Le Mare hydrothermal system. Anne Thompson (1992) examined alteration and conducted an x-ray diffraction study on 9 clay samples from the South Gossan zone (section 4.2.1, this report).

During October 1 to 18, 1992, 900.5 m (2,954.4 ft) of BQ core was drilled in six holes: one hole was drilled into the Culleet Creek zone. Three holes were drilled into a geophysical anomaly just east of it (Figures 4 and 6), and one hole was drilled in each of the South Lake and South Gossan zones (Figure 4).

Dave Heberlein (1993B) reported that, "the best targets generated by the field program were drill tested". It is assumed that Heberlein was referring to both the airborne surveys and follow-up ground work. The only hole that intersected sections containing significant copper concentrations was DDH 92-676-2:

Table 3
Significant Intersections in 1992 Minnova Diamond Drill Holes

Drill Hole	Location	Interval		Length		Copper > 500 ppm	Molybdenum > 50 ppm
		m.	ft.				
92-676-2	Culleet Creek zone	11.1-13.1	36.4-43.0	2.0	6.56	684	
		13.1-15.1	43.0-49.5	2.0	6.56	719	
		19.0-21.0	62.3-68.9	2.0	6.56	746	
		21.0-23.0	68.9-75.5	2.0	6.56	863	
		23.0-25.0	75.5-82.0	2.0	6.56	959	
		58.0-62.7	190.3-205.7	4.7	15.42	529	

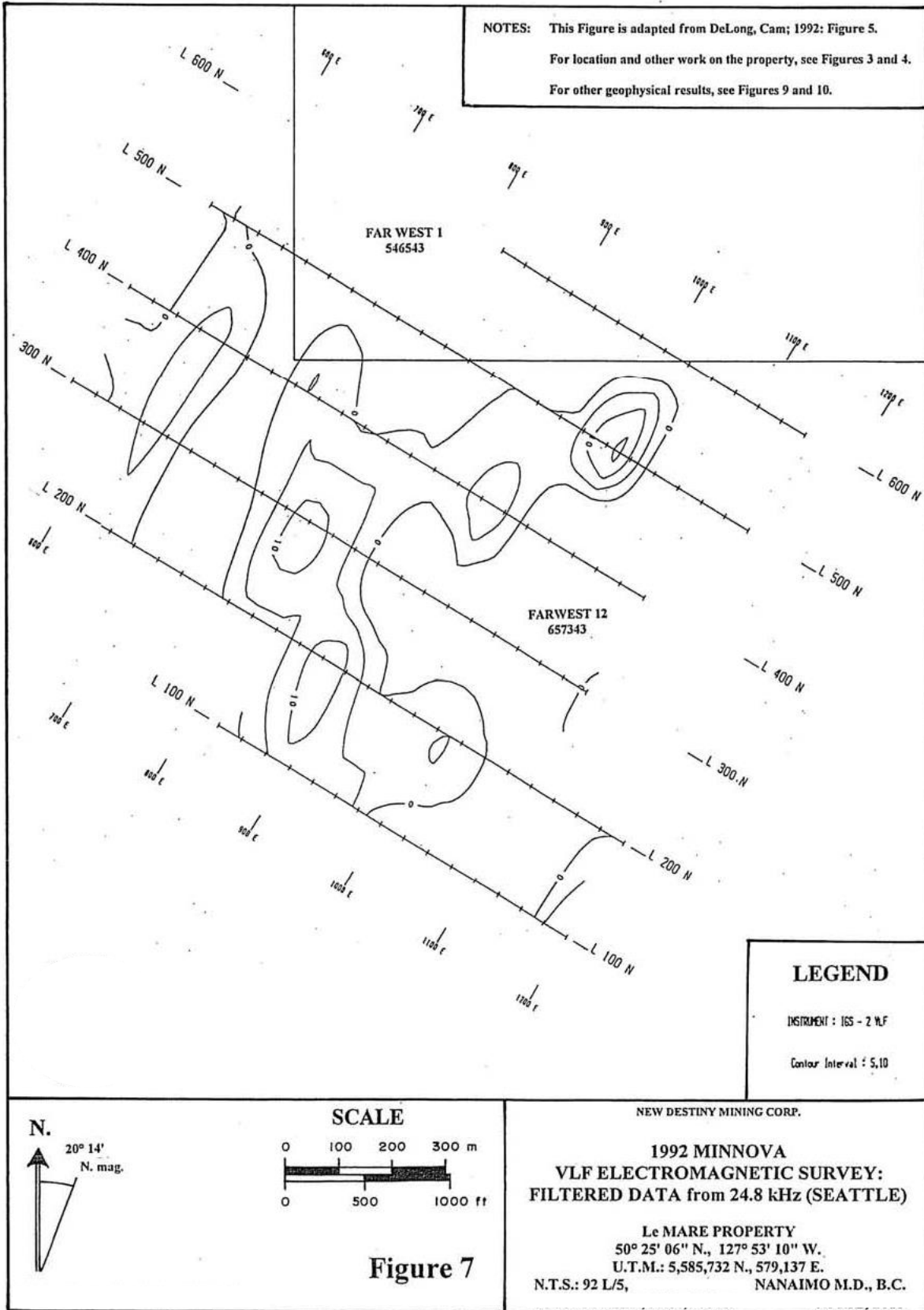
NOTES: This table is produced by the writer from the certificates of analysis attached to the report of Heberlein, Dave; 1993B.

For locations of 1992 drill holes, see Figures 4 and 6.

Diamond drill holes 92-676-1 (lost in poor ground), 92-676-1A, 92-676-3, and 92-676-5 were drilled into a geophysical anomaly located southeast of Harvey Cove and south of Culleet Creek (Heberlein, 1993B) about 150 m (492.1 ft) east of the Culleet Creek hydrothermal plume's margin (Figures 4, 6, and 15). This could account for Heberlein's (1993B) report of weak potassic alteration and copper mineralization encountered in these holes.

Drill hole 92-676-2 was drilled on the access road about 50 m (164 ft) east of the Gorby showing (Figure 6), well within the Culleet Creek plume. The results from that drill hole were summarized as follows:

92-676-2 ... was drilled to test the depth extent of disseminated chalcopyrite mineralization at the Gorby Zone. The hole penetrated a sequence of potassic to chlorite altered flow banded rhyolites, rhyolite breccias and felsic tuffs with rare intervals of basalt. Consistent fracture controlled chalcopyrite mineralization (to 3%) (qualitative visual estimate) occurs in the upper 26 m (85.3 ft) of the hole.



Quartz stockworking is well developed in the mineralized section. Wall rocks are pervasively silicified and potassium feldspar alteration envelopes occur. Up to 3% (qualitative visual estimate) chalcopyrite is present throughout this interval and Cu grades (concentrations) range up to 959 ppm.

Lower in the hole chlorite-calcite-hematite alteration is prevalent. Traces of chalcopyrite occur to a depth of 252.1 m (827.1 ft), but copper grades (concentrations) do not exceed 124 ppm.

Heberlein, Dave; 1993B: p. 13.

Drill hole 92-676-4 penetrated the South Gossan zone in the eastern part of an area that was reported to have hosted pervasive argillic and advanced argillic alteration over a mineralized potassic alteration zone. Results from that drill hole were summarized as follows:

92-676-4... was the only hole drilled into the South Gossan Zone. It penetrated a section dominated by highly vesicular rhyolite flows (silicified vesicular basalt flows?) and fragmental rocks. Alteration is moderate and consists of pervasive sericitization with minor silica flooding. Chlorite is also abundant, particularly near a basalt dyke at 91.0 m (298.6 ft).

Heberlein, Dave; 1993B: p. 14.

The writer is of the opinion that the 1992 Minnova crew mis-identified silicified mafic volcanic rocks as rhyolitic rocks (section 3.4.1, this report), the same mis-identification during 1992 core logging is probable.

1992 Continued

Quest Canada Exploration Services conducted a ground very-low-frequency electromagnetic survey on a 6-line grid on Gooding Ridge between Gooding Cove and the Culleet Creek zone to test a distinct airborne anomaly in that area. The surface anomaly was considered to be weak and of little interest (DeLong, 1992) (Figures 4 and 7).

1992 to 1997

No exploration was recorded and the 1991-era LeMare claim group lapsed.

1997: On February 6, 1997, David J. Pawliuk recorded the LEM 1 to 6 (353575 to 353580) 2-post claims. The LEM 1 and 2 claims were located on the Culleet Creek zone and the LEM 3 to 6 claims occupied the eastern part of the Southern Gossan zone as defined by Birkeland (1991) (Figure 4). During the 1997 prospecting program conducted by Fox Geological Services Inc., 10 rock samples were taken. None were significantly mineralized with either copper or molybdenum (Pawliuk, 1998). Enough assessment credit was applied to the LEM claims to keep various claims in good standing to February 6, 2001 to February 6, 2003.

During the summer of 1997, geologists from Phelps Dodge Corp. visited the Le Mare Lake area as part of the company's project No. 207. Rob Cameron submitted grab samples 62960 to 62965 taken around the Gorbey showing on the LEM 1 (353575) claim to Acme Labs. They were found to contain from 1,005.7 to 5,245.1 ppm copper and from 0.3 to 4.9 ppm molybdenum. The relation between Fox Geological Services Inc. and Phelps Dodge Corp., if any, is unknown to the writer.

1998 to 2006

No exploration was recorded and the LEM claim groups lapsed.

2006: From December 4 to 6, 2006, J.T. Shearer map-staked the FAR WEST 1 to 6 (546543, 546454, 546562, 546563, 546565, and 546689) claims to cover the slopes southwest of Le Mare Lake (Figure 3). Those claims formed the core-area of the current Le Mare property.

2007: J.T. Shearer enlarged the current Le Mare property-area by map-staking the FAR WEST 7 and 8 (563795 and 563802) claims south and southeast of the core-area respectively on July 29, 2007. The property-area was expanded farther to the north and east by Shearer's map-staking of the FAR WEST 9 to 11 and GEYSERITE (569848 to 569850 and 570078) claims from November 10 to 14, 2007. The property was optioned to Equus Energy Inc. of Vancouver, B.C.

Homegold Resources Ltd., a private exploration company controlled by J.T. Shearer, conducted a program of prospecting and soil sampling along several of the lower roads around Le Mare Lake focusing on previously defined anomalous areas (Shearer, 2007). A total of 131 soil and 4 rock samples were taken and analyzed by the induced plasma coupling (ICP) method for 30 elements. Gold concentrations were determined by fire assay and atomic adsorption techniques.

Upon contouring Shearer's 2007 and 2009 soil-survey data (Figures 21E to 22W) (sections 4.2.1, 5.3.2 and 7.0, this report), the writer found that Shearer's data more precisely defined soil copper and molybdenum anomalies and could be used to help define hydrothermal plumes in the northwestern part of the Le Mare hydrothermal system (Figures 15 and 21W).

2008: During the 2007 exploration program, chalky geyserite, a grey-white hydrated silicate ($\text{SiO}_2 \cdot n\text{H}_2\text{O}$), an ingredient in Portland cement, was discovered to occur in small amounts along a road south of Culleet Lake. By sometime during 2007 or 2008, J.T. Shearer had optioned the copper and molybdenum of the Le Mare property to Equus Energy Inc. and the geyserite on the same property to Electra Gold Ltd.

From October 25 to December 4, 2008, J.T. Shearer's exploration company Homegold Resources Ltd., conducted prospecting for geyserite along a disused logging road southwest of Culleet Creek and Lake, and near the South Gossan zone (Shearer, 2009). A total of 51 samples were taken from those areas. No significant concentrations of that industrial mineral were found.

On April 5, December 5 and 14, 2008, J.T. Shearer expanded the Le Mare property-area by mapstaking the MAHATTA 1, NORTHEAST LEMARE, and FAR WEST 13 (580535, 595599, and 596074) claims to the northeast and south of the established property area.

2009: The options of Equus Energy and Electra Gold with regard to the Le Mare property were terminated. On October 7, 2009, New Destiny Mining Corp. optioned the Le Mare property from J.T. Shearer.

Upon reviewing the exploration data, the writer became of the opinion that the Le Mare hydrothermal system occupied an area shaped like a lima bean and was not part of a linear, asymmetric, mineralized trend as assumed by previous explorationists of the area. J.T. Shearer map-staked the FARWEST 12 and 13 (657343 and 657363) claims to cover the projected southwestern extension of the hydrothermal system (Figure 3).

J.T. Shearer, Bryce Clark (President of New Destiny Mining Corp.), and John Ostler examined the Le Mare property on November 4, 2009. The 2009 exploration program was conducted from that time until December 15, 2009 by Homegold Resources Ltd. The program comprised prospecting, soil sampling, and some check-mapping in two areas: between the Culleet Creek zone and Gooding Cove, and in the South Gossan zone (sections 3.4, 4.2, and 5.1 to 5.4, 6.0 and 7.0, this report) (Figures 4, and 17 to 22W). A total of 235 soil and 33 rock samples were taken. All samples were analyzed for 33 elements by induced coupled plasma (ICP) techniques; high concentrations were determined by fire assay and atomic adsorption. Soil-copper anomalies between the Culleet Creek zone and Gooding Cove confirmed

the presence of mineralized hydrothermal plumes in that area, southwest of the linear trend that had previously been thought to have hosted all significant porphyry copper mineralization.

Historical Mineral Resource and Mineral Reserve Estimates,
and Production from the Le Mare Property-area

No historical estimates of mineral resources or reserves related to the Le Mare property, or historical production from the Le Mare property-area are known to the writer.

GEOLOGICAL SETTING

Regional Geology

Dave Heberlein (1993A) described the geology of northwestern Vancouver Island as follows: Northwestern Vancouver Island lies within Wrangellia; a part of the Insular belt of British Columbia. Oldest rocks in the region are Upper Triassic tholeiitic basalts of the Karmutsen Formation which form the basement to the overlying Jurassic and Cretaceous stratigraphy.

Middle Jurassic Bonanza Supergroup rocks outcrop over much of the western part of northern Vancouver Island. The basal part of the Bonanza Supergroup is a marine volcanic sequence consisting of amygdaloidal, pillowed basalts and andesite with interbedded tuffs and intraformational breccias. It grades upwards into a succession of andesitic to dacitic flows, tuffs, and breccias which are in turn overlain by a sub-aerial sequence of interbedded intraformational breccias and maroon subaerial basalt flows, dacites and rhyolites. Felsic rocks are abundant close to volcanic-intrusive centres and are often interbedded with volcanoclastic sediments.

The Bonanza volcanic sequence is unconformably overlain by or faulted against shallow marine clastic sedimentary rocks of the Cretaceous Long Arm Formation.

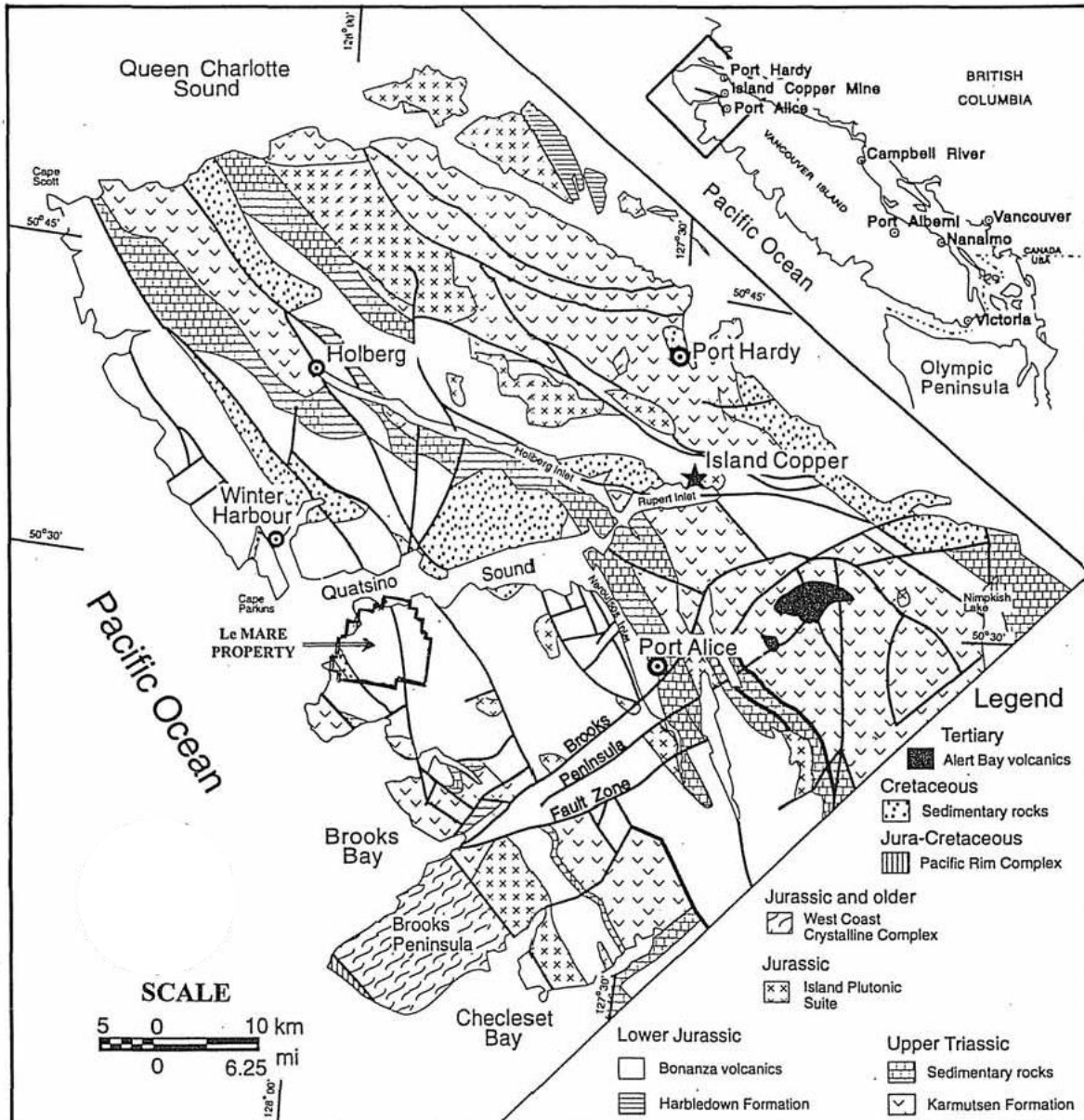
Intrusive rocks in the region are interpreted to be coeval with the Lower Jurassic Bonanza volcanic rocks. Known as the Island Intrusives, they consist mostly of granodiorites and monzonites. These intrusions are associated with porphyry and skarn mineralization throughout the central and north parts of Vancouver Island.

The Lemare claims lie within a fault bounded structural block named the Cape Scott block by Muller (1977). Brittle faulting and broad open folding are the main styles of deformation. Muller (1977) and Jeletzky (1970) attribute this to the thick, brittle section of Karmutsen basalt that forms the basement to the Jurassic rocks.

Heberlein, Dave; 1993A: pp. 4-5.

G.T. Nixon of the British Columbia Geological Survey conducted a regional mapping program throughout the northern part of Vancouver Island during the early 1990s that resulted in a regional geological map of the area (Nixon et al., 1994) (Figure 8).

The writer's tabulation of the geologic history of the region around the Le Mare property-area is as follows:



NOTE: This Figure is adapted from Nixon, G.T. et al.: 1994; Figure 1.



Figure 8

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**REGIONAL GEOLOGY from NIXON et al.
in B.C.E.M.&P.R. Pap. 1994-1**

Le MARE PROPERTY

50° 25' 06" N., 127° 53' 10" W.

U.T.M.: 5,585,732 N., 579,137 E.

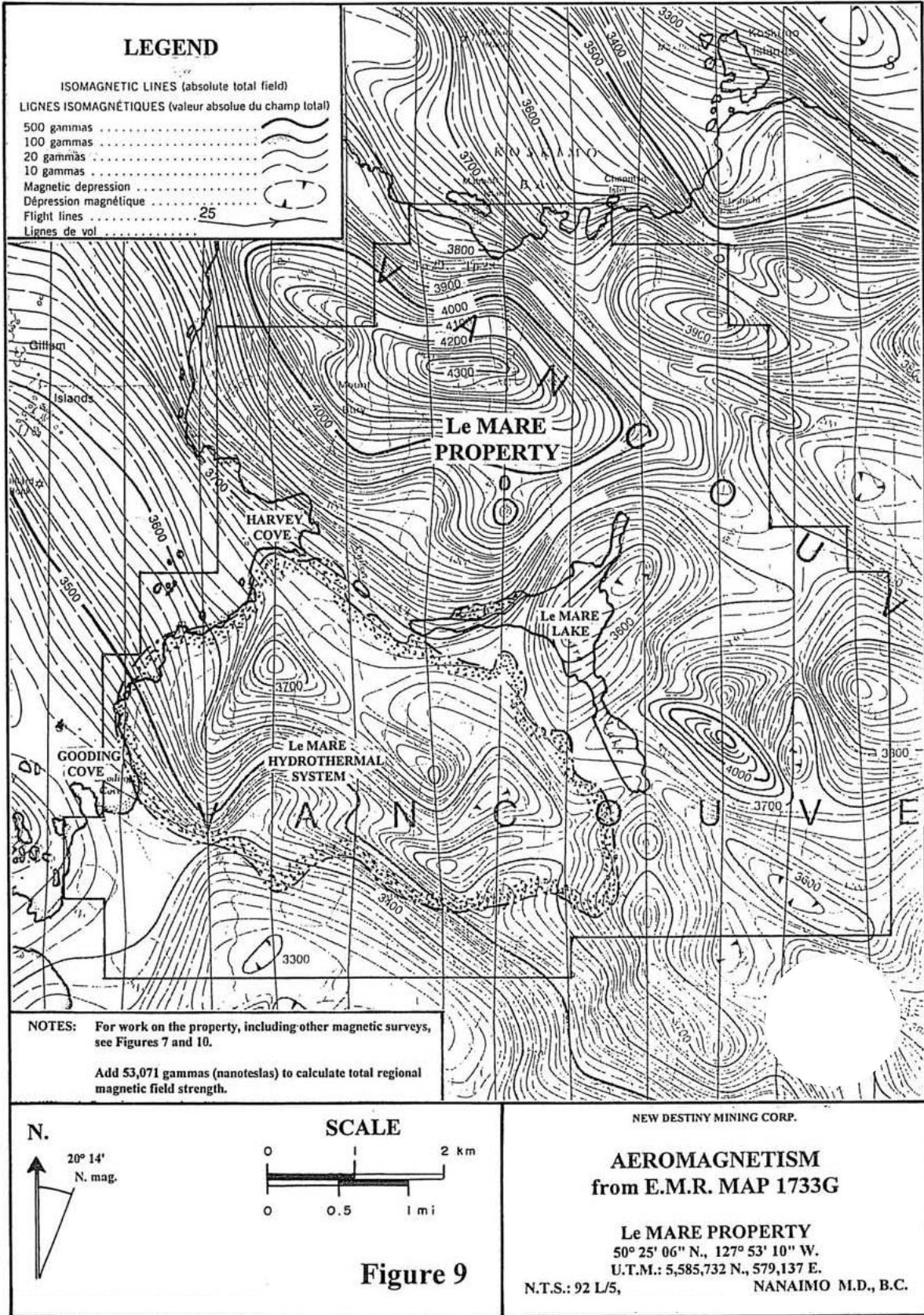
N.T.S.: 92 L/5.

NANAIMO M.D., B.C.

Table 4
Table of Geological Events and Lithological Units in the Le Mare Property-area

Time	Formation or Event
Recent 0.01-0 m.y.	Valley rejuvenation: Down cutting of stream gullies through till, development of soil profiles.
Pleistocene 1.6-0.01 my.	Glacial erosion and deposition: Removal of Tertiary-age regolith, deposition of till and related sediments at lower elevations, smoothing of the Tertiary-age land surface.
Late Miocene 7.6-7.9 m.y.	Tensional faulting: Deposition of the Alert Bay basaltic volcanic rocks
Eocene to Late Oligocene 32 - 59 my.	Northeasterly trending tensional faulting: Emplacement of the Sooke intrusions and Metchosin volcanic rocks MINERALIZATION: Emplacement of gold-bearing quartz veins
Late Cretaceous to Paleocene 75.0-57.0 m.y.	Laramide Orogeny: Mild folding and faulting, in central British Columbia. Northeastward tilting on the eastern side of the Vancouver Island area. Emplacement of the Nanaimo Formation sediments
Early to Middle Cretaceous (Valanginian to Cenomanian) 137.0 - 93.5 m.y.	Deposition of the Logram and Queen Charlotte Group clastic sedimentary rocks on the Late Mesozoic erosional surface.
Middle Jurassic to Early Cretaceous 163-137 m.y.	Uplift and erosion: Gentle westward tilting of the western part of the Vancouver Island area resulting in partial unroofing of the early Mesozoic stratigraphy
Late Jurassic to Late Cretaceous 144-88 m.y.	Columbian Orogeny: Emplacement of the Coast Intrusions east of the Vancouver Island area, thrusting and transcurrent faulting, deformation of Cache Creek rocks in a northeastward dipping subduction zone, accretion of Nicola Group rocks to North America
Middle Jurassic 166.0-159.7 m.y.	Nassian Orogeny: Final emplacement of the Island Intrusions accompanied by local folding and contact metamorphism in adjacent cover rocks and lower greenschist facies regional metamorphism. Regional faulting and tilting resulting in southwestward dipping monoclines followed by uplift and erosion.
Early to Middle Jurassic (Sinemurian to Bajocian) 197.0 - 166.0 m.y.	Subduction and calc-alkaline island arc volcanism and related clastic sedimentation: Deposition of the Bonanza Supergroup mafic to felsic volcanics and Island Intrusions MINERALIZATION: 175 m.y. Development of the Island Copper Complex calc-alkaline porphyry Cu-Au-Mo deposits Presumed time of development of the Le Mare hydrothermal system
Late Triassic (Karnian to Norian) 220.7- 209.6 m.y.	Deposition of the Vancouver Group in a fore-arc basin: Quatsino Formation reef-related limestone beneath Parson Bay Formation calcareous wacke and argillite
Middle Triassic (Ladnian to Karnian) 240.6-220.7 m.y.	Deposition of Karmutsen Group mafic volcanics on a spreading oceanic crust.
	m.y. = million years ago

NOTE: Data for this table was compiled by the writer from various sources including Muller (1977) and Douglas ed. (1970).



Regional and Property Geophysics

Regional Aeromagnetic Survey

In September, 1962, the Geological Survey of Canada conducted a fixed-wing airborne aeromagnetic survey over the northern part of Vancouver Island. Energy, Mines, and Resources Map 1733G covering N.T.S. map-area 92 L/5 was one of the aeromagnetic maps produced. The current Le Mare property-area is in the west-central part of that map-area (Figure 9).

The northeastern part of the property-area coincides with a regional northwesterly trending magnetic high that may be a reflection of mafic volcanic stratigraphy in that area. Peaks in this magnetic trend are located at the hill top east of the southern end of Le Mare Lake and near the peak of Mount Bury (Figures 3 and 9). Exposures of the Le Mare hydrothermal system are located on the southwestern flank of the aeromagnetic trend. Three local magnetic highs occur along the ridge that transects the hydrothermal system. A distinct magnetic low coincides with the phyllic-argillic alteration that covers much of the South Gossan zone (Figures 5, 9, and 15). The writer presumes that magnetic low to be an effect of magnetite destruction by that alteration.

During a preliminary investigation of the Le Mare Lake area in 1991, Keewatin Engineering manipulated data generated from E.M.R. Map 1733G to produce maps of enhanced total field and calculated gradient magnetic data (Figure 10) superimposed on the 1: 50,000-scale N.T.S. Map sheet 92 L/5 (Birkeland, 1991). The maps submitted for assessment were in colour and without legends; thus, the locations of magnetic highs and lows, and the magnetic gradient can only be assumed from the colour distribution. The writer knows of no report of how the magnetic data manipulation was accomplished.

A.O. Birkeland's (1991) conclusions regarding the results of this data were as follow:

The calculated gradient map (Figure 10) indicates the following:

- A northwest trending low magnetic trough corresponding to the major cross property LeMare alteration trend. This magnetic low is likely caused by the destruction of magnetite within the argillic alteration trend.
- Anomaly A is coincident with the South Gossan Zone and indicates that although magnetite destruction is present at a high level in the advanced argillic and phyllic zones which outcrop on surface, magnetite alteration exists at depth beneath the alteration cap.
- Anomaly B is located on the ridge west of Dumortierite Creek where the best anomalous soil geochemistry on the property occurs. It is interpreted that this area is underlain by a porphyry system with corresponding flanking magnetite alteration and associated Cu-Mo-Au mineralization.
- Anomaly C is the highest magnetic anomaly adjacent to the LeMare-Culleet alteration trend. This anomaly is on strike with east-west faults exposed in the South Gossan Zone and on trend with east-west structures and geochemical anomalies encountered on the east side of LeMare Lake (Trapper cabin area).
- Anomaly D occurs in a covered low-land in the vicinity of the gold geochem anomalies “down plunge” of the main South Gossan Zone alteration cap. This large positive anomaly within the northwest trending magnetic low indicates that a porphyry and associated magnetite-bearing Cu-Mo-Au system may be at depth beneath the valley till and has not been detected by conventional soil geochemistry completed to date.

Birkeland, A.O.; 1991: pp. 19-20.

Birkeland's “northwest trending low magnetic trough” is one of a series of such “troughs” that transect the volcanic stratigraphy in the Quatsino Sound area. It cuts through the area of soil-copper enrichment separating the North Lake zone from the main part of the zone of soil-copper enrichment (Figures 10 and 21W). The writer

interprets this magnetic feature to have been due to post-mineralization weathering along a west-northwesterly trending fault, possibly previously responsible for the location of the Le Mare hydrothermal system (Figure 11).

Anomaly 'A' as plotted on Birkeland's (1991) magnetic gradient map is 1 km (0.61 mi) north of the South Gossan zone and not coincident with it. Similarly, Anomaly 'B' is plotted 1 km (0.61 mi) northnortheast of its described location. The described locations of these two anomalies make more sense than their plotted locations. The plotted locations of anomalies 'C' and 'D' are much better matches to their descriptions.

Anomalies 'A', 'B', and 'D' are small, local magnetic features. Although quite intense, anomaly 'C' doesn't resemble any of the magnetic gradient features spatially related to the areas of alteration and soil-metal enrichment associated with the Le Mare hydrothermal system. During the 1992 field season, Minnova geologists visited the area of anomaly 'D' and could not associate it with a body of hydrothermal alteration in the Bonanza Supergroup mafic volcanic rocks. That anomaly may be related to local volcanic stratigraphy.

During the early 1990s, it was well-known that the porphyry deposits of the Island Copper Cluster located near Port McNeill were concentrated at dilational jogs along a west-northwest trending, steeply dipping regional fault (Figure 12). Birkeland (1991) seems to have focused on the regional structure and ignored the round to oval shapes of individual deposits. He became convinced that alteration and mineralization near Le Mare Lake was exposed as a 6-km (3.66-mi) long linear belt of copper enrichment flanked to the northeast by belts of gold, molybdenum, and zinc enrichment, and to the southwest by a belt of zinc enrichment. A desire to promote that interpretation may have influenced focus on small magnetic features around the southeastern arm of Le Mare Lake to the exclusion of larger ones within the area defined by the magnetic gradient between Gooding Cove and Le Mare Lake (Figures 9 and 10).

When the 1991 Stow soil-copper and molybdenum anomalies, the 1992 Minnova ground electromagnetic anomaly, the results of the 1991 Keewatin calculated gradient magnetics, and those of the 2007 and 2009 soil surveys are combined, they indicate that the Le Mare hydrothermal system covers a 5 X 3 km (3.05 X 1.83 mi) or 15 km² (5.6 mi²) oval-shaped surface-area and not an asymmetric linear belt (Figures 10,11,15,and 21E to 22W).

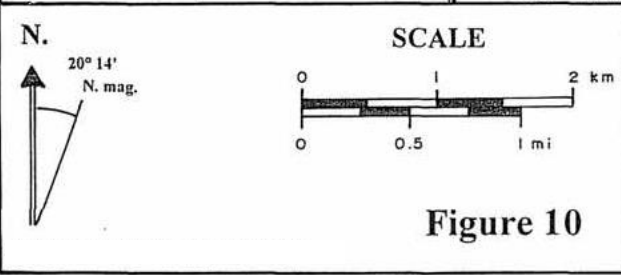
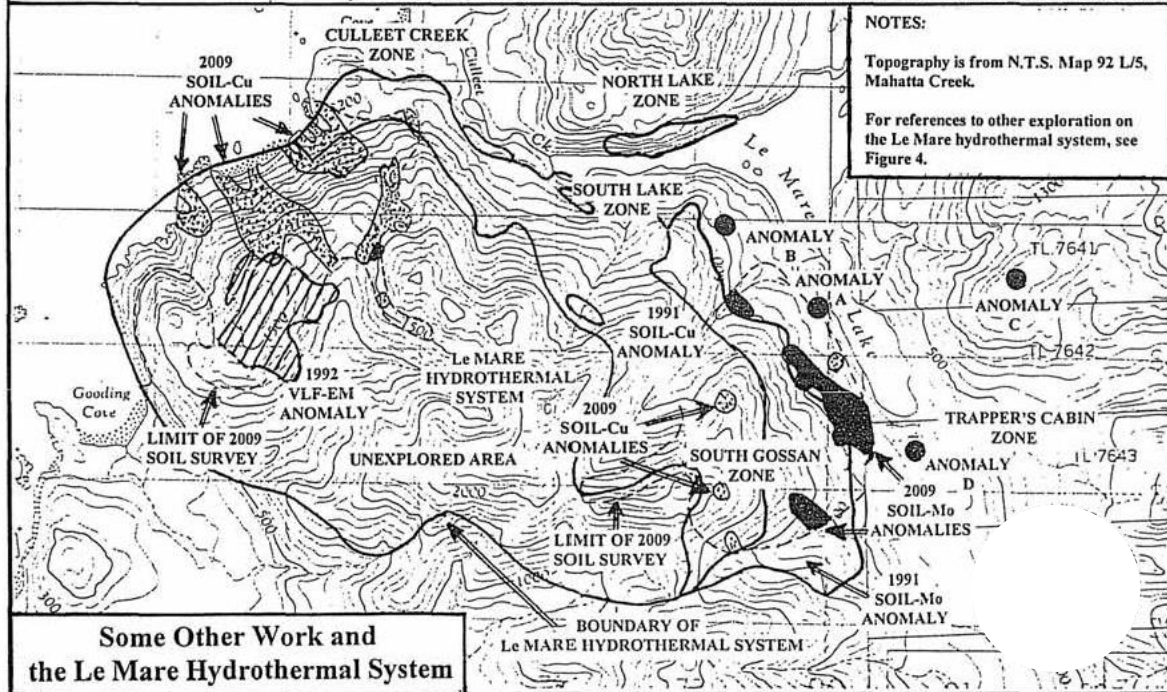
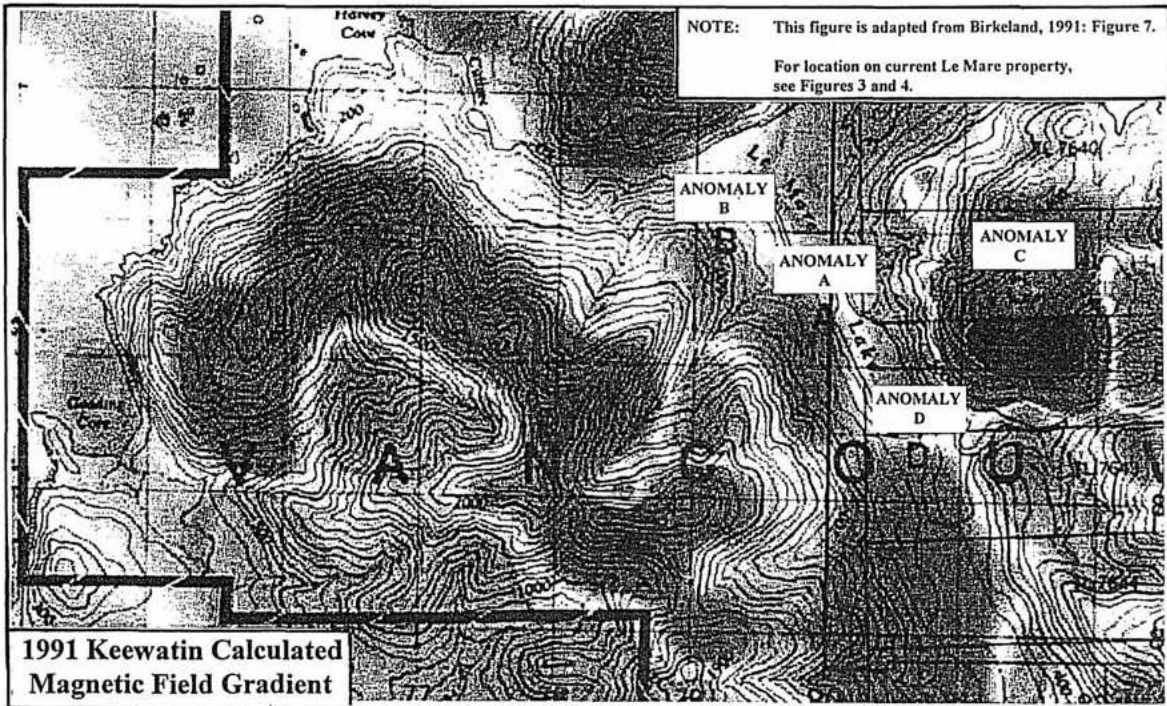
Regional Silt Geochemistry

A reconnaissance regional stream sediment sampling program was conducted during 1988 throughout the northern part of Vancouver Island, through a joint federal-provincial initiative resulting in the production of Geological Survey of Canada Open File 4020 (Matysek et al., 1988).

Research by Keewatin Engineering Inc. during March, 1991, including investigation of regional geochemical survey results, revealed that a belt similar to the Island Copper Belt was located between Kyuquot Sound and Quatsino Sound. It was named the Mahatta-Kashutl belt (Birkeland, 1991). Attributes of the two areas were sufficiently similar for Keewatin to stake and explore the 1991-era LeMare property (section 2.1, this report).

J.A. Perelló et al. (1995) reported that the porphyry deposits of the Island Copper Cluster were concentrated along dilational jogs in a west-northwesterly trending, steeply dipping, right lateral, transcurrent fault (Figure 12).

Selected silt-metal concentrations of silt samples taken from locations near the Le Mare property (Figures 3 and 11) were tabulated as follows:



NEW DESTINY MINING CORP.

RELATION of AEROMAGNETIC GRADIENT to the Le MARE HYDROTHERMAL SYSTEM

Le MARE PROPERTY
50° 25' 06" N., 127° 53' 10" W.
U.T.M.: 5,585,732 N., 579,137 E.
N.T.S.: 02 L/5. NANAIMO M.D., B.C.

Table 5
Selected Regional Silt-metal Concentrations

Sample Number	Water pH	Copper ppm	Lead ppm	Zinc ppm	Arsenic ppm	Moly. Ppm	Silver ppm	Gold ppb
883053	7.3	38	1	82	7	1	0.1	1
883082	7.1	41	13	240	10	1	0.1	1
883128	7.1	32	1	76	6	1	0.1	1
883129	7.0	44	1	86	6	1	0.1	1
883131	6.8	33	2	75	4	1	0.1	1
883237	6.7	34	3	87	12	1	0.1	107
883238	7.1	19	1	68	7	1	0.1	1
883262	7.2	34	9	230	14	1	0.1	2
883263	7.1	39	3	152	11	2	0.1	2
883264	7.0	42	5	155	11	1	0.1	18
883265	7.4	41	1	102	11	2	0.1	2
883266	7.4	43	3	135	11	1	0.1	1
883267	7.3	44	1	87	7	3	0.1	4

NOTE: For sample locations, see Figure 11.

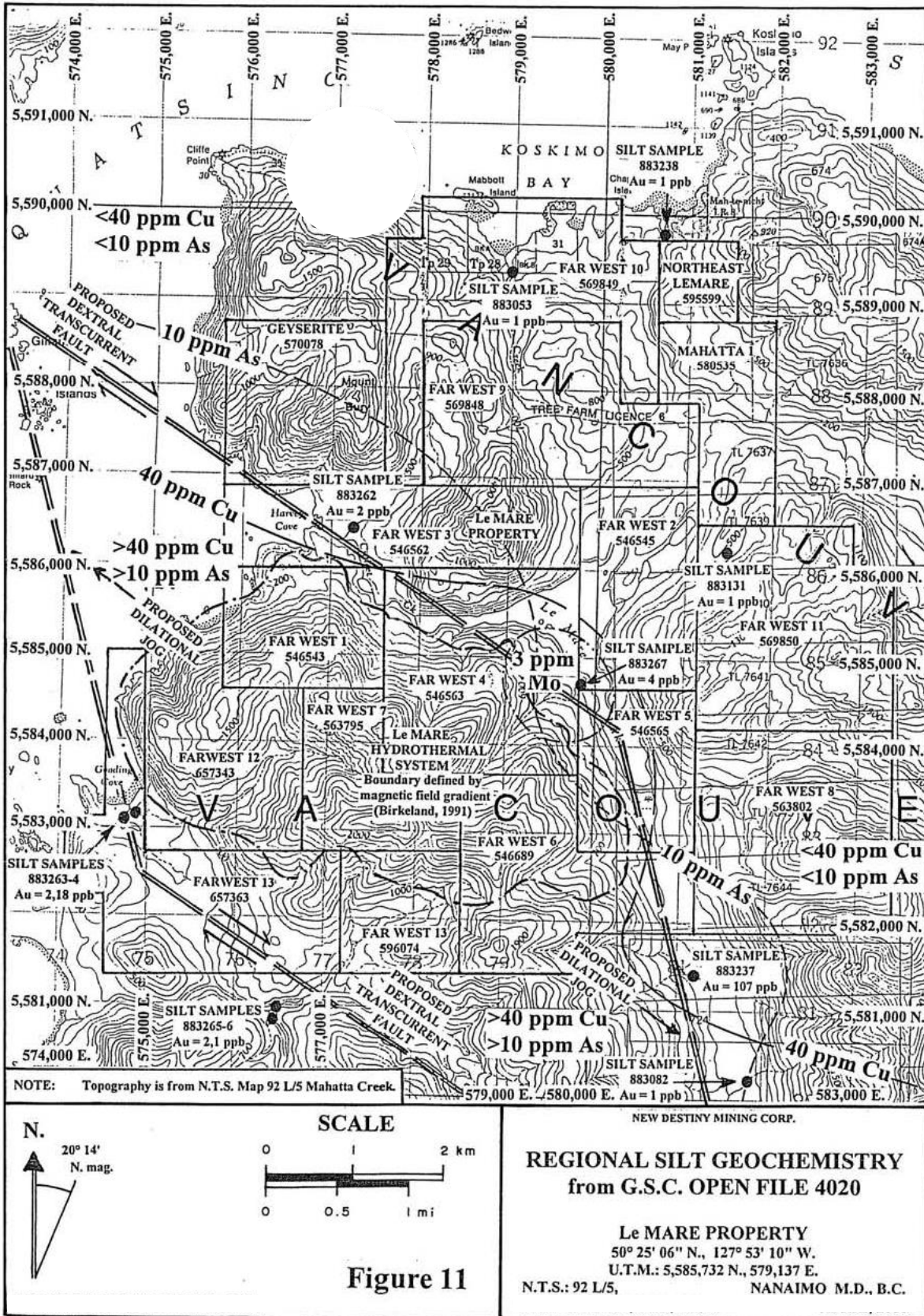
Regional silt survey results indicate that the Le Mare hydrothermal system may also occupy a dilational jog in a regional fault similar to those which controlled mineralization of the Island Copper Cluster (Figure 11).

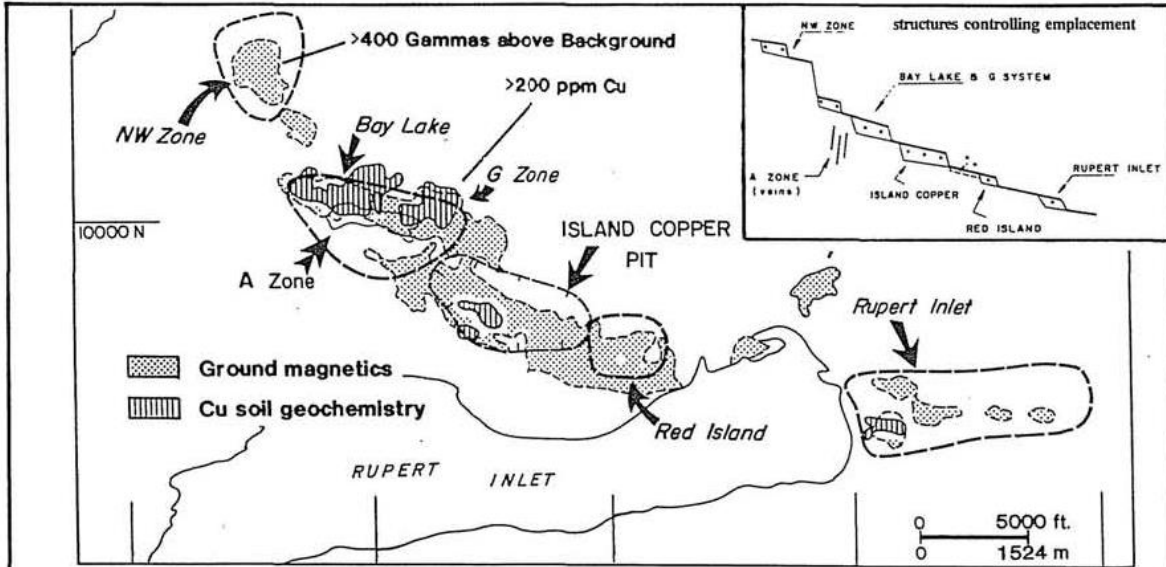
It is proposed that a steeply dipping right-lateral fault, trending at 306° may extend from beneath Quatsino Sound southeastward to Le Mare Lake where it terminates. A parallel structure may accommodate right-lateral displacement from Gooding Cove southeastward to beyond the head of Klatskino Inlet (about 12.5 km (7.6 mi)) southeast of the southeastern corner of the Le Mare property. A dilational jog between these two west-northwesterly trending faults may be defined by two steeply dipping faults that trend at about 338°. The easterly one may underlay the south arm of Le Mare Lake and Keith River; and the westerly one may extend from Gooding Cove north-northwestward to Gillam Islands beneath Quatsino Sound. The Le Mare hydrothermal system occupies an area bounded by these proposed faults (Figure 11).

Elevated silt-gold concentrations occur in six samples in the Le Mare property-area: 883237, 883262 to 65, and 883267, all of which are within 300 m (984 ft) the surface traces of the proposed faults. The 40ppm copper and 10-ppm arsenic contours separate areas of comparatively low silt-copper and arsenic concentrations to the north and east of Le Mare Lake with areas of higher concentrations to the south and west of it. The two contours roughly follow the northern and eastern boundaries of the proposed dilational jog, and could be the result of comparatively copper and arsenic-rich volcanic stratigraphy having been translated west-northwestward into contact with rocks with lower copper and arsenic contents along a regional dextral transcurrent fault system.

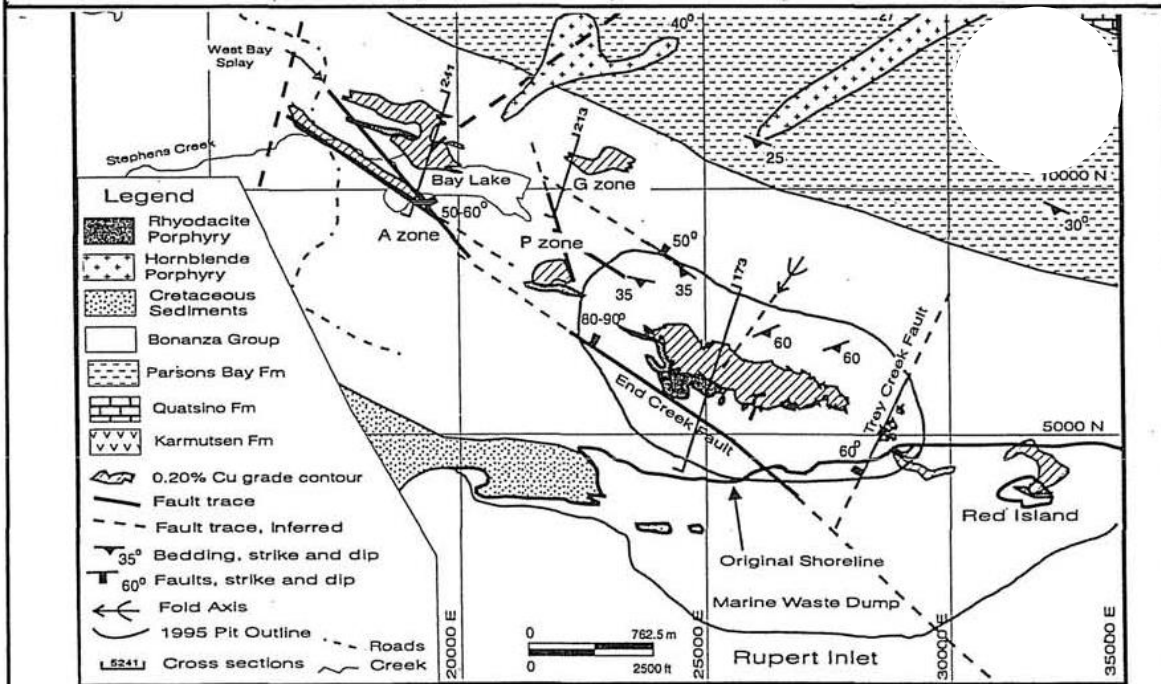
Silt sample 883267, taken near the mouth of Dumortiorite Creek and down-stream from the South Gossan Zone soil-molybdenum anomaly, contained 3 ppm molybdenum. That concentration was determined by the writer to be sub-anomalous in soils of the area (section 5.3.2, this report) (Table 11) (Figures 22E and 22W). The only other two silt samples with elevated molybdenum contents were samples 883263 and 883265 which were taken from streams that drain the southern part of the Le Mare hydrothermal system (Figure 11).

Regional silt-silver, lead, and zinc distributions are not very diagnostic of regional structures or of mineralized locations.





Location of porphyry centres of the Island Copper Cluster in relation to ground magnetic and copper-in-soil geochemical anomalies.



Geology of the Island Copper mine area. Rhyodacite porphyry contacts and 0.2% Cu boundaries are projected to the 920 elevation (sea level = 1000 feet) for the Red Island, Island Copper and P zone centres and to the 1140 elevation for the Bay Lake and G zone centres.



NOTE: These figures are adapted from Perelló, J.A. et al.; 1995: Figures 4, 5, and 24.

NEW DESTINY MINING CORP.

CONFIGURATION and GEOLOGY of the ISLAND COPPER CLUSTER DEPOSITS

Le MARE PROPERTY

50° 25' 06" N., 127° 53' 10" W.

U.T.M.: 5,585,732 N., 579,137 E.

N.T.S.: 92 L/5,

NANAIMO M.D., B.C.

Figure 12

Property Geology

Stratigraphy

Two mapping programs in the Le Mare property-area have been recorded for assessment: those of A. O. Birkeland (1991) for Stow Resources Ltd., and of J.T. Shearer (2010) for New Destiny Mining Corp., which formed part of the current (2009) work program. A third mapping program, conducted by Dave Heberlein (1993B) for Minnova during 1992, was not filed. Various aspects of all three mapping programs are somewhat deficient, and the mapping from any one of the programs is at variance with other mapping of the same area. Presently, there is no reliable comprehensive geological map of the Le Mare property-area.

Mappers of Birkeland's (1991) field crew offered very little lithological description on their geological map. Rocks were described mostly andesite or rhyolite with or without brecciation. Structural determinations were sparse, and when compared with the results of later mapping, many bedding-attitude determinations were revealed to be orientations of dominant cleavages.

Mappers on Heberlein's (1993B) field crew had difficulty distinguishing felsic volcanic rocks from silicified andesite and basalt. Consequently, their map depicted an unrealistic amount of felsic volcanic rocks. Almost no structural symbols were recorded on Heberlein's (1993B) geological map.

Mapping during the current (2009) work program was conducted mostly in the South Gossan zone area to check the inconsistencies and the level of reliability of previous mapping in order to determine if a subsequent, more thorough mapping program would be advantageous. Time was not spent on detailed structural and alteration studies.

The Le Mare property hosts mostly mafic volcanic rocks of the Bonanza Supergroup, including autobreccias, lahars, and minor amounts of tuff and other pyroclastic beds. Rhyolitic rocks comprise a minor amount of the stratigraphy in the property-area. A thin rock unit previously identified as quartzite, was observed by the writer to be a pyritic, rhyolitic tuff. It may be one of the most useful stratigraphic marker beds in the property-area.

A 50-m (152.4-ft) thick "dyke", occupied by a rock described on Shearer's (2010) map as aplite, was located at the divide at the head of the Dumortiorite Creek valley just south of the phyllitic-argillic alteration of the South Gossan zone. It pre-dates the Dumortiorite Creek fault and could be coeval with the development of the Le Mare hydrothermal system.

Perelló et al. (1995) described three intrusive phases responsible for emplacement of the Island Copper Cluster deposits: an "early" rhyodacite porphyry associated with potassic alteration, an "inter-mineral" rhyodacite associated with sericite-clay-chlorite alteration and molybdenum deposition, and a barren, "lateminal" rhyodacitic porphyry. The spatial association of the "aplite" with the sericite-clay-chlorite (phyllitic-argillic) alteration and soil-molybdenum anomalies of the South Gossan zone indicates that it may be an equivalent of the "inter-mineral" or "late-mineral" rhyodacite identified at the Island Copper deposits (Figures 13 and 15).

A prominent topographic knob, located at U.T.M: 5,584,800 N., 578,850 E. about 400 m (1,312.4 ft) northwest of Dumortiorite Creek, was found to host a rhyodacitic dome that was described by Shearer (2010) as follows:

... Adjacent to the road (northwest of Dumortiorite Creek) is a bench-like ridge overlooking the west arm of Le Mare Lake where a resistant weathered dome-like feature was examined. An intensely siliceous, brittle, silicified hetero-volcanic breccia is cut by numerous quartz-chalcedony-like veinlets. The breccia fragments include angular banded, lapilli rhyolite, dacite and sub-rounded altered andesite. The dome-like breccia measures roughly 200 X 200 m (656 X 656 ft). Other than the numerous quartz-chalcedony veinlets, no alteration or sulphide minerals were noted.

Shearer, J.T.; 2010: p. 17.

This rock-unit may be a volcanic vent filling above a rhyodacite porphyry like those exposed at the Island Copper mine (Figures 13 and 15). It is curious that this dome is located adjacent to the South-Gossan zone, argillic-phyllic alteration plume.

Structure and Metamorphism

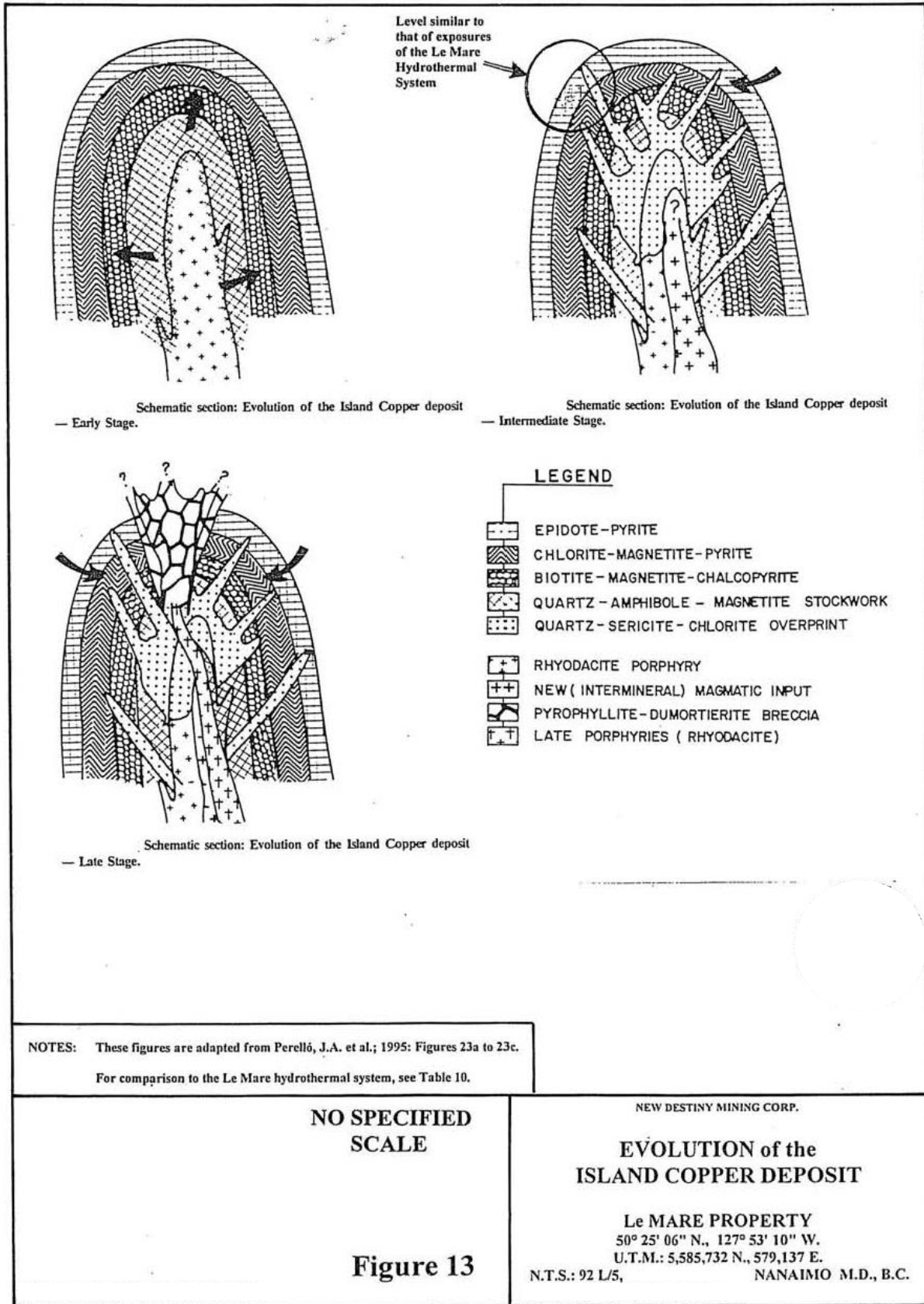
Regional mappers of the northern part of Vancouver Island have been in general agreement that folding of the Mesozoic and Cenozoic-age rocks exposed in that area has been minimal, and that block and transcurrent faulting have been the main mechanisms for stratigraphic displacement. J.E. Muller (1977) concluded that: Triassic-age rifting, westward tilting of the western part of Vancouver Island area during the Middle Jurassic-age Nassian orogeny, and eastward tilting east of the island's core area during the Late Cretaceous Laramide orogeny disrupted Vancouver Island stratigraphy into a series of tilted homoclines (Table 6). His conclusion has not been challenged.

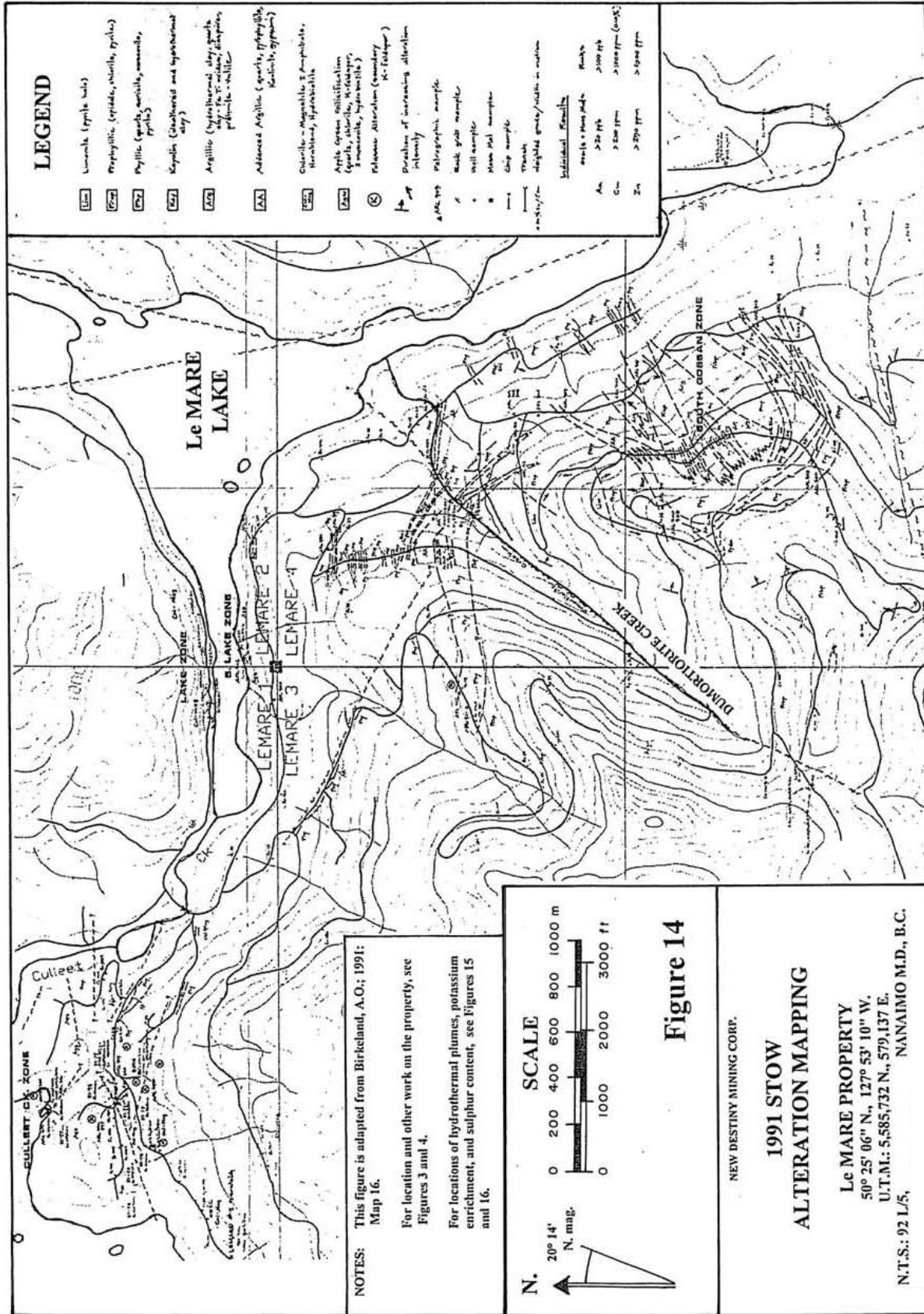
However, there is a structural complication in the Le Mare property-area. The mostly mafic volcanic stratigraphy near the hydrothermal system has been deformed into a series of open to close outcrop-scale folds that have a wide variety of axial-plane orientations. Development of this deformation before that of the Le Mare hydrothermal system and great diversity of fold axis orientations indicate that this deformation was related to local intrusion and not to regional deformation.

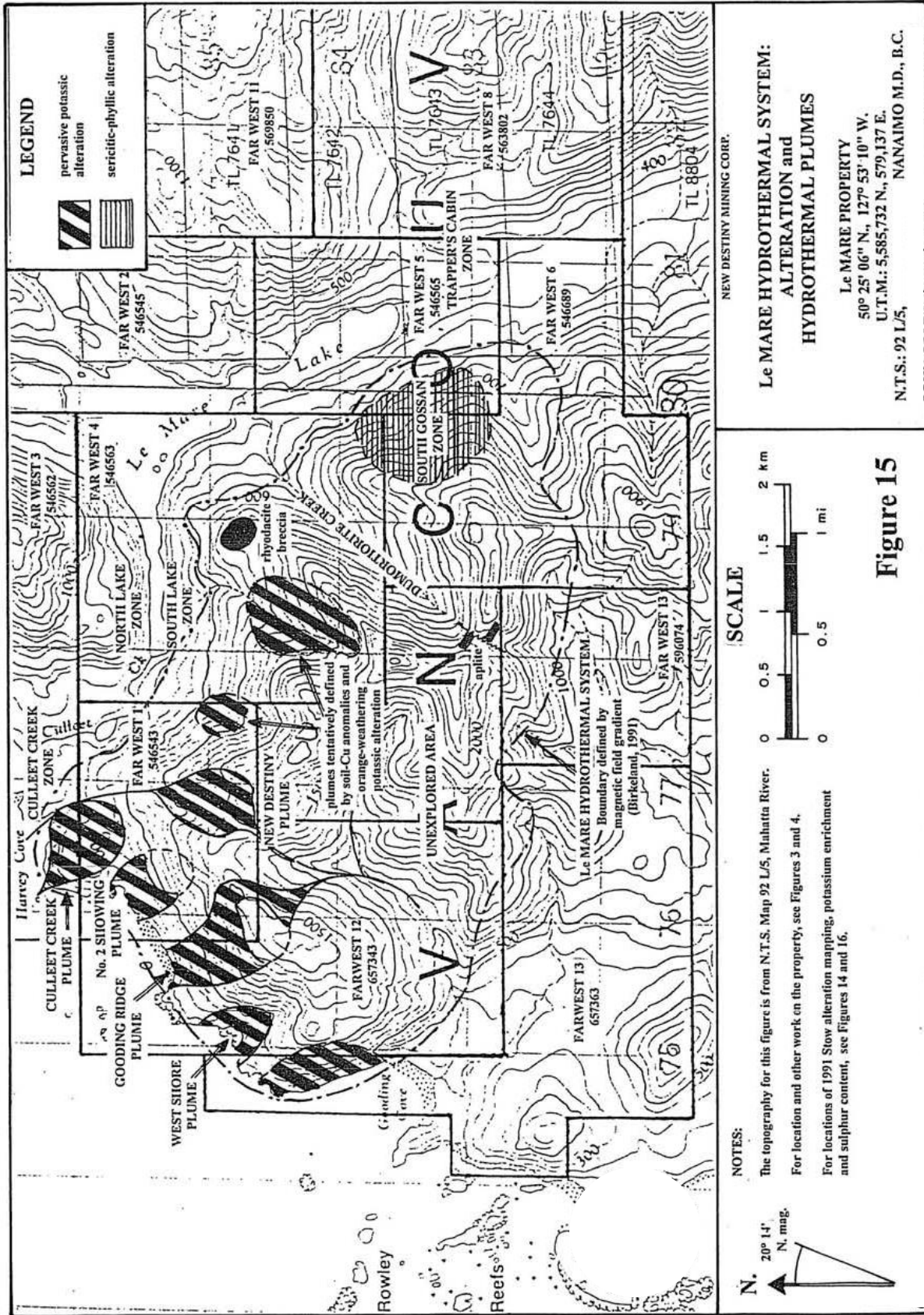
V.A. Preto (1979) concluded that such folding near the southern terminus of the Nicola batholith was related to emplacement of that intrusion. Similar folding mapped by the writer in mafic Nicola Group volcanics south of Merritt, British Columbia appeared to be the result of volcanic stratigraphy draping down over the margins of local plutonic cupolas. It was assumed that radial patterns of axial-plane orientations could be used to locate the locations of apices of such plutonic cupolas.

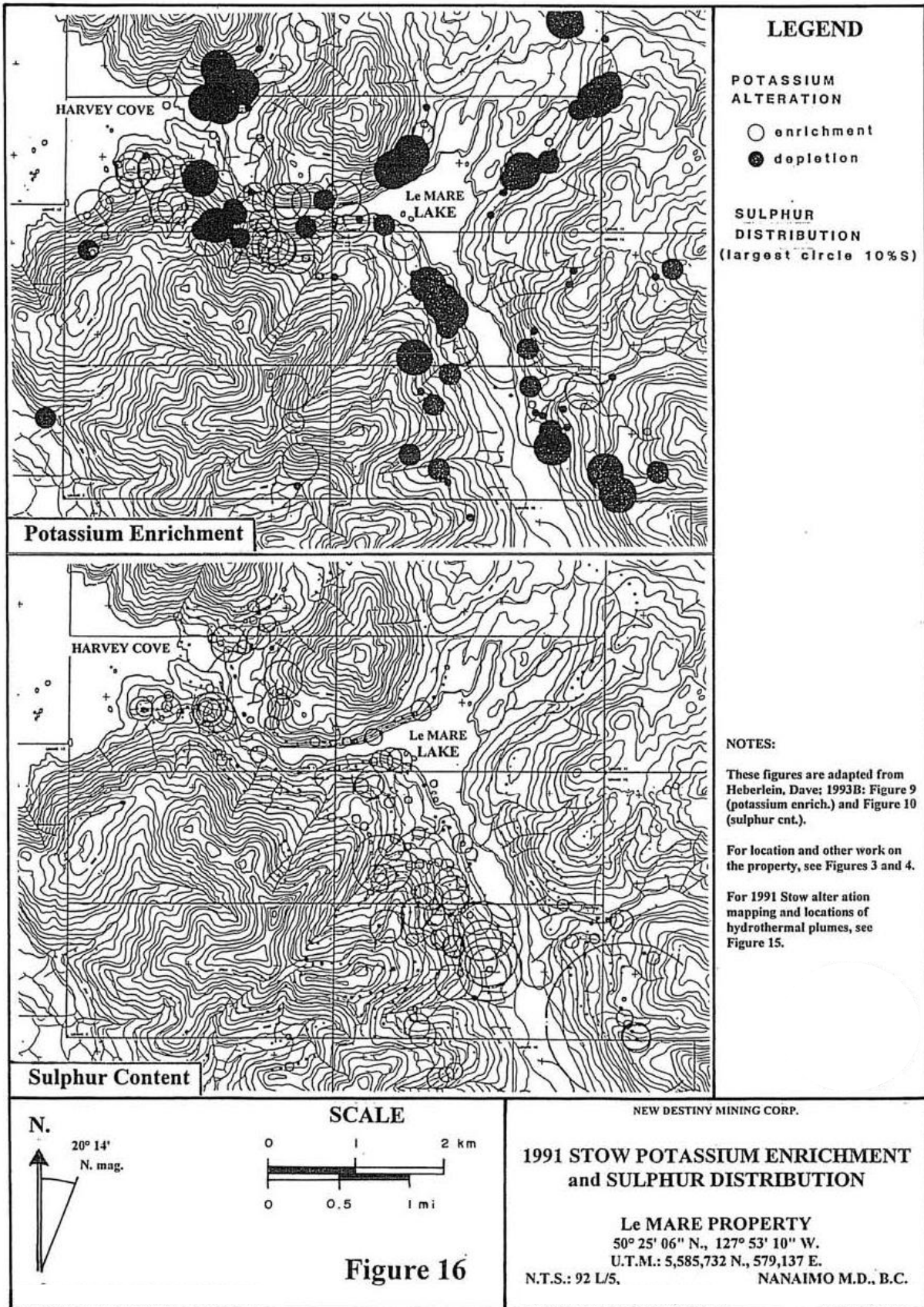
This style of folding indicates that the volcanic rock hosting the Le Mare hydrothermal system was buried at sufficient depth and sufficiently close to an intrusive contact for local heat, confining, and differential pressures to result in plastic, rather than brittle deformation. The existence of a near-surface contact of the volcanics with either of a coeval sub-volcanic intrusion or a rhyodacitic porphyry body is also supported by the exposure of the aplite rock unit at the head of Dumortiorite Creek.

Regional metamorphism around the Le Mare property-area does not exceed prehnite-pumpellyite or zeolite facies. It is difficult to discern around the Le Mare hydrothermal system due to pervasive, lower greenschist facies, thermal "contact" metamorphism that resulted in the formation of the axial plane cleavages in the drape-folds. Subsequently this was overprinted by pro-grade propylitic, potassic, and argillic-phyllic alteration. The writer presumes that this folding, thermal metamorphism, alteration and mineralization occurred during the Middle to Late Jurassic Period at about 175 million years ago, contemporaneous with development of the Island Copper Cluster deposits.









LEGEND

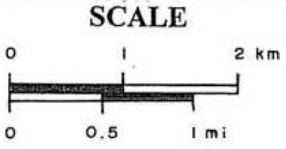
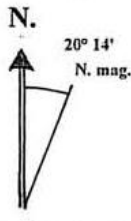
POTASSIUM ALTERATION
 ○ enrichment
 ● depletion

SULPHUR DISTRIBUTION
 (largest circle 10% S)

Potassium Enrichment

Sulphur Content

NOTES:
 These figures are adapted from Heberlein, Dave; 1993B: Figure 9 (potassium enrich.) and Figure 10 (sulphur cnt.).
 For location and other work on the property, see Figures 3 and 4.
 For 1991 Stow alteration mapping and locations of hydrothermal plumes, see Figure 15.



NEW DESTINY MINING CORP.
1991 STOW POTASSIUM ENRICHMENT and SULPHUR DISTRIBUTION
 Le MARE PROPERTY
 50° 25' 06" N., 127° 53' 10" W.
 U.T.M.: 5,585,732 N., 579,137 E.
 N.T.S.: 92 L/5, NANAIMO M.D., B.C.

Figure 16

MINERALIZATION

Copper

At the Le Mare hydrothermal system, copper mineralization is related to an early potassic alteration event; molybdenum enrichment is related to a later argillic-phyllitic event. High concentrations of copper and molybdenum occur together in significant amounts only where molybdenum enrichment has overprinted that of copper. The Le Mare hydrothermal system's potassic alteration zone has just been unroofed by erosion. At this level, copper mineralization occurs in discrete showings-areas located preferentially in the central parts of sub-vertical alteration plumes (previous) (Figure 15). Copper mineralization occurs mostly as chalcopyrite with minor amounts of bornite. In weathered rock, primary minerals are replaced by varying degrees by chalcocite, covellite, and black (copper-rich) limonite. In intensely weathered areas, sulphides have been oxidized to brick-red hematite and limonite; copper concentrations have been reduced to very low levels. This occurred above the Gooding Cove road in the Gooding Ridge plume where the writer's sample N4-1 contained 3 ppm copper and traces of molybdenum, gold and silver (Figure 15) (Table 9).

Of the five hydrothermal plumes located between Harvey and Gooding coves, the Culleet Creek plume is the only one that has been explored intensively during the early 1990s (Figures 4 and 15). A.O. Birkeland (1991) described copper mineralization of the Culleet Creek plume as follows:

Rocks in the vicinity of the Culleet Creek Zone exhibit a white weathering rind on surface (kaolinite after chlorite-K-spar). Numerous voids and boxwork textures with remnant secondary Cu mineralization is being leached by surface weathering and all values (concentrations) encountered near surface are likely depleted. This distinctive weathering characteristic (including chalcedonic quartz intergrowths) occurs over an area of approximately 500 m X 750 m (1,640.4 X 2,460.6 ft) (Figure 15). Two road borrow pits (Gorby and Boris showings) have fresher rock exposed in the pitwalls and roadfill debris. All rock types exposed in the pits are silicified and mineralized to various degrees. Modes of occurrences of copper mineralization are described as follows:

- chalcopyrite, chalcocite, minor bornite, covellite, and native copper in apple green silicified (AGS) zones
- associated with chalcedonic intergrowths, jasper and quartz veinlets and fractures, amygdules or disseminated in breccia matrix overprinting all rock types
- disseminated chalcopyrite in lesser silicified dark green chloritized volcanics

The 500 m X 750 m (1,640.4 X 2,460.6 ft.) alteration zone of AGS has been trenched with 8 plugger and blast hole trenches.

Birkeland, A.O.; 1991: p. 13.

Within all of the hydrothermal plumes examined by the writer, the early phase of potassic alteration comprises veinlets and disseminations of predominantly orthoclase, minor quartz, and sparse red-brown biotite which hosts chalcopyrite, with small amounts of bornite associated with pyrite, commonly with a chalcopyrite:pyrite ratio greater than 2:1. Orthoclase-rich, alteration passes gradually to a distal phase of silicification which, as A.O. Birkeland (1991) correctly observed, was accompanied by a gradual decrease to low copper concentrations with chalcopyrite being the only significant copper-bearing sulphide.

Orthoclase-quartz alteration is post-dated by quartz-jasper veinlets, pods, and disseminations that host vein-segregations and disseminations of chalcopyrite, bornite, and pyrite. These look similar to, but can be seen to cross-cut earlier orthoclase-quartz related mineralization in fresh rock at the Gorby showing. Generally, copper mineralization seems to be more abundant in quartz-jasper alteration than in the preceding orthoclase-quartz alteration.

Tabulated averages of Birkeland's (1991) sampling results weighted per linear metre, from the eight plugger and blast-hole trenches that Birkeland mentioned (previous quote). Grab samples were excluded. That tabulation is as follows:

Table 6
Results of Birkeland's 1991 Sampling in the Culleet Creek Zone
Weighted per Metre of Sampling

Location	Analysis Number Sequence	Total Sampling Length metres	feet	Copper ppm	Molybdenum ppm	Gold ppb	Silver ppm	Zinc ppm
Harvey Cove showing	125229-37 131488-500	22.0	72.2	1043	<2	<6	<0.4	102
Gorby showing	125357-61 125383-90 125403-07 131451-53	30.5	100.1	315	<1	<5	<0.2	84
Boris showing	125391-99	9.0	29.5	1134	<1	<5	0.5	30
91-T2	131457-61	5.0	16.4	93	<1	<5	<0.2	102
91-T3	131462-67	6.45	21.2	2665	4	<5	<0.4	70
91-T4	131468-70	3.0	9.8	660	<1.7	77	<0.3	77
91-T5	131471-73	3.0	9.8	577	3	17	<0.2	144
91-T6	131474-78	5.0	16.4	170	<1	<7	<0.2	167
91-T7	131479-83	4.8	15.7	687	<2.8	29	<0.2	50
91-T8	131484-87	4.3	14.1	133	<1	<5	<0.2	63
Average/m of Culleet Creek zone sampling		93.05	305.3	740	<1.5	<8.9	<4.7	87

NOTES: This table is produced from the data of A.O. Birkeland, A.O., 1991.
1991 grab samples have been excluded from this tabulation.
For locations of sampled areas, see Figures 4 and 6.

Average copper concentrations from the 1991 Stow Resources trenches varied from a low of 133 ppm to a high of 2,665 ppm (Table 8). Such variance is intrinsic to discontinuous copper mineralization near the top of the potassic alteration zone of any calc-alkalic porphyry system. A discussion of sample variance comprises part of section 7.0 of this report.

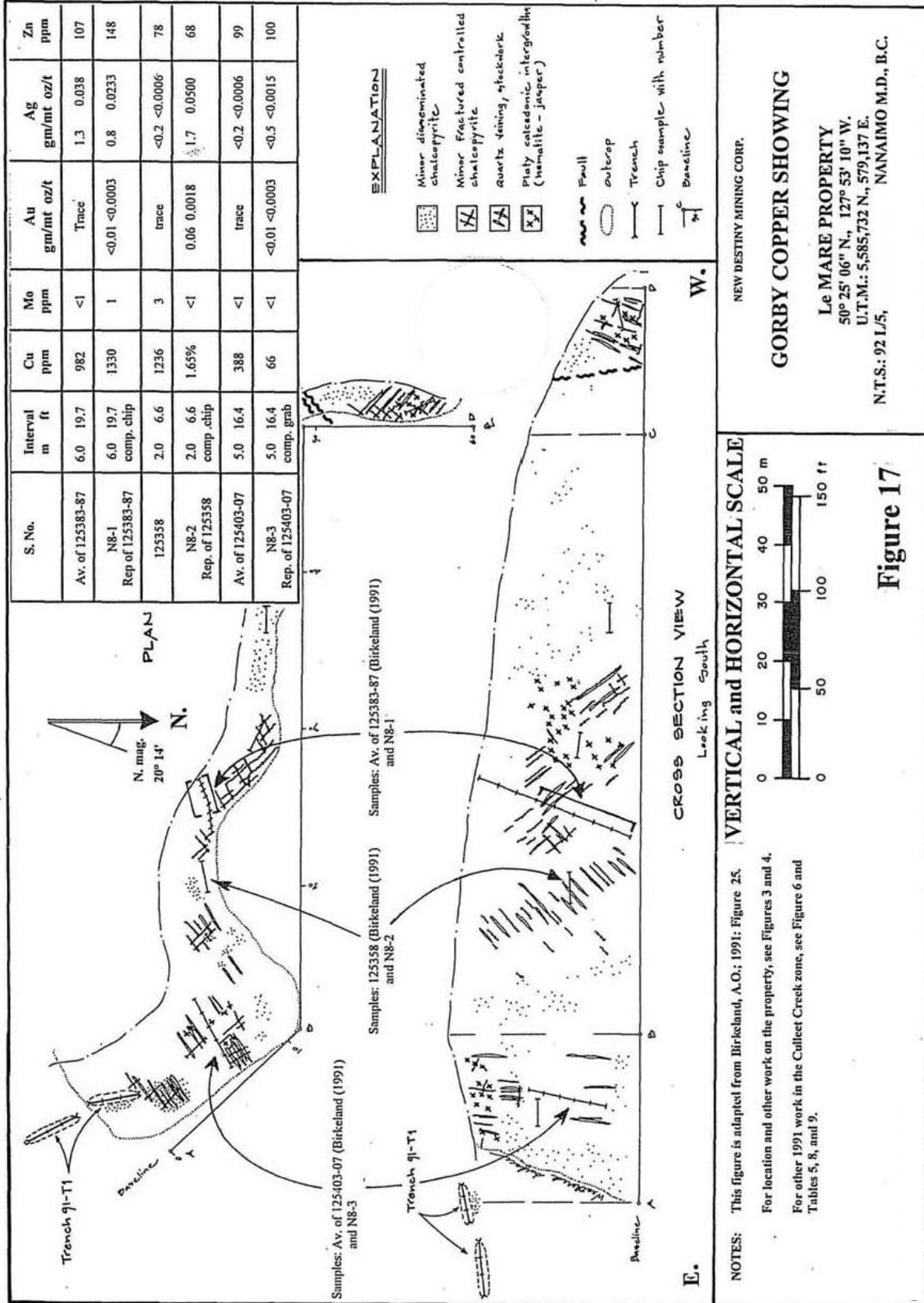
The Gorby occurrence is located on a spur road about 80 m (262 ft) north of the Gooding Cove road in the southern boundary-area of the FAR WEST 3 (546562) claim (Figures 3, 4, and 6). It is near the geographic centre of the Culleet Creek plume and hosts the most extensive exposure of fresh, mineralized rock in the plume. A road borrow pit was extended into a 50-m (164-ft) long side-hill cut during the 1991 Stow Resources program (Figures 4, 6, and 17). Although Birkeland (1991) did not describe specifically the mineralization at the Gorby showing, his comments regarding copper mineralization in fresh rock of the Culleet Creek zone match what the writer observed in the cut itself.

J.T. Shearer (2010) added to a description of the Gorby showing as follows:

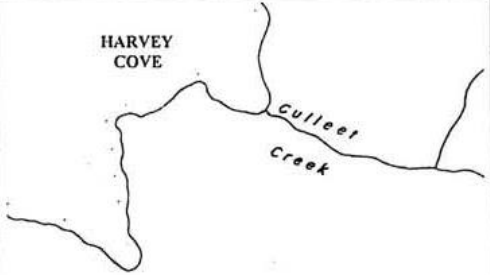
The Boris and Gorby copper showings were briefly examined and are well documented by (Birkeland, 1991) ... One of the key differences the writer noted at the Gorby showing was the increase (greater intensity) in quartz (and lesser calcite) veining hosted in the andesite (at that location). This was not observed in other andesitic rocks mapped - although minor (<0.05%) free chalcopyrite was occasionally noted. Also at the Gorby, quartz-filled stretched amygdaloidal andesitic flows are associated with disseminated chalcopyrite ...

J.T. Shearer; 2010: p. 17.

One of the 1992 Minnova Inc. diamond drill holes, No. 92- 676-2, penetrated the Culleet Creek potassic alteration plume at a location about 50 m (164 ft) east of the centre of the Gorby cut (Figure 6, Table 5). That hole went through five 2-m (6.56-ft) and one 4.7-m (15.4-ft) long intersections that contained from 500 to 959 ppm copper. Those copper concentrations were similar to many of the average concentrations that the writer calculated from Birkeland's (1991) trench sampling results (Tables 5 and 8), indicating that the discontinuous style of copper mineralization recorded from the trenches is present to a depth of at least 200m (656 ft). Sparse copper mineralization to depth in this plume may be related in part to its location at the outer edge of the hydrothermal system (Figure 15).



S. No.	Interval m ft	Cu ppm	Mo ppm	Au gm/mt oz/t	Ag gm/mt oz/t	Zn ppm
SA 004	10 32.8 composite chip?	982	<1	Trace	1.3 0.038	107
N2-1	composite chip	61	<1	<0.01 <0.0003	<0.5 <0.0015	106
N3-1	composite chip	320	<1	<0.01 <0.0003	0.6 0.0175	75
N10-1	composite chip	48	<1	<0.01 <0.0003	<0.5 <0.0015	338
N10-2	10 32.8 composite chip	898	4	<0.01 <0.0003	1.2 0.0350	154
N10-3	composite grab	6.57%	<1	0.05 0.0015	6.2 0.1809	74

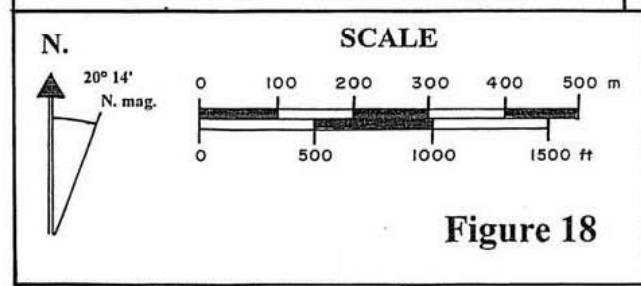
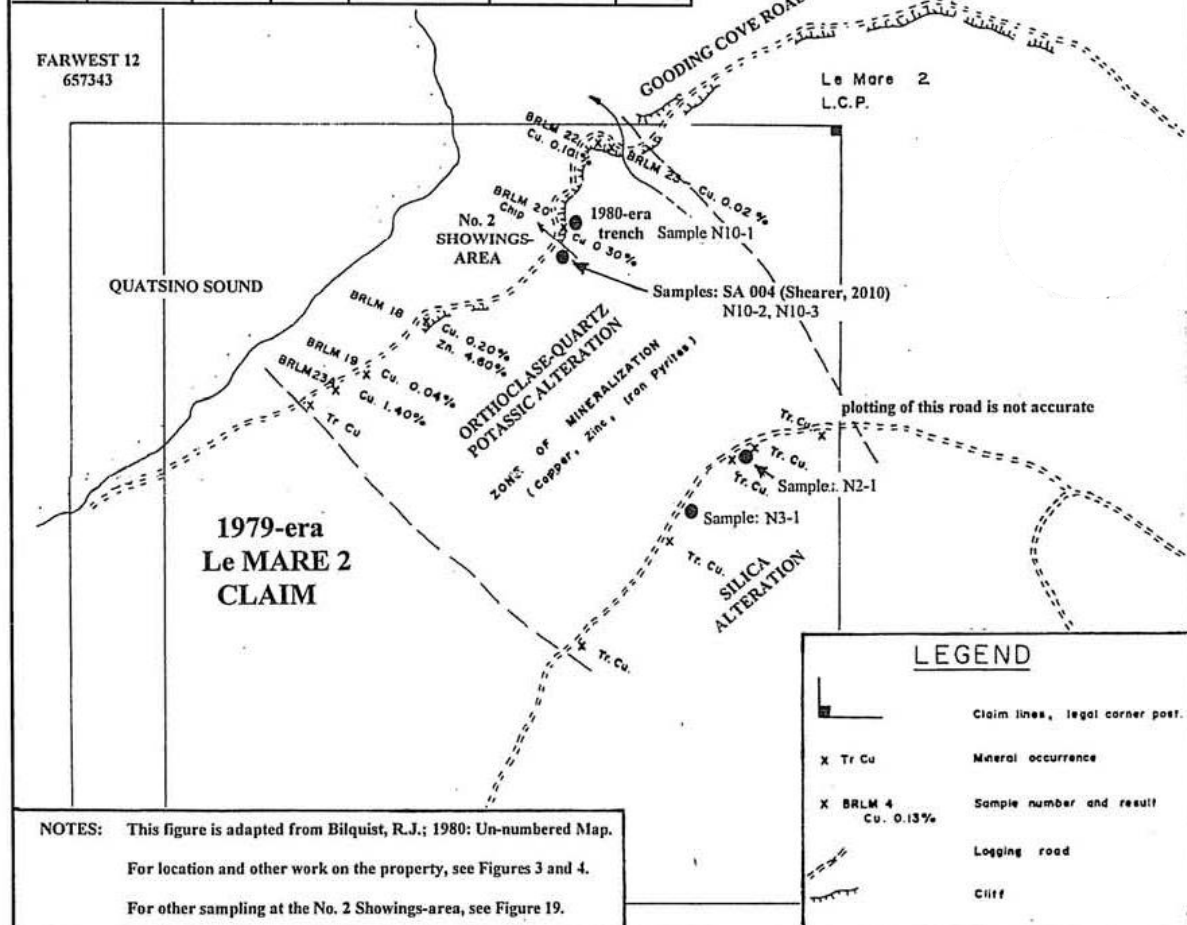


HARVEY COVE

CULLEET CREEK

FAR WEST 3
546562

FAR WEST 1
546543



NEW DESTINY MINING CORP.

**No. 2 COPPER SHOWINGS-AREA:
1980 BRINCO PROSPECTING**

Le MARE PROPERTY
50° 25' 06" N., 127° 53' 10" W.
U.T.M.: 5,585,732 N., 579,137 E.
N.T.S.: 92 L/5, NANAIMO M.D., B.C.

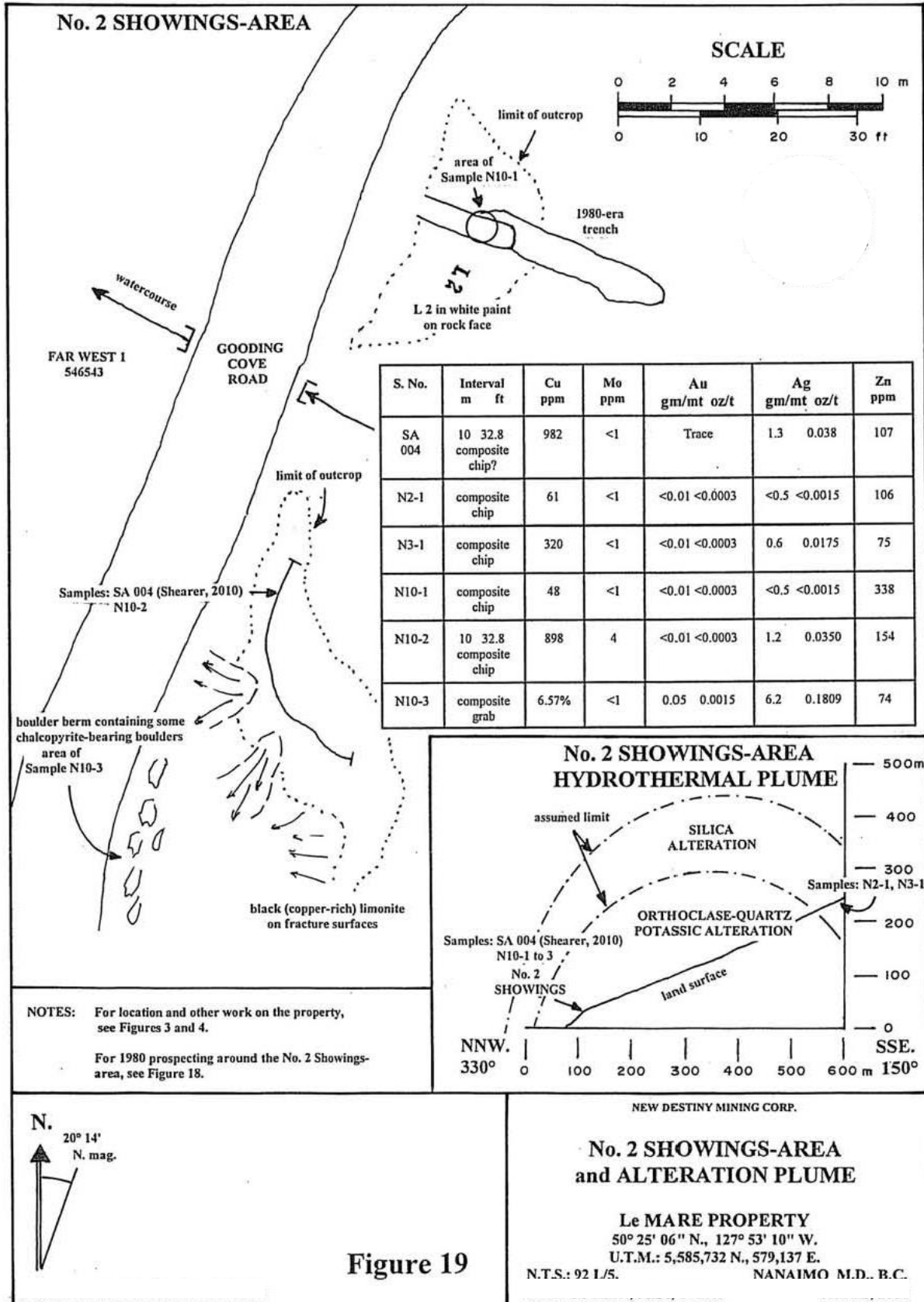


Figure 19

The No. 2 showings-area is located on the up-hill side of the Gooding Cove road in the northwestern part of the FAR WEST 1 (546543) claim (Figures 3 and 4). It is in the northwestern part of the potassic alteration zone of the No. 2 Showings-area plume (Figure 15).

During 1980, British Newfoundland Exploration Ltd. (BRINCO) conducted a prospecting program on the Le Mare No. 1 (later known as the North Lake zone) and the Le Mare No. 2 showings-areas (Figure 4). R.J. Bilquist (1980) recorded the results of BRINCO's work on the No. 2 Showings-area as follows:

Prospecting on the LE MARE NO.2 mineral claim resulted in the discovery of a zone of mineralization. This zone was traced along the road cut a distance of 600 m (1,968.5 ft) (Figure 18). The mineralization found included chalcopyrite, malachite, azurite, sphalerite, and iron pyrites. Mineralization appears to be related to faults and fractures and in places it is abruptly cut off at the boundaries of these. The rock appears to be mainly andesite flows and tuffs cut by an occasional andesite dike. Near chip sample BRLM 20, secondary potassium feldspar was seen as fracture fillings. Samples from here assayed between 0.20% and 1.40%. The copper mineralization was noted in another parallel road approximately 400 m (1,312.3 ft) to the southeast. No samples from here were assayed but from visual examinations it is assumed that the values (concentrations) would be similar.

Bilquist, R.J.; 1980: p. 6.

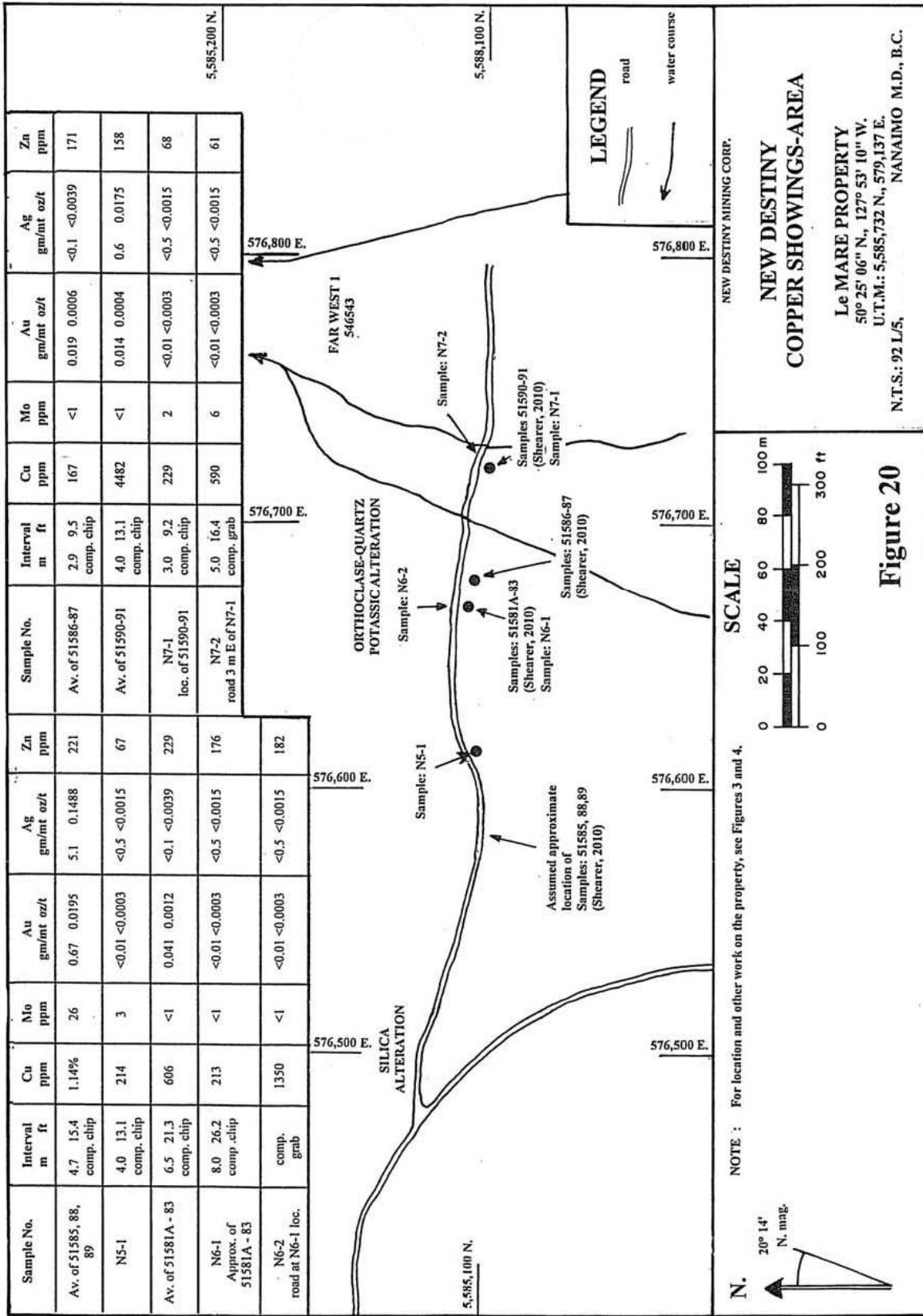
Bilquist's (1980) description of a 600-m (1,968.5 ft) section of the Gooding Cove road has been copied in various versions by subsequent writers who all have ascribed that description to the No. 2 showings-area itself. The location of Bilquist's sample No. BRLM 20 is the actual location of the showings-area (Figures 18 and 19).

Recognition of "secondary potassium feldspar" at the BRLM 20 sample site may have encouraged the BRINCO prospectors to work the slope above the No. 2 (BRLM 20) showings-area along a road where the writer took samples N2-1 and N3-1 (Figures 18 and 19) (Table 9). Although R.J. Bilquist (1980) did not report the presence of a hydrothermal plum, he did outline the potassic alteration zone of the No. 2 Showings-area plume fairly accurately (Figures 15 and 18).

Later, when A.O. Birkeland (1991) developed his theory of the presence of a linear, 6-km (3.7-mi) long, mineralized belt from Harvey Cove to the southeastern end of Le Mare Lake, he seems to have ignored the 1980 BRINCO discoveries southwest of his preferred zone. Being outside Birkeland's trend, the No. 2 showings-area and Gooding Ridge was explored little prior to the current (2009) exploration program.

At the No. 2 showing itself, there is an old trench dug into chloritic andesite hosting orthoclase-quartz and quartz-jasper (potassic) alteration similar to that in the Gorby cut. Any significant copper mineralization has been removed from the trench (Figure 19). A few metres southwest of the trench, the letters L 2 have been painted in white on the rock face, probably by either the 1991 or 1992 field crew.

There is a relatively new cut into the rock for road material about 18 m (59 ft) south of the trench along the road. Sample SA 004 was taken along a variably weathered 10-m (32.8-ft) long section near the base of the rock face (Figure 19). Composite chip samples contained 982 ppm and 898 ppm copper respectively. A boulder berm, dug out of the cut and left beside the road in front of it, comprised clasts of andesitic flow and flow-breccia rock. Comparatively 'permeable' parts of the matrix in several flow-breccia boulders contained blebs of chalcopyrite up to 5 cm (about 2 inches) in length. Grab sample N10-3 containing 6.57% copper, was comprised of pieces hammered off some of the boulders. On the rock face itself, very little chalcopyrite remained. Most of it had been replaced by black copper-rich limonite.



The New Destiny showings-area is near the western end of Le Mare Ridge in the south-central part of the FAR WEST 1 (546543) claim (Figures 3 and 4) in the potassic alteration zone of the plume (Figure 15). Dave Pawliuk, a member of the 2009 field crew and a former owner of the 1997-era LEM claims (Figure 4), discovered the showings on December 5, 2009 (Figure 20). J.T. Shearer's account of them was as follows:

The New Destiny Copper Zone ... is exposed along a new logging road hosted by rhyodacite and andesite with pervasive chlorite and hematite, locally brecciated alteration. Mineralization consists of up to 2% chalcopyrite and pyrite (Figure 20) ... , sample 51585, a chip over 0.9 m (3.0 ft) assayed 2.34% Cu, 1.97 g/tonne (0.057 oz/ton) gold and 9.0 g/tonne (0.26 oz/ton) silver.

Shearer, J.T.; 2010: p. 22.

The western part of the showings area hosts intensely chloritized and silicified dacitic rock near the base of a Tertiary-age weathering profile. This rock contains significant amounts of chalcopyrite and pyrite that have been partly weathered to hematite and limonite. D.J. Pawliuk's samples: 51585, 51588 and 51589, taken from felsic volcanic rocks near the western end of the showings-area contained an average of 1.14% copper (Figure 20). This high concentration may have been due in part to local copper concentration in "permeable" areas.

Rock with blebs of massive chalcopyrite-pyrite-bornite mineralization were sampled by the writer farther east in the showings-area (subsequent).

East of the dacite is medium-green silicified mafic andesite or basalt with sparse to moderately intense orthoclase-quartz alteration. Black (copper-rich) limonite and traces of azurite and malachite occur on fracture surfaces. Rusty blebs throughout this rock may be the result of weathering of pyrite and chalcopyrite to hematite and limonite. Sparsely disseminated chalcopyrite is present on fresh surfaces.

Averages of D.J. Pawliuk's samples 51581A to 51583 and 51590 and 51591, from about the same locations as samples N6-1 and N7-1, contained 606 and 4482 ppm respectively (Figure 20). Like at the other copper showings that the writer has examined in this part of the Le Mare hydrothermal system, there is great local variability in copper concentrations. The molybdenum content of samples from the New Destiny showings-area is low (Figures 17 to 20) (Table 8).

Potassic alteration and accompanying copper mineralization have been overprinted by argillic-phyllic alteration in the South Gossan zone, and in a small area around the Mo Road showing west of Culleet Lake. Locally, along discrete fracture systems in the South Gossan zone, argillic-phyllic alteration is in turn, overprinted by minor amounts of advanced argillic alteration. The effects of the overprinting alteration events have been to liberate copper deposited during the previous potassic alteration event and to redistribute it, probably upward, to rock that has now been eroded away. This is indicated by the lack of distinct soil-copper anomalies in the South Gossan zone (Figure 21E). J.T. Shearer (2010) summarized copper occurrence in and around the South Gossan zone as follows:

South Gossan Zone (SGZ)

Copper mineralization flanks the (argillic, phyllic, and advanced argillic) alteration zones occurs in volcanic wallrocks. Modes of occurrences are described as follows:

- Massive fine-grained chalcopyrite and bornite/chalcocite/covellite (may be Au bearing) veinlets and fractures radiating out from beneath the northeast plunge beneath the advanced alteration cap.
- Disseminated fine grained chalcopyrite associated with black chlorite-magnetite-hydrobiotite? in mafic volcanic (transitional potassic-phyllic "mafic porphyry") alteration.
- East of the SGZ and across the Le Mare Lake valley (Trapper Cabin area) (Figure 4) are fault controlled chalcopyrite and bornite occurrences in siliceous pyritic volcanics.

- To the west of the SGZ and in the headwaters of “Dumortiorite Creek”, carbonate veins up to .3m (1 ft) in width occur in propylitic alteration envelopes. The veins have been traced for a strike length of up to 15m (49.2 ft).

Shearer, J.T.; 2010: p. 18.

Copper-bearing veins radiating out from subsequent alteration could describe re-mobilized copper that has been flushed outward from the sloping margin of a vertically zoned argillic-phyllitic alteration plume. Shearer’s description of chalcopyrite in association with “transitional potassic-phyllitic” alteration could be a manifestation of local partial overprinting of early potassic by subsequent argillic-phyllitic alteration as mentioned previously.

Molybdenum

The distribution of molybdenum enrichment related to the Le Mare hydrothermal system has been defined mostly by soil-molybdenum anomalies. All of the largest and most intense of these anomalies are spatially associated with quartz-sericite-pyrite (phyllitic) alteration lower eastern flank of the argillic-phyllitic alteration plume (Figure 22E). Molybdenum enrichment is conspicuously absent in the copper showings that are related to potassic alteration. Discussions of the relevance of molybdenum enrichment to alteration and the distribution of soil-molybdenum anomalies form part of sections 4.2.1 and 5.3.2 of this report.

A small, roadside outcrop hosting visible molybdenite was located by the 1991 Stow mapping crew (Birkeland, 1991) on the main Gooding Cove road southwest of Culleet Lake. It was described as follows:

At the Mo Road showing to the west of Culleet Lake, sparse chalcopyrite and molybdenite mineralization has been noted in the road cut associated with advanced argillic and phyllic alteration ...

Birkeland, A.O.; 1991: p. 14.

The outcrop was less than 5 m (16.4 ft) long and was composed of white to yellow sericite with subsequent and veinlets and disseminations of clay and a white chalky mineral that Shearer identified as geyselite. Traces of fine-grained molybdenite and possibly chalcopyrite were disseminated throughout the rock.

The Mo Road outcrop is located at about U.T.M. co-ordinates: 5,585,884 N., 577,209 E. (50° 25' 12" N., 127°54' 47"W.) on the FAR WEST 1 (546543) claim. It is within a small area of phyllic alteration between the road and Culleet Lake (Figure 14). The most important aspect of this outcrop is that, as at the flank of the argillic-phyllitic alteration plume in the South Gossan zone and at the Island Copper mine deposit, molybdenite mineralization is demonstrated to be intimately associated with phyllic alteration in outcrop (sections 4.2.3 and 5.3.2, this report) (Table 10).

Comparison of the Island Copper and Le Mare Hydrothermal Systems

The Island Copper mine deposit covered an elongate 1,750 X 480 m (5,741 X1,575 ft) oval-shaped area. From the mine’s opening until 1994, a total of 345 million tonnes (380 million tons) of ore having average head grades of 0.41% copper, 0.017% molybdenum, 0.19 gm/mt (0.006 oz/ton) gold and 1.4 gm/mt (0.041 oz/ton) silver were produced.

The Le Mare hydrothermal system is exposed in an oval-shaped area with axes measuring about 5,000 X 3,000 m (16,404 X 9,843 ft). There is more than sufficient area within the exposed boundaries of the Le Mare hydrothermal system to accommodate a pit the size of the one required at the Island Copper mine deposit.

Many aspects of the Le Mare hydrothermal system are quite similar to those of the Island Copper mine deposit. Similarities and differences between the two systems are tabulated by the writer as follows:

Table 7
Comparison of the Island Copper and Le Mare Hydrothermal Systems

Aspect	Island Copper Hydrothermal System	Le Mare Hydrothermal System
Mineral occurrence class	Calc-alkalic porphyry Cu-Au-Mo	Calc-alkalic porphyry Cu-Mo (Au potential is not assessed)
Age	175 m.y - Middle Jurassic Period Aaelnian-Bajocian Stage	175 m.y - Middle Jurassic Period Aaelnian-Bajocian Stage
Host rocks	Bonanza Supergroup mafic to intermediate meta- volcanic and associated meta-sedimentary rocks	Bonanza Supergroup mafic to intermediate meta-volcanic and associated meta-sedimentary rocks
Controlling structures	End Creek Fault: west-northwest trending, right-lateral, sub-vertical, regional fault	proposed west-northwest trending, right lateral, sub-vertical, regional fault
Local structures	block faults, minor folds	block faults, drape folds
Localization	dilational jog along the regional structure	proposed dilational jog along a regional structure
Alteration	Early Potassic and Pro-grade Propylitic: 1. Inner potassic: qtz-actinolite-hb-Na.plag- +/- scapolite-apatite (low Cu + Mo contents) 2. Outer potassic: bio-mag-albite-kspar +/- amphiboles (>0.2% Cu) 3+4. Propylitic: chlorite-calcite-epidote-pyrite 3. (<0.3% Cu) 4. (<0.1% Cu) Intermediate phyllic-argillic: sericite kaolinite-illite-chlorite +/- pyrite (Mo and minor Cu mineralization) Late Advanced Argillic: (hosted in pyrophyllite-dumortiorite breccia) pyroph-qtz-sericite-kaoliniteclays-dumortiorite	Early Potassic plumes surrounded by Pro-grade Propylitic 1. Potassic plume: core of kspar-qtz +/-bio intruded by qtz-jasper all contained in silicic envelope (Cu showings in core areas) 2. Outer propylitic: chlorite-calciteepidote-pyrite (low Cu) Intermediate phyllic-argillic: sericitekaolinite-clays-chlorite at the South Gossan zone (asst. with soil-Mo anomalies) Late advanced argillic: (restricted to a few permeable faults) sericite-kaolinite-clays
Intrusion	1. Early mineral rhyodacite (altered and associated with potassic alt and most Cu mineralization) 2. Intra-mineral rhyodacite (altered and asst with most Mo and minor Cu mineralization) 3. Late-mineral rhyodacite (unaltered) and pyrophyllite breccia (post-mineral)	1. Rhyodacite breccia at Culleet Creek zone with qtz-jasper (late potassic) alteration 2. Altered + unaltered felsic dykes in the South Gossan zone 3. Rhyodacite northwest of Dumortiorite Creek- Unaltered aplite at the head of Dumortiorite Creek
Mineralization	1. Early Cu-Au+/-Mo asst with kalt 2. Late Mo-Cu+/-Au asst with argillicphyllic alt	1. Cu showings + soil anomalies asst with kalt 2. Mo Road showing and soil anomalies asst with phyllic alt

NOTE: Au = gold, Cu = copper, Mo = molybdenum, bio = biotite, hb = hornblende, kspar = potassium feldspar, mag = magnetite, plag = plagioclase feldspar, qtz = quartz, alt = alteration, kalt = potassic alteration, m.y. = millions of years ago.

The deposits of the Island Copper Cluster differ from typical calc-alkalic porphyry copper-molybdenum deposits in that, for the most part, they have gold contents similar to those of alkalic porphyry copper-gold deposits (Perelló et al., 1995).

PREVIOUS WORK 2011

Background

In 2011 a series of mapping surveys were completed in the Le Mare Lake area focusing in on an area located along the western section of the Farwest claim group. More specifically, in an area roughly bounded by: west of Le Mare Lake, south of Culleet Creek and east of Gooding Cove with surveys extending from near tide water to summit of 450 meters.

Previous geophysical VLF-EM surveys (1992) and soil geochemical surveys (2009) conducted in this area have outlined copper (gold) anomalous targets. A VLF-EM conductive signature was outlined along a northeast trending ridge (summit elev. 488 m) which is coincidental with a geochemical gold high. Three separate copper soil anomalies were outlined from the 2009 surveys. One of these anomalies is coincidental and responsible for the New Destiny copper zone discovered by backhoe trenching during March-April exploration in 2011.

Access to the mapping project site was via the Restless Creek mainline logging and branch roads. For mapping control, the author utilized a PC mapping tool referred to as the Yuma Tablet PC. The mapping tool has Windows 7 Professional operating system and installed with Microsoft Office Suite. For mapping, geospatial Arc GIS software was installed and a 1:20,000 scale topographical map, supplied by the logging company detailing all natural and man-made features, was uploaded. All geological rock outcrops encountered were entered into the Yuma, generally at 1:5000 scale. A hand-held GPS Garmin model was also utilized as backup for field mapping plus, a 1:20,000 hard copy base map was used to manually plot all the outcrops mapped. This traditional procedure was a precautionary measure taken as the author was a first time user of the Yuma. Once proficient, the author found the Yuma quite time saving.

The Yuma Tablet along with the base map were submitted to Coastal Resource Mapping Ltd., GIS specialists, where the mapping data was downloaded and computer generated geological maps produced for future field mapping surveys and updates.

Historical exploration surveys along the south end of the southeast arm of Le Mare have outlined hydrothermal alteration signatures related to porphyry mineral environment. Subsequent geochemical soil surveys have delineated a coincident copper-molybdenum anomaly, referred to as the 'South Gossan zone', that supports a porphyry type model. The copper mineralization (e.g. Gorby, New Destiny and other related showings) found in the area mapped noted-above (see Figure 1), is currently viewed in the technical report (J. Ostler, P.Geol., April 30, 2010) as at least 6 distinct 'hydrothermal-plume' copper-potential hosted systems and interpreted as been hosted in 'dilatational jog' (pull-apart structure) similar to the Island Copper cluster deposits.

Based on the mapping surveys and empirical field data presented in this summary report, the author herein presents an argument that supports evidence for the potential of an epithermal and or a volcanogenic-type, massive sulphide environment – a long side the porphyry copper model discussed in the technical report. Although no massive sulphide mineralization (e.g. float, etc.) as yet has been documented (to the author's knowledge), however the proxy to such potential mineralization can be found in the rocks mapped and interpreted as discussed below.

The regional tectonostratigraphic framework is represented by the northwest trending, Early to Middle Jurassic Bonanza volcanic arc. The Bonanza arc, evolved as part of the upper stratigraphic Bonanza Group, in a convergent-margin setting, built on basement comprising distinctive mid-Paleozoic arc volcanic rocks of Sicker and Buttle Lake groups and the Late Triassic Vancouver Group which includes, tholeiitic flood basalts of the Kurmutsen Formation and Quatsino (carbonate) Formation. Resurgence of arc magmatism in Early Jurassic time gave rise to the Bonanza arc. The arc was thought to have developed in response to eastward-directed subduction of Pacific Ocean lithosphere during Early to Middle Jurassic times.

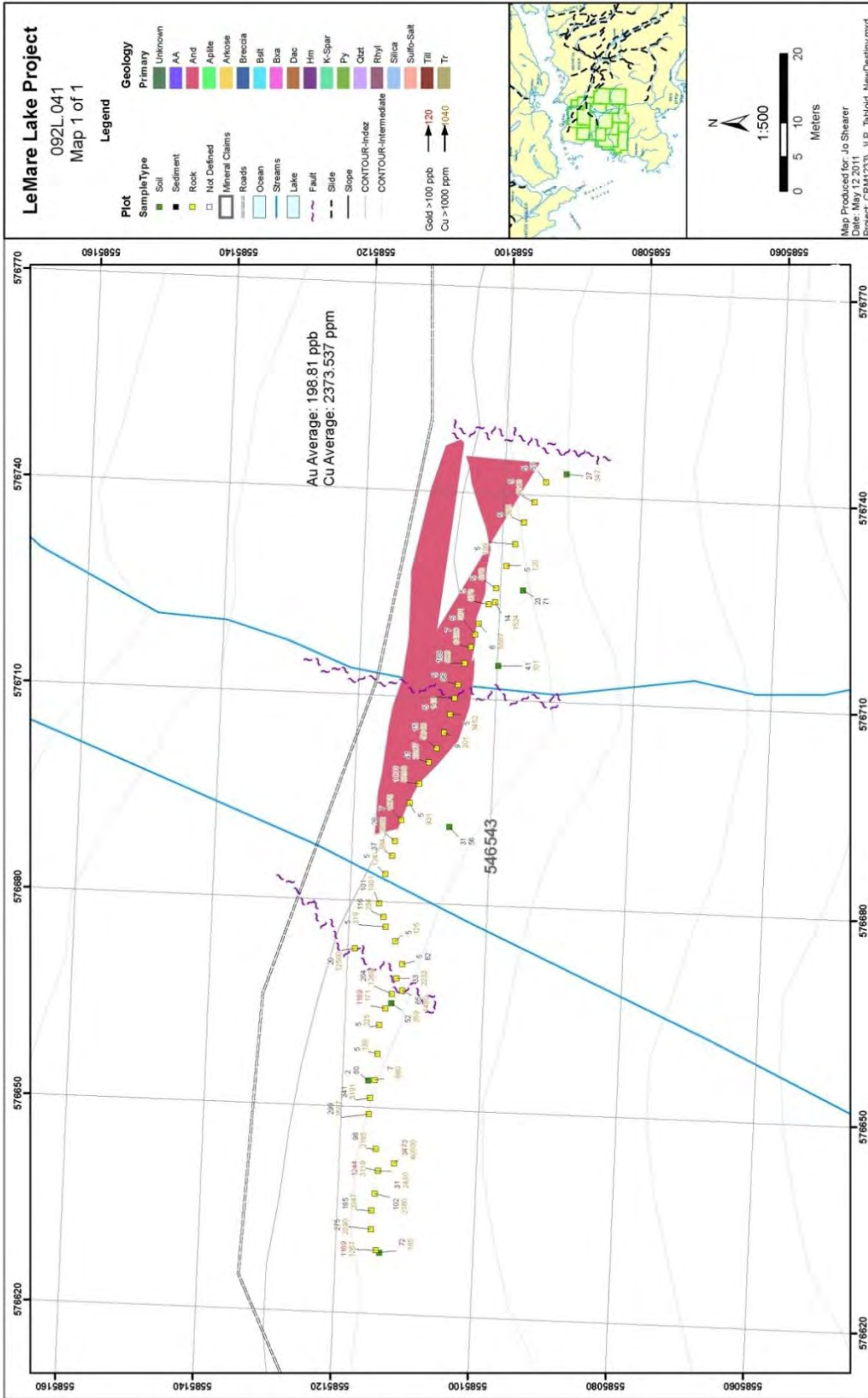


FIGURE 20B

The Bonanza Group also forms part of the Insular Belt, a morphological belt that defines the southern portion of the Wrangellia Terrane, underlying Vancouver Island and coastal sections of British Columbia. The accretion of the Wrangellia to the inboard terranes of the Coast and Intermontane Belts occurred as late as mid-Cretaceous or as early as Middle Jurassic. This accretionary deformational event resulted in the development of the Middle Jurassic Nassian Orogeny manifested by regional lower greenschist facies metamorphism, transpressional and tensional faulting and shallow west dipping monoclines. This was followed by uplift, erosion and unroofing caused by the Late Jurassic to Late Cretaceous Columbia Orogeny, with deposition of detritus derived from the Bonanza Group volcanic rocks to form Middle Jurassic and Early Cretaceous conglomerates.

The Bonanza Group is comprised of: Late Triassic, platformal carbonates and siliclastic sedimentary rocks of the Parson Bay Formation; upper Late Triassic to Early Jurassic volcanic-sedimentary unit, suggested to be a nascent arc and; Early to Middle Jurassic Le Mare Lake mature arc volcanics.

The Le Mare Lake volcanics constitute thick sequences of intercalated volcanic and marine sedimentary strata and mark episode of regionally extensive subaerial volcanism in the Bonanza Group. Wide variety of rock types in the Le Mare Lake proper include: black to grey-green or reddish grey, aphanitic to plagioclase-phyric, amygdaloidal flows of basaltic to andesitic composition; siliceous apple green to purplish banded andesite; grey to creamy pink and pale buff dacitic to rhyolitic flows; rhyodacitic to rhyolitic pyro-megaclastic flows and basaltic to rhyolitic volcanic breccia.

Culleet Creek Copper-enriched Volcanic Horizon

A volcanic horizon which is cut by Culleet Creek and which makes up part of the overall Le Mare Lake volcanic suite, is herein referred to as the Culleet Creek volcanic horizon by the author (Figure 1.) in order to distinguish it from the remaining Le Mare Lake volcano-lithostrata. This particular volcanic flow-horizon is bimodal and appears to be more enriched in copper than other volcanic horizons currently mapped in the Le Mare Lake proper to date.

The Culleet Creek volcanic andesitic horizon is temporal with rhyolitic pyroclastic flow horizons (see Figure 1). The andesite is characteristically aphanitic with alternating greyish green-maroon flow bands (Photos 1 & 2).



Photo 1 West Side Road, Off Restless Mainline

Within in this copper enriched andesitic horizon, the flow bands predominately trend northerly and dip between 40-60 degrees west as depicted in Photo 1 above. Within this road section some 20 meters up the road and to the right of this photo, are well mineralized, angular copper-epidote-bearing float, scattered along the ditch line, which suggest to be in-place. One of the better grab samples obtained by the author assayed **0.64% Cu and 77 ppb gold** (also Figure 1). Approximately 30-50 meters down the road and to the left of the photo, exposed along the stream bed, are intensely sheared, brecciated, creamy-kaolinitic altered rhyolite flows. The flows also carried mega-pyroclastic, thinly laminated rhyo-dacitic angular fragments which appear to floating and carried along in a grey siliceous, aphanitic matrix see Photo 3 below.



Photo 2 Restless Mainline

Photo above is from log landing-road cut, located about 200 meters higher in elevation than Gorby copper showing and about 300 meters lower from the sample collected in Photo 1. The exposure, characteristically displays siliceous (almost chert-like) dacitic to andesitic of geyish-green, marooned coloured flow banding. This section hosts limited chalcopyrite and malachite staining along fractures. A chip sample collected from the above photo assayed **0.45% copper and 20 ppb gold**.

Copper mineralization found along this exposed section is hosted within the same stratigraphic volcanic horizon as found in Photo 3 above. These 2 copper zones are temporal and are related to the copper mineralization found in the New Destiny and Gorby zones. Although the copper zones appear to occur in slightly different levels or horizons within the andesitic flow and vary in size and tenure, they suggest to be related to one and the same copper mineralizing event. Of the 4 zones found to date, New Destiny is the largest containing the highest copper and gold values associated with mineralization hosted along intense shearing and brecciation and pyroclastic-like andesitic fragments, over approximate andesitic flow- true thickness of at least 80-100 meters. The highest sample assay value collected from the New Destiny copper zone, based on the GPS sample position, appears to have been obtained by on the samplers, along a major shear-breccia structure. This sample contained **3.473 gm/t Au, 4.05% Cu, 15.2 gm/t Ag** along with epithermal signature-like minerals: **2,046 ppm As, 49.2 ppm Cd** and, **152 ppm Hg**.



Photo 3 Below the New Destiny Showing

Photo 3 displays large, mega-pyroclastic, rhyolite to rhyodacite angular fragments, also incorporated with the pyroclastic flow, is a large mafic (andesitic) clast. The matrix is composed of aphanitic, greyish, translucent silica.



Photo 4 Pyroclastic Rhyolite with Grey, Siliceous, Pyrite-rich Matrix (45 ppm Cu & <5 ppb Au)

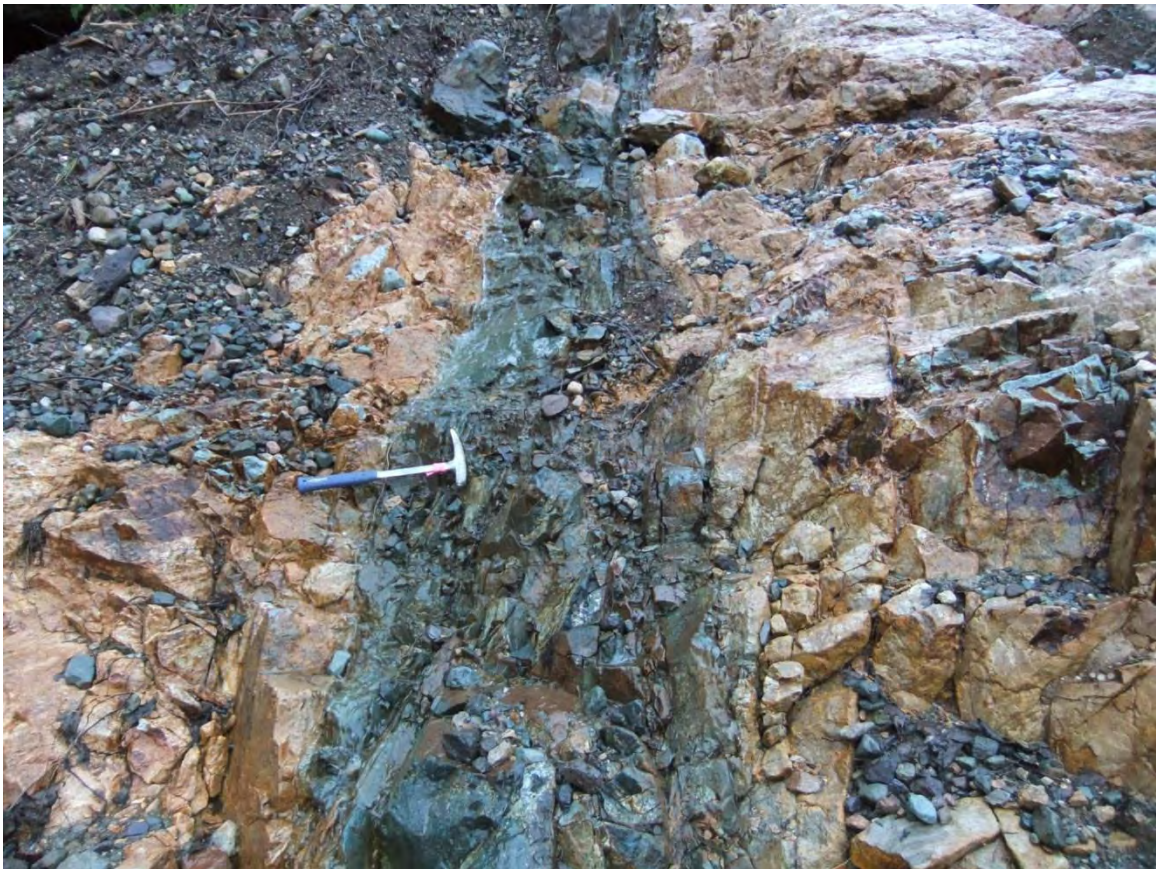


Photo 5 Intercalated Andesite Lens in Rhyolite Breccia Flow



Photo 6 West End of Le Mare Lake

The bottom road near west end of Le Mare Lake, is a section of faulted rhyolite (photo 6) displaying a major northeast striking – southeast steeply dipping, fault with 3-4 meter wide grey gouge (see Figure 2).

Photos 4-6 are taken of northwest trending pyroclastic rhyolite flows near the western end of Le Mare Lake. This is flow is interpreted as part of and temporal to the hydrothermal-geyserite alteration mapped along the road cut as shown in Figure 1.

The pyroclastic rhyolite and rhyodacitic flows mapped above are similar to the Pemberton Hills rhyolite. The Pemberton Hills rhyolite horizon is considered to be younger than the Le Mare Lake volcanics however, possible viscous flow-dome-like complexes similar to the Pemberton Hills should not be overlooked in the Le Mare Lake proper. Further mapping of the Culleet Creek andesite- rhyolite pyroclastic volcanic pile could vector to proximal environment(s) for epithermal and or, massive sulphide type mineralization.

DEFORMATION AND METAMORPHISM:

The Le Mare Lake volcanics were subjected to regional deformation (D1) during collision and accretion of the Wrangellia Terrane to west coast Intermontane Belts of British Columbia, between Middle Jurassic to mid-Cretaceous time. During the Nassian Orogeny (D1), the volcanic rocks would also have experienced regional lower

greenschist facies metamorphism. A second deformation phase (D2) would have occurred during the Late Jurassic to Late Cretaceous Columbian Orogeny as the result of on-going subduction of the Pacific Oceanic (Juan De Fuca) plate. This orogeny would have produced D2 greenschist overprinting and further tilting of the Le Mare Lake volcanic as shown in the following photos.

Photos 1 and 2 above show low grade greenschist facies volcanic flows moderately dipping to the west which were probably subjected to the initial deformation (D2) folding producing large open monoclines and subsequently further tilted by D2 deformation.



Photo 7 Just East of New Destiny Showing

Photo shows andesitic flows with open fold limb dipping to the northwest probably related to D2 folding. Above the yellow dashed-line are incipient pillow-like lavas.



Photo 8 Part of New Destiny Showing

Part of a section of the New Destiny copper zone across 20 meters displaying intense shearing and brecciation (between dotted lines) probably related to tranpressional deformation (D3). It is along this section that the high grade copper-gold-silver sample was collected.



Photo 9 Part of New Destiny Showing

Photo showing near the western end of the New Destiny some 130 meters west of photo 8. Fault above (marked in dashed line) is probably related to same fault-shear system in photo 8. The fault strikes northeast and dips shallow to southeast. This structure could also be interpreted as a possible thrust fault with HW riding over FW related to D3 deformation (see Figure 2).

Mineralization:

Presently, all of the copper mineralization examined by the author to date is hosted in the Le Mare Lake andesitic volcanic rocks, with the Culleet Creek volcanic horizon more copper enriched than others. Although the pyroclastic rhyolite flows can carry abundant siliceous, fine pyrite, the copper content is generally low. The copper mineralization found on all of the 4 copper zones noted above are predominately structurally controlled, occurring as thin fracture veinlets or as fracture healed, irregular quartz-chalcopyrite veins. Some disseminated or isolated blebs of copper can be found away from the structurally controlled veinlets. The copper-bearing quartz veins characteristically fill architecturally prepared structural sites such as in the case of the New Destiny zone and to a lesser extent at the Gorby. Where there is an increase in quartz veining, chalcopyrite and pyrite mineralization tend to be more abundant. This is evident in the New Destiny, especially along one narrow exposed section where there is highly siliceous quartz veining carrying abundant chalcopyrite and pyrite, as displayed by the photo below.

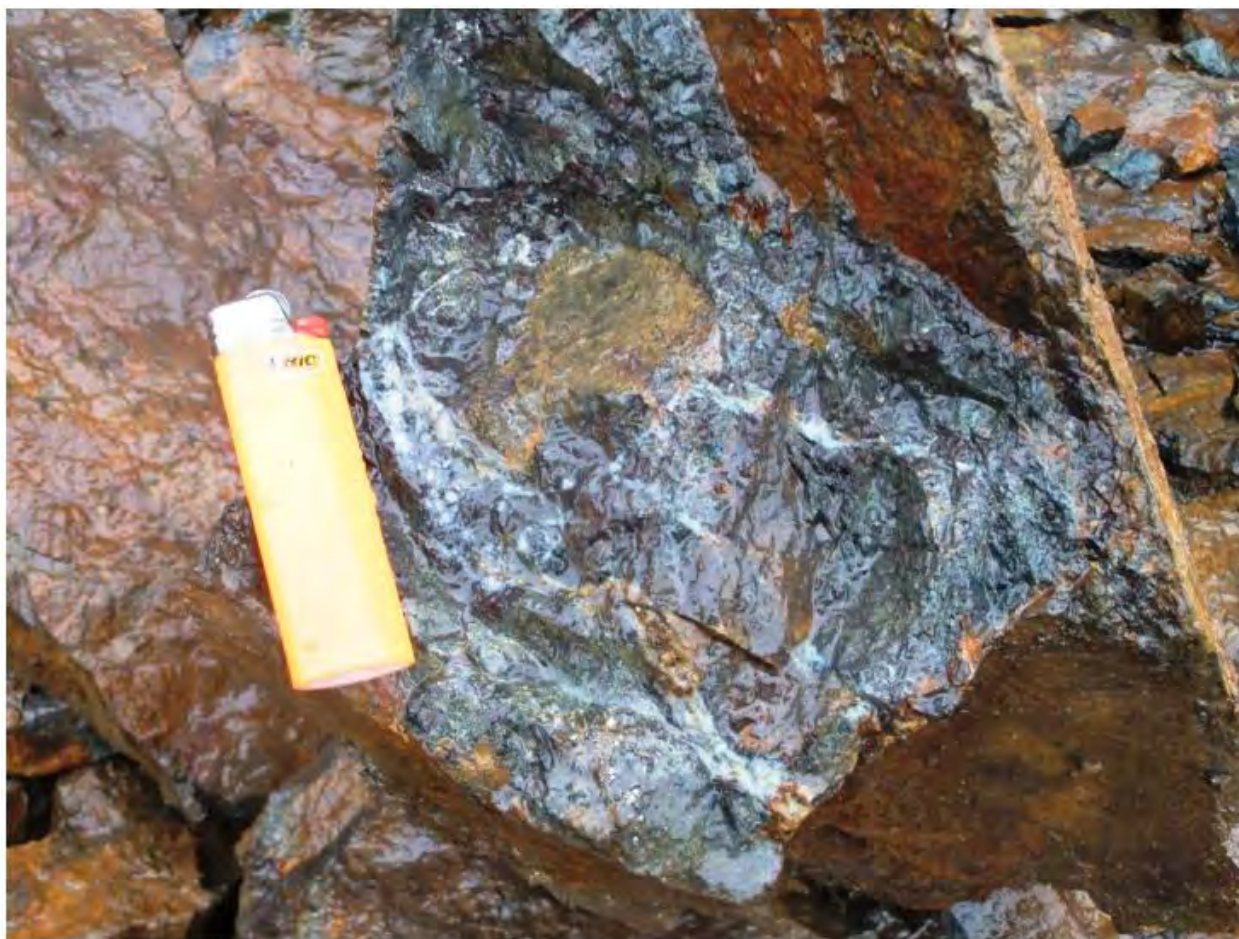
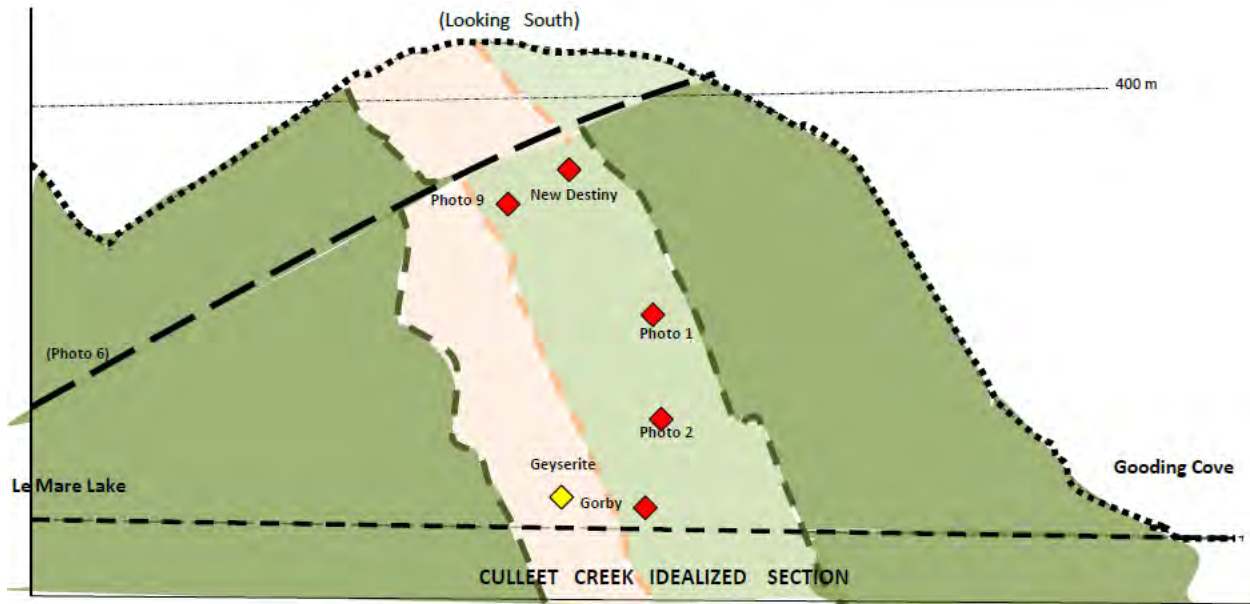


Photo 10 Quartz Veining at the New Destiny Showing

The copper(gold)-bearing andesite and the rhyolite and pyroclastic flows are temporal and suggest some pre-tomagmatic activity. The possibility that some of this mineralization was syngenetically deposited and the

possibility of defining a volcanogenic style mineralization of temporal epithermal environment on the Farwest property, is a concept that will require further mapping and prospecting.



(Schematic Cross Section not to scale)

FIGURE 20C

SOME PRELIMINARY CONSTRAINTS TO THE COPPER AND GOLD MINERALIZATION

Copper and gold mineralization is hosted along an andesitic volcanic horizon temporally and spatially related to rhyolitic and rhyolite pyroclastic flows. The mineralization is post deformational and appears in part, to be structurally related, and could also be considered as a volcanic-hosted orogenic style mineralization, with some of the mineralized-bearing fluids originating from a deeper seated (mesozonal) pluton.

The New Destiny Showing was discovered in 2010. In the 2011 program the showing was trenched with a tracked excavator and sampled in 3m intervals by chip samples, Figure 20. The results show over 200m averaging over 0.2% copper with significant gold.

Gold in soil anomalies are widespread, the largest is on the knoll southwest of the New Destiny showing which is 100m long NE-SW and 400m east-west. There may be a mineralized fault zone on the top of the knoll that is the source of the gold. Gold values range up to 947 ppb gold.

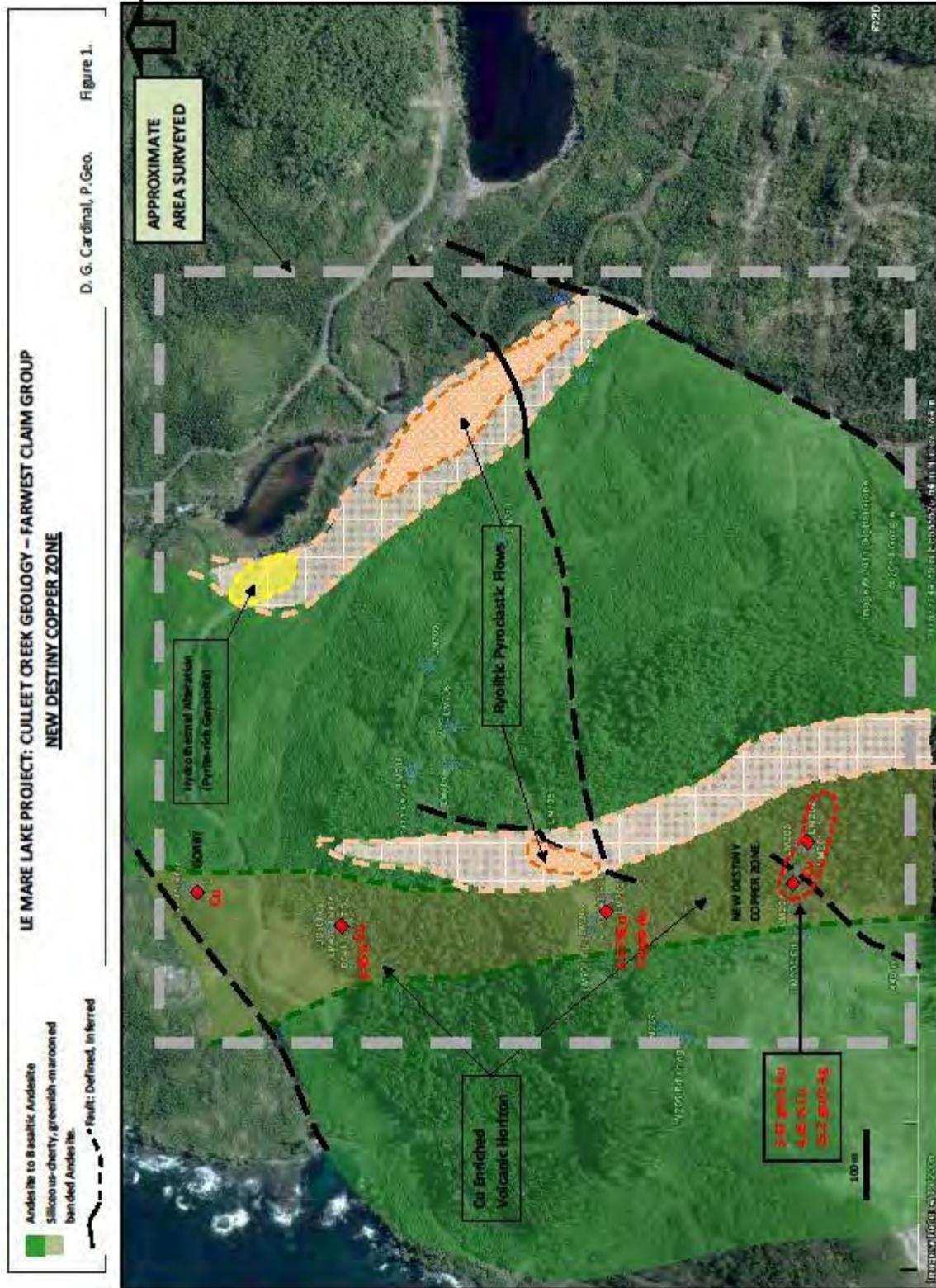


FIGURE 20D

Soil Surveys

Of the three major soil surveys conducted around the Le Mare hydrothermal system, the 1991 Stow survey (Birkeland, 1991) was the most extensive and most closely resembled a regional survey. Also A. O. Birkeland (1991) made the only calculation of soil-metal thresholds using sample populations. However, he adjusted some of his thresholds downward, which enlarged anomalies (Table 4).

Contoured soil-copper and molybdenum concentrations in soils from the 1991 Stow, 2007 Equus and 2009 New Destiny surveys (Birkeland, 1991; Shearer, 2007 and 2010, respectively) (Figures 21E to 22W). The most recent data was preferentially plotted in areas of overlap.

Most high soil-copper concentrations coincide with orange-weathering, orthoclase-quartz, potassic alteration. This co-occurrence facilitated identification of the hydrothermal plumes in the northwestern part of the La Mare hydrothermal system (Figures 15 and 21W). It supports the thesis that copper mineralization at depth, probably is related to potassic alteration like at the Island Copper mine deposit (Section 11.1 of this report) (Figure 13). At the South Gossan zone, where argillic-phyllitic alteration has overprinted on previous potassic alteration, soil-copper anomalies are small and weak (Figure 21E). This indicates that copper, previously deposited with potassic alteration, has been partly removed from the rocks of this area during this later alteration phase.

All significant soil-molybdenum anomalies are spatially associated with the lower eastern flank of the plume of argillic-phyllitic alteration adjacent to an area of quartz-sericite-pyrite (phyllitic) alteration (Figure 22DE) (section 10.2.1 of this report). None are associated with the earlier potassic hydrothermal plumes (Figure 22W).

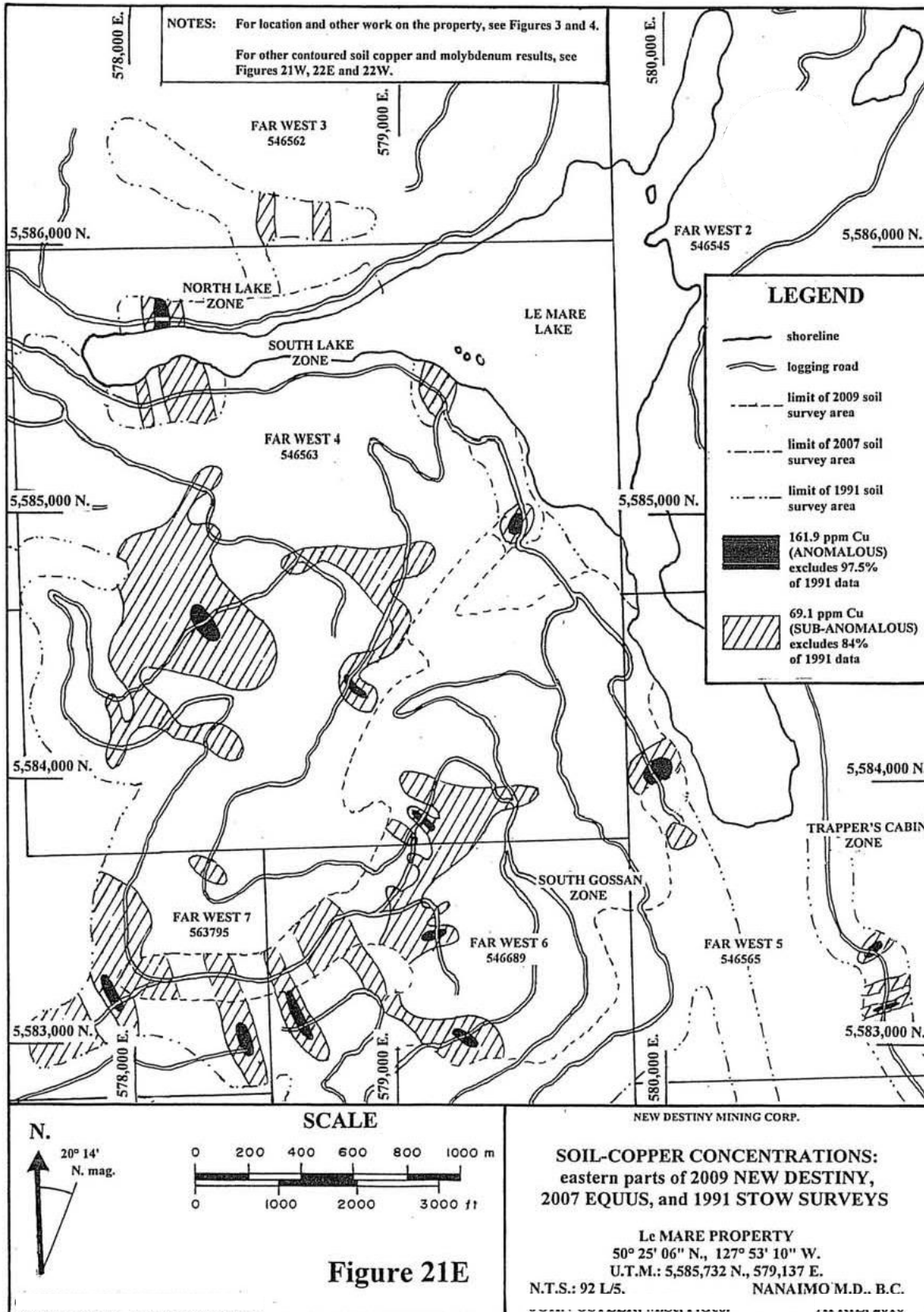


FIGURE 21A

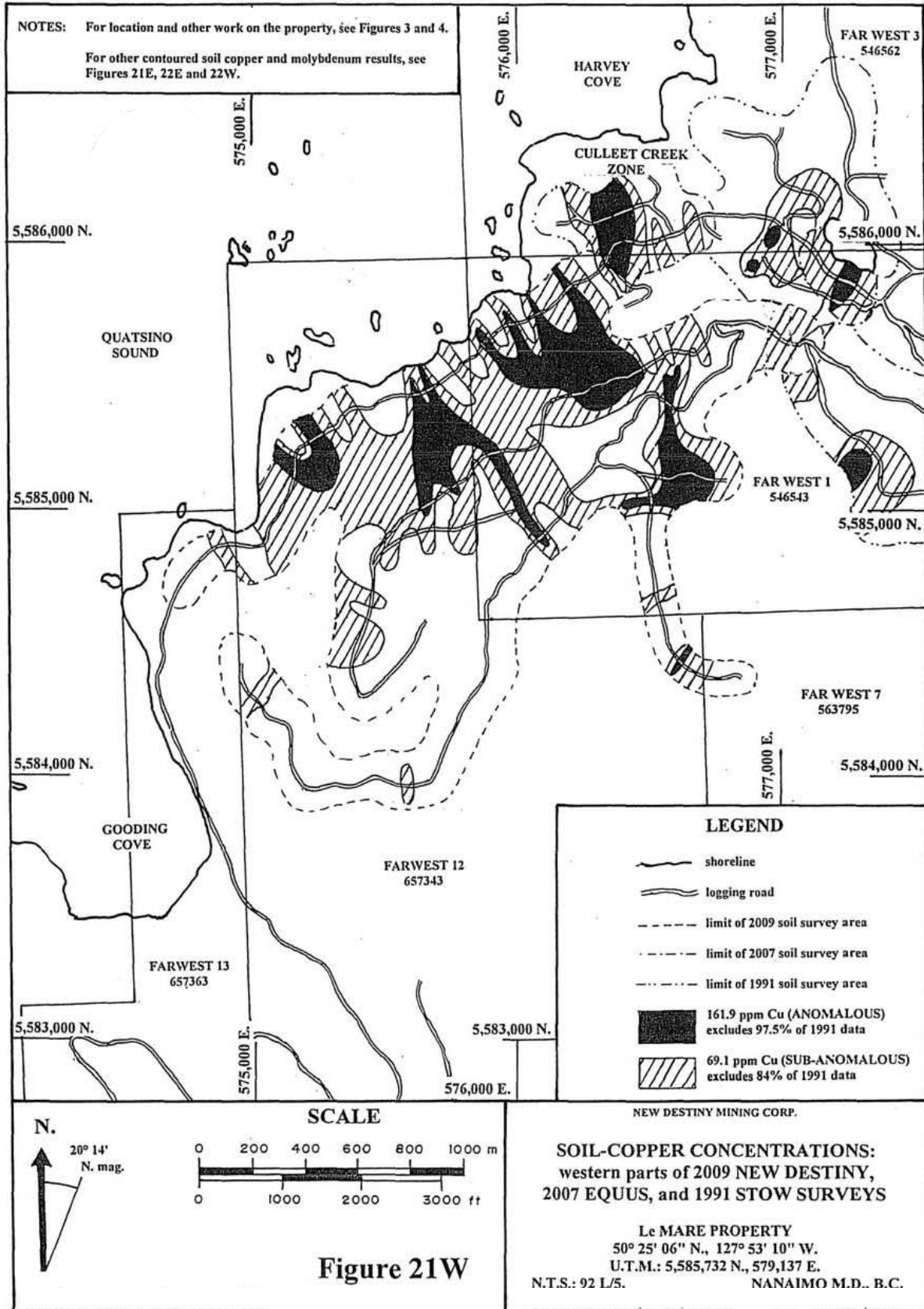


FIGURE 21B

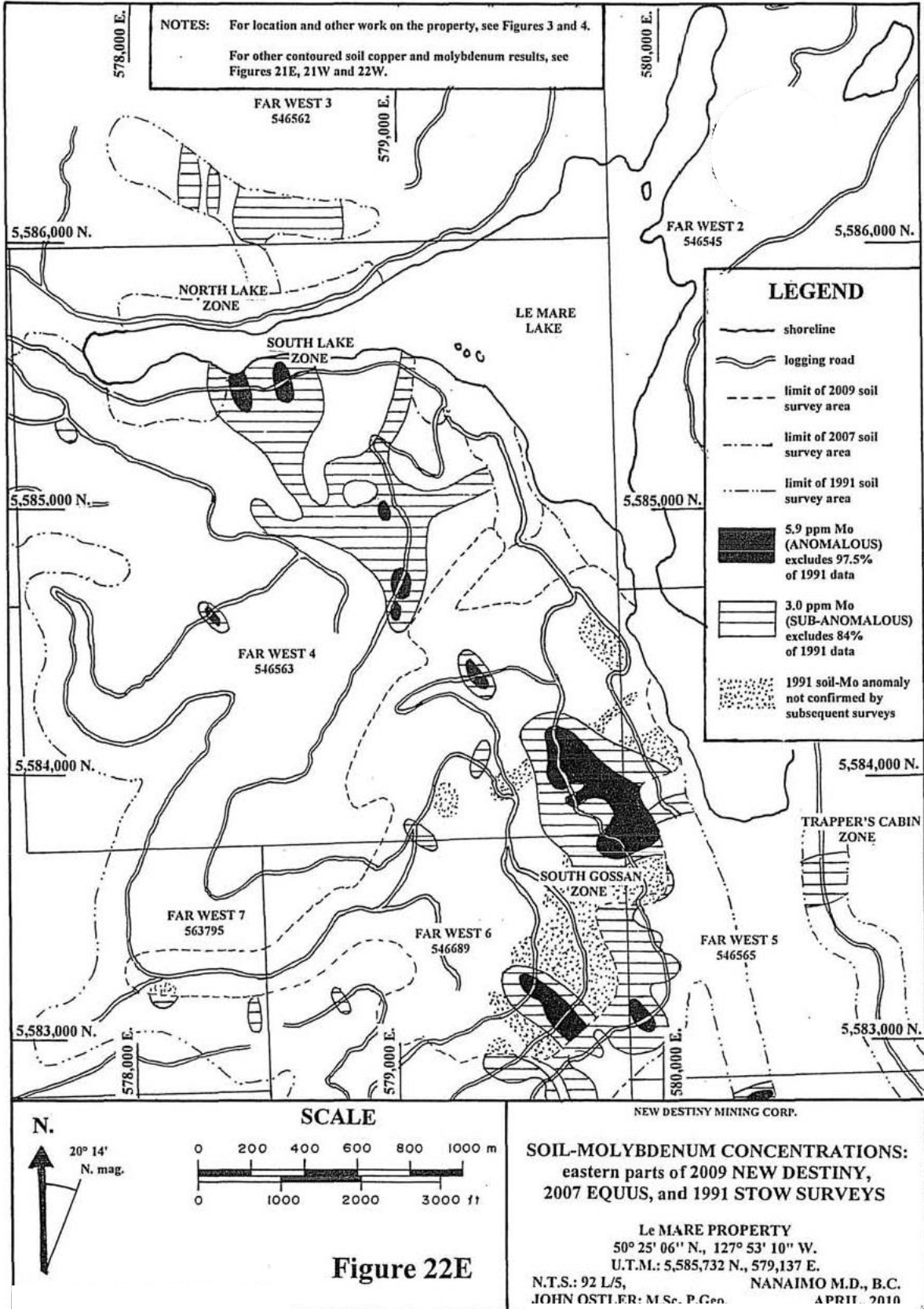


Figure 22A

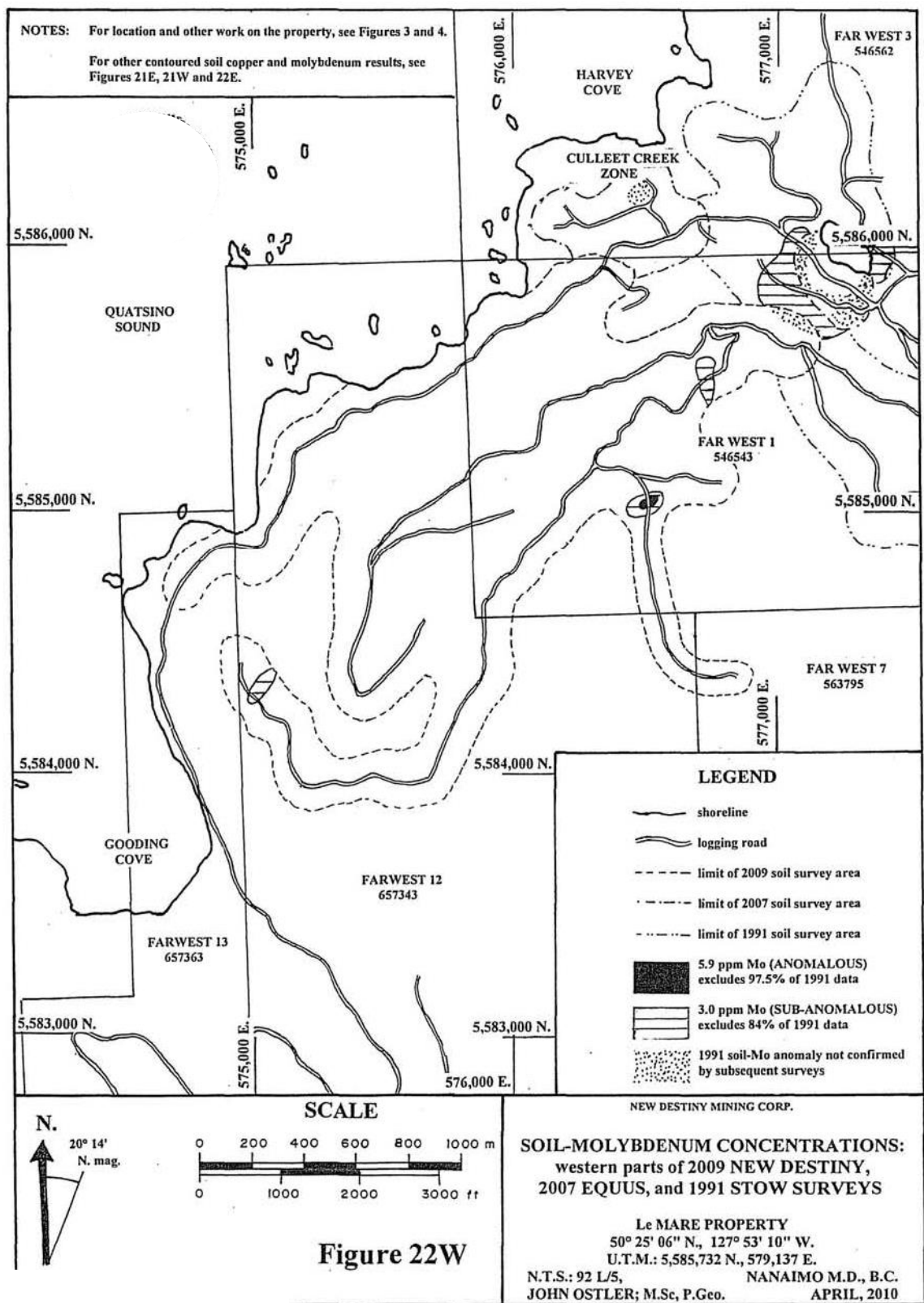


FIGURE 22B

EXPLORATION 2014

Stephen Wetherup, Mark Rein, Max Estaris and Rob Cameron completed three days of geological mapping on the LeMare Property, in northern Vancouver Island from July 22nd to July 24th, 2014. The purpose of the mapping was to determine if geology and alteration on the LeMare Property were indicative of a porphyry Cu-Au-Mo system occurring on the Property.

Access to the property was along logging roads many of which were heavily overgrown and some areas were just too far to reach on foot although most of the focus area (the South Gossan) was covered at lower elevations. A total of 16 samples were collected during the mapping for later Terraspec analysis and mapping data focused on rock types, structures, alteration minerals/type and intensity of the alteration.

The LeMare Property is largely underlain by Jurassic age, Wrangellian island arc Terrane Bonanza Group bi-modal volcanic rocks. The Bonanza group rocks are dominated by andesitic flow and volcanoclastic rocks with rare siltstone, wacke and rhyolite/dacite flows and tuffs.

Bonanza Group rocks generally strike southward and dip moderately westward which are folded locally to a SE strike and near vertical dip. A major NE trending fault is interpreted to occur along Dumortierite Creek and appears to down-drop the NW block of Bonanza Group rocks on the Property. This assumption is based on alteration in the Bonanza rocks which is distinct in each block and described below.

On the southwest corner of the Property a downthrown block of Cretaceous age, Longarm Formation basalt and shale/siltstone occurs and presumably overlies the Bonanza Group rocks. The Longarm Formation rocks are cut by numerous faults, mainly WNW striking, steep, dextral strike-slip faults, N striking steep normal block faults and NE striking oblique faults. The Longarm block is bounded by the WNW and NE faults and locally contains N striking qtz-cb-ep+/-py+/-apy veins and breccia zones.

Alteration and Mineralization

Bonanza group rocks are generally chlorite-pyrite (propylitic) altered. In the NW block of Bonanza rocks the chlorite-pyrite alteration is overprinted by silica (locally chalcedonic)-hematite+/-jasperoid locally (Gorby showing) and silica-clay-pyrite (advanced argillic?). At the Gorby showing minor amounts of chalcopyrite occur with the silica replacement. Several zones (beds?) of advanced argillic alteration comprised mainly silica-pyrite-clay which appear to be 25-50m thick. There are also rare zones of sericite-silica-pyrite along structural zones (possibly bedding planes as well) approximately 1-2 m wide and generally along LeMare Lake on the east side of the NW block.

The SE block of the Bonanza group rocks (South Gossan Zone) is also propylitically altered by chlorite-pyrite but on the eastern margin of the block by LeMare Lake the andesite is chlorite-epidote-pyrite-magnetite altered with abundant epidote-calcite+/-chalcopyrite (rare covellite/bornite) veins. This area coincides with a moderate magnetic high on the aeromagnetic data. Up slope from LeMare Lake the Bonanza volcanic rocks are chlorite-pyrite-epidote altered and are cut by numerous zones of sericite-pyrite-silica alteration which is generally structurally controlled but also appears along bedding planes or within permeable layers. These QSP zones contain pyrite veinlets and rare quartz (with no pyrite) veinlets locally.

North of LeMare Lake several K-feldspar altered fault zones occur within Bonanza andesite rocks and is the only observed potassic alteration on the property.

The Longarm formation is weakly chlorite-epidote alteration with local vuggy quartz-epidote-calcite-pyrite veins.

The Bonanza group rocks in the NW block on the property contains extremely few veins and any alteration more intense than the regional chlorite-pyrite propylitic alteration is very high level in character with advanced argillic

silica-pyrite or chalcedonic silica-hematite. Chalcopyrite mineralization associated with the silica-hematite is not likely to be porphyry related. Overall, this block of rocks does not appear to have any porphyry potential.

The Bonanza rocks SE of Dumortierite Creek (South Gossan Zone) are distinct as the propylitic alteration of the lower elevation andesite units near LeMare Lake and south of the lake contain abundant epidote and magnetite which was nearly absent north of the creek. And, there are many more QSP alteration zones within the otherwise propylitic rock. Overall, it appears that these rocks were lower in the hydrothermal system than the NW block.

The presence of numerous epidote-calcite-chalcopyrite/bornite veins in the magnetic area is encouraging in terms of porphyry potential. However, the lack of veining in the overlying rocks, lack of any appreciable intrusive rocks and the presence of the faulting that cuts the SE block 2km to the south, severely limits the exploration potential.

Furthermore, the geochemical data from historical work in the South Gossan shows very weak Cu-Au-Mo and a single drill holes located in the South Gossan also did not intersect porphyry alteration or mineralization.

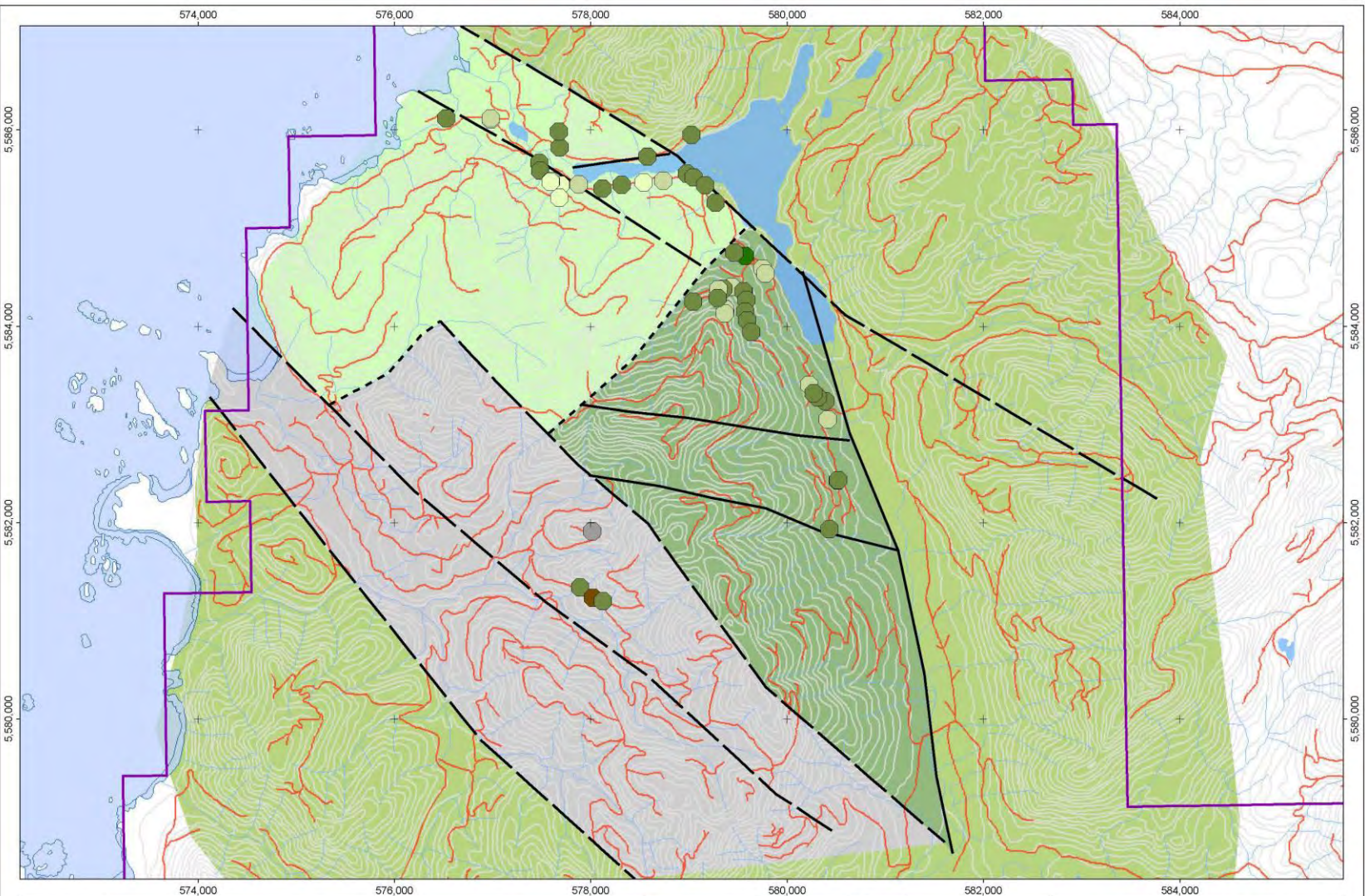


Figure 23

<p>Geology - Rock Type</p> <ul style="list-style-type: none"> Rhyolite Andesite Andesite tuff Basaltic andesite Longarm Basalt Siltstone 	<p>Faults:</p> <ul style="list-style-type: none"> Dextral Normal Ob-Norm 	<p>Geology Blocks:</p> <ul style="list-style-type: none"> Bonanza Gp - Upper Block Bonanza Gp - Undivided Bonanza Gp - Lower Block Longarm Fm. 	<ul style="list-style-type: none"> Roads - Unpaved Rivers/Streams Contours (ft) Lakes/Ponds Ocean Tenure Outline 	<p>Lemare Property, BC, Canada.</p>	<p>Scale</p> <p>0 200 400 600 800 1,000 m</p> <p>Projection: NAD 1983 UTM Zone 9N.</p>	<p>Lemare Property</p> <p>Geology Rock Type Map, Lemare Property, BC, Canada.</p> <p>Date: 23/9/14 Figure: #1 Office/Author:</p>
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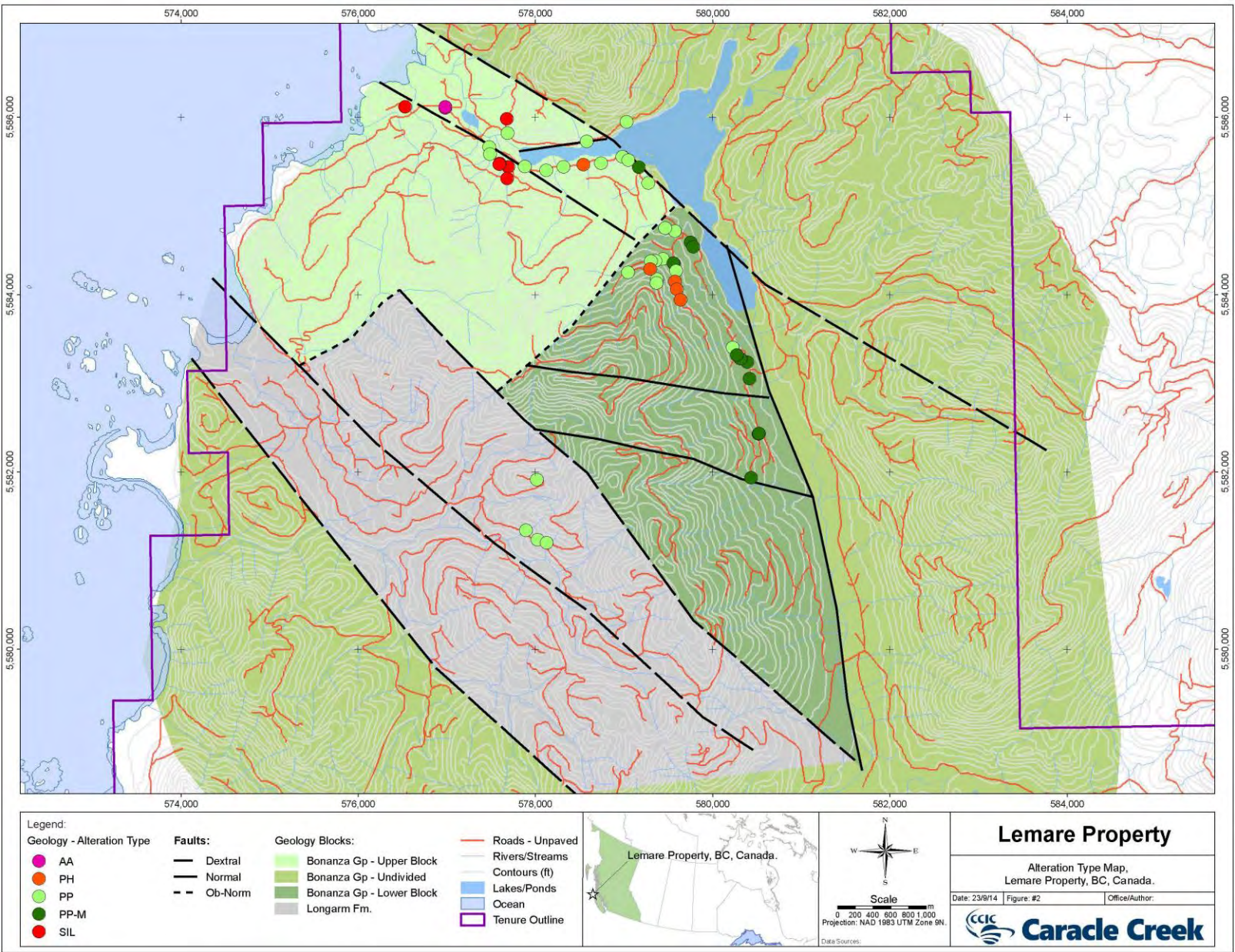


Figure 24

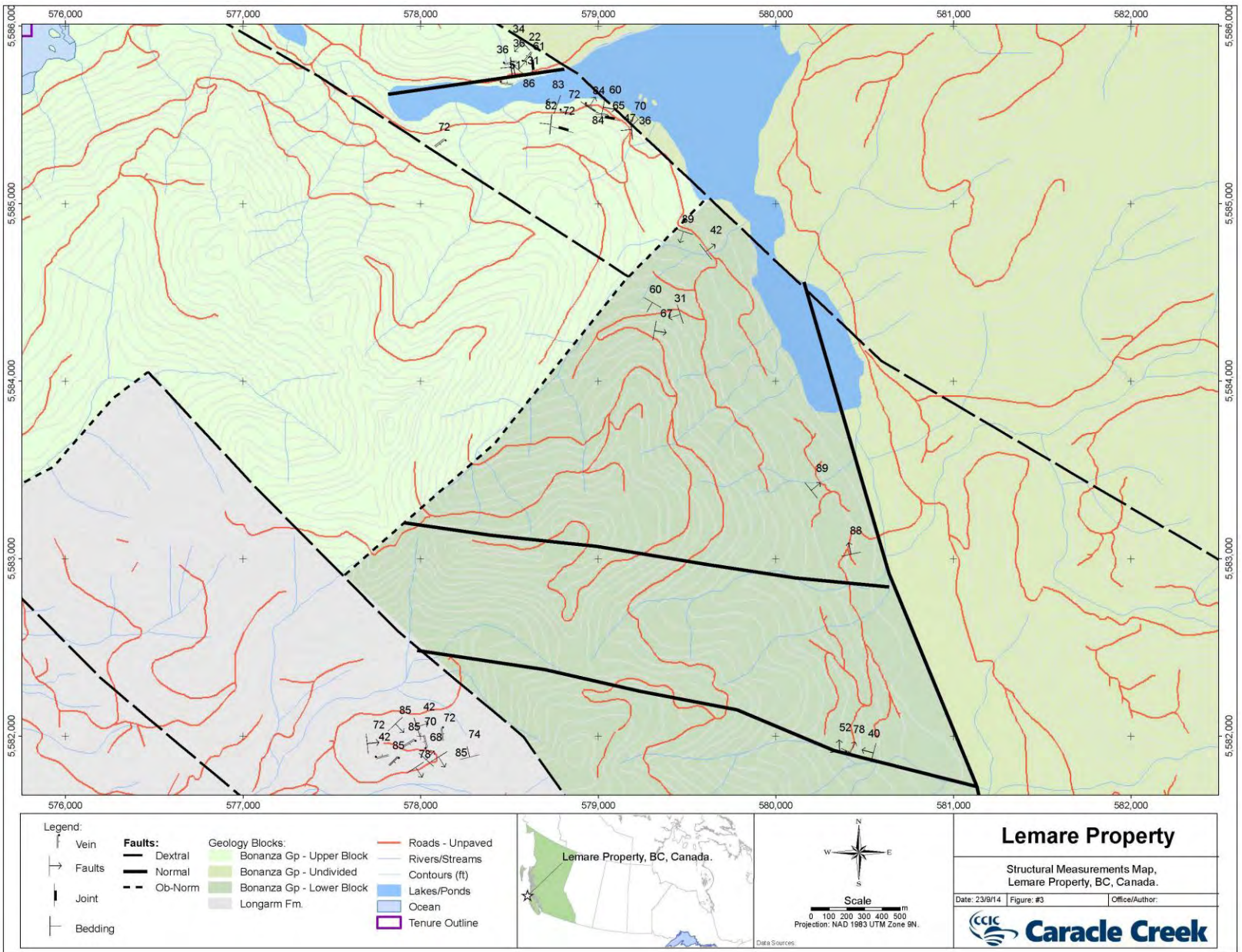


Figure 25

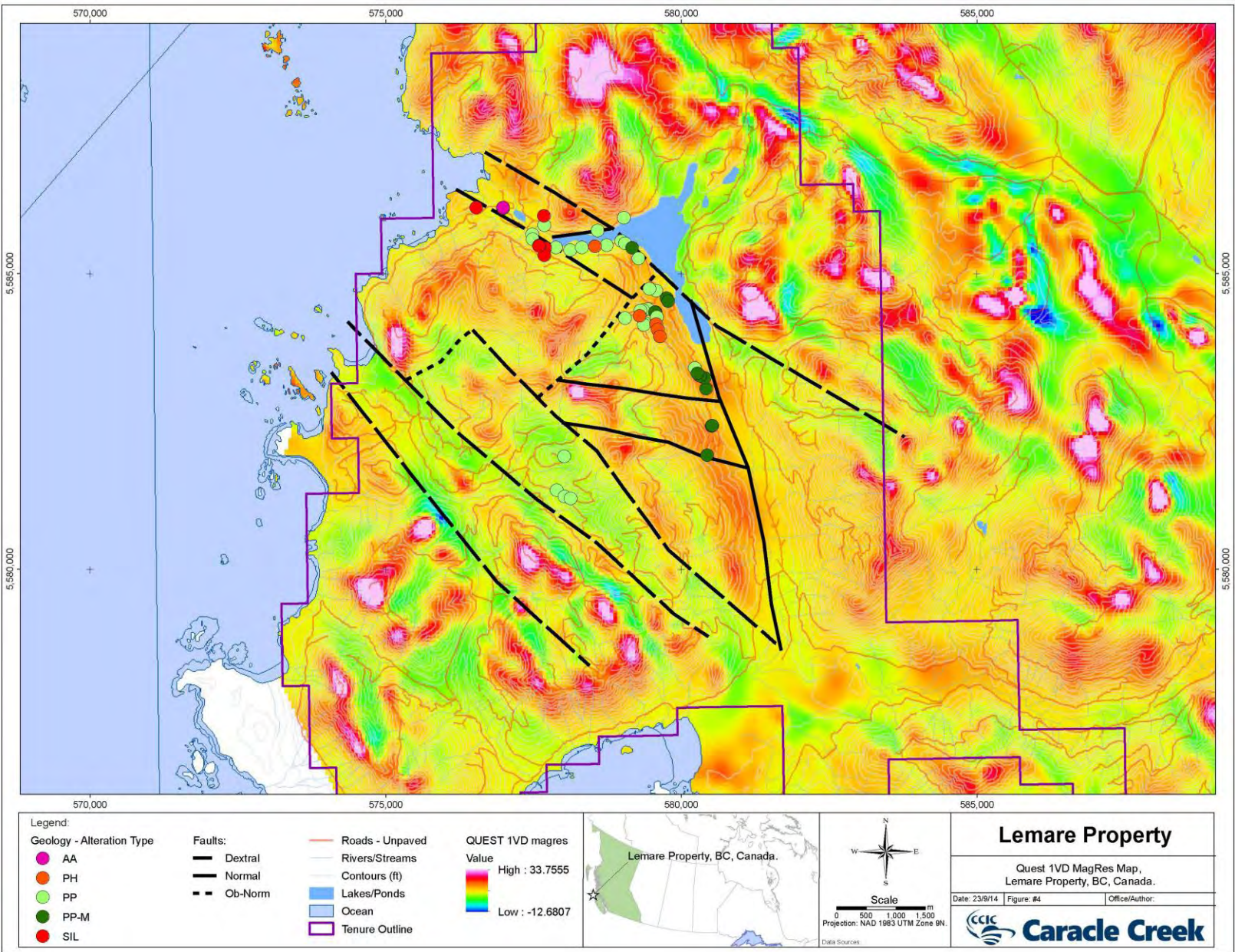


Figure 26

CONCLUSIONS and RECOMMENDATIONS

The Le Mare hydrothermal system has been just barely unroofed by erosion. The top of the potassic alteration zone is exposed along the crests of Le Mare and Gooding ridges, located between Le Mare Lake and Gooding Cove in the southwestern part of the Le Mare property. Local magnetic field gradient indicates that this system occupies a 5 X 3 km (3.05 X 1.83 mi) or 15 km² (5.6 mi²) oval-shaped area that may be hosted by a dilational jog in a regional right-lateral fault system. The proposed fault system is similar to the one that hosts the Island Copper Cluster deposits near Port McNeill, British Columbia.

At surface, copper mineralization occurs in discrete showings-areas, located preferentially in the central parts of sub-vertical hydrothermal plumes. These plumes have core-zones of orthoclase-quartz-biotite (potassic) alteration, enveloped in siliceous exteriors. The gradual decrease in the orthoclase:quartz ratio from about 4:1 in potassic core zones to less than 1:20 in siliceous envelopes, indicates that peripheral silicification is a distal phase of the core-zone potassic alteration and is not overprinted by it. Orthoclase-quartz-biotite alteration is succeeded by quartz-jasper alteration; both phases are mineralized with chalcopyrite, and minor quantities of bornite. This potassic alteration is accompanied by co-incident soil-copper and local magnetic anomalies. Discovering economically viable concentrations of copper mineralization within the Le Mare hydrothermal system depends on the successful identification of zones where these hydrothermal plumes and copper occurrences coalesce.

Molybdenum enrichment occurs in areas flanking phyllic alteration in a 600-m (1,968.5-ft) diameter argillic-phyllic alteration plume, covering a 0.28 km² (0.1 mi²) area in the eastern part of system in the South Gossan zone. Another, much less extensive plume of argillic-phyllic alteration is exposed between the Culleet Creek zone and Culleet Lake in the system's northwestern part. These two plumes cover less than 2% of the total exposure-area of the Le Mare hydrothermal system. Argillic-phyllic alteration post-dates and overprints potassic alteration.

Both sample results and the distribution of soil-copper and molybdenum anomalies, demonstrate that copper and molybdenum mineralization are associated with early potassic and subsequent argillic-phyllic alteration events respectively. They occur together in significant amounts only where molybdenum enrichment has overprinted that of copper.

Most aspects of the Le Mare hydrothermal system are similar with those of the Island Copper Cluster deposits. Geology, alteration, and mineralization at surface at the Le Mare hydrothermal system correspond with those attributes at the Island Copper mine above the main deposit. These similarities indicate that the Le Mare hydrothermal system may host a calc-alkalic porphyry copper-molybdenum deposit of the Island Copper Cluster type.

The Early Jurassic-age land surface above the Le Mare hydrothermal system, and whatever near surface hot-spring environment that it may have hosted, has been lost to erosion. Only a few narrow roots of a late, advanced argillic alteration occur in the argillic-phyllic alteration plume in the South Gossan zone. They attest to the former existence of hot spring development above the current erosional level.

Previously, the Le Mare hydrothermal system has been investigated for geysirite (SiO₂.nH₂O), an industrial mineral related to hot spring deposits. The level of exposure of the Le Mare hydrothermal system is beneath that favourable for the development of near-surface clays and industrial minerals. The chance of finding a commercially viable geysirite deposit in this area is low to nil.

Most exploration has been conducted in the northeastern part of the Le Mare hydrothermal system; its southeastern part remains sparsely explored to unexplored. Six BQ diamond drill holes penetrated the northeastern margin of the Le Mare system in 1992. One hole that penetrated the Culleet Creek potassic alteration plume, intersected five 2-m (6.56-ft) and one 4.7-m (15.42-ft) long intersections that contained from 500 to 959 ppm copper, which is similar to the tenor of copper mineralization in nearby trenches. Copper mineralization at surface is locally quite variable. The writer's samples range from 3 ppm to 6.57% copper.

Generally, the reproducibility of small-scale sampling is low. Such variability should be expected in mineralization located near the top of the potassic alteration zone of a porphyry copper-molybdenum deposit. Less than 1% of the surface area of the Le Mare hydrothermal system has been drilled.

Recommendations

It is recommended that a two-phase work program be conducted on the Le Mare hydrothermal system to explore for the presence of a calc-alkalic porphyry copper-molybdenum deposit of the Island Copper Cluster type.

The first phase of the recommended work program comprises geological mapping, prospecting, and soil survey. The services of an excavator, two pick-up trucks and an all-terrain vehicle (quad) will be required to provide efficient access to the work-area.

Geological mapping and prospecting should be conducted over the whole 15 km² (5.6 mi²) covered by the hydrothermal system. Focus of this work should be on: 1. identifying areas of mineralization that may be present, 2. recording the relationship among mineralization and the various alteration types that are known to be present in the system, 3. mapping the distribution of alteration types to determine the locations of potassic and argillic-phyllitic alteration plumes which are hosts of copper and molybdenum mineralization respectively, 4. discerning both the volcanic and intrusive stratigraphy, and the pattern of drape folding.

Soil samples should be collected along traverses using the many logging roads that are present throughout the area covered by hydrothermal system. In most parts of the project-area, logging roads are sufficiently close together to enable contouring of soil data between them. Contour traverses through the bush will be necessary in locations with widely spaced roads. Soil samples should be taken at 50-m (82-ft) intervals along traverses. During the current (2009) work program, soils were surveyed over 5 km² (1.9 mi²) of the 15 km² (5.6 mi²) area of the hydrothermal system. If the sampling density of the 2009 soil survey is maintained, then to survey the remaining 10 km² (3.7 mi²) will require about 580 samples to be taken along 27.6 km of line. A cost-estimate of this first-phase of the recommended work comprises Table 13.

If reasonable encouragement is generated by the results of the first-phase program, it should be followed by a second-phase program of 3-dimensional induced polarization and ground magnetic surveys. Those surveys should be conducted over the whole 15 km² (5.6 mi²) area of the hydrothermal system. Induced polarization surveys are conducted most efficiently along cut lines where wires of known lengths can be strung out from generators. A survey over the Le Mare hydrothermal system will require the cutting and surveying of about 75 km (45.75 mi) of line to produce a survey grid-area with lines spaced 200 m (656 ft) apart. Cutting lines through dense second-growth rain forest over the rugged terrain on the Le Mare property will be slow work. It is estimated that a 2-man crew will be able to cut only about 0.5 km (0.31 mi) of line per day. A cost estimate of this second-phase of the recommended work comprises Table 14.

The results of the two phases of the recommended work program should produce a three-dimensional assessment of the rock to a depth of about 300 m (984 ft) beneath the surface exposure of the Le Mare hydrothermal system. It is expected that this assessment will be sufficiently detailed to enable precise location of the best target-areas to drill for porphyry copper and molybdenum mineralization.

Table 8
Estimated Cost of the Recommended Second-phase Exploration Program

Item	Costs	Accumulated cost
Line-cutting Costs: 75 km of line @ \$2,250/km (all-in contractor price) Excavator for road opening; 90 hours @ \$150/hour Excavator mobilization	\$ 168,750 \$ 13,500 <u>\$ 5,000</u> \$ 187,250	\$ 187,250
Geophysical Survey Costs: 75 km of 3-dimensional induced polarization and ground magnetic surveys @ \$3,200/km (all-in contractor price including data manipulation and reporting)	\$ 240,000	\$ 240,000
Geological Support and Project Management: J.T. Shearer, senior geologist and project manager; 25 days @ 700/day 1 geologist; 25 days @ \$650/day each including field work, data manipulation and reporting for assessment	\$ 17,500 <u>\$ 16,250</u> \$ 33,750	\$ 33,750
Transport and Crew Costs for Geological Support and Management: 1-ton 4X4 pick-up truck; 20 days @ \$100/day Gasoline Hotel; 20 man-days @ \$100/day Meals in transit; 20 man-days @ \$60/day	\$ 2,000 \$ 2,500 \$ 2,000 <u>\$ 1,800</u> \$ 8,300	\$ 8,300
Communication Costs: Satellite phone rental; 4 weeks @ \$400/week 1 FM truck radio; 1 month @ \$750/month	\$ 1,600 <u>\$ 750</u> \$ 2,350	\$ 2,350
Reporting Costs and Office Expenses: Digital Map Drafting Physical and Electronic Assessment Report Production Costs	\$ 3,000 <u>\$ 800</u> \$ 3,800	\$ 3,800
Environmental and Compliance Costs: Top up of current environmental bond for road work and line cutting	\$ 10,000	<u>\$ 10,000</u>
Itemized Cost of Recommended Second-phase Induced Polarization and Ground Magnetic Survey Program		\$ 485,450
Harmonized goods and services tax (H.S.T.) (12% of previous items)		<u>\$ 58,254</u>
Itemized Budget		\$ 543,704
Contingency; 10% of itemized budget		<u>\$ 54,370</u>
Total Estimated cost of Recommended Second-phase Induced Polarization and Ground Magnetic Survey Program		\$ 598,074

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APPENDIX I

STATEMENT of QUALIFICATIONS

January 2, 2015

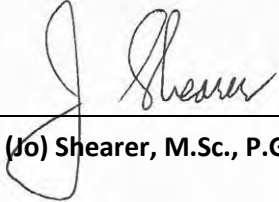
STATEMENT of QUALIFICATIONS

I J. T. (Jo) Shearer, of Unit 5 – 2330 Tyner St. Port Coquitlam, BC, V3C 2Z1, do hereby certify that:

1. I am an independent consulting geologist and principal of Homegold Resources Ltd.
2. My academic qualifications are:
 - Bachelor of Science, Honours Geology from the University of British Columbia, 1973
 - Associate of the Royal School of Mines (ARSM) from the Imperial College of Science and Technology in London, England in 1977 in Mineral Exploration
 - Master of Science from the University of London, 1977
3. My professional associations are:
 - Member of the Association of Professional Engineers and Geoscientists in the Province of British Columbia, Canada, Member #19,279 and the APGO in Ontario, Member 1867.
 - Fellow of the Geological Association of Canada, Fellow #F439
4. I am responsible for the preparation of all sections of the technical report entitled "Geochemical Assessment Report on the Le Mare Copper-Gold Property" dated January 2, 2015.
5. I have visited the property between February 24, 2011, May 15, 2011 and July 21-24, 2014. I have carried out mapping and sample collection and am familiar with the regional geology and geology of nearby properties. I have become familiar with the previous work conducted on the Le Mare Project by examining in detail the available reports and maps and have discussed previous work with person knowledgeable of the area.

January 2, 2015

Date



J.T. (Jo) Shearer, M.Sc., P.Geo.

Appendix II

Statement of Costs

January 2, 2015

Statement of Costs 2014

Professional Services

Wages	Total not including HST
J.T. Shearer, M.Sc., P.Geo., 4 days @ \$700/day, July 21-24, 2014	\$ 2,800.00
R. Savelieff, B.Sc., P.Geo., 4 days @ \$450/day July 21-24, 2014	1,800.00
Stephen Wetherup, P. Geo. 5 days @ \$700/day, July 21-25, 2014	3,500.00
Subtotal	\$ 8,100.00

Expenses

Transportation:

Truck 1, fully equipped 4x4, 5 truck days @ \$120/day	600.00
Truck 2, fully equipped 4x4, 4 truck days @ \$120/day	480.00
All Terrain Vehicle, 4 days @ \$100/day	400.00
Gas	810.00
Camp, 14 man days @ \$150/day	2,100.00
Ferry	410.00
Survey and Camp Supplies	350.00
Food & Meals, 14 man days @ \$45/day	630.00
Satellite Phone Rental	250.00
Truck Radios	300.00
Mark Rein, 5 days @ \$450/day, July 21-25, 2015	2,250.00
Max Estaris, 5 days @ \$500/day, July 21-25, 2015	2,500.00
Rob Cameron, 5 days @ \$400/day, July 21-25, 2015	2,000.00
Computer Work for Geological Maps, Figures 23 to 26, 1:5,000	600.00
Data Compilation and Interpretation	1,400.00
Word Processing and Reproduction	350.00
Subtotal	\$ 15,430.00

Grand Total \$ 23,530.00

Event #	5539864
Date Filed	January 26, 2015
Amount	\$23,000.00
PAC filed	\$4,050.80
Total Paid	\$27,050.80

APPENDIX III

ROCK SAMPLE DESCRIPTIONS

January 2, 2015

Station	Date	Easting	Northing	Rock type	1' Alteration	Alt Int1	2' Alteration	Alt Int2	Alt type	Veins	V Int.	Sample	Photos
sw352	22-Jul-14	578580	5585730	Andesite flows	chl-ep-py		3 k-spar		2 PP	ksp		3	yes
sw353	22-Jul-14	576980	5586109	Rhyolite-dacite	silica-clay		4		AA			87155	
sw355	22-Jul-14	577880	5585448	wacke-tuff	chl-py		3 ser-py		1 PP			87251	
sw356	22-Jul-14	578121	5585406	Andesite flows	chl-ep-py		3 ep		2 PP			87252	
sw357	22-Jul-14	578315	5585445	Andesite flows	chl-ep-py		3 ep		PP				
sw354	22-Jul-14	578737	5585486	Andesite tuff, wacke, flows	chl-ep-ht-py		3		PP				yes
sw358	22-Jul-14	578982	5585558	Andesite flows	chl-py		3 ser-py		2 PP			87253	
sw359	22-Jul-14	579049	5585519	Andesite flows-agglomerate	chl-ep-py		4 ser-py		2 PP			87254	
sw360	22-Jul-14	579168	5585438	Andesite flows	chl-ep-mt-py		3		PP-M				
sw361	22-Jul-14	579272	5585262	Andesite flows-agglomerate	chl-ep-py		4		PP				
sw362	23-Jul-14	578017	5581917	Longarm basalt and seds	chl-ep		2 chl-ep		1 PP	qtz-cal-py		1	
sw363	23-Jul-14	580516	5582433	Andesite	chl-ep-mt		3		PP-M	ep		3	
sw364	23-Jul-14	580434	5581937	Andesite	chl-ep-mt-py		4		PP-M	ep		4	
sw365	23-Jul-14	580416	5583057	Andesite (tuff)	chl-ep-mt-py		4 ep-cal-cpy		PP-M	ep-cal-py-c		3	
sw366	23-Jul-14	580225	5583408	Siltstone and andesite	chl-py		3		PP	ksp		2	
sw367	24-Jul-14	578539	5585467	Rhyolite-dacite	ser-clay-py		4		PH	qtz-py		1	87255
sw368	24-Jul-14	579559	5584357	Andesite tuff	chl-ep-mt-py		3		PP-M	ep		2	87256
sw369	24-Jul-14	579436	5584405	Andesite tuff	chl-py		3 sil-clay-py-ht		2 PP				
sw370	24-Jul-14	579352	5584384	Andesite breccia	chl-py		3		PP	py		2	
sw371	24-Jul-14	579304	5584382	Andesite tuff	chl-py		3		PP				
sw372	24-Jul-14	579044	5584252	Andesite	chl-sil-py-ht		4		PP				
sw373	24-Jul-14	579370	5584178	Andesite tuff	chl-py		4 sil-ser-py		2 PP	ksp-qtz-py		2	87258
sw374	24-Jul-14	579365	5584136	Andesite tuff	chl-py		3 sil-ser-py		2 PP				
sw375	24-Jul-14	579749	5584593	Andesite tuff	chl-py		3		PP-M				
sw376	24-Jul-14	579779	5584542	Andesite tuff	chl-py-ep		3		PP-M	ep-cpy		2	87259
sw377	24-Jul-14	579571	5584717	Basaltic andesite	chl-ep		3		PP				
sw378	24-Jul-14	579461	5584751	Andesite	chl-ep		3		PP				
87257	24-JUL-14	579295.7	5584290	Andesite			0		0 PH			87257	
R10	22-JUL-14	577683.3	5585824	andesite	prop		2 silica		2 PP				
R11	22-JUL-14	577675.3	5585984	andesite?	silica		4		0 SIL				
R8	22-JUL-14	577694.1	5585442	rhyolite?	silica		4		0 SIL				
R9	22-JUL-14	577681.5	5585313	rhyolite	silica		4		0 SIL				
Rc1	22-JUL-14	579027.9	5585952	andesite	prop		3 pot		1 PP				
Rc12	23-JUL-14	577891.3	5581344	andesite	prop		2		0 PP			87202, 87203	
Rc13	23-JUL-14	578024.8	5581238	siltstone			0		0 PP				
Rc14	23-JUL-14	578124.8	5581204	andesite	prop		2		0 PP				
Rc15	23-JUL-14	580521	5582438	andesite	prop		4 silica		4 PP-M				
Rc16	23-JUL-14	580390.8	5583241	andesite	prop		4 silica		2 PP-M				
RC17	23-JUL-14	580315.3	5583278	andesite	prop		4 silica		2 PP-M				
Rc18	23-JUL-14	580270.1	5583321	andesite	prop		4 silica		2 PP-M				
Rc19	24-JUL-14	579556.9	5584358	andesite	prop		3		0 PP-M				
Rc2	22-JUL-14	576530.1	5586120	andesite	silic		4		0 SIL				
Rc20	24-JUL-14	579587.8	5584271	andesite	prop		4 silica		1 PP			87204	
Rc21	24-JUL-14	579573.9	5584150	andesite	phyllic		2		0 PH			87205	
Rc22	24-JUL-14	579596.1	5584062	andesite?	phyllic		2		0 PH			87206	
Rc23	24-JUL-14	579637.1	5583945	andesite?	phyllic		2		0 PH				
Rc3	22-JUL-14	576980.5	5586121		aa		4		0 AA				
Rc4	22-JUL-14	577476.5	5585669	andesite	prop		3		0 PP			87201	
Rc5	22-JUL-14	577486.3	5585585	andesite	prop		2		0 PP				
Rc6	22-JUL-14	577594.4	5585481	rhyolite	silica		5		0 SIL				
Rc7	22-JUL-14	577592.9	5585472	rhyolite	silica		5		0 SIL				

Comments

Chl-ep-py altered andesite cut by k-spar veins that strike e-w and dip moderately northward.
Along Gooding Bay road at advanced argillic (silica) altered zone
Interbedded wacke and tuff, chl-py altered with some weak possibly sericitic alteration
Chl-ep altered andesite with epidote along fractures, ~ 50 m long outcrop along road
Chl-ep altered andesite to basaltic andesite; ~ 100 m long cliff face along road
Chl-ht altered andesite with abundant ht along fractures and bedding planes and siliceous; locally unit is brecciated with siliceous and hematite veins clasts in the chloritic andesite matrix.
Nearly continuous oc along road of andesite (chl-py) and very little ep; several zones of sericite-py alteration
Highly chl-py altered andesite agglomerate with some ser-py fractures
Chl-ht-py-mt altered andesite
Strongly chl-py altered andesite
Entire road up is weakly chloritized and ep altered basalt and shales Longarm Fm; few cryptocrystalline quartz-cal-py veins and breccia zones
Epidote veined chl-mt-ht-py altered andesite with epidote extending 3-4 cm from fractures
Magnetic andesite chl-mt-py altered with heavy ep veining
Chl-ep-mt altered andesite tuff with ep-py-cal-cpy veins
Hornfelsed siltstone-andesite tuff with chl-py alteration and some pink along fractures (ksp?)
Angular blocks of rhyolite or highly sil-ser-clay-py altered volcanic ~10-15% py and few qtz veinlets
Shearer sample 51567; chl-mt altered andesite tuff with epidote clots and veins

Possibly dacitic and very siliceous or hornfelsed tuff with chl-py-ht alteration
Highly pyritic andesite chl-py altered with zones of sil-ser-py alteration and lensy qtz=py and py veins; some fractures with pinkish selvages possibly ksp.
A small orbicular dacite lapilli tuff unit occurs here.
Slightly magnetic andesite, that is sheared
2-3% pyrite with clots and veins of ep-py-cpy

brecciated, cut by (5-10%) mm-size py vlt & minor qtz vlt; sil-ch; occur at FW of structure
NW corner of Le mare lake, brown weathering andesite, with bands of jasper, fractured, local flow banding weak chlorite flow banding 045/53 SE
tan weathering FG cherty rock, laminated
banded white silica and minor jasper, hackly weathered surface- brecciated
flow banded rhyolite, local breccia, tan weathered, alternating red jasper and white silica, flow banding 125/60 SW
on road N of lake. Epidote clots and chlorite and pyrite on fractures overprint of Kspar veins and clots in patches, local pseudobx.
Parson Bay Fm? Green to grey green, feldspar phyrlic, andesite, chlorite on fractures, trace fg pyrite along microveinlets, sample 87202, 87203
Parson bay? Long 70 m exposure of brown rusty sediments. 10 to 20 cm beds of siltstone separated by 5 cm fissile partings . Fe Crab spots to 15% beds 200/39 W
massive grey tan andesite, feldspar phyrlic
S side of Le mare in new cut. Fg aphanitic andesite? Rare mafic phenos- possible hornblende, silicified, very magnetic, hematite stained possible hornfels? Late veins and masses of epidote to 20cm,
grey silicified andesite, dusting of FG pyrite, <1%, 1 metre wide fracture zone with increased pyrite to 5% euhedral, trace chalco?, isolated coxcomb quartz. fractures zone 300/90
rusty silicified andesite, minor epidote, diss euhedral pyrite to 3% with local enrichments to 15% in narrow fracture zones to 30 cm wide fractures zone 250/90
rusty silicified andesite, trace epidote in microveinlets and open vugs, very pyritic 1 to 10%
black FG granular andesite? (mapped as basalt) , 3-5% fg pyrite, trace epidote, chlorite, blocky weathering, rusty, weak to mod magnetic
Gorby showing. Quarry on coast . Red burgundy jasper veinlets, masses and breccia matrix with pyrite and trace chalcopyrite. Andesite host
samp 87204, dark green, 5% diss pyrite, hard silicified, chlorite
samp 87205, very rusty, light grey bleached, pyritic andesite, pyrite 1-2% possible sericite
samp 87206, rusty, grey siliceous, pyritic, (3-5%), this < 1mm quartz microveinlets, possible sericite, no chlorite
as rc22,
on road near Gorby. Yellow tan massive white to grey silica breccia trace fine grained pyrite, ptorolith indeterminate. Advanced argillic? Or just silicified. Sample 87155
5m rhyolite dyke cutting andesite. Quartz eyes, flow banded parallel to contact, green chloritic andesite, dyke contact 230
brown weathering andesite, major fault exposed in creek bed ~1.5 m og grey gouge, fault 208/46 SE
continuous to RC7, massive silica, grey to burgundy, rare jasper, brecciated with late silica in lighter colour, laminated (flow banding?), flow banding 150/40 SW
see RC6

Legend

Plot SampleType

- Soil
- Sediment
- Rock
- Not Defined

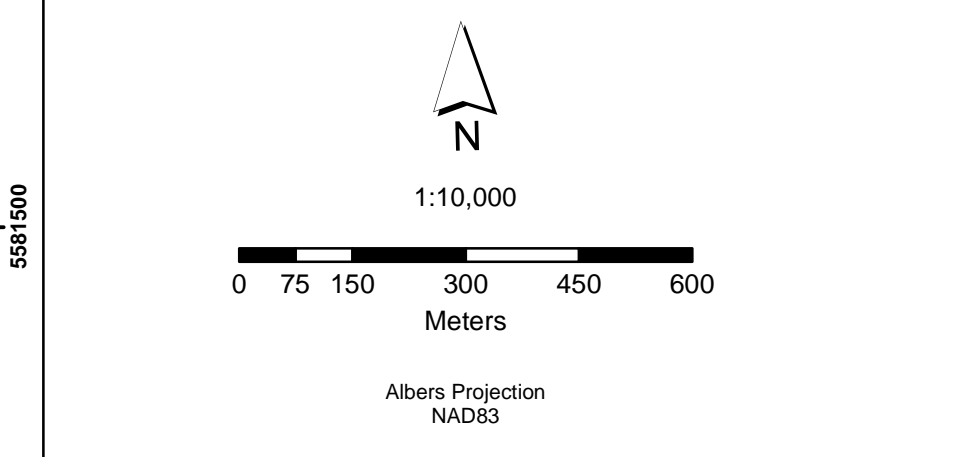
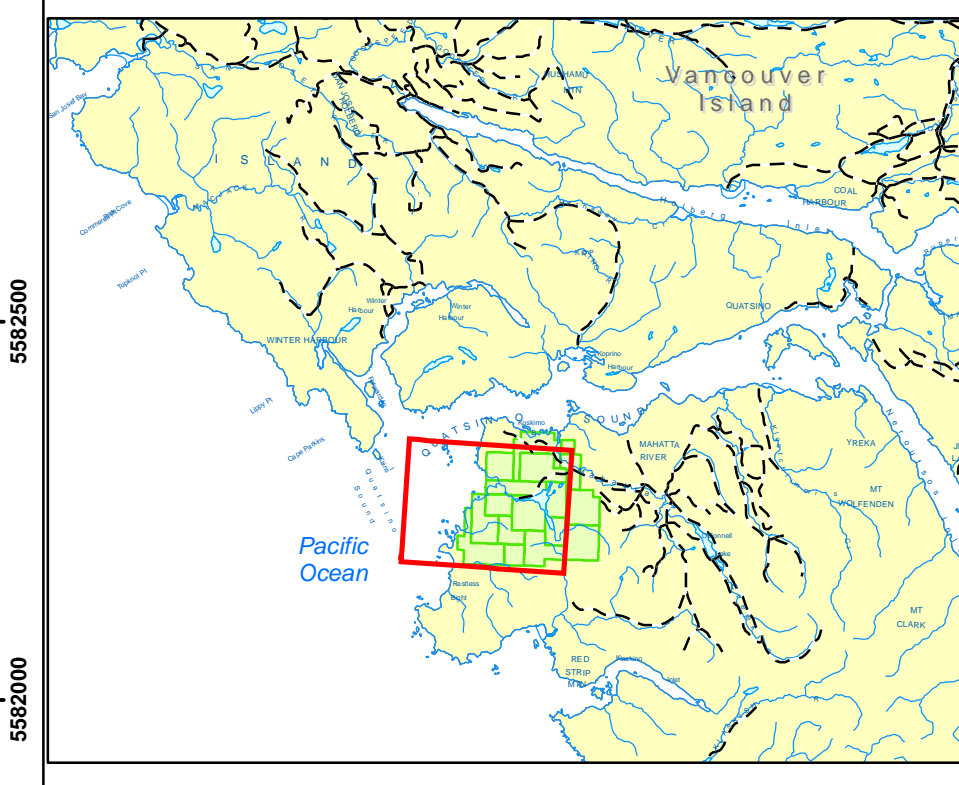
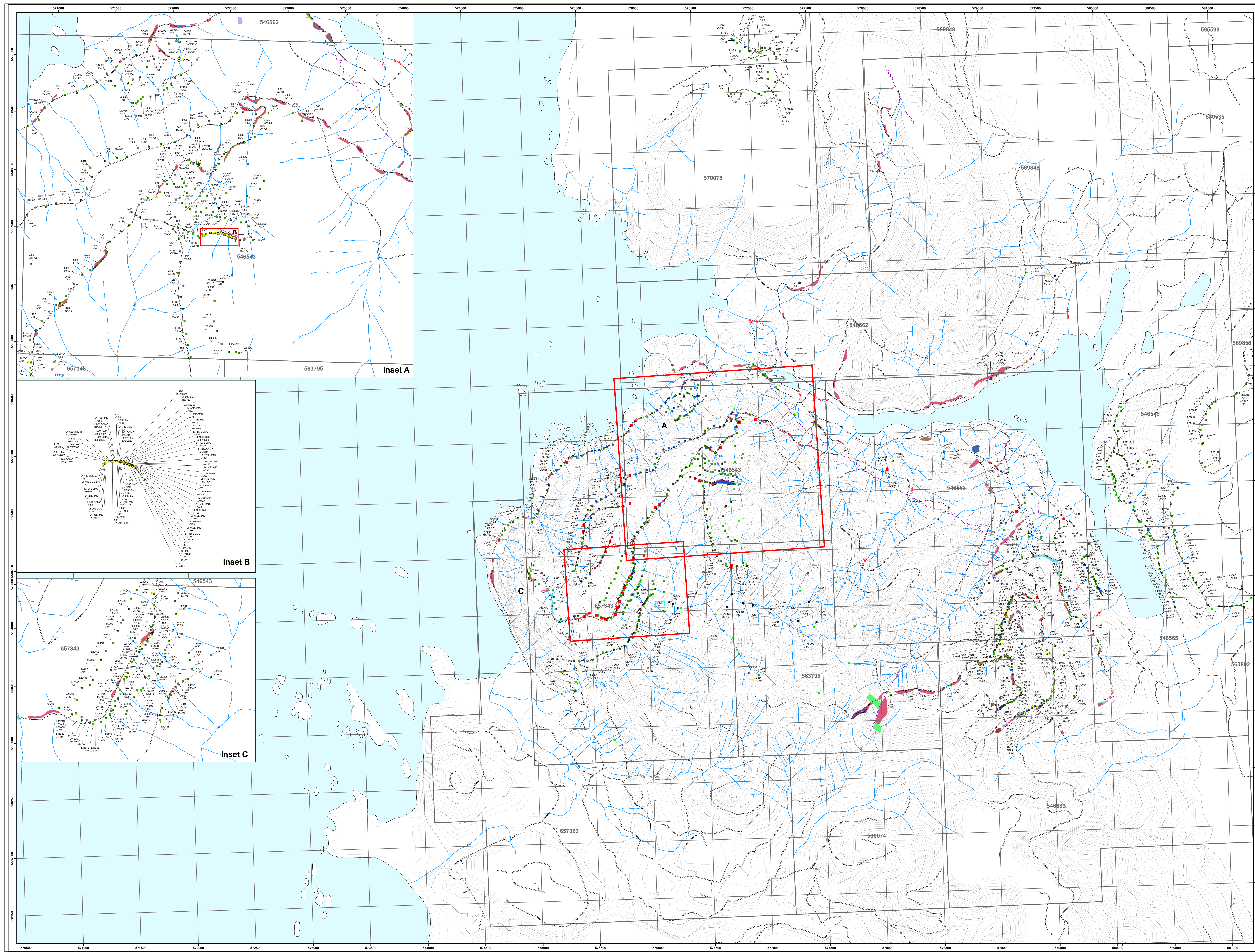
Sample Number → L111
 Gold >10 ppb → 35x/60
 Molybdenum >2 ppm
 Cu >10 ppm

The notation "*" indicates values not fulfilling the above requirement.

- Au > 50 ppb
- Waypoints
- Mineral Claims
- Ocean
- Lake
- Streams
- Fault
- Slide

Geology

- Unknown
- AA
- And
- Aplite
- Arkose
- Breccia
- Bst
- Bxa
- Dac
- Hm
- K-Spar
- Py
- Qtz
- Rhyl
- Silica
- Sulfo-Salt
- Till
- Tr



Legend

- Plot**
- Sample Number → L111
 - Gold >10 ppb → 25x/60
 - Silver >0.5 ppm → 25x/60
 - Cu >10 ppm → 25x/60
- The notation "x" indicates values not fulfilling the above requirement.

- SampleType**
- Soil
 - Sediment
 - Rock
 - Not Defined
 - Au > 50 ppb
 - Waypoints

- Mineral Claims
- Ocean
- Lake
- Streams
- Fault
- Slide

- Geology**
- Unknown
 - AA
 - And
 - Aplite
 - Arkose
 - Breccia
 - Bst
 - Bxa
 - Dac
 - Hm
 - K-Spar
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 - Rhyl
 - Silica
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 - Till
 - Tr

